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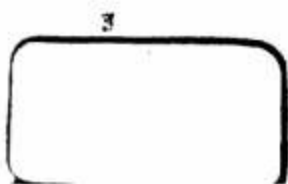
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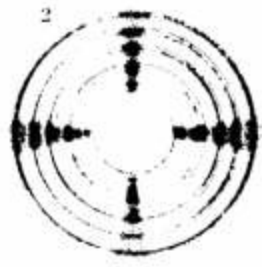
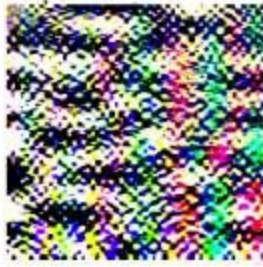
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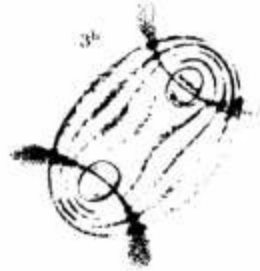
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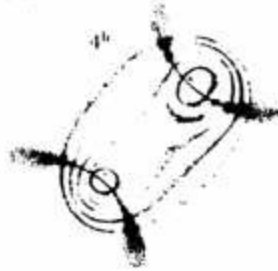
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1903.

PREFACE.

The remarkable advance in the Science of Mineralogy, during the years that have elapsed since this Text-Book was first issued in 1877, has made it necessary, in the preparation of a new edition, to rewrite the whole as well as to add much new matter and many new illustrations,

The work being designed chiefly to meet the wants of class or private instruction, this object has at once determined the choice of topics discussed, the order and fullness of treatment and the method of presentation.

In the chapter on Crystallography, the different types of crystal forms are described under the now accepted thirty-two groups classed according to their symmetry. The names given to these groups are based, so far as possible, upon the characteristic form of each, and are intended also to suggest the terms formerly applied in accordance with the principles of hemihedrism. The order adopted is that which alone seems suited to the demands of the elementary student, the special and mathematically simple groups of the isometric system being described first. Especial prominence is given to the "normal group" under the successive system, that is, to the group which is relatively of most common occurrence and which shows the highest degree of symmetry. The methods of Miller are followed as regards the indices of the different forms and the mathematical calculations.

In the chapters on Physical and Chemical Mineralogy, the plan of the former edition is retained of presenting somewhat fully the elementary principles of the science upon which the mineral characters depend; this is particularly true in the department of Optics. The effort has been made to give the student the means of becoming practically familiar with all the modern methods of investigation now commonly applied. Especial attention is, therefore, given to the optical properties of crystals as revealed by the microscope. Further, frequent references are introduced to important papers on the different subjects discussed, in order to direct the student's attention to the original literature.

The Descriptive part of the volume is essentially an abridgment of the Sixth Edition of Dana's System of Mineralogy, prepared by the author (1893). To this work (and future Appendices) the student is, therefore, referred for fuller descriptions of the crystallographic and optical properties of species, for

directions, however, the work has been expanded when the interests of the student have seemed to demand it; for example, in the statement of the characters of the various isomorphous groups. Attention is also called to the paragraph headed "ms." in the description of each common species, in which are given the distinguishing characters, particularly those which serve to separate it from other species with which it might be easily confounded.

The list of American localities of minerals, which appeared as an Appendix in the earlier edition, has been omitted, since in its present expanded form it requires more space than could well be given to it; further, its reproduction here is unnecessary since it is accessible to all interested not only in the System of Mineralogy but also in separate form. A full topical Index has been added, besides the usual Index of Species.

The obligations of the present volume to well-known works of other authors—particularly to those of Groth and Rosenbusch—are too obvious to require special mention. The author must, however, express his gratitude to his colleague, Prof. L. Y. Firsson, who has given him material aid in the part of the work dealing with the optical properties of minerals as examined under the microscope. He is also indebted to Prof. S. L. Penfield of New Haven and to Prof. H. A. Miers of Oxford, England, for various valuable suggestions.

New Haven, Conn., Aug. 1, 1898.

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INTRODUCTION.

1. The science of Mineralogy treats of those inorganic species called mineral*, which together in rock masses or in isolated form make up the material of the crust of the earth, and of other bodies in the universe so far as it is possible to study them in the form of meteorites.

S. Definition of a Mineral. — A Mineral is a body produced by the processes of inorganic nature, having a definite chemical composition and, if formed under favorable conditions, a certain characteristic molecular structure which is exhibited in its crystalline form and other physical properties.

This definition calls for some further explanation.

First of all, a mineral must be a homogeneous substance, even when minutely examined by the microscope; further, it must have a definite chemical composition, capable of being expressed by a chemical formula. Thus, much basalt appears to be homogeneous to the eye, but when examined under the microscope in thin sections it is seen to be made up of different substances, each having characters of its own.

• ■ ^ ■ AgaiTl, "ObFi*y««lJ," or'voleanic, glasses, though it may be essentially homogeneous, has a definite composition \ <f. corresponding to a specific chemical formula, and is hence classed as a rock, ' ^ ■ not as a mineral species. Further, several substances, such as hyalite, hyalome-j -(J lane, etc., which at one time passed as minerals, have since been relegated to petrology, because it has been shown that they are only local fossils of basalt, J retuiiiiiig.aH~&ppai«iilJ;Ju><noi;«ndeOA fot:in-dti« to-rapid cooling. ^ ^ . -^

Again, a mineral has in all cases a definite molecular structure, unless the conditions of formation have been such as to prevent this, which is rarely true. This molecular structure, as will be shown later, manifests itself in the physical characters and especially in the external crystalline form.

It is customary, as a matter of convenience, to limit the name mineral to those compounds which have been formed by the processes of nature alone, while compounds made in the laboratory or the smelting-furnace are at most called artificial minerals. Further, mineral substances which have been produced through the agency of organic life are not included among minerals, as the pearl of an oyster, the opal-silica (tabasheer) secreted by the bamboo, etc. Finally, mineral species are, as a rule, limited to solid substances; the only liquids included being metallic mercury and water. Petroleum, or mineral oil, is not properly a homogeneous substance, consisting rather of several hydrocarbon compounds; it is hence not a mineral species.

It is obvious from the above that minerals, in the somewhat restricted sense usually adopted, constitute only a part of what is often called the Mineral Kingdom.

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INTRODUCTION

3. Scope of Mineralogy.—In the following pages, the general subject of Mineralogy is treated under the following heads:

(1) Crystallography, —This comprises a discussion of crystals in general and especially of the crystalline forms of mineral species,

(2) Physical Mineralogy. —This includes a discussion of the physical characters of minerals, that is, those depending upon cohesion and elasticity, density, light, heat, electricity, and so on.

(3) Chemical Mineralogy. —Under this head are presented briefly the general principles of chemistry as applied to mineral species; their characters as chemical compounds are described, also the methods of investigating them from the chemical side by the blowpipe and other means.

(4) Descriptive Mineralogy, —This includes the classification of minerals and the description of each species with its varieties, especially in its relations to closely allied species, as regards crystalline form, physical and chemical characters, occurrence in nature, and other points.

4. Literature.—Reference is made to the Introduction to the Sixth Edition of Dana's System of Mineralogy, pp. xlv-lxi, for an extended list of independent works on Mineralogy up to 1892; the names are also given of the many scientific periodicals which contain original memoirs on mineralogical subjects. For the convenience of the student the titles of a few works, mostly of a general character, are given here. Further references to the literature of Mineralogy are introduced through the first half of this work, particularly at the end of the sections dealing with special subjects.

Crytallography and Phytikai Minertxlogy.

Early works* include those of RoinS de Mala, 1773; HnOy, 1832; Neumann, Krv-
■allonoinie, IUSS, uud Erystallographle, 183fi: Eiiipffer, 1823; Grassmann,
Kiystallouomfo 1639: N'aniiiHmi. 1839 and later; Quetietedt. 1816 (also 1878); Miller,
18S9 and 1808' Omilich, IHri6; Eopp, 1803; vod Lnns, 1B66; Bravtiia, EtudeB CrUt., Paris,
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Iii Phybical Minrraloot tbe moat important general works are those ofSchraaf (1868),
Qrotli {18T6-18B3I. Mallard (1884), Llebisch il89ii. mentioned in tbe above list; alio
jRosenbiiacb. Mikr. Physiographfe. etc. (1893). In addition to Ibese (to wbMi might be
added tbe namei of some general works od PliyBicsi memoirs of especial importance on
the different subjects are enumerated iu many cases at tbe end of the respective sections
ot this work.

e given Id pp. ll-)xi of Daits's System of Uiner-

General Mineralogy/.

Of the man; works, a knowledge of which is deemed by one who wishes a full acquaintance with the historical development of Mineralogy, the following are particularly important. Very early works include those of Theophrastus, Pliny, Liuweus, Wallerius, Cronstedt, Werner, Bergmann, Klaproth.

Within the nineteenth century: Haüy's Treatise, 1801, 1882; Jameson, 1816, 1830; Werner's Letztete Mineral-System, 1817; Clendinning's Mineralogy, 1818, 1833; Leohard's Handbuch, 1821, 1830; Holis's Min., 1832; Haidinger's Introduction of Mineralogy, 1834; Drdt-haidinger's Charakteristik, 1820, 1833, 1833; Beudantic's Treatise, 1834, 1852; Phillips's Hlu., 1837; Shepard's Hlu., 1833-85 and later editions; Eobell's Oruddege. I, 1888; Mohs's Hlu., 1819; Brellhaupt's Min., 1818-1847; Haidinger's Handbuch, 1845; Niemann's Hlu., 1846 and later; Haumimius Handbuch, 1847; Daubroy's Hlu., 1844-1847 (also 1856-1859); Brooke & Miller, 1852; J. D. Dana's System of 1887, 1844, 1800, 1854, 1851.

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Abbbetiatioks.

Lm pi. Plane of the optic asea. B. Hardness.

Ix. Bxt. Acute biaectrix (p. 208). Ob*. Observatloos oo occurrence, etc.

Jx,. Oljtuse bisectrii (p. 808). O.F. Oxidizing Flame (p. 2B7).

J.B, Before llie Blowpipe (p. fl66).! Pyr. Pyrogoosiica or blowpipe And

Jomp. CompoalLion. allied characters.

as. Differences, or distinctive char O, Ratio, Oxygen Ratio (p. 249),

ftciets. R.P. Reducing FUmE (p. 357),

V. SpeciUc Gravity. Var. Varieties.

The sign A is used to Indicate the angle betneeo two faces of a crystal, as

»(100 A 110) = 44'30'.

*See the bibliography given by RoseDbuecb.

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PART L CRYSTALLOGEAPHT,

GENERAL MORPHOLOGICAL RELATIONS OF CRYSTALS.

ft. Crystallograpliy.—The subject of Grystallograph; inclndee the description of the characters of crystals in general; of the various forma of crjatala and their division into groups and Byetoms; of the methods of studying crystals, incinding the determination of

the mathematical relations of their faces, and the measurement of the angles between them; finally, a description of compound or twin crystals, of irregularities in crystals, of crystalline aggregates, and of polymorphous crystals.

Allied to Crystallography is the subject of Crystallogeny, which describes the methods of making crystals which may be applied in the laboratory, and discusses the theories of their origin in nature. This department is only briefly touched upon in the present work.

6. Definition of a Crystal—A crystal* is the regular polyhedral form, bounded by smooth surfaces, which is assumed by a chemical compound, under the action of its intermolecular forces, when passing, under suitable conditions, from the state of a liquid or gas to that of a solid.

As expressed in the foregoing definition, a crystal is characterized, first, by its definite internal molecular structure, and, second, by its external form. A crystal is the normal form of a mineral species, as of all solid chemical compounds; but the conditions suitable for the formation of a crystal of ideal perfection in symmetry of form and smoothness of surface are never realized. Farther, many species usually occur not in distinct crystals, but in massive form, and in some exceptional cases the definite molecular structure is absent.

7. Molecular Structure in General—By definite molecular structure is meant the special arrangement which the physical units, called molecules, assume under the action of the forces exerted between them during the formation of the solid. Some remarks are given in a later article (p. 18 et seq.) in regard to

* In its original significance the term crystal is which the ancient philosophers believed to be white)-

which the ancient philosopher termed, from *Kratos* and *Xoia*, ice.

The relation between atomic and molecular structure is explained under the chapter on Chemical Mineralogy. The molecules here spoken of are the physical molecules.

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the kinds of molecular arrangement theoretically possible, and their relation to

the symmetry of the different groups and systems of crystals.

The definite molecular structure is the essential character of a crystal, and the external form is only one of the ways, although the most important, in which this structure is manifested. Thus it is found that all similar directions in a crystal, or a fragment of a crystal, have like physical characters,* as of elasticity, cohesion, action on light, etc. This

is clearly shown by the cleavage, or natural tendency to fracture in certain directions, yielding more or less smooth surfaces; as the cubic cleavage of galena, or the rhombohedral cleavage of calcite. It is evident, therefore, that a small crystal differs from a large one only in size, and that a fragment of a crystal is itself essentially a crystal in all its physical relations, though showing no crystalline faces.

Further, the external form without the corresponding molecular structure does not make a solid a crystal. A model of glass or wood, on the one hand, is not a crystal, though having its external form, because there is no relation between form and structure. Also, an octahedron of malachite, having the form of the crystal of cuprite from which it has been derived by chemical alteration, is not a crystal of malachite.

On the other hand, if the natural external faces are wanting, the solid is not called a crystal. A cleavage octahedron of fluorite and a cleavage rhombohedron of calcite are not properly crystals, because the surfaces have been yielded by fracture and not by the natural molecular growth of the crystal.

8. Crystalline and Amorphous.—When a mineral shows no external crystalline form, it is said to be massive. It may, however, have a definite molecular structure, and then it is said to be crystalline. If this structure, as shown by the cleavage, or by optical means, is the same in all parallel directions through the mass, it is described as a single individual. If it varies from grain to grain, or fiber to fiber, it is said to be a crystalline aggregate, since it is in fact made up of a multitude of individuals.

Thus in a granular mass of galena or calcite, it may be possible to separate the fragments from one another, each with its characteristic cubic, or rhombohedral, cleavage. Even if the individuals are so small that they cannot be separated, yet the cleavage, and hence the crystalline structure, may be evident from the spangling of a freshly broken surface, as with fine-grained statuary marble. Or, again, this aggregate structure may be so fine that the crystalline structure can only be resolved by optical methods with the aid of the microscope. In all these cases, the structure is said to be crystalline.

If optical means show a more or less distinct crystalline structure, which, however, cannot be resolved into individuals, the mass is said to be crypto-crystalline; this is true of some massive varieties of quartz.

If the definite molecular structure is entirely wanting, and all directions in the mass are sensibly the same, the substance is said to be amorphous. This is true of a piece of glass, and nearly so of opal. The amorphous state is rare among minerals.

* This subject is further elucidated in the chapter devoted to Physical Mineralogy, where it is shown that, with respect to many, but not all, of the physical characters, the average of this proposition is true, viz., that unlike directions in a crystal have in general unlike properties.

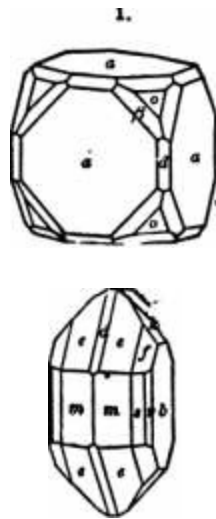
The crystallization of the various forms of crystalline aggregates is postponed to the end of the present chapter.

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There is a tendency to go over into the amorphous condition by molecular rearrangement. A transparent amorphous mass of «reueic Irloxiide (Al_2O_3). (formed by fusion, sometimes opaque and crystalline after a time. Similarly the joints of a railroad bridge may eventually become crystalline units thus lose some of their original strength because of the molecular rearrangement made possible by the vibrations caused by the frequent jar of the train. The microscopic study of rocks reveals many cases in which an analogous change in molecular structure has taken place in a solid mass, as caused, for example, by great pressure.

9. **Crystalline Form.**—A crystal is bounded by smooth plane surfaces, called faces or planes,* showing in their arrangement a certain characteristic symmetry, and related to each other by definite mathematical laws.

Thus, without inquiring, at the moment, into the exact meaning of the term symmetry as applied to crystals, and the kinds of symmetry possible, which will be explained in detail later, it is apparent that the accompanying figures, 1-3, show the external form spoken of. They represent, therefore, certain definite types. ^



Zircon.

Chrysolite.

10. **Variation of Form and Surface.**—Actual crystals deviate, within certain limits, from the ideal forms.

First, there may be variation in the size of like faces, thus producing what are defined later as distorted forms. In the second place, the faces are rarely absolutely smooth and brilliant; commonly they lack perfect polish, and they may even be rough or more or less

covered with fine parallel lines (called striations), or show minute elevations, depressions or other peculiarities. Both the above subjects are discussed in detail in another place.

peculiarities. Both the above subjects are discussed in detail in another place.

It may be noted in passing that the characters of natural faces,

just alluded to, in general make it easy to distinguish between them and a face artificially ground, on the one hand, like the facet of a cut gem; or, on the other hand, the splintery, uneven surface yielded by cleavage,

11. Constancy of Angle in the Same Species.—The crystals of any species are essentially constant in the angle of inclination between like faces, wherever they are found, and whether products of nature or of the laboratory. These angles, therefore, form one of the distinguishing characters of a species. Apatite.

Thus, in Fig. 4, of apatite, the angle between the adjacent faces x and

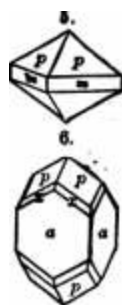
y is the same in all crystals of the species, whether the direction, rather than the definite

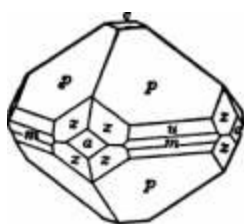
angle.

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($130^\circ 18'$) is the same for any two like faces, similarly situated with reference to each other. Further, this angle is constant for the species, differing but little on crystals from different localities. Moreover, the angles between all the faces on crystals of the same species (cf. Figs. 5-9 of zircon below) are more or less closely connected together by certain definite mathematical laws.

12. Diversity of Form, or Habit. —While in the crystals of a given species there is constancy of angle between like faces, the forms of the crystals may be exceedingly diverse. The accompanying figures (5-9) are examples of a few of the forms of the species zircon. There is hardly any limit to the number of faces which may occur, and as their relative size changes, the habit, as it is called, may vary indefinitely. Yet for the crystals of each species, the angles between like faces are essentially constant.





Crystals of Kircon.

13. Diversity of Si». —Crystals occur of all sizes, from the merest microscopic point to a yard or more in diameter. It is important to understand, however, that in a minute crystal the development is as complete as with a large one. Indeed the highest perfection of form and transparency is found only in crystals of small size.

A single crystal of quartz, now in Milan, is three and a quarter feet long and five and a half inches in diameter, and its weight is estimated at eight hundred and seventy pounds. A single cavity in a vein of quartz near the Tiefen Glacier, in Switzerland, discovered in 1857, contained smoky quartz crystals weighing in the aggregate about 20,000 pounds; a considerable number of the single crystals measured a weight of 300 to 350 pounds, or even more. A gigantic crystal from Acworth, New Hampshire, measured four feet in length and two and a half in circumference; another, from Orafto, was over four feet long, and three and a half inches in one of its diameters, and weighed about two and a half tons.

14. Symmetry in General —The faces of a crystal are arranged according to certain laws of symmetry, and this symmetry is the natural basis of the division of crystals into groups and systems. The symmetry may be defined relatively to (1) a plane of symmetry, (2) an axis of symmetry, and (3) a center of symmetry.

These different kinds of symmetry may, or may not, be combined in the same crystal. It will be shown later that there is one group of crystals of which have neither center, axis, nor plane of symmetry; another where there is only a center of symmetry. On the other hand, some groups have all these elements of symmetry represented.

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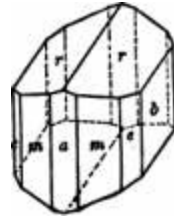
16. Planes of Symmetry.—A solid is said to be geometrically symmetrical with reference to a plane of symmetry when for each face, edge, or solid angle there is another similar face, edge, or angle which has a like position with reference to this plane. Thus it is obvious that the crystal of amphibole, shown in Fig. 10, is symmetrical with reference to a central plane of symmetry, parallel to the face b, passing vertically through the edge formed by the faces r, r and through the middle of the face a.

In the ideal crystal this symmetry is right symmetry in the geometrical sense, where every point on the one side of the plane of symmetry has a corresponding point at equal

In the ideal crystal this symmetry is right symmetry in the geometrical sense, where every point on the one side of the plane of symmetry has a corresponding point at equal

distances on the other side, measured on a line normal to it. In other words, in the ideal geometrical symmetry, one half of the crystal is the exact mirror-image of the other half.

A crystal may have as many as nine planes of symmetry, three of one set and six of another, as is illustrated by the cube (Fig. H). Here the planes of the first set pass through the crystal parallel to the cubic faces; they are shown in Fig. Z. The planes of the second set join the opposite cubic edges. On the other hand, some crystals have no plane of symmetry.



AnipuilKjIe.

16. Axes of Symmetry.—If a solid can be revolved through a certain number of degrees about some line as an axis, with the result that it again occupies precisely the same position in space as at first, it is said to have an axis of symmetry. There are four different kinds of axes of symmetry among crystals; they are defined according to the angular revolution needed in each case, that is, by the number of times which the crystal repeats itself in a complete revolution of 360° .

(n), A crystal is said to have an axis of binary, or twofold, symmetry when a revolution of 180° produces the result named above; in other words, when it repeats itself twice in a complete revolution. This is true of the crystal shown in Fig. 13 with respect to the vertical axis (and indeed each of the horizontal axes also).

(6) A crystal has an axis of trigonal, or threefold, symmetry when a revolution of 120° is needed; that is, when it repeats itself three times in a complete revolution. The vertical axis of the crystal shown in Fig. 14 is an axis of trigonal symmetry.

- The relation between the Ideal geometrical symmetry and the actual crystallographic symmetry is discussed in Art. IB.

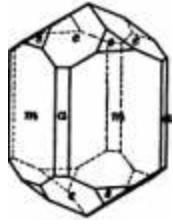
This is the cube of the normal group of the Isometric system.

;CoogIC

CRYSTALLOGRAPHY.

(c) A crystal has an axis of tetragonal, or fourfold, symmetry when a revolution of 90° is called for; in other words, when it repeats itself four times

15.



Tellurite. Kettle.

The vertical axis in t

Mimetite. 3 crystal shown in Fig. 15 ia

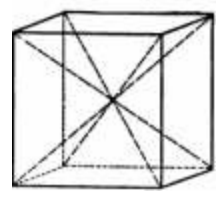
Folycrsm.

in a complete revolution. such an axis.

(d) Finally, a crystal has an axis of hexagonal, or sixfold, symmetry when a revolution of 60° is called for; in other words, when it repeats itself six times in a complete revolution, this is illustrated by Fig. 16.

The cube illustrates three of the four possible kinds of symmetry with respect to lines of symmetry. It has six axes of binary symmetry joining the middle points of opposite edges (Fig. 17). It has four axes of trigonal symmetry, joining the opposite solid angles (Fig. 18). It has, finally, three axes of tetragonal symmetry joining the middle points of opposite faces (Fig. 19).

18.

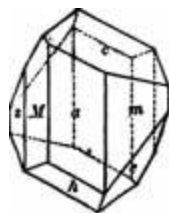


17. Center of Symmetry.—Most crystals, besides planes and axes of symmetry, have also a center of symmetry. On the other hand, a crystal, though possessing neither plane nor axis of symmetry, may yet be

symmetrical with reference to a point, its center. This last is true of the triclinic crystal shown in Fig. 20, in which it follows that every face, edge, and solid angle has a face, edge, and angle similar to it in the opposite half of the crystal.

There is another method of viewing it; the symmetry in this case, which is adopted by some authors. If the crystal be thought of as divided into two similar halves by a plane parallel to any one of its faces, and one half be revolved 180° about an axis normal to

this face, this half would be brought into position in which it would be the mirror-image of the remaining half. This symmetry is hence described as C_2 with reference to an axis of binary symmetry and a plane normal to it, both taken together. This method is not followed here since, though having theoretical advantages. It is likely to confuse the student meeting the problem of crystallography for the first time.



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* This is actually the cube of the normal group of the isometric system.

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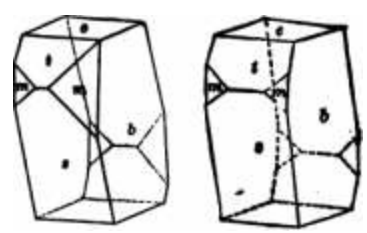
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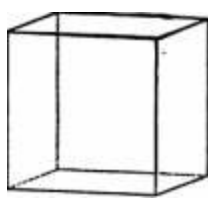
18. Definition of Geometrical to Crystallographic Symmetry.— Since the symmetry; in the arrangement of the faces of a crystal is an expression of the internal molecular structure, which in SI. 22.

general is alike in all parallel directions, the relative size of the faces and their distance from the plane or axis of symmetry are of no moment, their angular position alone is essential. Hence Fig. 21 has actually a vertical plane of symmetry (parallel to b) as Fig. 22 if the faces have exactly the same angular position as in that, although the strict geometrical definition * could not be applied to it.

Also in a normal cube (Fig. 33) the Heulandite.

three central planes parallel to each pair of cubic faces are like planes of symmetry, as stated in Art. 16, But a crystal is still crystallographically a cube, though differing widely from the requirements of the strict geometrical definition, as shown in Figs. 34, 35, if only it can be proved, «.:. by cleavage, the



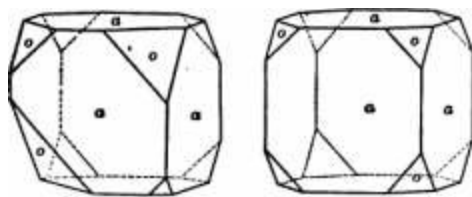


physical nature of the faces, or by optical means, that the three pairs of faces are like faces, independently of their size, or, in other words, that the molecular structure is the same in the three directions normal to them.

Further, in the case of a normal cube, a face of an octahedron on any solid angle requires, as explained beyond, similar faces on the other angles. It is a necessary, however, that

these eight faces should be of equal size, for in the crystallographic sense Fig. 26 is just as symmetrical with reference to the planes named as Fig. 27. 19. On the other hand, the molecular and hence the crystallographic symmetry is not always that which the geometrical form would suggest. Thus, deferring for the moment the consideration of pseudo-symmetry, an illustration of the fact stated is afforded by the cube.

* It is to be noted that the crystals of sulfur always show the geometrically ideal form. In which like faces, edges, and angles have the same shape, size, and position. In other words, the ideal crystal is uniformly represented as having the symmetry called for by the strict geometrical definition.



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It has already been implied and will be fully explained later that while the cube of the normal group of the isometric system has the symmetry described in Art. 15, 16, a cube of the same geometrical form but belonging molecularly, for example, to the tetrahedral group, has no planes of symmetry parallel to the faces, only the six diagonal planes; further, though the four axes shown in Fig. 18 are still axes of trigonal symmetry, the cubic axes (Fig. 19) are axes of binary symmetry only, and there are no axes of symmetry corresponding to those represented in Fig. 17. Other more complex cases will be described later.

Further, a crystal having interfacial angles of 90° is not necessarily a cube; in other words, the angular relations of the faces do not show in this case whether the figure is bounded by six like faces; or whether only two are alike and the third unlike; or, finally, whether there are three pairs of unlike faces. The question must be decided, in such cases, by the molecular structure as indicated by the physical nature of the surfaces, by the cleavage,

or by other physical characters, as [pyro-electricity, those connected with light phenomena, etc.

Still again, the student will learn later that the decision reached in regard to the symmetry to which a crystal belongs, based upon the distribution of the faces, is only preliminary and approximate, and before being finally accepted it must be confirmed, first, by accurate measurements, and, second, by a minute study of the physical characters as just insisted upon.

The method based upon the physical characters, which is the most convenient and admits of the widest application, is the skillful eliciting of the symmetry of the crystal by some appropriate solvent. By this means there are, in general, produced upon its minute depressions the shape of which always conforms to the symmetry in the arrangement of the molecules. This process, which is practically one involving the dissection of the crystal structure, is more particularly discussed in the chapter on Physical Mineralogy.

20. Pseudo-Symmetry. The crystals of certain species approximate closely in angle, and therefore in apparent symmetry, to the requirements of a system higher in symmetry than that to which they actually belong: they are then said to exhibit pseudo-symmetry. Numerous examples are given under the different systems. Thus the micas have been shown to be truly monoclinic in crystallization, though in angle they seem to be in some cases rhombohedral, in others orthorhombic.

It will be shown later that compound, or twin, crystals may also simulate by their regular grouping a higher grade of symmetry than that which belongs to the single crystal. Such crystals also exhibit pseudo-symmetry and are specifically called mimetic. Thus aragonite is an example of an orthorhombic species, whose crystals often imitate by twinning those of the hexagonal system.* Again, a highly complex twinned crystal of the monoclinic species, phillipsite, may have nearly the form of a rhombic dodecahedron of the isometric system. This kind of pseudo-symmetry also occurs among the groups of a single system, since a crystal belonging to a group of low symmetry may by twinning gain the geometrical symmetry of the corresponding form of the normal group. This is illustrated by a twinned crystal of scheelite like that figured (Fig. 378) in the chapter on twin crystals.

Pseudo-symmetry of still another kind, where there is an imitation of the symmetry of another system of lower grade, is particularly common in crystals of the isometric system (e.g., gold, copper). The result is reached in

* The terms pseudo-hexagonal, etc., used in this and similar cases explain themselves.

Bncn cases b; an abnormal development or "distortion" in the direction of certain axes of Bynimetry. This subject is discussed and illustrated on • later page.

21. Fowible Oroaps of Symmetry.—The theoretical consideration of the different kinds of symmetry possible among crystals built up of like molecales, as explained in Arts. 30-32, has led to the conclusion that there are thirty-two (32) types in all, difleriiiig with respect to the combination of the different symmetry elements just described. Of these thirty-two natural groups among crystals baaed upon their symmetry, seven groups include by far tue larger number of crystallized minerals. Besides these, some thirteen or fourteen others are distinctly represented, though several of these are of rare occurrence. Further, eight or nine others, making in all twenty-nine or thirty, are known among crystallized salts made in the laboratoiy. The characters of each of the thirty-two groups are given under the discussion of the several crystalline systems.

' 22. Cryatallographic Axes.—In the description of the form of a crystal, especially as regards the position of its faces, it is found convenient to a^time, after the methods of analytical geometry, definite lengths of certain lines passing throuh the center of the ideal crystal, as a basis of reference, (See further Art. 38 et teq.)

These lines are called the cryatallographic axes. Their direction is to a greater or less extent fixed by the symmetry of the crystals, for an axis of symmetry is in almost all cases* a possible crystallographic axis. Further, their unit lengths are fixed sometimes by the symmetry, sometimes by the faces assumed as fundamental, i.e., the itnit forms in the sense defined later. The dotted lines'shown in Fig. 19 are the crystallographic axes to which the cnbic faces are referred.

23. Syatema of Crystallization.^The thirty-two possible crystalline groups, distinguished from one another by their symmetry, are classified in this work under six systems, each characterized by the relative lengths and inclinations of the assumed crystallographic axes. These are as follows:

I. Isometric System. Three equal axes at right angles to each other.

II. Tetragonal System. Three axes at right angles to each other, two of them —the lateral axes—equal, the third—the vertical axis—longer or shorter.

III. Hexagon'AL Sxsteu. Four axes, three equal lateral axes in one plane intersecting at angles of 60° , and a vertical axis at right angles to this plane and longer or shorter.

IV. Orthorhombic System. Three axes at right angles to each other, but all of different lengths.

V. MoNOCLiNic System. Three axes unequal in length, and having one of their intersections oblique, the two other intersections equal to 90° .

essentially connected together by their common optical characters, and in general separated * from those of the other systems in the same way.

In the paragraphs immediately following, a synopsis is given of the symmetry of the normal group of each of the different Systems, and also that of one subordinate group of the hexagonal system, which is of so great importance that it is also often conveniently treated as a sub-system even when, as in this work, the forms are referred to the same axes as those of the strictly hexagonal type—a usage not adopted by all authors.

25. Symmetry of the System.—With respect to the symmetry of the form, which finds practical expression, as before stated, in the axial relations, the normal groups under the different systems are characterized as follows.

I. Isometric System. Three like axial planes of symmetry (principal planes) parallel to the cubic faces, and fixing by their intersection the crystallographic axes; six like diagonal planes of symmetry, passing through each opposite pair of cubic edges, and hence parallel to the faces of the rhombic dodecahedron.

Further, three like axes of tetragonal symmetry, the cubic or crystallographic axes; four like axes of trigonal symmetry, the octahedral axes; and six like axes of binary symmetry, the dodecahedral axes. There is also obviously a center of symmetry. These relations are illustrated by Fig. 28; also by Fig. 41; further by Figs. 70 to 110.

*Crystals of the tetragonal and hexagonal systems are alike in being optically uniaxial; but the crystals of all the other systems have distinguished optical characters.

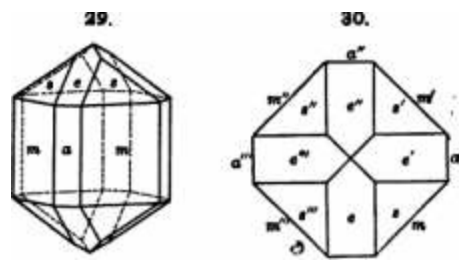
Two systems of symmetry are said to be like when they divide the identical crystal into halves which are identical to each other; otherwise, they are said to be unlike. Axes of symmetry are said to be like or unlike. If a plane of symmetry includes two of the crystallographic axes, it is called an axial plane of symmetry. If the plane includes two or more like axes of symmetry. It is called a principal plane of symmetry; also an axis of symmetry in which two or more like planes of symmetry meet is a principal axis of symmetry.

{In describing the symmetry of the different groups, here and later, the center of symmetry is ordinarily not mentioned when its presence or absence is obvious.

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II. Tetragonal System. Three axial planes of symmetry: of these two are like planes intersecting at 90° in a line which is the vertical crystallographic axis, and the third plane (a principal plane) is normal to them and

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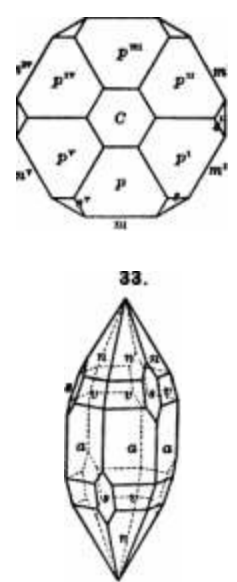
hence contains the lateral axes. There are also two diagonal planes of symmetry, intersecting in the vertical axis and meeting the two axial planes at angles of 45° .

Further, there is one axis of tetragonal symmetry, a principal axis; this is the vertical crystallographic axis. There are also in a plane normal to this four axes of binary symmetry—like two and two—those of each pair at right angles to each other. Fig. 29 shows a typical tetragonal crystal, and Fig. 30 a basal projection of it, that is, a projection on the principal plane of symmetry normal to the vertical axis. See also Fig. 42 and Figs. 149-151.

iii. Hexagonal System. In the Hexagonal Division there are four axial planes of symmetry; of these three are like planes meeting at angles of 60° , their intersection-line being the vertical crystallographic axis; the fourth plane (a principal plane) is at right angles to these. There are also three other diagonal planes of symmetry meeting the three of the first set in the vertical axis, and making with them angles of 30° .

^^^

Beryl



Beryl.

Corundum.

Further, there is one principal axis of hexagonal symmetry; this is the vertical crystallographic axis; at right angles to it there are also six binary axes. The last are in two

sets of three each. Fig. 31 shows a typical hexagonal crystal, and Fig. 32 a basal projection of the same. See also Fig. 43 and Figs. 195-209.

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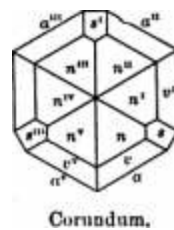
CBYSTA.LLOGBA.PHT.

STSteid there are three

lo 'i the vertical axis.

In the Trigonal or Rhombohedral Division of this s like planes of symmetry intersecting at angles of (Further, the forme belonging here have a vertical principal axis of trigonal symmetry, and three horizontal axes of binary symmetry, diagonal in position to the crystallographic axes. Fig. 33 shows a typical rhombohedral crystal, and Fig. 34 a basal projection. See also Figs. 326-252.

34. 36. 36.



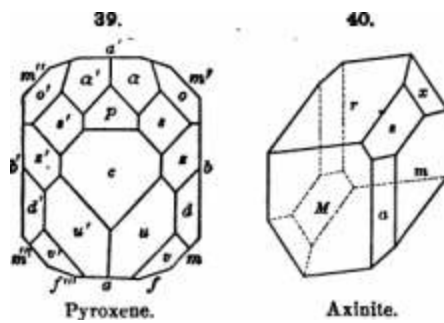
Chrysolite. I System. Three unlike plat

Cli'ysolfte.

IV. Orthorhombic System. Three unlike planes of symmetry meeting at 90° , and fixing by their intersection-lines the position of the crystallographic axes. Further, three unlike axes of binary symmetry coinciding with the last-named axes. Fig. 35 shows a typical orthorhombic crystal, and Fig. 36 a basal projection. See also Fig. 44 and Figs. 211-303.

V, Monoclinic System. One plane of symmetry which contains two of the crystallographic axes. Also one axis of binary symmetry, normal to this plane and coinciding with the third crystallographic axis. See Figs. 37-39; also Fig 45 and Figs, 312-327,

Pyroxene.



VI. Triclinic System. No plane and no axis of symmetry, hut symmetry solely with respect to the central point. Fig. 40 and Fig. 46 show typical triclinie crystals. See also Figs. 333-341,

26. The relations of the normal groups of the different systems are further illustrated both as regards the cryetallographic axes and symmetry by the accompanying figures, 41-46. The exterior form is here that bounded by faces each of which is parallel to a plane through two of the crystallographic axes indicated by the central broken lines. Further, there is shown, within this, the combination of faces each of which joins the extremities of the unit lengths of the axes.

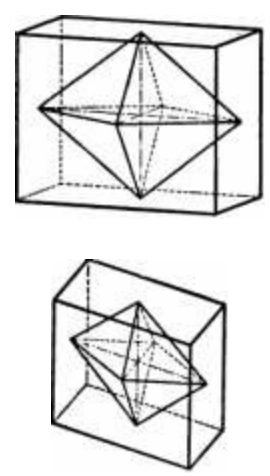
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GENERAL MORPHOLOQfCAL RELATIONS OP CRYSTALS.

17

The full understanding of the subject will not be gained ootil after a Btudjr of the forms of each system in deuil. Keverthelesa the student will do well to make himself familiar at the outset with the fundamental relation! here illustrated.

41. *a. 43.



It will be shown later that the symmetry of the different groups can be most clearly and easily exhibited by the nse of the spherical projection explained in Art. 39 ei aeq.

VI. Mod«li—Ulaas (or trnn'parcDt cellulotd) modeli IllustratiDg the dlSeTciit systema,

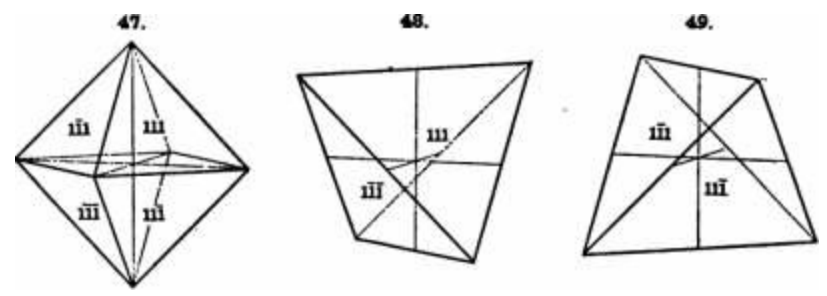
having the forms shown in Figs. 41-46. will be very useful to the student, especially in determining the relationship as regards symmetry. They show the outlines of one or more simple forms. Most of them are made in great variety and perfection of form; these are indispensable to the student in mastering the principles of crystallography.

38. So-called Holohedral and Hemihedral Forms.—It will appear later that each crystal form of the normal group in a given system embraces all the faces which have a like geometrical position with reference to the crystallographic axes; such a form is said to be holohedral (from *holos* and *hedra*, face). On the other hand, under the groups of lower symmetry, a certain form, while necessarily having all the faces which the symmetry allows, may yet have but half as many as the corresponding form of the normal group; these half-faced forms are sometimes called on this account hemihedral. Furthermore, it will be seen that, in such cases, to the given holohedral form there correspond two similar and complementary hemihedral forms, called respectively plus and minus (or right and left), which together embrace all of its faces.

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A single example will help to make the above statement intelligible. In the normal group of the cubic system, the octahedron (Fig. 47) is a "holohedral" form with all the possible faces—eight in number—which meet the axes at equal distances. In the tetrahedral group of the same system, the forms are referred to the same crystallographic axes, but the symmetry defined in Art. 19 (and more fully later) allows for but four similar faces in the position described. These yield a four-faced, or "tetrahedron," form, the tetrahedron. Figures 48 and 49 show the plus and minus tetrahedron, which together embrace all the faces of the octahedron, Fig. 47.



In certain groups of still lower symmetry; a given crystal form may have but one-quarter of the faces belonging to the corresponding normal form, and, after the same method, such a form is sometimes called tetartohedral.

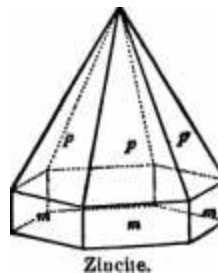
The development of the various possible kinds of hemihedral (and tetartohedral) forms under a given system has played a prominent part in the crystallography of the past, but

it leads to mncn complexity and is distinctly less simple than the direct statement of the symmetry in each case. The latter method is systematically followed in this work, and the subject of hemihedrism is dismissed with the brief (and incomplete) statements of this and the following paragraphs.

29. Hemimorphic Forms.—In several of the systems, forms occur under the groups of lower symmetry than that of the normal group which are characterized by this: that there is no transverse plane of symmetry, but the faces present are only those belonging to one extremity of an axis of symmetry (and crystallographic axis). Such forms are conveniently called hemimorphic forms. A simple example under the hexagonal system is given in Fig. 50. It is obvious that hemimorphic forms have no center of symmetry.

30. Molecular Networks.—Much light has recently been thrown upon the relations existing between the different types of crystals, on the one hand, and of these to the physical properties of crystals, on the other, by the consideration of the various possible methods of grouping of the molecules of which the crystals are supposed to be built up. This subject, very early treated by Hauy and others (including J. D. Dana), was discussed at length by Frankenheim and later by Bravais. More recently it has been extended and elaborated by Sohncke, Wulff, Schrieffer, Fedorow, Barlow, and others.* All solid bodies, as stated in Art. 7, are believed to be made up of definite

- See the literature (following Art. 83.



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physical units, called the physical, or crystal, molecules. Of the form of the molecules

nothing is definitely known, and though theory has something to say about their size, it is enough here to understand that they are infinitely small, so small that the surface of a solid— e.g. of a crystal—may appear to the touch and to the eye, even when assisted by a powerful microscope, as perfectly smooth.

The molecules are further believed to be not in contact but separated from one another—if in contact, it would be impossible to explain the motion to which the sensible heat of the body is due, or the transmission of radiation (radiant heat and light) through the mass by the wave motion of the ether, which is believed to penetrate the body.

When a body passes from the state of a liquid or a gas to that of a solid, under such conditions as to allow perfectly free action to the forces acting between the molecules, the result is a crystal of some definite type as regards symmetry. The simplest hypothesis which can be made assumes that the form of the crystal is determined by the way in which the molecules group themselves together in a position of equilibrium under the action of the inter-molecular forces.

As, however, the forces between the molecules vary in magnitude and direction from one type of crystal to another, the resultant grouping of the molecules must also vary, particularly as regards the distance between them and the angles between the planes in which they lie. This may be simply represented by a series of geometrical diagrams, showing the hypothetical groupings of

61. B3.

points which are strictly to be regarded as the centers of gravity of the molecules themselves. Such a grouping is named a network, or point-system, and it is said to be regular when it is the same for all parallel lines and planes, however they be taken. For the fundamental observed fact, true in all simple crystals, that they have like physical properties in all parallel directions, leads to the conclusion that the grouping of the molecules must be the same about each one of them (or at least about each unit group of them), and further the same in all parallel lines and planes.

The subject may be illustrated by Figs. 51, 52 for two typical cases, which are easily understood. In Fig. 51 the most special case is represented where the points are grouped at equal distances, in planes at right angles to each other. The structure in this case obviously corresponds in symmetry to the

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oabe described in Arts. 15 and 16, or, in other words, to the normal group of the isometric system. Again, in Fig. 53, the general case is shown where the molecules are unequally grouped in the three directions, and, further these directions are oblique. The symmetry

is here that of the normal group of the triclinic system.

If, in each of these cases, the figure be bounded by the simplest possible arrangement of eight points, the result is an elementary parallelepiped, which obviously defines the molecular structure of the whole. In the grouping of these parallelepipeds together, as described, it is obvious that in whatever direction a line be drawn through them, the points (molecules) will be spaced alike along it, and the grouping about any one of these points will be the same as about any other.

31. Certain important conclusions can be deduced from a consideration of such regular molecular networks as have been spoken of, which will be enumerated here though it is impossible to attempt a full explanation.

(I) The prominent crystalline faces must be such as include the largest number of points, that is, those in which the points are nearest together.

Thus in Fig. 53, which represents a section of a network conforming in symmetry to the structure of a normal orthorhombic crystal, the common crystal-faces would be expected to be those having the position ib , aa , mm , then

"j-v, ..., \

nn , and so on. This is found to be true in the study of crystals, for the common forms are, in nearly all cases, those whose position bears some simple relation to the assumed axes; forms whose position is complex are usually present only as small faces on the simple predominating forms, that is, as modifications of them. So-called vicinal forms, that is, forms taking the place

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of the simple fundamental forms to which they approximate closely in an equal position, are exceptional.

(3) When a variety of faces occur on the same crystal, the numerical relation existing between them (that which fixes their position) must be rational and, as stated in (1), a simple numerical ratio is to be expected in the common cases. This, as explained later, is found by experience to be a fundamental law of all crystals. Thus in Fig. 53, starting with a face meeting the section in turn, nn would be a common face, and for it the ratio is $l:l$ in the directions b and a ; mm would be also common with the ratio $3:1$.

(3) If a crystal shows the natural easy fracture, called cleavage, due to a minimum of cohesion, the cleavage surface must be a surface of relatively great molecular crowding, that is, one of the common or fundamental faces. This follows (and thus gives a partial, though not complete, explanation of cleavage) since it admits of easy proof that that plane in which the points are closest together is farthest separated from the next molecular

plane. Thus in Fig. 53 compare the distance separating two adjoining planes parallel to *bb* or *aa*; then two parallel to *mm*, *ll*, *nn*, etc. Illustrations of the above will be found under the special discussion of the subject of cleavage.

S2. Kindt of Molecular Groupings.—The discussion on the basis just described shows that there are fourteen possible types of arrangement of the molecules. These agree as to their symmetry with the seven groups defined in Art. 25 as representing respectively 6*.

the normal groups of the six systems a b o

with also that of the trigonal (or $\sqrt{3}$ -o' \sim o • fclli[^] • ^ ^ S ^ --- ^ the rhombohedral) division of the $\sqrt{3}$ [FNXjirT iW[^] ^ ^ v" hexagonal system. Of the fourteen, i i ■*

three groupings belong to the iso- Lo

metric system (these are shown, for * ')

sake of illustration, in Fig. 54, a,b,c, from Groth); two to the tetragonal; one each to the hexagonal and the rhombohedral; four to the orthorhombic system; two to the monoclinic, and one to the triclinic.

In its simplest form, as above outlined, the theory fails to explain the existence of the groups under the several systems of a symmetry lower than that of the normal group. It has been shown, however, by Sohncke and later by Fedorow, Schöenflies and Barlow, that the theory admits of extension. The idea supposed by Sohncke is this: that, instead of the simple form shown, the network may consist of a double system, one of which may be conceived of as having a position relative to the other (1) as if pushed to one side, or (2) as if rotated about an axis, or finally (Z) as if both rotated as in (3) and displaced as in (1). The complexity of the subject makes it impossible to develop it here. It must suffice to say that with this extension Sohncke concludes that there are 65 possible groups. This number has been further extended to 230 by the other authors named, but it still remains true that these fall into 33 distinct types as regards symmetry, and thus all the observed groups of forms among crystals, described under the several systems, have a theoretical explanation.

Literature. —A complete understanding of this subject can only be gained by a careful study of the many papers devoted to it, a partial list of which is added below. Further references particularly to the early literature are given in Sohncke's work (see below). An excellent and very clear summary of the whole subject is given by Groth in the third edition of his *Physikalische Krystallographie*, 1898.

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Wiilfiig. For title see p. 3.

GENERAL MATHEMATICAL RELATIONS OF

CRYSTALS.

33. Axial Ratio, Axial Plane.—The crystallographic axes have been defined

{Art. 22) as certain lines, usually determined by the symmetry, which are used in the description of the faces of crystals, and in the determination of their position and angular inclination, with these objects in view, certain lengths

55. of these axes are assumed as units to which the

occurring faces are referred.

The axes are, in general, lettered a, b, c , to correspond to the scheme in Fig. 55. To aid the memory, the letters may be further distinguished; as i (vertical axis); $d, 5^{\wedge}$ (shorter and longer lateral axes), etc.

If two of the axes are equal, they are designated a, a, d ; if three, a, a, ij . In one system, the hexagonal, there are four axes, lettered a, a, a, i .

Further, in the systems other than the isometric, one of the lateral axes is taken as the unit to which the other axes are referred; hence the lengths of the axes express strictly the axial ratio. Thus for sulphur (orthorhombic, see Fig. 57) the axial ratio is

$d:l\% 6 = 0.8131 : 1 : 1.9034$. For rutile (tetragonal) it is

$a : i = 1 : 0.64415$, or, simply, $6 = 0.64415$.

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23

The plane of any two of the axes is called an axial plane, and the space divided by the three axial planes is an octant, since the total space about the center is thus divided by the three axes into eight parts. In the hexagonal system, however, where there are three lateral axes, the space about the center is divided into 12 parts, or sectors.

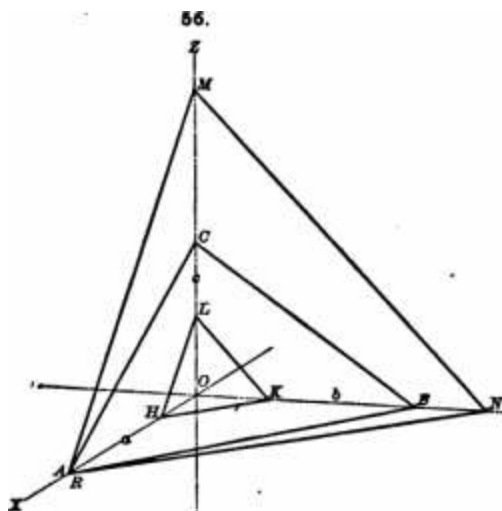
34. Parameters, Symbol.—The parameters of a plane are its intercepts on the assumed axes. The symbol expresses, often in abbreviated form, the relation of these intercepts to certain lengths of the axes taken as units,

For example, in Fig. 56 let the lines OX, OY, OZ be taken as the directions of the crystallographic axes, and let OA, OB, OC represent the unit lengths, designated (always in the same order) by the letters a, b, c . Then the parameters for the plane (1) HKL are OH, OK, OL ; for the plane (2) RNM they are OR, OA, OM . But in terms of the unit lengths

these axes, respectively,

$$(1) \frac{a}{h} + \frac{b}{k} + \frac{c}{l} = \frac{1}{n}$$

These two expressions are identical, since the two planes HKL, MNR are parallel and hence crystallographically the same. Obviously each of the above expressions may be changed into the other by multiplying (or dividing) by 4,



It will be noted that in (1) the numerators of the fractional numbers expressing the relation to the axes are all unity; while in (3) the number referring to one of the lateral axes (n) is made unity. The significance of this distinction will appear at once.

The general expression for any plane referred to these axes, written after the same method, will be

$$(1) \frac{a}{h} + \frac{b}{k} + \frac{c}{l} = \frac{1}{n}$$

$$\text{or (2) } la : nb : mc$$

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Here in (1) the numbers, or indices, hkl (in the case above, 433) constitute the symbol after the method of Miller (1839; earlier developed by Whewell and Orasfmann).

The second form (2) is the symbol essentially as early written by Weiss. This last was contracted by Kanmann to mPn imOn in the isometric system), the axes being omitted from the expression and the order reversed; the same with the omission of the P (or O), m-n, is adopted in Dana's System of Mineralogy, in the last edition (1892) of which work, however, the Miller symbols are given the preference.

In the hexagonal system there are assumed four axes, three of them lateral axes. Corresponding to this, in the symbols after the method of Miller as adapted by Bravais, there are four indices, hkit. The relation of these to the axes is the same as in the other

cases, as explained under the hexagonal system.

The following are other examples of planes with the symbols written after the two methods given. It will be seen that the respective expressions under (1) and (2) are identical.

HlUor** SfinbaL Muimum'i BrmboL

:j-« 821 or la: WsSe 2Por2

: le 818 " lo : S* : Ic i>8orl.2

la:<t>b:2o 3i^» or

la: 2b:a>o ooPSoi

If the axial values are measured behind for the axis a, to the left for J, or below for c, they are called negative, and a minus sign is placed over the corresponding number of the Miller symbols; as.

' 1 ■

It is soraetimPB slaieii that Nanmann"s armbols are the more ewy of comprelienaiOD b«csii5e more readily referred to the aiee. and tbU is In a measure true. If the Bludeut, however, wlllscuatom hhiiBelf m thick of the Miller symbols iu the form given above, that ie, always us the deiinomiftors of the fractional values ot the axes whose numerators am unity, he will never have any trouble in seeing the position of a given plane relatively to the axes. He must remember that the order is always that given ftbo?e. A, k. and (referring respectively lo the axes a, b, and e-, moreover, he will note that a zero, o, alw^iB roeaus that ihe given plane is parallel to the mis to whicli it refers, since t- = cd.

With the symbols of Naumann, the m, written first, always refers to the vertical axti, while the n, which follows, and is always greater than unity, refers to one of the ialeral axes. tUe other being made unity. To which lateral axis the n belongs is often indicated by a mark over the n (n, or n, or n), or attached to the Pas explained under the different syBtema. When m = 1, it Is omitted before the Por o (but not so when the Pis dropped); --' = I. it Is omitted in all c- -

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Of Ihete the Bjmbou of 'WelsB are easenllally thone alreaiiy given (under 2. p, 28) which, ■bbrieriated {anci inrerted In order), were adopted by Nnuamnu, The ajmbols of Ury bsTe been eiteusivel]' used by the Freuch school of ininerulogista, A. \ery full ezplanatlos of til the different systems, o8 of that recently devised by himself, la given in GoldBchmldt'i jDilez (IBH6-1B9I). Transformatlun eoualtons for the important cases, are giveu by Orolb

(Pbye. Kryet.J, Mallard (Cric, vol. 1), Lfebisch (Kryst.), and others ; see p. 2.

35. Law of Rational Indices.—The study of crystals has established the general law that the ratios between the intercepts on the axes for any face of a crystal to those of any other face can always be expressed by rational numbers. These ratios may be 1:2, 2:1, 2:3, 1:0 (as ∞), etc., but never $1:4^2$, etc. Hence the values of hkl in the Miller symbol must always be either whole numbers or zero, and similarly the m and n of Naumann's symbol may be whole numbers or fractions, or infinity.

If the form whose intercepts on the axes a, b, c determine their assumed unit lengths—the unit form as it is called—is well chosen, these numerical values of the indices are in most cases very simple. In the Miller symbols, 0 and the numbers from 1 to 6 are most common.

The above law, which has been established as the result of experience, in fact follows from the consideration of the molecular structure as hinted at in an earlier paragraph (Art. 31).

The law of rational indices finds an illustration later under the isometric system. It is stated there that three of the five regular solids of geometry, viz., the cube, octahedron, and the regular triangular pyramid (crystallographically the tetrahedron) all occur as crystal forms; the regular pentagonal dodecahedron and icosahedron, on the contrary, are impossible forms. This is true because their intercepts on the axes for such forms would be irrational; thus for the regular dodecahedron the ratio would be $1 : \sqrt{5}$. ■

There are, it is true, two forms respectively twelve-sided and twenty-sided which approximate to these regular solids, but their faces in the latter case are not regular polygons, ■ but in the second they are not all regular triangles. In the latter case it will be seen that the twenty faces in fact belong to two distinct forms, eight of one and twelve of the other.

36. Form.—A form in crystallography includes all the faces which have a like position relative to the planes, or axes, of symmetry. The full meaning of this will be appreciated after a study of the several

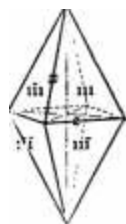
members. It will be seen that in the most general case, that of a form having the symbol (hkl) , whose planes meet the assumed unit axes at unequal lengths, there must be forty-eight like faces in the isometric system * (see Fig. 101), twenty-four in the hexagonal (Fig. 301), sixteen in the tetragonal (Fig. 166), eight in the orthorhombic (Fig. 57), four in the monoclinic, and two in the triclinic. In the first four systems the faces named yield an enclosed solid, and hence the form is called a closed form; in the remaining two systems this is not true, and such forms in these and other cases are called open forms. Fig. 275 shows a crystal bounded by three pairs of unlike faces; each pair is hence an open form.

Figs. 58-61 show open forms.

The unit or fundamental form is one where parameters correspond to the assumed unit lengths of the axes. Fig. 7 shows the unit pyramini of sulphur whose symbol is (111); it has eight similar faces, the position of which determines the ratio of the axes given in Art. 33.

* The normal group is referred to in each case.

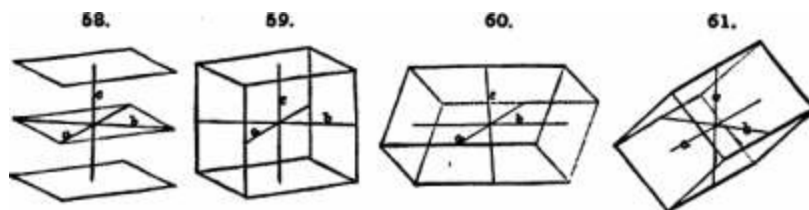
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The forms in the laometric system have special individual names, given later. In other systems certain general names are employed which may be briefly mentioned here. A form whose faces are parallel to two of the axes* is called a pinacoid (from pinax, & baird. It is shown in Fig. 58. One whose faces are parallel to the vertical axis but meet both the lateral axes is called a prism, as Fig. 6B. If the faces are parallel to one lateral axis only. It is a dome (Figs. 60, 61). If the faces meet all three axes, the form is a pyramid (Fig 57); this is usually given even if there is only one face belonging to the form.

In Fig. 6S. a (100), b (010), c (001) are pinacoids; m (110), n (120) are prisms; p (101), also A (011), q (021) are domes; all these are open forms. Finally, r (111), s (121) are pyramids, and in this case they are closed forms. The relation existing in each of these cases between the symbol and the position of the faces to the axes should be carefully studied.



(00)

As shown in the above cases, the symbol of a form is usually included in parentheses, as (111), (100); or it may be in brackets [111] or [111]. If the symbol is written without parentheses, as 111, it usually refers to a single face of the form only. Note also that with the Miller symbols, each face of a given form has its own individual symbol.

37. Zone, —A zone includes a series of faces on a crystal whose intersection-lines are

mutually parallel to each other and to a common line drawn through the center of the crystal, called the zone-axis. This parallelism means simply that the parameters of the given faces have a constant ratio for two of the axes. Some simple numerical relation exists, in every case, between all the faces in a zone, which is expressed by the zonal equation. The faces a, m, s, b (Fig. 62) are in one zone; also, b, k, l, c , etc.

If a face of a crystal falls simultaneously in two zones, it follows that its position is fixed and can be determined from the two zonal equations, without the measurement of angles. Further, it can be proved that the face corresponding to the intersection of two zones is always a possible crystal face, that is, one having rational values for the indices which define its position.

In many cases the zonal relation is obvious at sight, but it can always be determined, as shown in Arts. 43, 44, by an easy calculation.

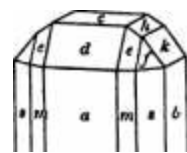
38. Horizontal Projection. — In addition to the usual perspective figures of crystals, projections on the basal plane (or more generally the plane normal to the prismatic zone) are very conveniently used. These give in fact a map of the crystal as viewed from above looking in the direction of the axis of the prismatic zone. Figs. 30, 32, 34 give simple examples; also Fig. 63 a projection of Fig. 62, both repeated from p. 16. In these the successive faces may be indicated by accents, as in Fig. 63, passing around in the

, prism and (101) a

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On the construction-



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direction of the axes a, b, a' , that is, counter-clockwise, of these projections see the 6a.

Appendix A.

39. Spherical Projection. — The study of actual crystals, as regards the angular and zonal relations of their faces, is much facilitated by the use of the spherical projection. In this the position of each face is represented by point called its pole, when a normal drawn to it from the center and produced meets the surface of the sphere. The

eymboh aftur Miller are immediately connected with this projection, and by means of it all ordinary calculations can be performed in a very simple manner. Fig. 65 shows a spherical projection of the orthorhombic crystal. Fig. 62.

If the center of the crystal, the point of intersection of the crystal's axes, is taken as the center of a sphere, auxiliary lines be drawn from it to the successive faces of the crystal, the lines which

meet the surface of the sphere, will be, as

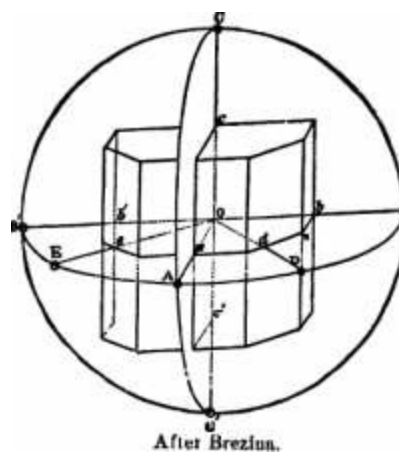
before the end, the poles of the crystal faces. For example, in Fig. 64, the common center of the crystal and sphere is O , the normal to the face ABC the surface of the sphere at B , of B' at B' . of d and e on D and E respectively, and so on. These poles evidently determine the position of the faces in each case.

It is obvious that the pole of the face b' (010) opposite $6(010)$ will be the opposite extremity of the diameter of the sphere, and so in general for (120) and $(120'$, etc. It is seen also that the poles, or normal points, of faces which lie in the same zone, that is, faces whose intersection lines are parallel, are in the same great circle, for instance $B(010), J(110), \dots, (100), \dots, (110)$, and so on. It is customary in the use of the sphere to represent it as projected upon a horizontal plane, usually the normal to the prismatic zone, so that, as in Fig. 65, the poles of the prismatic faces lie in the circumference of the circle, and those of the other faces within it. The eye being supposed to be situated at the opposite extremity of the diameter of the sphere normal to this plane, the great circles then appear either as arcs of circles, or as straight lines, i.e., diameters.

It will be further obvious from Fig. 84 that there are equal angles, between the normals of two faces, and between the normals of two faces, which is the supplement of the actual interior angle between the two faces; and this is true in general.

40. Construction of the Spherical Projection.—Since in the method ordinarily followed the poles of the prismatic faces lie in the circumference of the circle, their position is fixed at once by the angles laid off, e.g. from 100 , with a protractor. Further, the distances of the poles of all faces measured from the center of the circle (which, when the vertical axis is at right angles to those in the lateral plane, is the pole of the zone 001) are proportional to the tangents of half the angles. For example, to construct the spherical projection of Fig. 83, first draw a circle, and lay off on the circumference, from a point taken as 100 , the angular distances characteristic of the species (in this case):

am. 100 a $HO = 34^\circ 58'$; at, 100 a $120 = 48^\circ 68'$; oi, 100 a $010 = 90^\circ$.



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The positions of the poles of the faces a (100), m (110), n (120), h (010) are thus fixed. The poles of the other faces of these forms are obviously fixed. a', m', n', h', i', m'', n'', h'', i'', by

the symmetry. Again, to find the

position of the pole of d (101), which lies on the

great circle, or zone, a (100), e (001).

a' (100) (for which $k = \frac{h}{r}$ = Hr. mdev d

(001) A d (101) = sr ittr, the distance

is perpendicular to the great circle of

the zone (001) or $25^\circ 41'$, the distance of

the radius eu. Similarly for A (101)

on the zone-circle c (001),

the distance cA (001 A (101) = $80' 48''$

and eA (001 A (101) = $49' 88''$, the distances

of the poles are proportional to the sines of the

of half the angle respectively. So

also from the angle cA (001 A (111) =

$51' 16''$ and if (001 A (121) = $59' 50''$

ibc distances on the OorreapoDdJur

ione-cireU-8, e (001) m (110) aud e

(OUI) 1(120), may be delermined. In

priicilce, lioweever, tthese Inst steps are

Huiieci ssiiry ; since If thie circular arc

through b (010). d (101). 6" (010) Is

drawn, it elves the zone-circle for all

Ibe fiices for which $A =$ (; similarly

\wedge ... thai through a (100). h (Oil), a! (100)

"* \wedge give tbe zone circle for the plaues for

which $k = 1$, while thnl through a

(100), J; (031), a' (100) gives thiu zone-circle for the planes having $jt = 3i$. The
tutersecUoo-

poluts belweeu tiiese last arcs and ihut first drawn Axes Ibe)>o£itioDSof «(lll),/(12l),

pnch of which sutisties the two relutious. Further, through these same poinis must i>ass

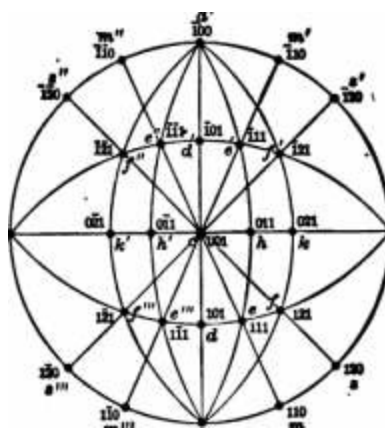
lliezone.circlee(001), in (110). for which $ft = t$, mid c (001). * (120). for which $t = 2A$, tbua

giving a check '

It is< Ihe intersection of either n

100, 031, iOO 010, 401. OiO 001. 480, OOi. In geaernl any fnce. hkl. iniisl lie In the three
zoue-clrcles

100. OW. Too. 010. AW. OiO, 001, AAO, OOl.



41. AnglM between Face*.—The angles most conveniently used with the Miller symbols, and those given in this work, are the normal angles, that is, the angles between the poles or normals to the faces, measured on arcs of great circles joining the poles as shown on the spherical projection, These normal angles are the supplements of the actual interfacial angles, as has been explained.

The relations between these normal angles, for example in a given zone, is much simpler than those existing between the actual interfacial angles. Thus it is always true that, for a series of faces in the same zone, the normal angle between two consecutive faces is equal to the sum of the angles of faces falling between. Thus (Figs. 62. (W)) the normal angle of ab (100 A 010) is the sum of am (100 A HO), m (110 a 130), and la (130 a O10). This relation holds true in all the systems.

Furthermore. It will be seen that, supposing acn' (Fig- 65) a plane of symmetry as in the orthorhombic system, the angle 100 A 110 (Fig. 62), is half the angle 110 A 110 (m'). Similarly 010 a 120 (6t) is half the angle 120 a 120 (m') ; again. 100 a 111 (ne) is the complement of half the angle 111 A 111 (aa) and 010 a 111 (&:) the complement of half the angle 111 a 111 (*^') ■

Here, as throughout this work, the symbol A is used to represent the angle between two faces, usually designated by letters.

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48. Use of the Spherical Projection to Exhibit the Symmetry.—The symmetry of any one of the crystalline groups may be readily exhibited by the help of the spherical projection, following the notation introduced by Gadow (1871, see p. 22).

The axes of binary, trigonal, tetragonal, hexagonal symmetry are represented respectively by the following signs: $\wedge A \wedge \#$. Further, a plane of symmetry is represented by a full line (zone-circle), while a dotted line indicates that the plane of symmetry is wanting. The position of the crystallographic axes is shown by arrows at the extremities of the lines. The pole of a face in the upper half of the crystal (above the plane of projection) is represented by a cross; one below by a circle. If two like faces fall in a

vertical zone a double sign is need, a cross within the circle. Figs. 69,111,136, etc., eive jllustrations. 43. Qener&l Bel&tlona between Planet m the Same Zone.—It may be demonstrated that if on a crystal two faces P {kkl) and B (pgr) lie in the same zone, then the follovring equation must hold good :

$$Ma \cos XQ + vb \cos YQ + wc \cos ZQ = 0,$$

where $a = kr - Ig$, $v = Ip - kr$, $w = hq - kp$.

The letters n, v, w are called the symbol of the zone or great circle PR. Every face (xyz) of this zone must satisfy the equation

$$ax + by + cz = 0.$$

If now (avw) be the symbol of one zone, and (efg) of another intersecting it, then the 'point of intersection will always be the pole of a possible crystal face. Its indices (hkt) must obviously satisfy two equations similar to (1). These indices are hence equal to

$$A = gv - fw, * = ew - gu, i = fa - ev.$$

The application of this principle is extremely simple, and its importance cannot be overestimated.

The zone-symbol B can be always obtained by rearranging the symbols of the two faces B la oHer. repeating the first two indices and then multiplying according to the following scheme:

I

xx:

$$HcDce \ i = kr - lq \ ^ = lp - !ir: w = ft? - tp.$$

44 BxftKVIM of Zoom kind Zonal Bvlatloni.—The following are cases in which the zonal equilibrium is not once. In Fig. 62. p. 27. the faces $a(100)$, $m(110)$, $.(130)$, $6(010)$, form a vertical zone with mutually parallel rulings, since the $^$ are alike in position. In so far as (this: that they are all parallel to the vertical axis; that is, for all faces in this zone it must be true that $l = 0$.

Also in the faces $a(100)$, $d(101)$, $e(001)$ are in zone, all hoini; parallel to a lateral axis i ; hence for them and all others in this zone $i = 0$. Also for (010) , $*(031)$, $A(001)$, $<(001)$ are in a zone, all being parallel to the axis a , so that $h = 0$.

Also the faces (181) , $\langle(1\rangle)$, $\langle*(100)$, $^(")$, $^(")$ are in a zone, since they have a common ratio for i . In fact, with them, obviously, $h = l$.

The faces (111) , (110) are also in a zone, and again (001) , (121) , (120) , though interrelated, can not happen between (111) in the one case, and (110) and (120) in the other. For each of the zones it is true that there is a common ratio of the lateral axes. That is, of h k in the symbols. For the (111) zone it may be shown that $A = h$; for the second, that $A = k$.

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$$u = 3, v = h, w = 0: \frac{h}{k} = \frac{u}{v} = \frac{w}{0}$$

The symbol of a face that lies in two zones, as stated above, must satisfy the zonal equation of each; these symbols are hence easily obtained either by combining the equations or by a scheme of multiplication like that given above.

"mole, in Fig. 66, of sulphur, the face lettered x is in the zone (1) with $6(010)$ and

0 in zone (2) with $6(111)$

(113) , also (111)

(2) with $6(111)$ and $6(010)$. These zones give, respectively :

$$(1) \quad 0 \ 1 \ 0 \ 0 \ 1 \quad (2) \quad 1 \ 1 \ 1 \ 1$$

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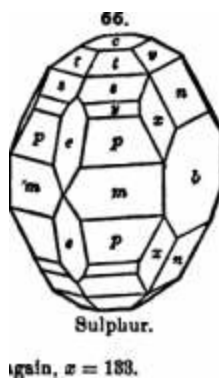
$$1 \ 1 \ 1 \ 1 \ 0 \ 110 \ 1$$

$$u = 8, v = 0, w = h: \frac{h}{k} = \frac{u}{v} = \frac{w}{0} = \frac{g}{1}$$

Hence for (1) the zonal equation is $h = 1$; for (2) $k = 1$. Combining these, we obtain $A = 1$, $k = 3$, $h = 3$. The symbol of the face x is, therefore, 138 . This result is given by multiplying the zonal indices 010 , 801 , together after the same method, thus:

$$0 \ 1 \ 1 \ 0 \ 1$$

■ XXX



This method of calculation belongs to all the diSereut s^lems. In the hexagonal system, in which there are four indices, one of the three referring to the lateral axes (usually the third) is omitted when the zonal relations are applied. See Art. 160.

46. Method of Calculation.—In general the angles between the poles can be calculated by the methods of spherical trigonometry from the triangles shown in the sphere of projection (Fig. 65)—which for the most part are right-angled. Certain fundamental relations connect the axes with the elemental angles of the projection ; the most important of these are given under the individual systems. Some general relations only are explained here.

46. Relation between the Indices of a Plane and the Angle made by it with the Axes.—When the assumed axes are at right angles to each other they coincide with the normals to the pinacoid faces (100, 010, 001), and consequently meet the spherical

surface at their poles. When the axial angles are not 90°, this is no longer true. In all cases, however, the following relation holds good between the cosines of the angles made by a plane, HEL, with the axes:

$$\cos \angle HOP = \frac{h}{a}$$

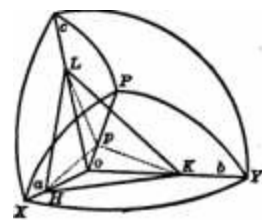
$$\cos \angle EOP = \frac{e}{b}$$

This is equivalent to

$$\cos \angle HOP = \frac{h}{a}$$

$$\cos \angle EOP = \frac{e}{b}$$

$$= \cos \angle POZ.$$



$$-\cos PX = T \cos FY = 7 \text{ COB PZ.}$$

,ab,GoOgIc

This equation is fundamental, and several of the relations given beyond are deduced from it.

... etlie Albo if ibe plaue E

OH -a. OK = 6, OL = e. The line HK, HL, EL give also the intersections of the planes 110, 101, 010 on the three axial planes, and their poles are hence at the points fixed by Donaula to these lines drawn from O. It will be obvious from this figure, then, that the following relations hold true:

$$\tan (100 \wedge 110) = -\frac{a}{b};$$

$$\tan (001 \wedge 101) = -\frac{a}{c};$$

$$\tan (001 \wedge 010) = \frac{a}{c}. \text{ These values are often used later.}$$

47. Cotangent and Tangent Relations.—If the angles between the poles of three faces in a zone are known, the angle between any one of them and the pole of a fourth face can be calculated by a formula called the cotangent formula. Conversely, if the angular position of this fourth face is given, the ratio of its indices can be calculated.

Let P, Q, S, R be the poles of four faces in a zone, taken in such an order that $PQ < PR$, and let the indices of these faces be respectively : P Q R S

hkl pqr vte xj/t Then it may be proved that

If one of these fractions reduces to an integer form, —, then one of

the others must be taken in its place.

This formula is chiefly used in the monoclinic and triclinic systems; and some special cases are referred to under those systems.

The cotangent relation becomes much simplified for a rectangular zone, that is, a zone between a pinacoid and a face in the zone of the other pinacoids at right angles to it. Thus if Pa, P6, Po, Qa, Q6, Qe represent respectively the angles between two faces in the same rectangular zone, viz., P (hkl) and Q (h'k'l') and the pinacoids a (100), b (010), c (001), the following relations hold good :

A line Pa k l

$$\frac{\tan Q_a}{q} = \frac{r}{r'}$$

$$\frac{k \tan P_b}{q} = \frac{l}{r'}$$

$$\frac{k}{q} = \frac{l}{r'} \cdot \frac{\tan P_c}{\tan Q_b}$$

* In the application of this principle it is essential (that the planes should be taken in the proper order, as shown above ; to accomplish this it is often necessary to use the indices and corresponding angles, not of (UO, but the face opposite (hkCi, etc.

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the above equations for I .

found pgo, or domes MH and jiOr or Qkl & Ojr, between two planar planes, we have

$$\tan(100 \text{ A } hkQ) \cdot \tan(100 \text{ A } pqO) \sim A - q' \tan(001 \text{ A } AOQ) \cdot \tan(001 \text{ A } pOr) \cdot i' \cdot p'$$

$$\tan(001 \text{ A } o^*o) \cdot r \cdot \tan(001 \text{ A } o^?r) \sim i \cdot 3 \cdot$$

These equations are the ones ordinarily employed to determine the azimuth of a crystal plane or dome.

The most common and important application of this tangent principle is where the positions of the unit faces 110, 101, 010 are known, in which case the relation becomes

$$\tan(100 \text{ A } ftO) \cdot \tan(010 \text{ A } hkO) \cdot h$$

$$\tan(100 \text{ A } HO) \cdot \tan(010 \text{ A } 110) \sim i'$$

$$\tan(001 \text{ A } AOf) \cdot \tan(001 \text{ A } Okl) \cdot k$$

$$\tan(001 \text{ A } 101) \cdot \tan(001 \text{ A } Oil) \sim r$$

Angles between the base, 001, and 102, 208, 303, 301, etc., are the tangent of the angle between 001 and 101. Again, the

tangent of the angle 100 A 130 is twice the tangent of 100 A HO [here t = 3], and one-half the tangent of 010 A HO.

4S. Formulas for Spherical Trigonometry.—For convenience, some of the more important formulas for the solution of spherical triangles are here added.

In right-angled spherical triangles

$$\sin A = \sin B$$

$\sin b$

$\tan 6, r. * \gg \gg$

$\cos J = \frac{a^2 + b^2 - c^2}{2ab}, \cos E = \frac{a^2 + c^2 - b^2}{2ac}$.

$\tan A \tan A$

In oblique-angled spherical triangles familiar relations are (1) $\sin J : \sin E = \sin a : \sin$

$\sin b \sin c \sin A; A + a \sin A \cos B; C + c \sin C \sin B \sin C \sin A$.

In the calculation it is often more especially arranged for logarithmic to mathematical formulae.

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ISOHETBIC SYSTEM.

ISOMETRIC SYSTEM.

49. The Isohetic System embraces all the forms which are referred to three equal axes at right angles to each other. Each of these axes is designated by the letter a.

There are five groups here included, of which the normal group,* possessing the highest degree of symmetry for the system and, indeed, for all crystals, is by far the most important. Two of the other groups, the pyritohedral and tetrahedral groups, also have numerous representatives among minerals.

1. NORMAL GROUP (1). GALENA TYPE.

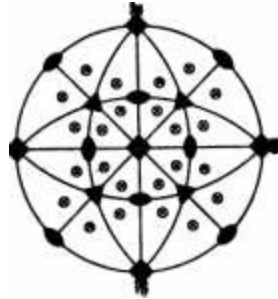
50. Symmetry.—Of each of the types of solids enumerated in the following table, as belonging to this group, as of all their combinations, it is true that there are three like principal planes of symmetry, whose intersections fix the position of the crystallographic axes (see Fig. 12, p. 9). There are also six other auxiliary planes of symmetry; these are situated diagonally to the others, each two equally inclined (45°) to the adjacent planes of chief symmetry, that is, to the axial planes.

Further, the crystals of this group have three principal axes of tetragonal symmetry, the cubic or crystallographic axes; four axes of trigonal symmetry, like octahedral axes; six axes of binary symmetry, the dodecahedral axes (see Art. 16, in the following paragraph). These axes are shown in Figs. 17, 18, 19, p. 10.

The accompanying spherical projection (Fig. 10, with the principles explained in Art. 42, shows the distribution of the faces of the general form, hkl, and hence represents clearly the symmetry of the group. Compare also the projection given later, fig. 110, p. 41.

51. Forms.—The various possible forms belonging to this group, and possessing the symmetry defined, may be grouped under seven types of solids.* These are enumerated in the following table, commencing with the most simple. The symbols are given in accordance with both the systems of Miller and Naumann; also the full expression showing the general position of the planes with relation to the axes. The last, however, are reduced to the form, corresponding to (2) in Art. 34, which shows how the Naumann symbols are

if constructed in accordance



* It is the most common and hence by far the most important group in the system; also, more fundamentally, because the forms here included possess the highest grade or symmetry possible in the system. There are five forms in this system, each geometrically a cube, but only that of this normal group actually has the full symmetry of a regular molecular structure which its geometrical shape suggests. If a crystal is said to belong to the isometric system, without further qualification, it is to be understood that it is included here. Similar remarks apply to the normal groups of the other systems.

The symmetry of the normal groups of the different systems has been already briefly explained in Art. 50.

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Miller. Notation.

1. Cube (100) $a : a : a$ or t

2. Octahedron (111) $a : a : a$ or t

3. Dodecahedron (110) $a : a : \sqrt{2}a$ or t

4. Tetrahedron (111) $a : a : a$ or t

5. Rhombohedron (100) $a : a : c$, etc

6. Triclinic (hkl) $a : b : c$ or m

w, (331) 3; (221) 2; (333) j, etc.

6. Trapezohedron (A//) a : ma : ma mOm or m-m

as, (311) 3-3; (211) %■%; (332) f-f, etc.

7. Hexoctahedron {hkl) a:na:ma mOn or m-n

as, (421) 4-2; (321) 3-), etc. In the general expression of Miller's symbols, $A > B > C$. In the notation of Naumiun, $\frac{a}{b} > 1$, Atteolio is called to the letters nlnfonly used in this work and in Duut'a SjrMem ol Hloerslog; (1893) to designate certain of the isometric forms.* They are:

Cube: a, ,

Octahedron: o.

Dodecahedron: d.

Tetrahedron: t = 310, $\frac{a}{b} = 2$; / = 81D, i-8; g = 9S6, $\frac{a}{b} = 410, \frac{a}{b}$.

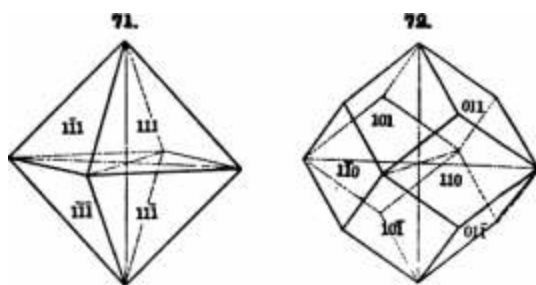
Trisoculiedrons: ; = 231, 2; q = S&\. 8; r = 832, 4; p = 411, 4.

Trapezohedrons: m = 311, 8-3; n = 211, $\frac{a}{b} = 2$; fl = &£, |(-

Hexoculiedrons: ■ = 321, 8-|; (= 421, i-2.

63. Cube.—The cube, whose general symbol is (100), is shown in Fig. 70. It is bounded by six similar faces, each parallel to two of the axes. Each face is a square, and the interfacial angles are all 90° . The faces of the cube are parallel to the principal or axial planes of symmetry. The lines joining the opposite solid angles of the cube are called the octahedral or trigonal inter-axes; those joining the middle points of opposite edges are the dodecahedral interaxes (see Figs. 17, 18, p. 10).

53. Octahedron.—The octahedron, shown in Fig. 71, has the general symbol (111). It is bounded by eight similar faces, each meeting the three axes at



equal distances. Each face is an equilateral triangle with plane angles of 60° . The normal interfacial angle, $(111 \wedge 111)$, is $70^\circ 31' 44''$.

54. Dodecahedron.—The rhombic dodecahedron, shown in Fig. 72, has the general symbol (110). It is bounded by twelve faces, each of which meets two

* The usage followed here (as also in the other systems) is in most cases that of Miller.
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of the axes at equal distances and is parallel to the third axis. Each face is a rhomb with plane angles of $70\frac{1}{2}^\circ$ and 109° . The real or interior interfacial angle is 120° , or the angle between two adjacent poles, that is, the normal interfacial angle, is 60° . The faces of the dodecahedron are parallel to the six auxiliary, or diagonal, planes of symmetry.

It will be remembered that, while the forms described are designated respectively by the symbols (100), (111), and (110), each face of any one of the forms has its own symbol. Thus for the cube the six faces have the symbols

100, 010, 001, 100, 010, 001

For the octahedron the symbols of the eight faces are :

Above 111, 111, 111, 111;

Below 111, 111, 111, 111. For the dodecahedron, the symbols of the twelve faces are :

110, 110, 110, 110,

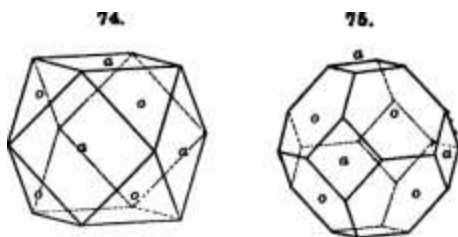
101, 101, 101, 101,

011, 011, 011, 011.

These should be carefully studied with reference to the figures (and to models), and also to the Spherical projection (Fig. 110). The student should become thoroughly familiar with these individual symbols and the relations to the axes which they express, so that he can give at once the symbol of any face required.

65. Combination of the Cube, Octahedron, and Dodecahedron.—Figs. 73, 74, 75 represent combinations of the cube and octahedron ; Figs. 76, 79, of the cube and dodecahedron ; Figs. 77, 78, of the octahedron and dodecahedron ; finally, Figs. 80, 81 show combinations of the three forms. The predominating

W S A



form, as the cube in Fig. 73, the octahedron in Fig 75, etc., is usually said to be modified by the faces of the other forms. In Fig. 74 the cube and octahedron are said to be "in equilibrium," since the faces of the octahedron meet at the middle points of the edges of the cube.

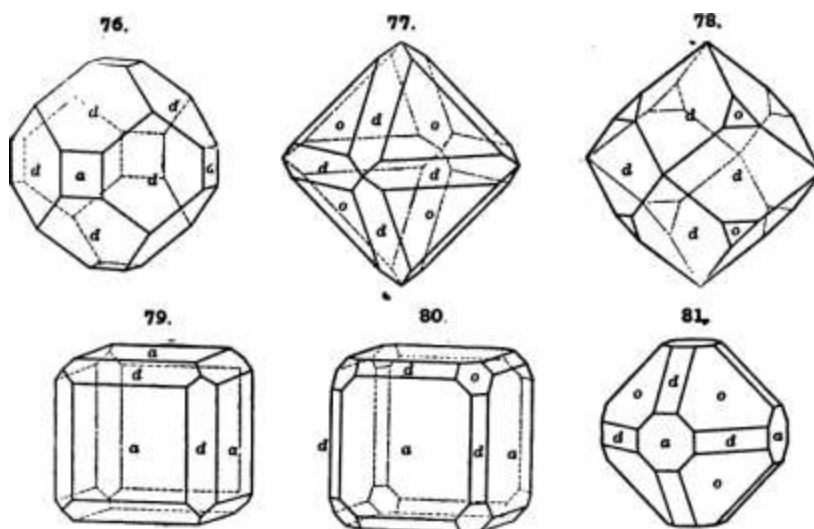
It should be carefully noticed, further, that the octahedral faces replace the solid angles of the cube, as regular triangles equally inclined to the adjacent cubic faces, as shown in Fig. 73. Again, the square cubic faces replace the six solid angles of the octahedron, being equally inclined to the adjacent octahedral faces (Fig. 75). The faces of the dodecahedron truncate the twelve

edges of the cube, as shown in Fig. 79. They also truncate the twelve edges of the octahedron (Fig. 77). Further, in Fig. 76 the cubic faces replace the six tetrahedral solid angles of the dodecahedron, while the octahedral faces replace its eight trihedral solid angles (Fig. 78).

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Similar edges of the cube, as shown in Fig. 79. They also truncate the twelve edges of the octahedron (Fig. 77). Further, in Fig. 76 the cubic faces replace the six tetrahedral solid angles of the dodecahedron, while the octahedral faces replace its eight trihedral solid angles (Fig. 78).



The normal interfacial angles for adjacent faces are as follows : Cube on octahedron, $m, 100 \wedge 111 = 54^\circ 44' 8''$.

Cube on dodecahedron, $ad, 100 \wedge 110 = 45^\circ 0' 0''$.

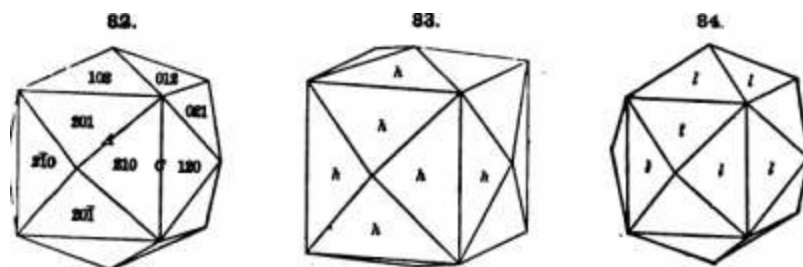
Octahedron on dodecahedron, $od, 111 A 110 = 35^\circ 15' 52''$, t_6 . As explained in Art. K, actual crystals always deviate more or less from the ideal solids. The deviation of the actual from the ideal is due to the fact that the faces of the crystal are not perfectly flat. Such deviations, however, do not vitiate the geometrical definition of right symmetry relatively to the principal axes. The six auxiliary planes mentioned on p. 33. still they do conform to the conditions of crystallographic symmetry, regarding like angular positions for similar faces. Again, it will be noted that in a combination of forms many of the faces do not actually meet the axes within the crystal, as, for example, the octahedral face of Fig. 73. It is still true, however, that this face would meet the axes at equal distances if produced: and since the axial ratio is the essential point in the case of each form, and the intercepts of the axes are of no importance, it is not necessary that the faces of the different forms in a crystal should be referred to the same actual axial lengths. The above remarks will be seen to apply also to all the other forms and combinations of forms described in the pages following.

67. Tetrahedron.—The tetrahedron (Figs, 82, 83, 84) is bounded by twenty-four faces, each of which is an isosceles triangle. Four of these faces together occupy the position of one face of the cube (hexahedron) whence the name commonly applied to this form. The general symbol is (hkO) , hence each face is parallel to one of the axes while it meets the other two axes at unequal distances. There are two kinds of edges, lettered A and C in Fig, 83 ; the interfacial angle of either edge is sufficient to determine the symbol of a given form (see below). The angles of some of the common forms are given on a later page (p. 42).

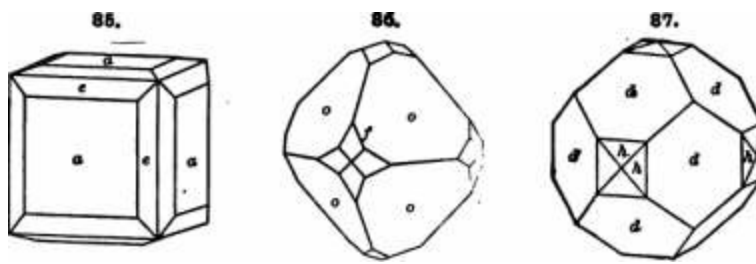
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ISOMETRIC SYSTEM.

There may be an indefinite number of tetrahedrons, as the intercepts of the two axes, and hence of h to i : varies ; for example (310) , (210) , (320) , etc. The form (110) is shown in Fig. 82 ; (410) in , and (530) in Fig. 84, All the tetrahedrons fall in a zone with a cubic and a dodecahedral face. As h increases relatively to i : the form approaches



the cube (in which $h:i = x : 1$ or $1:0$), while as it diminishes and becomes more and more nearly equal to h in value it approaches toward the dodecahedron ; for which $A = A$. Compare Fig. 83 and Fig. 84; also Fig. 110. The special symbols belonging to each face of the tetrahedron should be carefully noted.



The faces of the tetrahezedron bevel * the twelve similar edges of the cube, as in Fig. 85; they replace the solid angles of the octahedron by four faces inclined on the edges (Fig. 86), and also the tetrahedral solid angles of the dodecahedron by four faces inclined on the faces (Fig. 87).

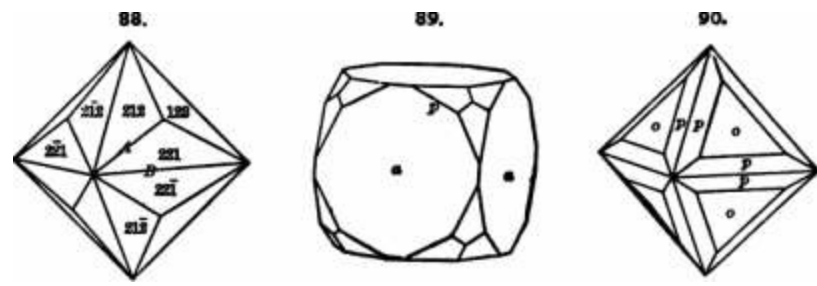
58. TrisootahedroE.—The trisoctahedron (Fig. 88). or, more definitely, the trigonal trisoctahedron, is bounded by twenty-four similar faces; each of these is an isosceles triangle, and three together occupy the position of an octahedral face, whence the common name. Further, to distinguish it from the trapezo-hedron or tetragonal trisoctahedron, it is sometimes called the trigonal trisoctahedron. There are two kinds of edges, lettered A and B in Fig. 88, and the interfacial angle corresponding to either is sufficient for the determination of the special symbol.

* The word bmtl (b used when two like faces replace the edge of a form and hence are inclined at equal angles to its adjacent similar faces.

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The general symbol is $h(hhl)$; common forms are (231) , (331) , etc $E \ll cl \gg$



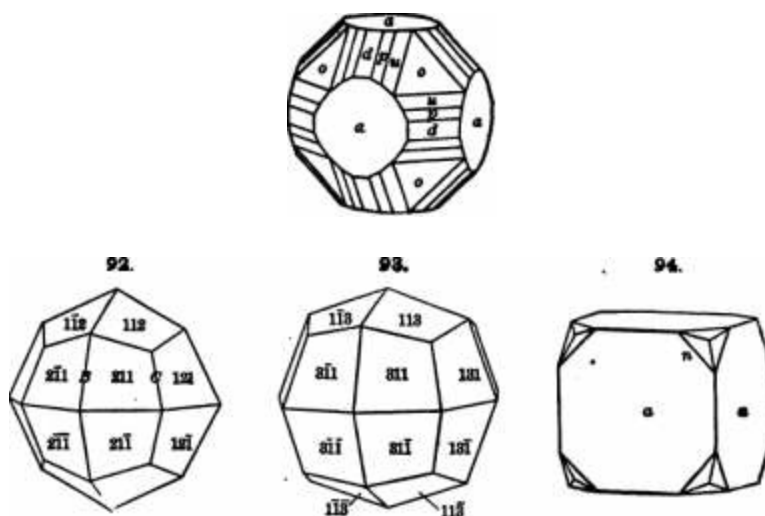
face of the trisoctahedron meets two of the axes at a distance less than unity and the third at the unit length, or (which is an identical expression *) it meets two of the axes at the unit length and the third at a distance greater than unity. The symbols belonging to each face should be carefully noted.

The normal interfacial angles for some of the more

common forms are given on a later page.

69. Trapezohedron.—The trapezohedron f (Figs. 92, 93) is bounded by twenty-four similar faces, each of them a quadrilateral or trapezium. It also bears in appearance a certain relation to the octahedron, whence the name, sometimes employed, of tetragonal trisoctahedron. There are two kinds of edges, lettered B and C, in Fig. 92. The Oaiana.

general symbol is hll ; common forms are (311) , $(3\bar{1}1)$, (333) , etc. Of the faces, each cuts an axis at a distance less than unity, and the other two at the unit length, or (again, an identical expression) one of them intersects an axis at the



unit length and the other two at distances greater than unity. The symbols belonging to each face should be carefully noted. The normal interfacial

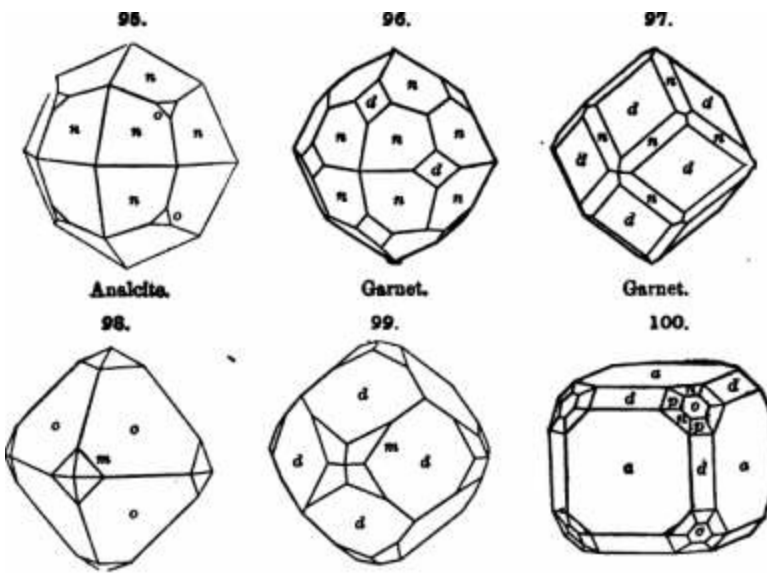
" Since $ia:ib:ie=la:lb:2o$. The student should read again carefully the explanations in Art. U.

It will be seen later that the name for these forms is trapezohedron, conspicuously the trapezohedron.

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angles for some of the common forms are given on a later page. Another name for this form is icositetrahedron.

60. The combination of these forms with the cube, octahedron, etc., should be carefully studied. It will be seen (Fig. 89) that the faces of the trisoctahedron replace the solid angles of the cube as three faces equally inclined on the edges. The faces of the trapezohedron appear as three equal triangles equally inclined to the cube (Fig. 94).



Spinet

Magnetite.

AmKlgam.

Again, the faces of the trisoctahedron bevel the edges of the octahedron (Fig. 90, also Fig 91, with $p(2^1)$ and $m(554)$), while those of the trapezohedra are triangles inclined to the faces at the extremities of the cubic axes (Fig. 98), Still again, the faces of the trapezohedron (211) truncate the edges of the dodecahedron (110) , as shown in Fig, 97; this can be proved to follow at once from the zonal relations (Arts. 43, ii), cf. also Fig. 110. The position of the faces of the form (311) , in combination with o , is shown in Fig. 98; with $'$ in Fig. 99. Fig, 100 shows both the trisoctahedron $p(221)$ and the trapezohedron $n(211)$ with a , o , and d .

It should be added that the trapezohedron $n(211)$ is a common form both alone and in combination; $m(311)$ is common in combination. The trisoctahedron alone is rarely met with, though in combination (Figs. 90, 91, 100) it is not uncommon.

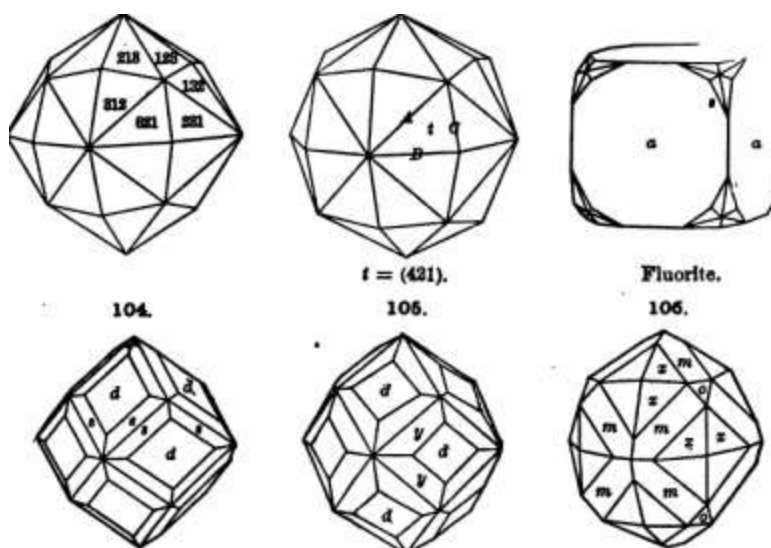
61. Hexoctahedron D,—The hexoctahedron. Figs. 101, 102, is the general form in this system; it is bounded by forty-eight similar faces, each of which is a scalene triangle, and each intersects the three axes at unequal distances. The general symbol is (kl) ; common forms are (321) , shown in Fig. 101, and (4^1) , in Fig. 103. The symbols of the individual faces, as shown in Fig. 101 and more fully in the projection (Fig. 110), should be carefully studied.

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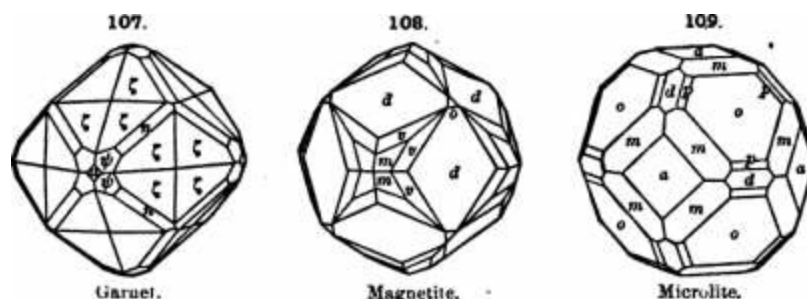
CRTSTALLOGUAPHT.

The hexoctahedron D has three kinds of edges lettered A, B, C (longer, middle, shorter) in Fig. 103; the angles of two of these edges are needed to fix the symbol unless the zonal

relations can be made use of. In Fig. 104 the faces of the hexoctahedron berel the dodecahedral edges, and hence for this



form $h = i : -1$; the form t has the special symbol (321) . The hexoctahedron alone is a very rare form, but it is seen in combination with the cube (Fig. 103, Anorite) as six small faces replacing each solid angle. Fig. 104 is common with garnet; Fig. 105 shows a combination observed in native copper ($y = 18 \blacksquare 10 \blacksquare 5$), and Fig. 106 with native gold ($.1; = 18-10-1$). The angles of some common hexoctahedrons are given on p. 42.



62. Some further examples of isometric forms are given in Figs. 107, 108, 109. In Fig. 107, \wedge is the trapezohedron (723) ; C is the hexoctahedron

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18oMBTBIo STSTKM.

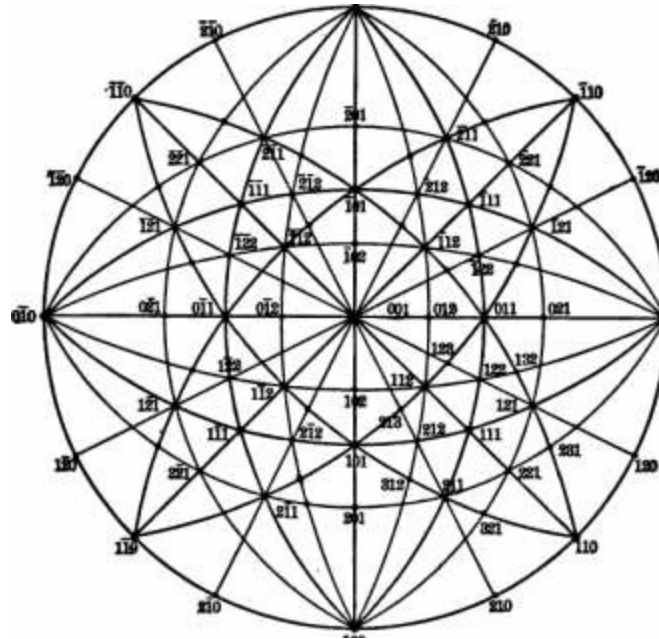
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($64 \blacksquare 63-1$), this last being called a rhomboform, since it deviates but slightly in angular position from the simple form ordinarily occurring ($d, 110$): hence the complex indices. In Fig. 108, v is the hexoctahedron (531) . In Fig. 109, $m = (311)$, $o = (221)$, etc.

63. Pseudo-symmetry in the Isometric System.—Isometric forms, by development in the direction of one of the cubic axes, simulate tetragonal forms. More common, and of greater interest, are forms simulating those of orthorhombic symmetry by extension, or

flattening, in the direction of an octahedral axis. Both these cases are illustrated Uter. OonTersely, certain rhombohedral forms resemble an isometric octahedron in angle and complex twinning.

64. Spherioal Projeotioo.—The spherical projeotion, Fig. 110, shows the



positions of the poles of the faces of the cube (100), octahedron (111), and dodecahedron (110); also the tetrahexahedron (210), the trisoctahedron (321), the trapezohedron (211), and the heioctuhedron (321).

The student should study this projection carefully, noting the symmetry indicated by the zone-circles 100, 001, 100, and 100, 010, 100; also by 110, 001, 110; 110, 001, 110; 010, 101, 010; 010, 101, 011. Note further that the faces of a given form are symmetrically arranged about a cubic face, as 001; a dodecahedral face, as 101; an octahedral face, as 111.

Note further the symbols that belong to the individual faces of each form, comparing the projection with the figures which precede.

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Finally, note the prominent zone-circles, for example, the zone between two cubic faces including a dodecahedral face and the faces of all possible tetrahedra. Again, the zones from a cubic face (as 001) through an octahedral (as $h^* 111$) through the trisoctahedron, as 118, 112, 228, and the trapezohedrons 883, 3S1, 8B1, etc. Also the zone from one dodecahedral face, as 110, to another, as 101, passing through 331, all, 812, etc. At the same time compare these zones with the same zones shown on the figures already described.

66. Angle* of Common Iometrie Formi.* Atkahexahbdronb,

Edge A Edge C

Cf. Fig. 88. 810 A 201, etc 310 A 120, et 1» 46 61 65(

86 001

410, v4 810, i-S 080, »4 810. i 8 680, H

sao, ^

480, H 640. f-f

Tbiboohdbbohs.

821,8 602,1 831,8 773.1 441.4 Tbafezohsdbon b .

Cf. Fig. 82. 411,4-4 733.1-1 811, S-8 032, t-l 811. 2-2 822. H

80 37

86 031

48 40

40 11)

00 181

03 36}

Edge A ai A 312. etc. 17 201 27 16 88 331 87 01{ 40 S» 48 201

03 17| 4fl 2S{

Angle on a (100, *-.) 14 3)

811 A ail, etc 27 16 80 43J 85 Of 40 40 48 111

23 S71 16 161 12 4of

EdgeB 31 A 32l. et 50 28| se 0«| 81 80} 36 811 23 00|

55 00}

50 28)

48 301

S3 831

19 45

80 07(83 411 86 S31

Angle on a (100, U) GO 14} 48 11 47 71 46 801 46 71 45 52

Angle on a (100. i-i) 19 28} 28 0 25 14} 39 3»} 30 15(48 18(

Angle on *(U1,1)

40 83} 48 6}

41 28 80 14 87 37 36 481 86 4} SO 461

Angle on o (111, 1) 10 11 10 47^ 19 28} 83 0 23 601 25 141

Angle on o(111.1> 86 16} S3 44 30 39} 36 141 19 38} tl 86}

Hbxootahbdrons.

Cf. Fig. 102. 431. 4-3 18-10-5. V-l 18-10-1,18-1 881,0-1 SSI, 8-1 488, »4 481, 4-f

Edge A 821 A 312, etc. J 17 45} 19 121 85 67} 27 SO} 81 47} 16 0}

Edge B

20 121

37 17}

0 83}

IB 87}

.81 0}

48 86}

33 87}

I, etc. 321 A 331, etc.

85 57

80 06

81 61} 27 801 31 47} 10 0} 16 601

Angle on o (100. i-i) 39 13}

81 50} 39 101

82 18} 86 42 43 1} 88 10}

o(111.1)

28 61

35 57}

35 411

28 83}

23 12}

15 13}

' A fuller list is given in the Introduction to Dana's System of Mineralogy, pp. 28-32.
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ISOTROPIC SUBSTANCE.

8. PYRITOHEDRAL GROUP (2). PYRITE TYPE.

66. Typical Form and Symmetry. The typical forms of the pyritohedral group are the pyritohedron, or pentagonal dodecahedron, Figs. 113, 113, and the diploid, or dyakis-dodecahedron, Fig. 118. The symmetry of these forms, as of the group as a whole, is as follows: There are but three

111. planes of symmetry; these are parallel to the

cubic faces and coincide with the planes of the cubic axes. The three crystallographic axes are axes of binary symmetry only; there are also four axes of trigonal symmetry coinciding with the octahedral axes.

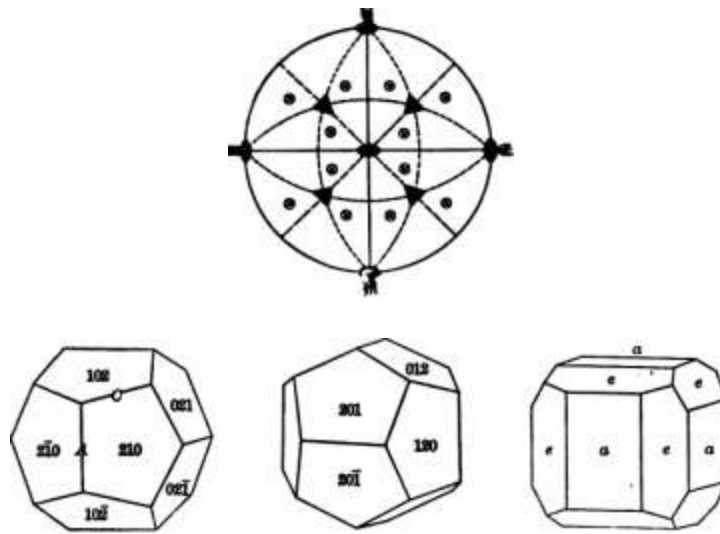
The spherical projection in Fig. 111 shows the distribution of the faces of the general form (hkl) and thus exhibits the symmetry of the group. This should be carefully compared with the corresponding projection (Fig. 69) for the normal group, so that the lower grade

of symmetry here present be thoroughly understood. In studying the forms described and illustrated in the following pages, this matter of symmetry, especially in relation to that of the normal group, should be continually before the mind.

It will be observed that the faces of both the pyritohedron (Fig. 113) and the diploid (Fig. 118) are arranged in parallel pairs, and on this account these forms have been sometimes called parallel hemihedron. Further, those authors who prefer to describe these forms as cases of hemibedrism call this type parallel-faced hemihedrism or pentagonal hemihedrism.

67. Pyritohedron.—The pyritohedron (Fig. 112) is so named because it is a typical form with the common species, pyrite. It is a solid bounded by twelve faces, each of which is a pentagon, but with one edge {A, Fig. 112) longer than the other four similar edges (C). It is often called a pentagonal dodecahedron, and indeed it resembles closely the regular dodecahedron of geometry, in which the faces are regular pentagons. This latter form is, however, as already noted (Art. Sti), an impossible form in crystallography.

Fig. 113. 114.



The general symbol is ka (kkO) or like that of the tetrahedron of the normal group. Hence each face is parallel to one of the axes and meets the other two axes at unequal distances. Common forms are (410) , (310) , (210) , (320) , etc. Besides the pyritohedron, as (210) , there is also the comple-

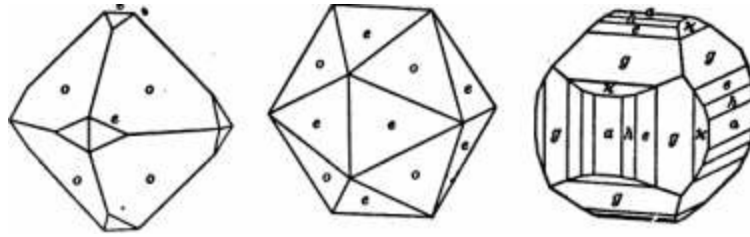
ct) by $GoOgIc$

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mentary minus form* shown in Fig. 113; the symbol is here (ISO). Other forms are (250) , (230) , (130) , etc.

The pyritohedron and the diploid together embrace twenty-four faces, having the same position as the twenty-four like faces of the tetrahedron of the normal group.

68. CombiSAtioat.— The faces of the pyritohedron replace the edges of the

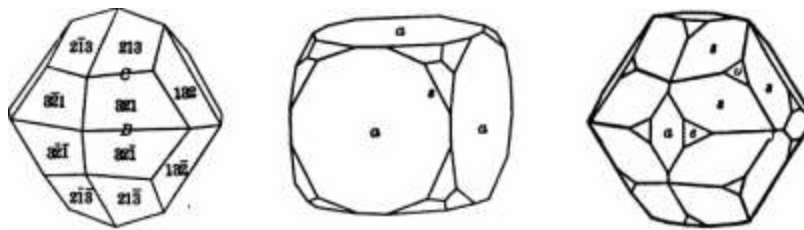


cube, but make unequal angles with two adjacent cubic faces; on the other hand, when the pyritohedron is modified by the cube, its faces truncate the longer edges of the pentagons. Cf. Fig. 114.

Fig. 115 shows ^ne combination of the pyritohedron and octahedron, and in Fig. 116 these two forms are eqjially developed. The resulting combin»tion bears a close similarity to the icosahedron, or regular twenty-facc^d solid, of geometry (see Art. 35). Here, however, of the twenty faces, the eigiit octahedral are equilateral triangles, the twelve others belonging to the pyritohedron are isosceles triaugles. Fig. 117 shows a nnumber of pyritohedroTia with the cube (a), namely, A (410), e (310), ff (320), and the minns form K (4.50).

69. Diploid. —The diploid is bounded by twenty-four similar faces, each meeting the axes at unequal distances; its general symbol is hence (M/), and common forms are (321), (421), etc. The form (321) is shown in Fig. 118;

118. 119. . 120

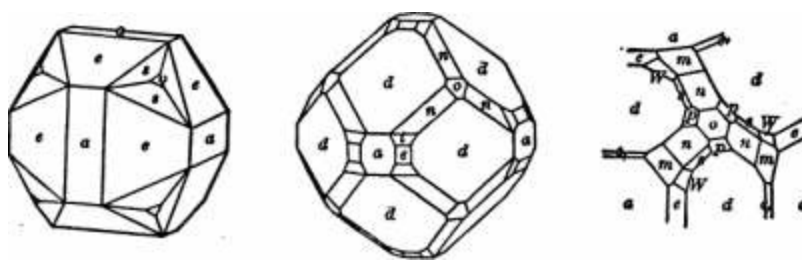


the symbols of its faces, aa given, should be carefully studied. As seen in the figure, the faces are quadrilaterals or trapeziums; moreover, they are grou]ied in pairs, hence the common name diploid. It is also called a dyakis-dodecahedron.

;,CoOgIc

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The complementary minus form bears to Fig. 118 the same relation as the minus to the plus pyritohedron. Its faces have the symbols 313, 231,123, in the front octant, and similarly with the proper negative signs iti the others. Tie pliia and minus forms together obviously embrace ail the faces of the lieioctahedroD of the normal group.



Pyrite.

Pyrite.

Pyrite.

In Fig. 119 the plus diploid is shown in combination with the cube. Here the three faces replace each of its solid angles. This combination forms resembles that of Fig. 89, but the three faces are here unequally inclined upon two adjacent cubic faces. Other combinations of the diploid with the cube, octahedron, and pyritohedron are given in Figs. 120 and 131.

70. Other Pyrite.—If the pyritohedral type of symmetry be applied to planes each parallel to two of the axes, it is seen that this symmetry calls for six of these, and the resulting form is obviously a cube. This cube cannot be

geometrically distinguished from the cube of the normal

group, but it has its own characteristic molecular sym-

metry. Corresponding to this it is common to find cubes of pyrite with fine lines (striations) parallel to the alternate edges, as indicated in Fig. 124. These are due to the partial development of pyritohedral faces (210). On a normal cube such striations, if present, must be parallel to both sets of edges on each cubic face. Similarly to the cube, the remaining forms of this pyritohedral group, namely, (111), (112), (113), (114), have the same geometrical form, respectively, as the octahedron, dodecahedron, the trisoctahedrons and tritrapezohedrons of the normal group. In molecular structure, however, these forms are distinct, each having the symmetry described in Art 66.

71. Other combinations of pyritohedral forms are shown in Figs. 132, 123, both of the species pyrite. Fig. 123 is dodecahedral in habit, with the diploid $\{ \llcorner \}$, the trapezohedron \gg (211), also a (100), o (111), e (310). In Fig. 133, a single angle of a pyrite crystal is represented with a (100), o (111), d (110); the two pyritohedrons e (210) and e, (120); the trisoctahedron y (221); the trapezohedrons \ll (311). H (311); the diploids a (321) ff (851).

This species illustrates well the complexity that may be observed among the crystals of a given mineral. Not only is there wide variation in habit, but the occurring forms are also very numerous. Thus some thirty-five pyritohedrons (— and —) have been noted and a like number of diploids; also five trisoctahedrons and eleven trapezohedrons.

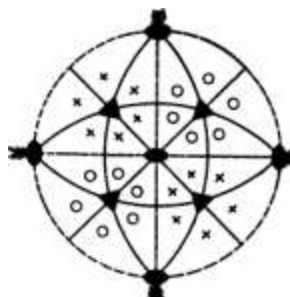
72. AaglM. —The following tables contain the angles of some common forms:

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3. TBTRAHEDRAL GROUP (3). TETRAHEDRITE TYPE.

73. Typical Forms and Symmetry.—The typical form of this group, and that from which it derives its name, is the tetrahedron, shown in Figs. 126, 137. There are also three other distinct forms, shown in Figs. 133,134, 1.S5.

The symmetry of these forms is that which is characteristic of the entire



X planes of symmetry, parallel faces of a rhomhic dodeca-

136. group. There are si

respectively to the

hedron, but no planes of symmetry parallel to the cubic faces. The three cubic axes are axes of binary symmetry only, and the four octahedral axes are axes of trigonal symmetry. There is no center of symmetry. f The spherical projection (Fig. 1251 showR the distribution of the faces of the genei'al form (hkl) and thus exhibits the symmetry of the group. It will be Been at once that the like faces are all grouped in the nlternale octants, and this will be seen to be characteristic of all the forms peculiar to this group. The relation between

the symmetry here described and that of the normal group must be cnrefnly studied.

In distinction from the pyritohedral forms whose faces were in parallel

pairs, the faces of the tetranedron and the analogous solids are inclined to>

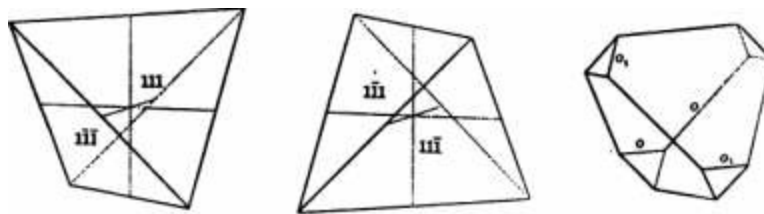
each other, and hence they are sometimes Bjmken of as inclined hemiAedron», and the type of 8o-called bemihedriem here illustrated is thea called inclined or tetrahedr&l hemihedriam.

74. Tetrahedron.—The tetrahedron,* as its name indicates, is a four^foed solid, bounded by planes meeting the axes at equal distances. Its general symbol is (111), and the_ four faces of the plus form (Fig. 126) have the symbols 111, ill. 111, HI. Theae are four of the faces of the octahedron of the normal group (Fig. 71), and those four which belong to the alternate octants as required by the symmetry already defined.

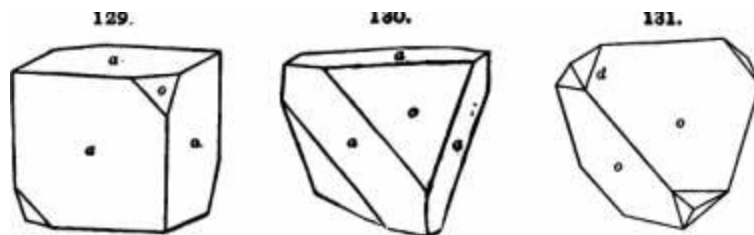
Each of the four faces of the tetrahedron is an equilateral triangle; the (normal) interfaci^ angle is $109^{\circ} 29' 16''$. The tetrahedron is the regular triangular pyramid of geometry, but crystallographically it must be so placed that the axes join the middle points of opposite edges, and one axis is Tertical

IM.

138.



There are two possible tetrahedrons: thepZu« tetrahedron (111), designated by the letter o, which has already been described, and the minvs tetrahedron, having the same_ geometrical form and symmetry, but the symbols of its four faces jire 111, 111, 111, 111. This second form is shown in Fig. 127; it is nsually designated by the letter o,. 'l.hese two forms are, as stated above, identical in geometrical shape, but they may be distingaished in many cases by the tests which serve to reveal the molecular structure, particularly the etch in, ^-figures. It is probable that the plus and minus tetrahedrons of



sphalerite (see that species) have a constant difference in this particular, which makes it possible to distinguish them on crystals from different localities and of different habit.

If both tetrahedrons are present together, the form in Fig. 138 results. This is geometrically an octanedron when they are equally developed, ■ but

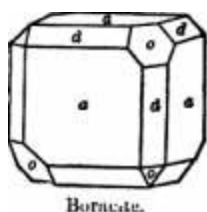
TlilB is nae of Out five regulnr Biiliils of geometry, which include also (lie cube, octa- bedro'. the regular pentagoiial dodecahedron, and the fcoBshedroD; tte last two are Im- powible forma among CTTStals.

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crystallogmphically it is always only a combination of two anlike forms, the plus and minus tetrahedrons, which can be diBtinguished as already noted.

The tetrahedron in combination with the cube replaces the alternate solid angles as in Fig. 129. The cube modifying the tetrahedron truncates itg edges as shown in Fig. lilO. The normal angle between adjacent cubic and tetrahedral faces is $54^{\circ} 44'$. In Fig. 131 the dodecahedron is shown modifying the pins tetrahedron, while in Fig. I'd'i the cube ia the predominating form with the plus and minus tetrahedrons and dodecahedron.

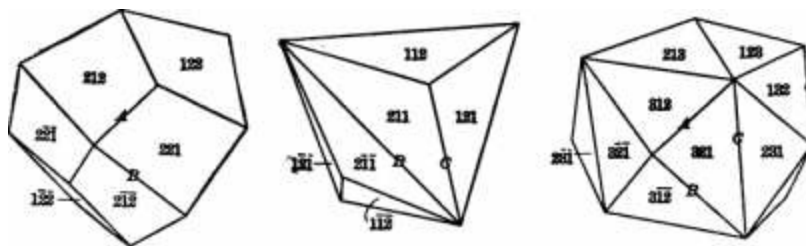
76. Other Typical Forms. —There are three other dis-} tinct types of solids iu this gronp, having the general J-/ 3jmbola{hkl),{hll),&ud{hi:I). Thefirst of these is shown ■^^^ in Fig. 133, here the symbol is (221). There are twelve faces, each » quadrilateral, belonging to this form, distributed as determined by the tetrahedral type of symmetry. They correspond to twelve of the faces of the trisoctahedron, namely, all those falling inal- ternate octants, This type of solid is sometimes called a tetragonai tristetra- hedro'n, or a deltoid dodecahedron. It does not occur alone among crystals, but its faces are observed modifying other forms.



133.

134.

136.



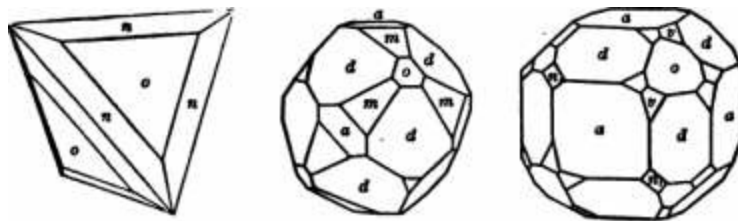
There is also a complementary minuti form, corresponding to the plus form, related to it in precisely the same way as the minus to the plus tetrahedron. Its tVFelvé faces are those of the trisoctahedron which belong to the other set of. alternate octants.

Another form, shown in Fig. 134, has the general symbol $\{hll\}$, here (211) ; it is bounded by twelve like triangular faces, distributed after the type demanded by tetrahedral symmetry, and corresponding consequently to the faces of the alternate octants of the form $\{hlj\}$ —the tritrihedron—of the normal group. This type of solid is sometimes called a trigonal tritetrahedron or trigondodecahedron.* It is observed both alone and in combination; it is much more common than the form $\{!ihl\}$. There is here again a complementary minus form. Fig. 136 shows the plus form $h(2U)$ with the plus tetrahedron, and Fig. 137, the form $m(311)$ with $a(100)$, $o(111)$, and $d(110)$. In Fig. 138, the minus form $w(211)$ is present.

*It is to be noted that the tritrihedron has faces which resemble those of the tritrihedron (Cubic trisoctahedron), although it is related not to the tritrihedron (Cubic trisoctahedron). On the other hand, the faces of the trigonal tritetrahedron resemble those of the tritrihedron. It is in fact related to the trapezo-

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The fourth independent type of solids in this $\{hxl\}$ is shown in Fig. 135. It has the general symbol $\{hxl\}$, here $(33L)$, and is bounded by twenty-four faces distributed according to tetrahedral symmetry, that is, embracing all the faces of the alternate octants of the forty-eight-faced hexoctahedron. This form is sometimes called a hexakistetrahedron. The complementary minus form $\{hkl\}$ embraces the remaining faces of the hexoctahedron. The plus, 136. 137. 138.



Sphalerite.

hexakistetrahedron, $i(531)$, is shown in Fig. 138 with the cube, octahedron, and dodecahedron, also the minus trigonal tritetrahedron $n(\text{II})$.

76. If the tetrahedral symmetry be applied in the case of planes each parallel to the two axes, it will be seen that there must be six such faces. They form a cube similar in geometrical appearance to the cube both of the normal and pyritohedral groups, but differing in its molecular structure, as can be readily proved, for example, by pyroelectricity. Similarly in the case of the planes having the symbol (110) , there must be twelve faces forming a rhombic dodecahedron bearing the same relation to the like geometrical form of the normal group. The same is true again of the planes having the position expressed by the general symbol $\{hk0\}$; there must be twenty-four of them and they together form a tetrahexahedron.

In this group, therefore, there are also seven types of forms, but only four of them are geometrically distinct from the corresponding forms of the normal group.

77. Angles.—The following tables contain the angles of some common forms:

TKT&AOKAI, TXIBTKTI&AHKDBONS.

iCoogIc

o BTBTALLOOaAPHr.

Hbxakibtetba bed bods.

Jl^geA EdgeB Edge C

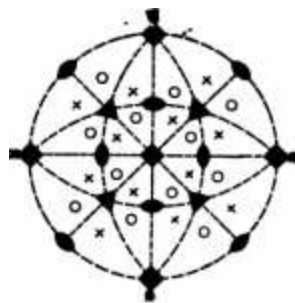
Cr. Fig. 185. 821 A 812, etc 8S1 A 8i2, etc. 891 a 281, etc

6ai 27' 89f' 6r 7i' 27" 891'

dil 21 47i e» 41 21 47i

482 10 si 83 4i 10 o)

481 82 l4 67 22f IS oo|



4. PLAGIHEDBAL GROUP (4). CUPRITE TYPE.

78. Typical Form and Symmetry.— The fourth group under the Uometric system is called the plagihedral or gyroidal 139.

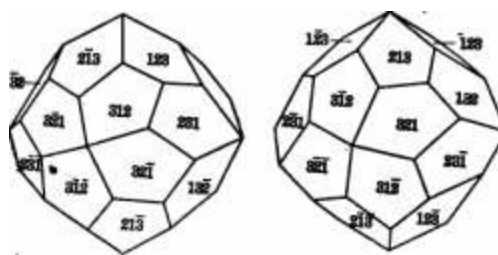
group because the faces of the general form {hkl} are arranged in spiral order. This is shown on the spherical projection, Fig. 139, and also in Figs. 140,141, which represent the single typical form of the group. These two complementary solids together embrace all the faces of the heioctahedron. They are distinguished from one another by being called respectively right-handed and left-handed pentagonal icositetrahedrons. The other forms of the group are geometrically like those of the normal group.

The symmetry characteristic of the group in general is as follows :

There are no planes of symmetry and no center of symmetry. There are, however, three axes of quaternary symmetry normal to the cubic faces, four axes of trigonal symmetry normal to the octahedral faces, and six axes of binary symmetry normal to the faces of the dodecahedron. ^ In other words, it has all the s of symmetry of the normal group while without planes or center of symmetry.

79. It is to be noted that the two forms shown in Figs. 140,141 are alike geometrically, hut are not superpcaable; in other words, they are related to one another as is a right- to a left-hand glove. They are hence said to be enantiomorpjious, and, as explained elsewhere, the crystals belonging here may be expected to show circular light polarization. It will be seen thai the complementary plus and minus forms of the preceding groups, unlike those here, may be superposed by being rotated 90" about one of the crystallographic axea. This distinction between plus and minus forms, and between rignt- and left-handed enantiomorphous forme, exists also in the case of the groups of Beveral of the other systems.

This group is rare among minerals; it is represented by cuprite, sal

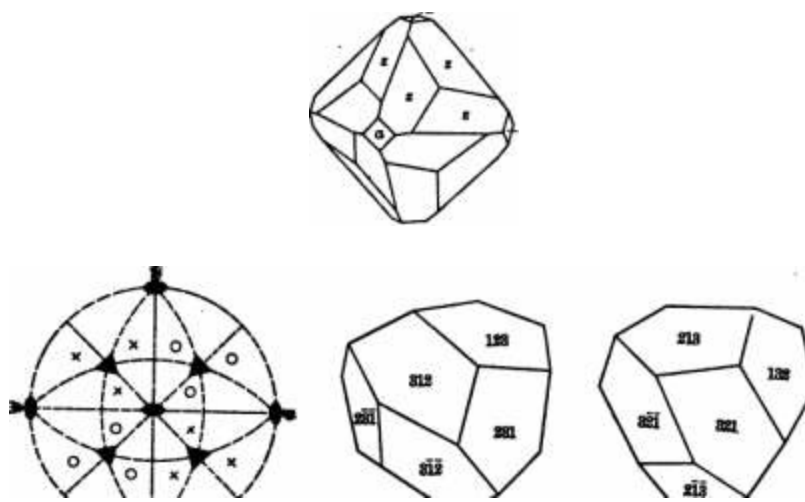


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ammoniac, sylvite, and halite. It is neually shown by the distribution of the snrall rnoUifyiDg faces, or by the form of the etchine 1«L

figures. Fig. 14a shows a crystal of cuprite from Cornwall (I'mtt) with the form a (13-10-13).

5. TETARTOHEDEAL GROUP (5). ULLMAN-NITE TYPE. SO. Symmetrr and Typical Forms.—The fifth remaining possible group under the isometric system is illustrated by Fig. 144, which represents the twelve-faced solid corresponding to the general symbol (/ikl). The dis-tribution of its faces is ahowm in the projection, Fig, 143. This form is sometimes called a tetrahedral-pentugonal dodecahedron. It is seen to have one-fourth as many faces as the form {kkl) in the normal group, hence there are four similar solids which together embrace all the faces of the hexoctahedron. These four solids, which are distinguished as right-handed (+ and —) and left-handed (-|- and —), are enantiomorphous, lilfe those of Figs. 140 and 141, and hence the suits crystallizing here may be expected to alio show circular polarization. The remaining forms of the group are

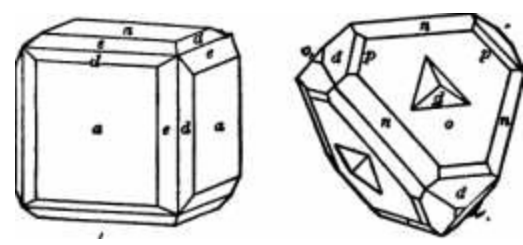


v^.

(besides the cube and rhombic dodecahedron) the tetrahedrons, the pyrito. hedrons, the tetragonal and trigonal tristetrahedron; geometrically they are like the solids of the

same names already described. This group has plane of symmetry and / no center of symmetry. There are three axes of binary symmetry normal to the cubic faces, and four axes of trigonal symmetry normal to the faces of the tetrahedron.

This group is illustrated by artificial crystals of barium nitrate, strontium nitrate, sodium chlorate, etc. Further, the species sillimanite, which shows sometimes pyritohedral (Fig. 146) and again tetrahedral form (Fig. 147), both having the same composition, must be regarded as belonging here.



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It is not difficult to see that the problems arising in the isometric system can be solved at once by the trigonometric relations in the sphere of projection (Fig. 110, p. 41; without the use of any formulas.

It will be remembered that the angle between a cubic face, $rs 100$. and the adjacent face of a tetrahedron, $SIO, 310, 82Q$, etc. can be obtained at once, since the angle U

112 k

this angle is equal to $g, ^{-}$, or in general —.

$$\tan(AW A 100) = \hat{}$$

Since all the forms of a face symbol under different species have the same angles, the tables of angles already given are very useful.

These and similar angles may be calculated immediately from the sphere, or often more simply by the formulas given in the following article.

19. Formulas.—(1) The distance of the pole of any face $P\{l>t\}$ from the cubic faces is given by the following equations. Here P_a is the distance between $(\&U)$ and (100) ; P_b is the distance between $\{fikt\}$ and (010) ; and P_c that between (hkl) and (001) .

These equations admit of much simplification in the various special cases, for (AiO) , (Mi) , etc.:

*.IW ** in '*

$$r^2 = f^2(A' + i^* + i') (p^2 + j^2 + r^2)$$

(8) The calculation of the supplement interfacial or normal angles for the several forms may be accomplished as follows:

Triclinic.—The angles A and B are, as before, the supplements of the interfacial angles of the edges lettered as in Fig. 88.

For the tetragonal-triclinic (Fig. 185), \cos

Trigonal bipyramid (Fig. 82). B and C are the supplement angles of the edges u lettered in the figure.

For the trigonal bipyramid (Fig. 186), \cos Tetrahedron (Fig. 82).

$$A' + A'' = \sqrt{A' + f^2}$$

For the pyrohedron (Fig. 113). $\cos A = \sqrt{}$; $\cos C = \sqrt{}$.

Hexagonal (Fig. 102).

$$A' + A'' = \sqrt{A' + A^* - f^2}, p = 2hk + l'$$

$$\sqrt{h^2 + k^2 + i^2} = \sqrt{h^2 + k^2 + i^2 - A} + 4 + (* \blacksquare)$$

For the hexagonal bipyramid (Fig. 186). $\cos B = \sqrt{} \sqrt{}$.

11. TETRAGONAL SYSTEM.

83. The Tetragonal System includes all the forms which are referred to three rectangular axes of which the two lateral axes are equal to each other and the third, the vertical axis, is either shorter or longer. The lateral axes are designated by the letter a ; the vertical axis by c (see Fig. 149) The length of the vertical axis expresses properly the axial ratio of $a : c$, a being uniformly taken as equal to unity.

Seven groups are embraced in this system. Of these the normal group is common and important among minerals; two others have several representatives, and another a single one only. It may be noted that in four of the groups the vertical axis is an axis of tetragonal symmetry; in the remaining three it is an AXIS of binary symmetry only.

1. NORMAL GROUP (6). ZIRCON TYPE.

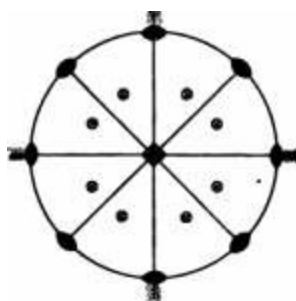
In addition to the normal group of the tetragonal principal plane of symmetry, the plane σ_h .

84. Symmetry.—The form system (of Figs. 149 to 171) has of the lateral axes a, a ; further, at right angles to this, and meeting each other at angles of 45° in the vertical axis, c , two pairs of planes of symmetry, like two-and-two. One of these sets, the axial planes, pass through the crystallographic axes, a, a, c , and are hence parallel to the faces lettered a ; the others are diagonal to them, or parallel to the faces m .

Further, the vertical axis, c , is a principal axis of tetragonal symmetry; there are also four axes of binary symmetry, like two-and-two; one set coincides with the lateral axes a, a ; the others are diagonal to them.

The distribution of the faces of the general form, hkl , belonging to this group, is shown in the spherical projection. Fig. 148.

86. Forms.—The various possible forms under the normal group of this system are as follows:



..(100) ..(110)

1. Base or basal pinacoid

2. Diametral prism, or prism)

of the second order \

3. Unit prism, or prism)

of the first order j" - ■ ■ ■

4. Ditetragonal prism (hkO)

as. (310) /-3; (210) t.2; (320) i-J, etc.

5. Pyramids of the diametral) ,,,,,,

or second order \ ' '

as, (203) l-i; e (101) It"; (201) 2-i, etc.

,ab,GoOgIc

CBYeTALLOQRAPHT.

Hlller. Naiim-nii

6. Pyramids of the cnit,) ,;^j» a:a:mc mP o. m

or first order, [^ ' '

as, (223) I; (111) 1; (221) 2, etc.

7. Ditetragonal pyramids, 1 ,j^ ^ ,, a-.na'.mc mPn or m-n

or Zirconoids,) * ' '

aa, (421) 4-3; (331) S-J; (133) 1-3, etc.

66. Basal Pinaoid or Base.—The hose is that form which includes the two similar faces which are parallel to the plane of the lateral aiea. These faces have the eymbola 001 and 001 respectiTely ; it is an " open form," as they do not inclose a space, consequently this form can occur only in combiniition with other forms. Cf. Figs. 149-152, etc. This form ia iilwajfe lettered c in this work.

87. Prinna.—Prisms, in systems other than the isometric, have been defined to be forma

whose faces are parallel to the vertical axis {S} of the crystal, while they meet the two lateral axes ; in this system the four-faced form whose planes are parallel both to the vertical and a lateral axis is also called a prism. There are hence three types of prisma here included,

88 Diametral Prism,—The diametral prism shown * in combination with the base in Fig. 149 includes the four faces which are parallel at once to the vertical and to a lateral axis; it has, therefore, the general symbol (100). It is a square prism, that is, the angle between any two adjacent faces is 90° .

The diametral prism is often called the prism of the second order; it is uniformly designated by the letter a, and its faces, taken in order, have the symbols 100, 100, 100, 100.

It will be seen that the combination of this form with the base is the analogue of the cube of the isometric system. It has four similar vertical edges and eight similar lateral edges. It has also eight similar solid angles. 149. 100. 100. 100.

$4 = \text{---}$

89. Unit Prism.—The unit prism includes the four faces which, while parallel to the vertical axis, meet the lateral axes at equal distances; its

* 111 Figs. 149-152 the diametral prism of the form are made to combine with the base in Fig. 156. It must be noted, however, that in the case of actual crystals of these forms, while the tetragonal system is usually indicated by the unlike physical character of the faces as compared with the face a, m. etc., in the vertical prismatic zone, no inference can be drawn as to the relative length of the vertical axis. This must be determined only when a pyramid is present: It is fixed for the species when a particular pyramid is fundamental or unit form, as explained later.

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general symbol is consequently (110). Like the preceding form, it is a square prism, with interfacial angles of 90° . It is shown in combination with the base in Fig. 150. It is often called the prism of the first order, and is uniformly designated by the letter v. The symbols of its faces, taken in order, are 110, 110, 110, 110.

The faces of the unit prism truncate the edges of the diametral prism and vice versa. When both are equally developed, as in Fig. 151, the result is a regular eight-sided prism, which, however, it must be remembered, is a combination of two distinct forms.

It is evident that the two prisms described do not differ geometrically from one another, and furthermore, in a given case, the symmetry of this group allows either to be made the

unit, and the other the diametral, prism according to the position assumed for the lateral axes. If on crystals of a given species both forms occur together equally developed (or, on the other hand, separately on different crystals) and without other faces than the base, there is no means of telling them apart unless by minor characteristics, as striations or other marking on the surface, etchings, etc.

90. Ditetragonal Prism.—The ditetragonal prism is the form which is bounded by eight similar faces, each one of which is parallel to the vertical axis while meeting the two lateral axes at unequal distances. It has the general symbol $H\{kO\}$. It is shown in Fig. 152, where $(hkO) = (210)$. The successive faces have the symbols $310, 120, 120, 510, 310, 130, 120, 210$.

In Fig. 164 a combination is shown of this form ($y = 310$) with the diametral prism, the edges of which it bevels. In Fig. 168 it bevels $\{h = 310\}$ the edges of the unit prism m . In Fig. 169 it is combined ($\{ = 310$) with both the square prisms.

91. Pyramids.—There are three types of pyramids in this group, corresponding, respectively, to the three prisms which have just been described. As already stated, the name pyramid is given (in systems other than the isometric) to a form whose planes meet all three of the axes; in this system the form whose planes meet the axis i and one lateral axis while parallel to the other is also a pyramid. The pyramids of this group are strictly double pyramids.

92. Diametral Pyramid.—The diametral pyramid, or pyramid of the second order, is the form. Fig 1b'A, whose faces are parallel to one of the lateral axes, while meeting the other two axes. The general symbol is $\{hOPf\}$. These faces replace the basal edges of the diametral prism (Fig. 154), and the solid angles of the unit prism (cf. Fig. 155). It is a square pyramid (also called a square octahedron), since its basal section is a square, and the inter-facial angles over the four terminal edges, above and below, are equal. The successive faces of the form (101) are as follows: Above 101, Oil, 101, Oil; below 101, Oil, 101, Oil.

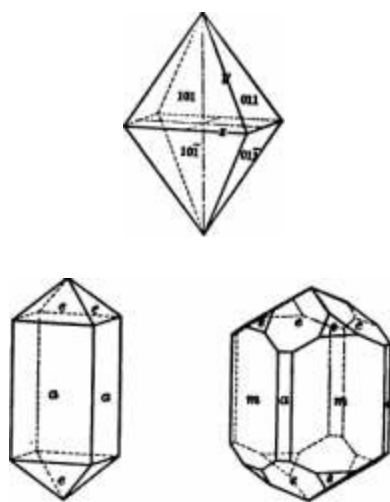
If the ratio of the intercepts on the lateral and vertical axes is the assumed axial ratio of the species, the symbol is (101), and the form is designated by the letter e. This ratio can be deduced from the measurement of either one of the interfacial angles [y or z) over the terminal or basal edges, as explained later. In the case of a given species, a number of diametral pyramids may occur, varying in the ratio of the axes n and i. Hence there is possible an indefinite number of such forms whose symbols may be, for example, (104), (103), (102), (101), (302), (201), (301), etc. Those mentioned first come nearest to the base (001), those last to the diametral prism (100); the base is therefore the limit of these pyramids (hOl) when $A = 0$, and the diametral

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prism (100) when $\alpha = 00$ and $I = 1$; or, what is the same thing, when $h = I$ and $I = 0$. Fig. 165 shows the three diametral pyramids u (105), e (101), q (30L).

153. 164. 166.



Ruttle. 93. Unit Pyramid.—A unit pyramid, or pyramid of the first order, is a form whose eight similar faces intersect the two lateral axes at equal distances and also intersect the vertical axis. It has the general symbol (hhl). Like the diametral pyramid, it is a square pyramid (or square octahedron) with equal interfacial angles over the terminal edges, and the faces replace the lateral, or basal, edges of the unit prism. If the ratio of the vertical to

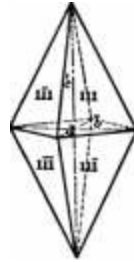
the lateral axis for a given unit pyramid is the assumed axial ratio for the species, the form is called the $\frac{w}{h}$ form, and it has the symbol (111) as in Fig. 156. Its faces mentioned in order as before are: Above $111, 111, 111, 111$; below $111, 111, 111, 111$.

166.

167.

168.

199.



$(4^{\wedge} / ^{\wedge}$

$^{\wedge} ^{\wedge}$

zircon. Zircon. Apophyllite

Obviously the angles of the unit pyramid, and hence its geometrical aspect, vary widely with the length of the vertical axis. For Fig. 166 (octahedrite) $i = 1.78$; for Fig. 161 $p = (111)$ and $(f = 0.64)$.

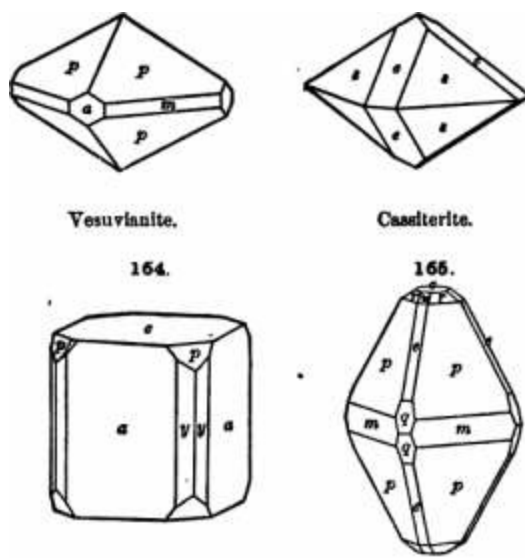
For a given species there may be a number of unit pyramids, varying in position according to the ratio of the vertical to the lateral axis. Their symbols, passing from the base (001) to the unit prism (110) , may thus be $(115), (113), (323), (U), (332), (291), (441)$, etc. In the general symbol of these forms $iAhJ$, as h diminishes, the form approximates more and more nearly to the base (001) , for which $A = 0$; as t increases, the form passes toward the unit prism, for which $h = \frac{1}{2}$ if $Z = 1$, that is, for which $A = 1$ if $Z = 0$. In Fig. 158 two pyramids of this order are shown, p (HI) and \ll (331).

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The faces of the unit pyramids replace the terminal edges of the unit prism (Figs. 157, 160) and the solid angles of the dihexagonal prism (Fig. 159).

$^{\wedge} ^{\wedge}$ isir::?'

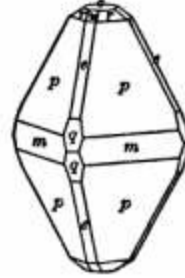
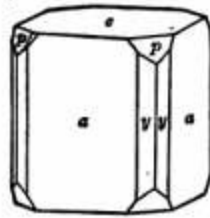


Vesuvianite.

Cassiterite.

164.

165.



Apopbytile,

Oclabedrite.

The application of the zonal relations proves that a dihemitral pyramid truncating the pyramidal edges of a given unit pyramid has the same ratio as it has for $\frac{1}{2}$ to 1. Thus (101) truncates the terminal edge of (111); (201) of (111), etc. Again, if a unit pyramid truncates the pyramidal edges of a given dihemitral pyramid, its ratio for $\frac{1}{2}$ to 1 is half that of the other form; that is, (101) truncates the pyramidal edges of (111); (201) of (111), etc. These relations are exhibited by Fig. 165, and the basal and spherical projections (Figs. 170, 171) corresponding to it. Here e (101) and y (105) truncate the terminal edges of p (111) and r (115), respectively, while p (111) truncates the edges of y (201).

94. Ditetragonal Pyramid, or Zirconoid.—The ditetragonal pyramid, or double eight-sided pyramid, is the form each of whose sixteen similar faces meet the three axes at unequal distances. This is the most general case of the symbol (Jikl), where A, k, l are all unequal and no one is equal to 0. That there are sixteen faces in a single form is evident. Thus, for example, for the form (213) the face 212 is similar to 122, the two lateral axes being equal (not, however, to 231). Hence there are two like faces in each octant. Similarly the symbols of all the faces in the successive octants are therefore, as follows:

Above 212 122 122 512 252 152 122 218 Below 213 125 132 513 132 153 318

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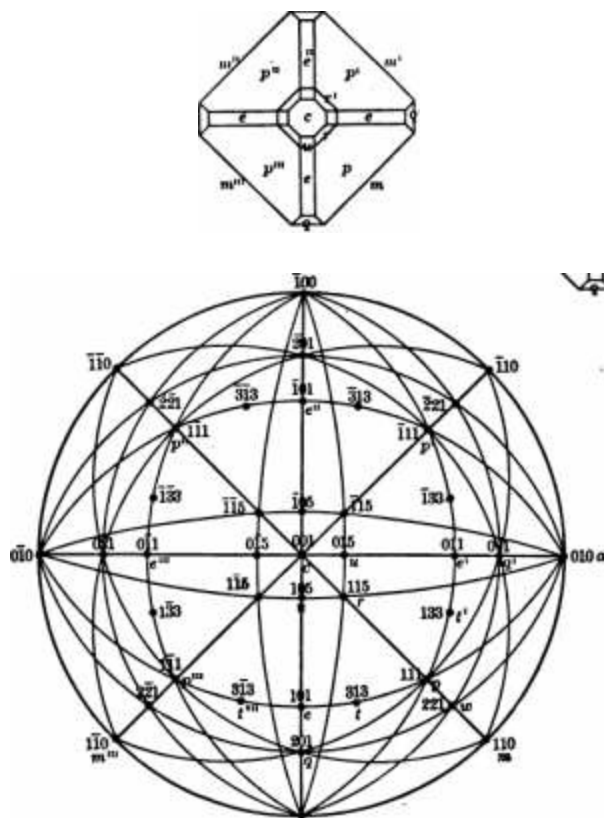
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This form is common with the species zircon, and is hence often called a zirconoid. It is shown in Fig. 166. It is not observed alone, though sometimes, as in Figs 167 (x = 311) and 168 (z = 321), it is the predominating form. In Fig. 169 two zirconoids occur, namely, I (313) and z (3'11). See also Fig. 6, 8, 9 of zircon on p. 8.

ZircoD. CiiMllerilc. Ruttle

96. In addition to the perspective figures already given, a basal projection (Fig 170) is added of the crystal of octahedrite $j^{\wedge}q$.

already referred to (Fig. 165); also a spherical projection of the same (Fig. 171) with the faces of the form (313) added. These exhibit well the general relations of this normal group of the tetragonal system. The symmetry here is to be noted, first, with respect to the similar zones $100, 001, 100$ and $010, 001, 010$; also, second, that of the other pair of similar zones, $110, 001, 110$, and $110, 001, 110$.



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2. HEMIMORPHIC GROUP (7).

96. Bynunetiy, —This group differs from the normal group only in having 172, no plane of symmetry through the plane of the

transverse axes; hence the forms are hemimorphic as defined in Art. 29. It is not known to be represented among minerals, and is sufficiently illustrated by the spherical projection (Fig. 172). Here the two basal planes are distinct forms, 001 and $00\bar{1}$; the prisms do not differ geometrically from those of the normal group, though distinguished by their molecular structure; further, the pyramids are no longer double

pyramids, but each form is represented by one half of FigB. 163, 156, 166 (cf Fig. 50, p 18). There are hence six distinct pyramidal forms, corresponding to B upper and lower halves of the unit pyramid, the diagonal pyramid and the ditetragonal pyramid.

3. PYRAMIDAL GROUP (8). SCHEELITE TYPE.

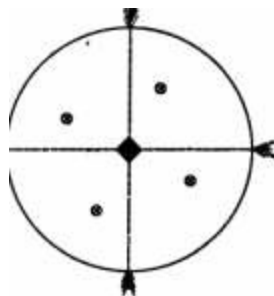
97. Typical Forms and Symmetry. —The forms here included have one plane of symmetry only, that of the transverse axes, and one axis of tetragonal symmetry (the vertical axis) normal to it. The distinct forms are the tetragonal prism (hkG) and pyramid (hkl) of the third order, shown in Figs. 174, 175,

The distribution of the faces of the general form (hkl) on the spherical projection. Fig. 173, exhibits the symmetry of the group. Comparing this, as well as the figures immediately following, with those of the normal group, it is seen that this group differs from it in the absence of the vertical planes of symmetry and the horizontal axes of symmetry. Further, half the faces, belonging to each octant, of the normal form (like) shown in Fig. 166 only are present, and these are the faces situated in a vertical zone, from 001 to 00l.

98. Prism and Pyramid of the Third Order.—The typical forms of the group, as above stated, are a square prism and a square pyramid, which are distinguished respectively from the square prisms a (100) and m (110), shown in Figs. 141) and 150, and from the square pyramids [M] and (hhl) of Figs. 153 and 156 by the name "third order."

There are two complementary forms in each case, designated left and right, which together include all the faces of the ditetragonal prism (Fig. 152) and ditetragonal pyramid (Fig. 166) of the normal group

The faces of the two complementary prisms, as (210), are: Left: 210, l30, 3i0, 150. Right: 120, 510, l30, 2l0.



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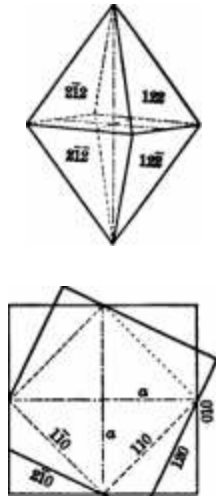
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The faces of the corresponding pyramids, as (SIS), are:

Left: Right.

above 313, l33, 3l3, iSS; below 2l2, l22, 313, l33. above ISS, 3lS, IS2, 2lS; below 1S3, 3l3, 133, 2l3.

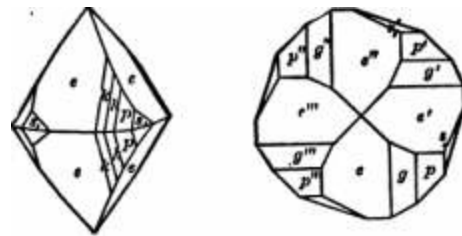
Fig. 176 gives a transverse section of the prism (100) and m (110), also the prism of the third series (ISO). Fig. 175 shows the right pyramid (1S3) corresponding to the same prism.



99. Other Forms.—The other forms of this group, that is, the base c (001); the other square prisms, a (100) and m (110): also the square pyramids (AO/)

and (hhl) are geometrically like the corresponding forms of the normal group already described.

100. To this group belong the important species scheelite; also the isomorphous species stolzite and powellite, unless it be that they are rather to be classed with wulfenite (p. 61). Fig. 177 shows a typical crystal of scheelite, and Fig. 178 a basal section of one similar; these illustrate well the charac-



^^

S'ySLS I'P^i (? .3,, ^v^, .. m. Ifig.

Here the forms are e (101), p (111), and the third-h (313), g (212), s, (131). Fig. 179 represents a meionite crystal with r (111) and the third-order pyramid z (311) See also Pies. 181, 182, in which the third-order pnam is shown.

The forms of this group are sometimes described (see Art. 28) a pyramidal hemitokedrim ; hence the name here given.

a showing

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-TSTRAOOMAL STSTEU.

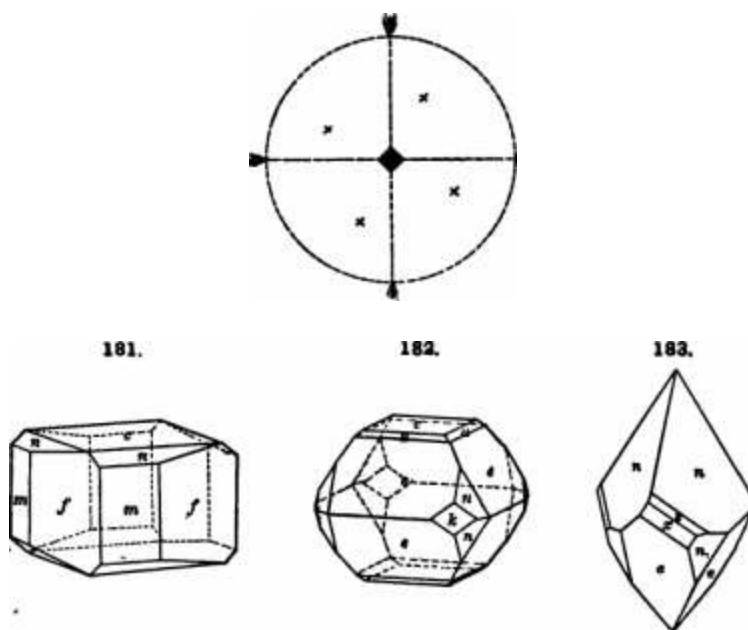
i. PYRAMIDAL-HEMIMORPHIC GROUP (9). WULFENITE TYPE.

101. Symmetry. —The fourth group of the tetragonal System is closely related to the group just described. It has the

same vertical axis of tetragonal symmetry, but there is no transverse plane of symmetry. The forms are, therefore, hemimorphic in the distribution of the faces (cf. Fig. 180). The species wulfenite of the Scheelite Group among mineral species probably belongs here, although the crystals

do not always show the difference between the pyramidal faces, above and below, which would characterize distinct complementary forms. Figs. 181, 183 could, therefore, serve as illustrations of the preceding group, but in Fig. 183 a characteristic distinction is exhibited. In these figures the forms are u (102), e (101), n (111); also v (230), A (210), t (432), x (311).

183.



Figs. 181-188. Wulfenite. 5. SPHENOIDAL GROUP (10). CHALCOPYRITE TYPE.

The typical forms of this group are the sphenoid and the tetragonal scalenohedron.

102. Typical Form and Symmetry.

the sphenoid (Fig. 185) and the tetragonal scalenohedron (Fig. 186).

and all the combinations of this group are 184.

characterized by the presence of two vertical planes of symmetry; these are diagonal to the crystallographic axes and intersect at angles of 90° in the vertical axis, which is an axis of binary symmetry only. Further, the two horizontal

This symmetry is exhibited in the distribution of the faces of the general form (hkl) in the spherical projection (Fig. 184). It is seen here that the faces are present in the alternate octants only, and it will be remembered that this same statement was made of the tetrahedral group under the isometric system. There is hence a close analogy between these

two groups.

The symmetry of this group should be carefully compared with that of the first and third groups of this system already described.

103. Sphenoid.—The sphenoid, shown in Fig. 185, is a four-faced solid, resembling a tetrahedron, but each face is an isosceles (not an equilateral) triangle. The general symbol of the plus unit sphenoid is (111), and the faces have the symbols: 111, III, III, III. The complementary minus sphenoid has the symbol (III), and these two forms include all the faces of the unit pyramid (111) of the normal group. When they occur together, if equally developed, the resulting solid, though having two unlike sets of faces, cannot be distinguished geometrically from the square pyramid (111).

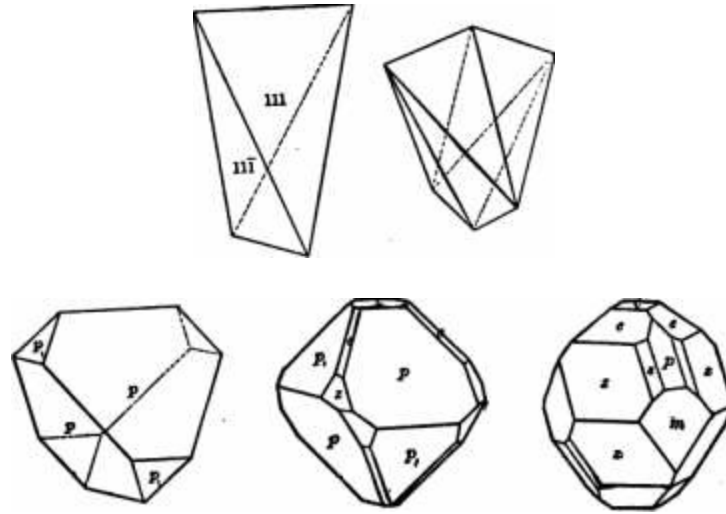
In the species chalcopyrite, which belongs to this group, the deviation in angle and axial ratio from the isometric system is very small, and hence the unit sphenoid cannot by the eye be distinguished from a tetrahedron (compare Fig. 187 with Fig. 128, p. 47). For this species $c/a = 0.85$ (instead of 1, as in the isometric system), and the normal sphenoidal angle is $108^\circ 40'$, instead of $109^\circ 28'$, the angle of the tetrahedron. Hence a crystal with both the plus and minus sphenoids equally developed closely resembles a regular octahedron.

In Fig. 188 the diametral pyramids e (101) and z (201) are also present, also the base c (001).

104. Tetragonal Scalenohedron.—The sphenoidal symmetry yields another distinct type of form, that shown in Fig. 186. It is bounded by eight similar scalene triangles, and hence is called a

tetragonal scalenohedron; the general symbol is $(Jikl)$. The faces of the complementary plus and minus forms embrace all the faces of the ditetragonal pyramid. This form appears in combination in chalcopyrite, but is not observed independently. In Fig. 189 the form s (531) is the plus tetragonal scalenohedron.

106, Other Forms.—The other forms of the group, namely, the two square prisms, the ditetragonal prism, and the two square pyramids $\{hhl\}$ and $\{hol\}$, are geometrically like those of the normal group. The lower symmetry in the molecular structure is only revealed by special investigation, as by etching.

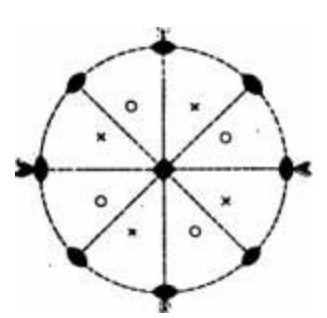


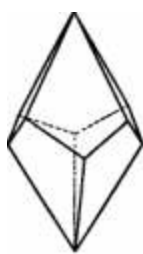
Figs. 187-189, Chalcopyrite.,

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6. TRAPEZOHEDRAL GROUP (11).

106. The trapezohedral group is analogous to the plagioclinal group under the triclinic system; it is characterized by the absence of any plane or center of symmetry: the vertical axis, however, is an axis of tetragonal symmetry, and perpendicular to this there are four axes of binary symmetry. The distribution of the faces of the general form $(/ikl)$ is shown in the spherical projection,





nder this system

Fig. 190, and Fig. 191 gives the resulting solid, a tetragonal Trapezohedron. The complementary right- and left-handed forms embrace all the faces of the ditetragonal pyramid of the normal group. These two forms are [^]antimorphous, and the salts belonging to this group show circular polarization. Phosgenite (p. 364) probably belongs to this group (Ooldschmidt).

7. TETAUTOHEDRAL GROUP (12).

107. Symmetry.—The seventh and last possible group has no plane nor center of symmetry, but the vertical axis is an axis of binary symmetry. The distribution of the faces of the general form

{*likl*} is shown on the sphere of projection (Fig. 192), and the solid resulting is a spherioid of the third order. There are also three other possible forms complementary to this*, and the four are respectively distinguished as right (—) and left (|— and —), These four together embrace all the sixteen faces of the ditetragonal pyramid. The other characteristic forms of this group are the prism of the third order *IhtOj*, the plus and minus sphenoids of the first order (111), and also those of the second order (101). This group has no known representative.

Mathematical Relation of

to TETRAHEDRAL SYMMETRY.

in. C₂h at Asm. —It differs from the (lacking of the symmetry of the few groups of this system with all of them the position of the vertical axis is fixed, in groups 1, 8, however, where there are two sets of vertical planes of symmetry. either set may be made the *h* planes and the other the *g* planes. The choice between these two possible positions of the lateral axes is decided principally by the habit of the occurring crystals and the relations of the given species to others of similar form.

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100. Definition of the *h* and *g* planes.—The following relations serve to connect the *h* and *g* planes. The angle of the vertical axis is, when $a = 1$, will be found to be

$\cos \theta = \frac{1}{\sqrt{3}}$ and $(001 \text{ and } 111)$:

$$\tan(\angle AHI) \times l^i = i-$$

For faces in the usual rectangular zone the tangent principle applies. The usual cases (cf. Fig. 171) are:

$$\tan(\angle AHO) \cdot \tan(\angle AHO) = i$$

$$\tan(\angle AHO) \cdot \tan(\angle AHO) = i$$

$$\tan(\angle AHO) = p \cdot \tan(\angle AHO) = p$$

110. Other Calculations.—It will be noted that in the spherical projection (Fig. 171) all those spherical triangles which are formed by great circles which meet the prismatic zone-circle 100, 010. Again, all those formed by great circles drawn between 100 and 010, and crossing respectively the zone-circles 100, 001, 010, or 010, 001, 010. Also, all those formed by great circles drawn between 110 and 010 and crossing the zone-circle: 110, 001, 110, or between 110 and 111) and crossing the zone-circle 001, 010.

These spherical triangles may hence be readily used to calculate any angles desired; for example, the angles between the faces of any crystal, such as hkl (say 111), and the prisms 100, 010, 001. The dihedral angles (r and f. Fig. 166) of the ditetragonal pyramid, 218 A 813 (or 313 A 813, etc.). and 213 A 122 (i.e. 813 and 133, etc.), can also be obtained in the same way. The zonal relations give the symbols of the poles on the zones 001, 100 and 001, 110 for the given case. For example, the zone-circle 110, 813, 133, 110 meets 010, 001, 110 at the pole 223, and the calculated angle: 313 and 223 is half the angle 813 and 133. If a large number of similar angles are to be calculated, it is more convenient to use a formula, as is given below.

111. Formulas.—It is sometimes convenient to have the normal interfacial angle* expressed directly in terms of the axis and the indices h, k, and l. Thus:

(1) The distance of the pole of any face P(hkl) from the poles a(100) = P_a, b(010) = P_b, c(001) = P_c are given by the following equation:

$\frac{1}{P^2} = \frac{1}{a^2} + \frac{1}{b^2} + \frac{1}{c^2}$ These may also be expressed in the form

$$\frac{1}{P^2} = \frac{1}{a^2} + \frac{1}{b^2} + \frac{1}{c^2}$$

(2) For the distance between the poles of any two faces (hkl), (pqr), we have $\cos PQ = \frac{hp + kq + lr}{\sqrt{h^2 + k^2 + l^2} \sqrt{p^2 + q^2 + r^2}}$

The above equations take simpler form for special cases often occurring; for example, for hkl and the angle at the edge y of Fig. 166.

119. Frismstie Anglss.—The angles for the commonly occurring diitetragonal prisms aro as follows:

Angle on Angle on Angle un

m (110, /) a (100, t-i") m (110. 7>

SO" 67J' S30. i-\ SO" 571' 14' ^'

ae 34. 820.1-1 38 4li 11 IS]

IS 20 480. j-i S6 52} 6 71

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HEXAGONAL BTSTEH.

m. HEXAGONAL SYSTEM.

113. The Hexagonal Stbtem includee all the forme whicb are referred to four axes, three equal lateral axes in a common plane intereecting at angles of 60° , and a fourth, vertical axis, at right angles to them.

Two sections are here included, each embracing a number of distinct groups related among themBelves. They are called the Hexagonal Division und tlie Trigonal (or Rkombohedral) Division. The symmetry of the former,, about the vertical axis, belougs to the hexagoual type, that of the latter to the trigonal type.

Miller (18&2> referred *X\ the rorins of ihe hexaeoriBl ayBtcm to tliree eqtiil axes pant11e> In llie fucea of Ibe finiilainental rbombubedroii, and heuce loUrsccling nt equal augles, uot 90° . Tbis metbod (furthet expUiui'd !□ An. IBt) lias tLe disitdvantiLge of fiiHiu): lo brlDK out the relutloDsbfp between llie noruinl bexa^ouuil mid telragoaiil types, botb cliaracterizea by a prioclpnl tixh of Bymtnelry, whlcli (on the system liere ailopled) Is Ibe vertical crystal I ograpble ails. It furtljer gives diSerent aymbols 10 faceR whicb nre crysiallo-graphicnllly ideoiical. It ia more uaturn) to employ the ibree rbombiiherfral avea for trigonal forms ooly. as ilonu b^ Orolh (IBM}, wlio iDcliidca these groups iu a Trigonal Sgifcm; but tliis also hna some diaadvanings.

114. Qroupi.—There are fire possible groups in the Hexagonal Division., Of these the normal grotip ia much the most important, and two others are also of importance among crystallied minerals.

In the Trigonal Division there are seven groups; of these the rhombo-hedral group or that of the Oalcite Type is by far the most common, and three others are also of importtiice.

119. Axes and Symbols.—The position of the four axes taken is shown in Fig. 193; the three Upright axes are called u, v, w , and the vertical axis is z . Further, when it is desirable to distinguish u, v, w ,

between the lateral axes they may be designated u, v, w . The general position of any plane on the method of Bravais (who adapted the system of Miller to this system) may be expressed in a manner analogous to that applicable in the other systems, viz.:

1111,

$$u \cdot v \cdot w = 1$$

The corresponding indices for a given plane are then h, k, l, i ; these always refer to the axes named in the above scheme:

It is found convenient to consider the axis z , as negative in front and positive behind, hence the general symbol is hkl . Further, as following from the angular relation of the three lateral axes, it can be readily shown to be always true that the algebraic sum of the indices h, k, l, i is equal to zero:

$$h + k + l + i = 0.$$

The general expression for any plane in accordance with the system of Bravais is

$$h : k : l : i$$

; Coefficient

CRYSTALLOGRAPHY.

Here it is always true that $h + k + l + i = 0$. The shortened form for the above expression as adopted by Kauffman is hkl .

(1) $h : k : l$:-

This is equivalent (after multiplying by 8) to

hkl

Here h, k, l, i are $\pm 1, 2, 3, 4, 6, 8$. The symbol is hence written hkl , or hkl .

194.

A. Hexagonal Division.

1. NORMAL GROUP (13). BERYL TYPE.

116. Symmetry.—Crystals belonging to the normal group of the Hexagonal Division have one principal plane of symmetry, the plane of the lateral axes ; also, normal to this and meeting in the vertical axis at angles of 30° , six other

planes of symmetry, like three-and-three. Those of one set pass through the lateral axes a, a, a (Figs. 195, 198) and are —

- hence parallel to the faces of the form (1010) lettered m, as in Fig. 208. The others are diagonal to the first set (Figs. 196, 197) and are parallel to the faces of the form (1130) lettered n, Fig. 208.

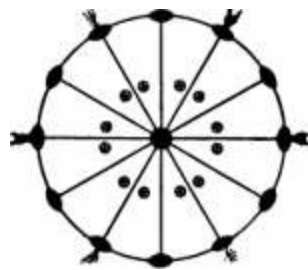
Further, these crystals have one principal axis of hexagonal, or sixfold, symmetry, the vertical crystallographic axis ; also six horizontal axes of binary symmetry; three of these coincide with the lateral crystallographic axes, the others are diagonal to them, The symmetry of this group is exhibited in the accompanying spherical projection, Fig. 194, and by the figures on the following pages from 195 to 209.

The analogy between this group and the normal group of the tetragonal system is obvious at once and will be better appreciated as greater familiarity is gained with the individual forms and their combinations.

117. Forms.—The possible forms in this group are as follows :

Hiller-Bravais.

1. Base (0001) ∞
2. Unit prism, or prism of the first order ∞



(1010)

, Diagonal prism, or prism of the second order ∞

(1010) c mP or I, m

ZaiZai

: CO e OS $\frac{1}{2}$ or t-3, a

Hlller-BnTdB NaumauD.

4. Dihexagonal prism {hkiO) na :pa : — a -.aic <aPn or t-n

as, (2l30) {a : 3a : - a : «) c oo Z' j or t-}

5. Pyramids of the unit, 1 i}, {xT,n „„„„„ „„„„„™p„„ „

oi firet order f *'' ('*' ^ '' * a : ^ a : - a : mc mPorm

M, (lOll) a-.-na: -a:c P or 1; also 203l (a :qd o : - a :2c) 2P or2

as, (1122) an : 2fl : - « :c i^ or 1-2

as, (2 l3l) ia:3a:~a: 3c SPi or 3-(

In the above $h > k$, and $A + k = -i$.

118 Baa0 —The base, or baetil pinacoid, includes the two faces, 0001 and 000l, parallel to the plane of the lateral axes It is uniformly designated by tlie letter c; see Figs. 19£*e(seq.

119. Prisms. Unit Prism.—There are three types of prisms, or forms in which the faces are parallel to the vertical aiaa.

The unit prism, or priam of the first order, Fig. 195, includes six faces, each one of which is parallel Co the vertical axis and meets two adjacent lateral sie^ at equal distances, while it is parallel to the third lateral axis. It has hence the general symbol (lOlO) and is uniformly designated by the letter m; its six faces, taiien in order (see Figs. 195 and 309), are;

lOlO, OlTO, lIOO, lOlO, OlLO, lIOO.

120. Diagonal Priim.—The diagonal prism, or prism of the second order. Fig 19C. hap six faces, each one of which is parallel to the vertical axis, and meets the three lateral axes, the two alternate at the unit distance, the other It one-half this distance; or, which is the same thing, it meets the last-named axis at the unit distance, the dthers at double this distance.* The general symbolL is (llSO) and it is uniformly designated by the letter a; the six faces (see Figs 196 and 309) in order are :

ll50. l2l0, 3110, nsO, l5l0, 2llo.

The unit prism and the diagonal prism are not to be distinguished geo* metrically, each being a regular hexagonal prism with normal interfacial angles of 60° . They are related to each other in the same way as the two square prisma tn (110) and o (IOC) of the

tetragonal system.

The relation in position between the unit prism (and pyramids) on the one hand and the diagonal prism (and pyramids) on the other will be understood better from Fig. 198, representing a cross-section parallel to the base c.

12L Dihexagonal Prism.—The dihexagonal prism, Fig. 197, is a twelve-sided prism bounded by twelve faces, each one of which is parallel to the vertical axis, and also meets two adjacent lateral axes at unequal distances the ratio of which always lies between 1 : 1 and 1 : 2 (see p. 66). This prism has two unlike edges, lettered z and y, as shown in Fig. 197. The general symbol is (hkl) , and the faces of a given form, as $(\frac{1}{2}10)$, are:

$110, 120, 1\bar{1}0, 530, 520, 310, \bar{1}\bar{1}0, 130, 1\bar{2}0, 250, 321, 310.$

* Since $101 : 10\bar{1} = -101 : 101$, the faces are $101, 10\bar{1}$.

$101, 10\bar{1}$

$\frac{1}{2}10$

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1 > 1

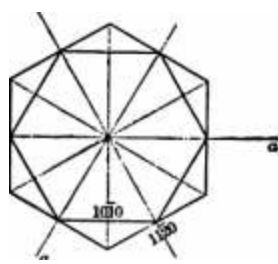
122. Pyramids. Unit Pyramids.—Corresponding to the three types of prisms just mentioned, there are three types of pyramids.

A unit pyramid, or pyramid of the first order. Fig. 199, is a form bounded by twelve similar triangular faces—six above and six below—which have the same position relative to the lateral axis as 183.

The faces of the unit pyramid, while they also intersect the vertical axis. The general symbol is hence (hkl) . The faces of a given form, as (101) , are:

Above $101, 011, 110, 10\bar{1}, 0\bar{1}1, 110.$

Below $101, 01\bar{1}, 110, 10\bar{1}, 0\bar{1}\bar{1}, 110.$ On a given species there may be a number n of unit pyramids, differing in the ratio of the lateral to the vertical axis, and thus forming a zone between the base (0001) and the faces of the unit prism (1010) . Their symbols, passing from the base (0001) to the unit prism (1010) , would be, for example, $1014, 1012, 3033, 1011, 3052, 2051$, etc. In Fig. 202, the faces o and s are unit pyramids and they have the symbols respectively (101) and (203) , here $(101) = 1014$. In Fig. 205, p is the unit pyramid (101) ;



here $i = 0.50$. As shown in these cases the faces of the unit pyramids replace the edges of the unit prism. On the other hand, they replace the solid angles of the diagonal prism a (1120) as shown in Fig. 20J.

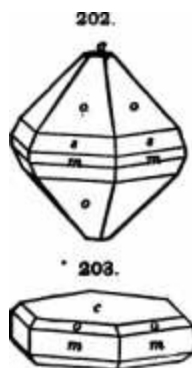
■, :cc., CoogIC

123. Diagonal Pyramidi. — The diagonal pyramid, or pyramid of the Kond order (Fig. 2UU), is a double six-sided pyramid including the twelve similar faces which have the same position relative to the lateral axes as the faces of the diagonal prism, and which also intersect the vertical axis. They have the general symbol $(\frac{h}{a} \frac{k}{b} \frac{l}{c})$. The symbols of the faces of the form (1122) are:

Above 1133, 12T2, 2112, 1!22, 15l2, 2112. Below 1153, ViU, 2llS, TT23, 13lS, 2115.

The faces of the diagonal pyramid replace the edges between the faces of the diagonal prism and the base. Further, they replace the solid angles of the unit prism vi (1010). There may be on a single crystal a number of diagonal pyramids forming a zone between the base c (0001) and the faces of the diagonal

ffism a (1150), as, naming them in order: 1134, 1152, 2213, 1121, etc In fig 205, o, s are the diagonal pyramids (1132) and (1131).



^^^

BaDkrit & CoDDflllste. Beryl

124. Sibezagonal Pyramid.— The dihexagonal pyramid. Fig. 201, is a double twelve-sided pyramid, having the twenty-four similar faces embraced under the general symbol $\{hkU\}$. It is bounded by twenty-four similar faces, each meeting the vertical axis and having a ratio for the intercepts on two adjacent lateral axes between 1 : 1 and 1 : 2 (cf. the general

symbol (2) given in Art. 115). Thus the form ($\bar{1}31$) includes the following twelve faces in the upper half of the crystal:

21. $\bar{1}$, 12. $\bar{1}$, $13\bar{1}$, 3311 , $3\bar{3}11$, $312\bar{1}$, $\bar{1}31$, $153\bar{1}$, $132\bar{1}$, $231\bar{1}$, $3S1\bar{1}$, $3\bar{1}31$.

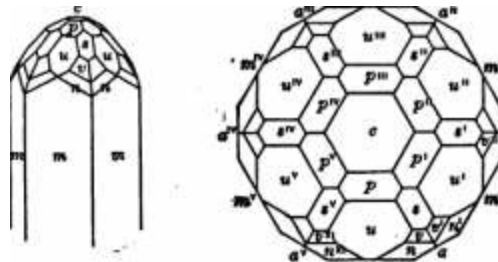
And similarly below with / (here $\bar{1}$) negative, $213\bar{1}$, etc. The dihexagonal pyramid is often called a berylloid because a common form with the species beryl, see Fig 206. w is the berylloid ($11 \cdot 3\bar{1}5\bar{3}$).

125 Combinations.—Fig. 207 of beryl shows a combination of the base c ; (0001) and prism m (1010) with the unit pyramids j (1011) and k ($203\bar{1}$); the diagonal pyramids s ($11\bar{3}1$) and the berylloids v ($213\bar{1}$) and n ($3\bar{1}1\bar{1}$). Both the last forms lie in a zone between j and s , for which it is true that $k \sim j$.

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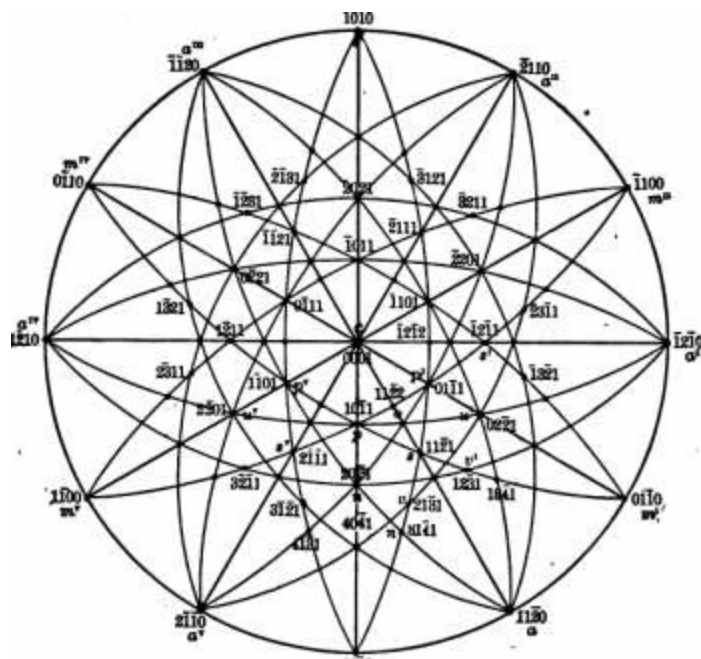
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The basal projection of a similar crystal shown in Fig. 208 is very instructive as exhibiting the symmetry of the normal hexagonal group. This is $6/m$



Crystal. Beryl Beryl.

of the spherical projection in Fig. 209 of a like crystal with also the form o ($11S2$).



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HSXAOONAL BYSTBU.

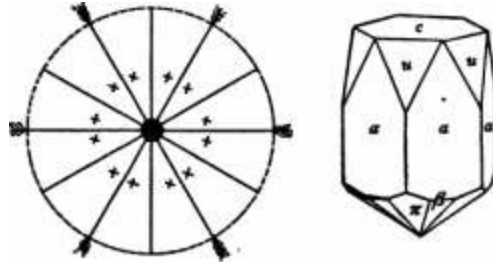
2. HEMIMORPHIC GROUP (U) IODYRITE TYPE.

126. Symmetry,—This group differs from the normal group only in having
 10 transverse plane of principal symmetry and no horizontal axis of binary
 symmetry. It has, however, the same two sets of planes of symmetry meeting
 at right angles in the vertical
 axis which is an axis of binary
 symmetry. By means of these there is no
 center of symmetry. The
 symmetry is exhibited in
 the spherical projection, Fig.

210.

127. Form. — The forms belonging to this group are the two basal planes, 0001 and 000 $\bar{1}$, here distinct forms, the plus (upper) and minus (lower) pyramids of each of the three types; also the three prisms, which last do not differ geometrically from the prisms of the normal group. An example of this group is found in iodyrite, or silver iodide, Fig-211; here $m = (4051)$, $t = (40\bar{1}5)$, $\sqrt{3} \wedge (9-9\bar{1}8'50)$. Greenockite and wurtzite, also zincite

(Fig 50, p. 18) are classed here, but there is some reason for believing that these species belong, with tourmaline, to the corresponding group under the trigonal (rhombohedral) division.

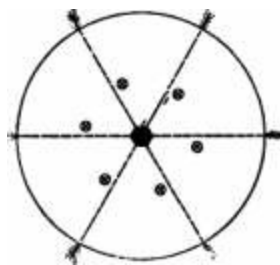


3 PYRAMIDAL GROUP (15). APATITE TYPE.

128. Typical Forms and Symmetry.—This group is important because it includes the common species of the Apatite Group, apatite, pyromorphite, mimetite, vanadinite. The typical form is the hexagonal prism (AAiO) and the hexagonal pyramid [hkil), each designated as of the third order. These forms are shown in Figs. 313 and 214. They and the other forms of the group have only one plane of symmetry, the plane of the horizontal axes, and also one axis of hexagonal symmetry (the vertical axis).

The symmetry is exhibited in the spherical projection (Fig. 212). It is seen here, as in the figures of crystals given, that, like the pyramidal group under the tetragonal system, the faces of the general form {hkil) present are half those belonging to each sextant, and further that those above and below fall in same vertical zone.

129. Prism and Pyramid of the Third Order.—The prism of the third order (Fig. 313) has six like faces embraced under the general symbol (hklO), and the form is a regular hexagonal prism with angles of 60° , not to be distinguished



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CRYSTALLOGRAPHY.

geometrically, if alone, from the other hexagonal prisms; cf. Figs. 195, 196, p. 68 The six faces of the right-handed form (2130) have the symbols 2150, 1330, 3^{10} , 5130, 13a0, 33l0. The faces of the complementary left-handed form have the symbols:

1230, 33l0, 3l20, T-330, 23l0, 3l30. The two forms together embrace all the faces of the dihexagonal prism

(Fig. jy;).

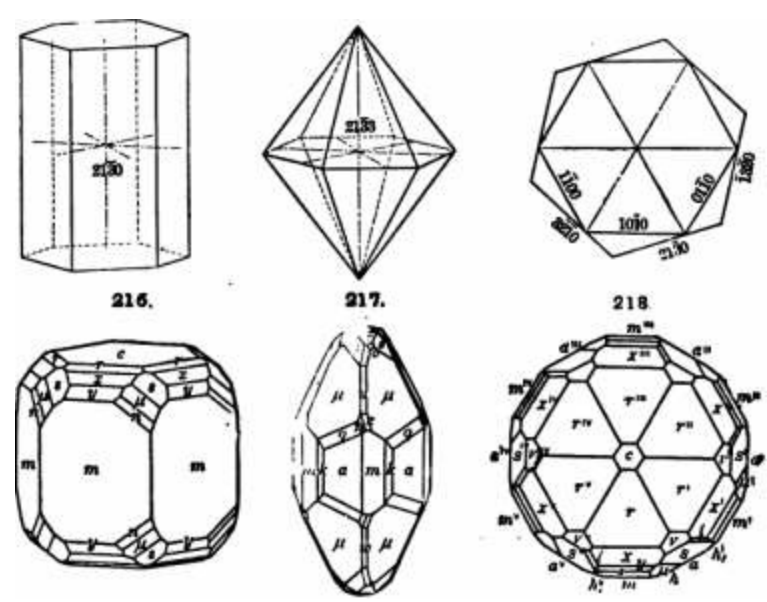
riu pyi'iiniiJ is also & regular liexagonal pyramid of the third order, and in iU- rclulioiis to the other hexagonal pyramids of the group (Figs. t99,tioo) it is analoguieb to the aquiie pyramid of the third order met wjlli in the correspond-iog group of the tetiugoual system (see Art. 98). The faces of the right* handed form ($\wedge l33$) are :

Above 2133, !333, 3213, 5133, l323, 32l3.

Below 21113, 1323, 5213, 3133, 1323, 3313.

There is also a com piemen tar; left-handed form, which with this embraces all the faces of the dihexagonal pyramid. The cross section of Fi^\wedge . 215 shows in outline the position ot the unit prism, and also that of the right-handed prism of the third orde.

213. 314. ai6.



Apaiite. Apatite. A|iHiie.

The prism and pyramid just described do not often appear on crystals as predominating forms, though this is sometimes the case; for example, Fig. 217 shows a crystal of apatite in which the prominent pyramid /j is a pyramid of the third order (2131). Commonly tiiese faces are present modifying other fundamental forms, and their character is obvious from their position relatively, for example, to the unit prism m (1010).

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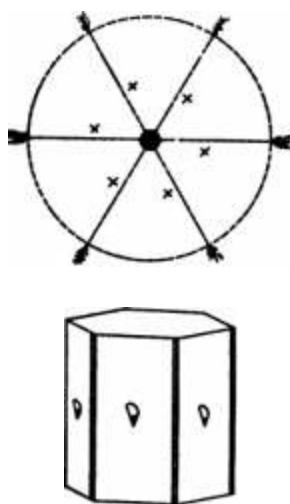
HEX&OONAL SYSTEM.

130. Other Fonoa. —The remainiug forms of the group are geometrically like those of (lie nurnal group, viz., lli« base (UOUI) ; the uuit prism (lulUj ; the diagonal prism (ll3o); the unit pyramids (hOhl); and the diagonal pyramids (h-ff-ilfl). That their molecular structure ia the same is readily proved, for example by etching. In this way it was shown that pyromorphite ■lid miuiecite belong in che same group with apatite (Baumhauer), though jsrystals with the typical forms have not been observed.

131. Typical crystals of apatite are given in Figs. 316,217, and S18 (a basal section). They show the third-order prisms A (ilSUj, h (2l3o), h, (1^3o), and the third order pyramids, right, M (^l3l), « (SlJl), q (437l), o CAlii), left, t (123-.i); also the unit pyramids r (lOia), x (lOll), y (2u3l), z (So5l); the diagonal pyramids v (LlS'.i), a (ll3l); finally, the prisms m (lOU), a (llSO), and the base c (0001).

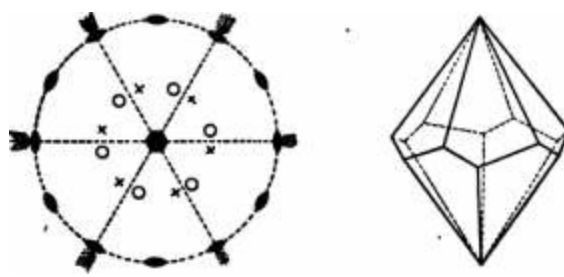
4. PYRAMIDAL-HEMIMORPHIC GROUP (16). NEPHELITE TYPE. 133. Symmetry.—A fourth group under the hexagonal system, the pyramidal-hemimorphic group, is like that just described, except that the forms are hemimorphic. The single horizontal plane of symmetry is absent, but the vertical axis is still an axis of hexagonal symmetry. This symmetry is shown in the spherical projection of Fig. 219, The typical form would be like the upper half of Fig. 214 of the pyramid of the third order. The species nephelite is shown by the character of the etching-figures (Fig. 220, Groth after Baumhauer) to belong here.

319. aao.



Nephelite. 5. TRAPEZOHEDRAL GROUP (17). 133. Symmetry. —The last group of this division is the trapezohedral group. It has no plane of symmetry, but the vertical axis is an axis of hexagonal

aai. 223.



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CaYBTALLOaOAPHT.

Bjmmetry, and there are, further, six horizontal axes of binary sjmmetrj. There is no center of symmetry. The distribution of the faces of the typiciil form (hkil) is shown in the spnerical projection (Fig. 321). The Cypicitl forinE. the right and left hexagonal trapezohedroua (see Fig. 222), are etiautioniorphoua, and the few cryatallized salts fulling in this group show circular polarization. The symbols of the right form {-llfi) are as follows :

Above 2153. 133;!, 3313, 5l33, 1^23, 3213.

Below 1235, S313, 3l25, 1533, 33l3, 3l33.

B. Trigonal or Rhombohedral Division.

134. Oeneral Character.—As stated on p. 65, the groups of this division are characterized by a vertical axis of trigonal, or threefold, symmetrT. There are seven groups liere included of which the group of the Caloite I'ype is by far the most important.

13&. TrigonoType Qroup.—The first group (18), that which has strictly the — nighest grade of symmetry, has

no known representatives among crystals, natural or artificial. It has, besides the vertical axis of trigonal symmetry, three horizontal axes of binary symmetry. There are four planes of aym-metry, one horizontal, and three others intersecting at angles of 60° in a ertical axis. The characteristic forms are the trigonal priamnnd pyramid and ditrigonal prism and pvramid. The symmetry is exhibited in Fig. 223. The typical form (Fig. 224) is a double ditrigonal pyramid with terminal edges alike in alternate sets of three each. This form may be compared to a scalenohedron twinned about the vertical axis.

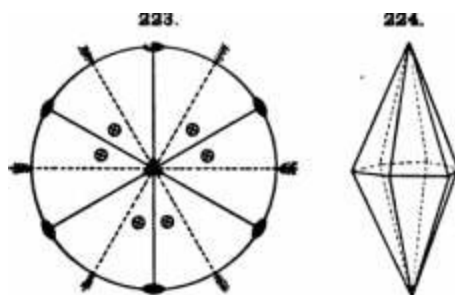
2. RHOMBOHEDRAL GROUP (19). CALCITE TYPE.

136. Typical Forma and Symmetry. —The typical forms of the rhombohedral (group are the rhombohedron (Fig. 226) and the scalenohedrjon (Fig. 242). These forms, with the spherical projection. Figs. 236.

225 and 252, illustrate the symmetry characteristic of the group. By comparing Fig. 253

with Fig. 209, p. 70, It will be seen that all the faces in half the sectants are present. This group is hence analogons to the tetrahedral group of the isometric system, and the sphenoidal group of the tetragonal system.

In this group there are three planes of symmetry only; these are diagonal to the crystallographic axes and intersect at angles of 60° in the vertical crystallographic axis. This axis is with these forms an axis of trigonal symmetry; there are, further, three horizontal axes of binary symmetry. Compare Fig. 235, also Fig. 226 ei seq.



,ab,GoOgIc

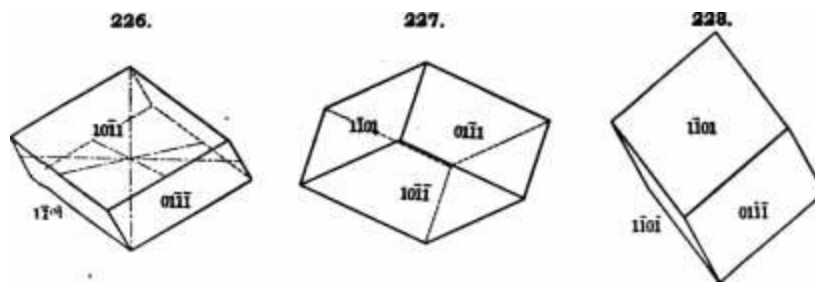
137. Rhombohedron.—Geometrically described, the rhombohedron is a solid bounded by six like faces, each a rhomb. It has six like lateral edges forming a zig-zag line about the crystal, and six like terminal edges, three above and three alternate position below. The vertical axis joins the two trihedral angles, and the lateral axes join the middle points of the opposite sides, as shown in Fig. 226.

The general symbol of the rhombohedron is $\{IMkl\}$, and the successive faces of the unit form $\{lOll\}$ have the symbols :

Above, $lOll$, lOl , Oll ; below, Oll , lOl , lOl .

The geometrical shape of the rhombohedron varies widely as the angles change, and consequently the relative length of the vertical axis i (expressed in terms of the lateral axis, $a = 1$), As the vertical axis diminishes, the rhombohedrons become more and more obtuse or flattened; and as it increases they become more and more acute. A cube placed with an octahedral axis is obviously the limiting case between the obtuse and acute forms where the interfacial angle is 90° . In Fig. 336 of calcite the normal rhombohedral angle is $74^\circ 55'$ and $b = 0.854$, while for Fig. 228 of hematite this angle is 94° and $h = 1.366$. Further, Figs. 239-334 show other rhombohedrons of calcite, namely, e ($OliS$), o (0554), f ($032l$), j ($4oll$), and v ($13-0-1B \blacksquare 1$), p ($16'0-15-1$), here the vertical axes are in the ratio of $i, f, 2, 4, 13, 16$, to that of the fundamental (cleavage) rhombohedron of Fig. 226, whose angle determines the value of i .

22fi. 227.



13B. Plus and Minus Rhombohedrons.—To every plus rhombohedron there may be an inverse and complementary form, identical geometrically, but bounded by faces falling in the alternate sextant. Thus the minus form of the unit rhombohedron (0111) shown in Fig. 237 has the faces:

Above, 1101, 1011, 1101; below, 1101, 0111, 1011.

The position of these in the spherical projection (Fig. 252) should be carefully studied; see also Fig. 262, of the figures already referred to. Figs. 238, 233, 333 are plus, and Figs. 237, 229, 230, 231 minus, rhombohedrons; Fig. 234 shows both forms.

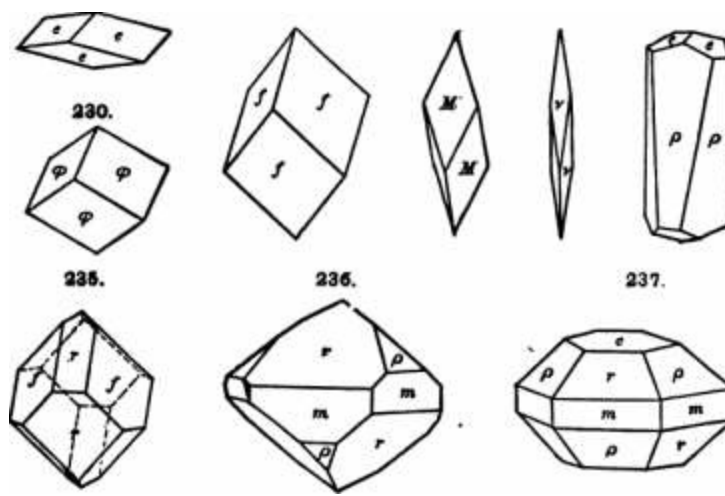
It will be seen that the two complementary plus and minus rhombohedrons of given axial length, that is, of given angle, e.g., 1011 (+ E) and 0111 (— 7?), together embrace all the like faces of the double six-sided pyramid. When these two rhombohedrons are equally developed the form is geometrically identical with this pyramid. This is illustrated by Fig. 237 of gmelinite r (1011), p (0111) and by Figs. 266, 267, p. 83, of quartz, r (1011), z (0111).*

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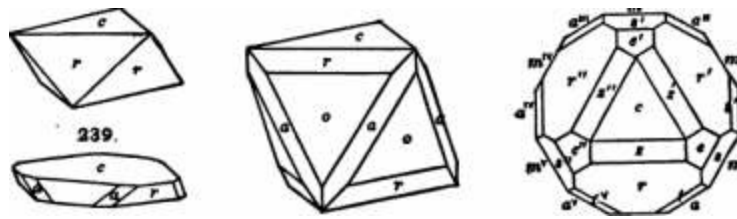
CRYSTALLOGRAPHY.

In each case the form, which is geometrically a hexagonal pyramid (in Fig. 237 with c and m), is in fact a combination of the two unit rhombohedrons, plus and minus. Commonly a difference in size between the two forms may be observed, as in Figs. 336 and 363, where the form taken as the plus rhombohedron predominates. But even if this distinction cannot be established, the two rhombohedrons can always be distinguished by etching, or, as in the case of quartz, by pyro-electrical phenomena.

329. 231. 332. 333. 334.



FigB. 329-235, Calcite Figs. 286, 387. Qmellnite. 139 Of the two series, or zones, of rhombohedrons the faces of the plus rhombohedrons replace the edges between the base (0001) and the unit prism (1010). Also the faces of the minus rhombohedrons replace the alternate edges of the same forms, that is, the edges between (0001) and (0110), (compare Figs. 236, 237, etc.). Fig. 338 shows the rhombohedron in combination with &c 338. 340. 341.



Fig* 296. 239. Hematite. Coquimbite. Eudialyte.

base Fig. 239 the same with the prism (1110). When the angle between the two forms happens to approximate to $70^{\circ} 32'$ the crystal simulates the aspect of a regular octahedron. This is illustrated by Fig. 240; here $\cos = 69^{\circ} 42'$

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also $\cos = 71^{\circ} 33'$, and the crystal resembles closely as octahedron with truncated edges (of Fig 77, p. 36).

140. There is a very simple relation between the plus and minus rhombohedrons which it is important to remember. The form of one series which truncates the terminal edges of a given form of the other has half the length of the vertical axis, and this ratio is expressed in the values of the indices

of the two forms Thus (0112), or $-5R$, truncates the terminal edges of the plus unit rhombohedron (1011), or R ; (1114), or $+R$, truncates the

terminal edges of (0112), or $-R$, (1015) of (2035). Again (1011), or $+R$,

truncates the edges of (025l), 01 - 2R; (4041), or + 4R, of 022l, or - 3B, etc. This is illustrated by Fig. 235 with the forms r (10ll) and /{023l}. Also in Fig. 241 a basal projection, z (10T4) truncates the edges of e (Oll2), e{ol!'}otr(1oll); r (IIUl) of s (023l).

141. Scaleno-hedron.—The scaleno-hedron, shown in Fig. 242, is the general form for this group corresponding to the symbol hkl. It is a 342, solid, bounded by twelve faces, each a scalene triangle. It has

roughly the shape of a double six-sided pyramid, but there are two sets of terminal edges, one more obtuse than the other, and the lateral edges form a zigzag edge around like that of the rhombohedron. Like the rhombohedron, the scaleno-hedrons may be either plus or minus according as to whether the faces are symmetrical to the zone-circle 1010,0001, or to 010, 0001. The former plus forms correspond in position to the plus rhombohedrons and conversely.

The plus scaleno-hedron (2l5l), Fig. 242) has the following symbols for the several faces:

Above 2l3l, 33ll, 5311, 1331, 1321, 3l5l.

Below 1231, 132l, 5131, 3T3i, 254l, 3511. $g^{\wedge}, \wedge, \wedge l^{\wedge} \wedge j, \dots$

For the complementary minus scaleno-hedron (123l) the symbols of the faces are:

Above 1331, 132l, 3l21, 5T31, 23ll, 33!l.

Below 2311, 3211, 123l, 133l, 3l2l, 2l31.

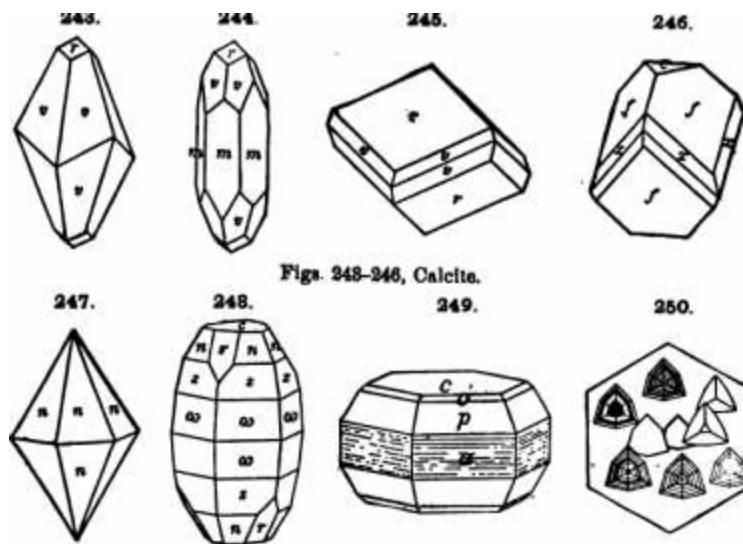
142. Definition of B«al«iioh*droii« to Blioubolitdroiii.—It was noted above that the scaleno-

hedron in general has a series of zigzag lateral edges. The rhombohedron. It is a 342, solid, bounded by twelve faces, each a scalene triangle. It has roughly the shape of a double six-sided pyramid, but there are two sets of terminal edges, one more obtuse than the other, and the lateral edges form a zigzag edge around like that of the rhombohedron. Like the rhombohedron, the scaleno-hedrons may be either plus or minus according as to whether the faces are symmetrical to the zone-circle 1010,0001, or to 010, 0001. The former plus forms correspond in position to the plus rhombohedrons and conversely.

The definition of the scaleno-hedron may be defined with reference to its parent rhombohedron. For example, in Fig. 245 the scaleno-hedron n (2l3l) has twice the vertical axis of / (0221). Hence the symbols devised by Naumann to express this relation, written in general mRn (in Dana's System m-), where the n expresses the multiple value of the vertical axis corresponding to the rhombohedron mR. The symbol

*o CRTSTALLOGRAPHT.

Of • la on this system R8 or 1«. and of * - 3R,. or - 2". If niPn U the symbol of a scaleno-
aeOTOD on ihe beiagoosl type uid m,Rn, that referred hi thU way, ft tney be ahown that



Figs. 248-246, Calcite.

Figs. S47 348, Coruodum.

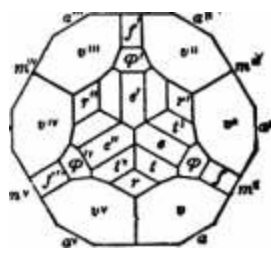
Figs. 24S, eOO, Spaagolite.*

g forniB of the normal group of the 251.

143. Other Forms.—The i

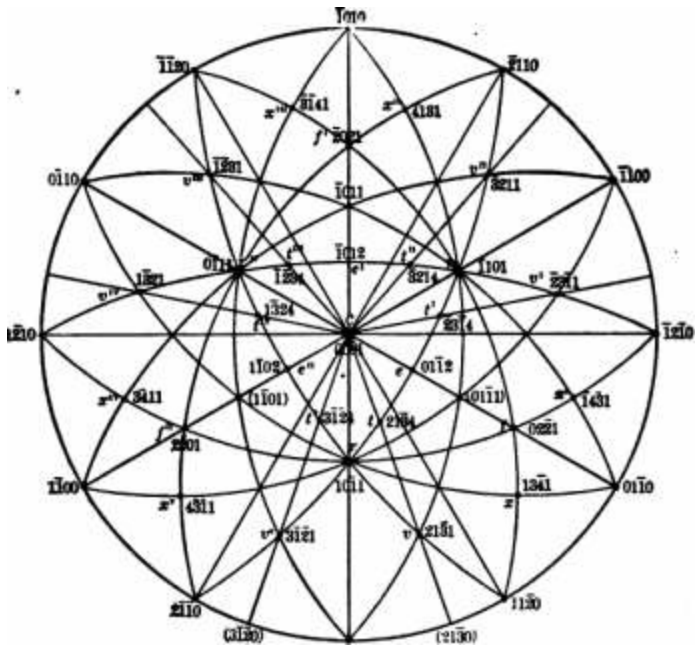
Tbomhohedral difisiou are geometrically like those of the correBponding group of the hexagonal syatem—viz., the base c (0001); the prisma vt (1010), a (1130), (//XriO); also the diagonal pyramids,'ae (U51). Some of these forma are flhown in the accompanyiug figiirsB. For further illustrations reference may be made to typical rhomhohedral species, as ralcite, hematite, etc. ' With respect to the diagonal pyramid, it is interesting to note that if it occurs alone (as in Fig. 247, n = 2243) it is impossible to say, on geometrical grounds, whether it has the trigonal symmetry of the rhomhohedral type or the hexagonal symmetry of the hexagonal type. In the ~ latter ease, the form might be made a unit pyramid by exchanging the axial and diagonal planes of symmetry. The true symmetry, however, is often indicated, as with corundum, by the occurrence on other crystals of rhomhohedral faces, as r (1011) in Fig. 248 (here i = 2241, a> = 14-14-58-3). Even if these are absent {Fig. 249), the etchiiig-figiires (Fig. 250) will often serve to reveal the true trigonal molecular symmetry; here o = (1124), p = (1122).

144. A basal projection of a somewhat complex crystal of calcite is given in Fig. 251, and ,■* spherical projection for the same species in Fig. 3.72; both



show well the symmetry in the distribution of the faces. Here the forms are: prisma, a (1130), m (1010); rhombohedrums, plus, r (1011), mistia « (0115), ^ (0554),/{oi!2l}j acalenohedrona, plus, v (2131), t (2134), minus, «(13il).

363.

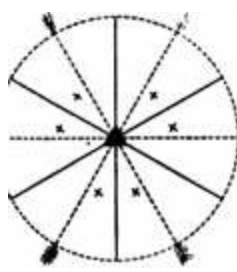


3. RHOMBOHEDKAL-HEMIMORPniC GROUP (20). TOURMALINE TYPE. 146. Symmetry.—A number of prominent rhombohedral species, i tourmaline, pyrargyrite, pronstite, belong to a hemimorphic group under this

2B3.

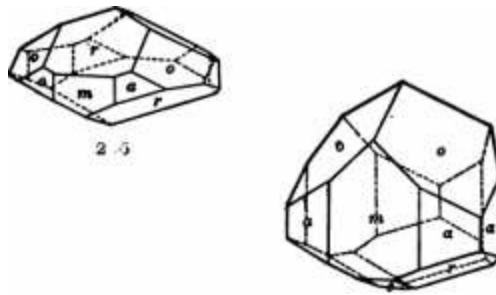
division. For them the sj-mmetry in the grouping of the faces differs at the two extremities of the vertical axis. The forms have the three diagonal planes of symmetry meeting at angles of CO" in the vertical aais, whieli is an aieg of trigoual symmetry. There are, however, no horizontal axes of symmetry, as in the rhombo-hedral grnnp, and there is no center of symmetry. Cf. Fig. 25.1.

146. Typio^ Forms. — In this group the basal planes (0001) and (000l) are distinct forms. The other characteristic forms are the two trigonal prisms7«(1010) and w), (01 iO) of the nnit aeries; also the four trigonal unit pyramids, corresponding respectively



to the three upper and three lower faces of a plus rhombohedron, and the three upper and three lower faces of the minus rhombohedron; also the bellimorphic diagonal hexagonal pyramid; finally, the four ditrigonal pyramids, corresponding to the upper and lower faces respectively of the plus and minus scalenohedron. Figs. 254-357 illustrate these forms. Fig. 25f is a basal section with r , (OII) and e , (IOLS) below.

264. 266. 367.



Figs. 254-257, Tourmaline, $m(111)$, $o(111)$, $a(111)$, & (111) , $r(111)$, $o(111)$, $c(111)$, $t(111)$.

4. TRI-RHOMBOHEDRAL GROUP (21). PHENACITE TYPE.

147. Symmetry.—This group, illustrated by the species diopside, phenacite, willemite, dolomite, ilmenite, etc., is an important

one. It is characterized by the absence of

planes of symmetry, but the vertical axis is

still an axis of trigonal symmetry, and there is

a center of symmetry. Cf. Fig. 258,

148. Typical Forms.—The distinctive forms

of the group are the rhombohedron of the second

order and the hexagonal prism and rhombo

hedron, each of the third order. The group is

thus characterized by three distinct rhombohedra \ / \

drons (each -|- and —), and hence the name "••, ^." \ ,. •'

given to it. ^ ■ -- -- ^

The complementary plus and minus rhombohedra of the second order together embrace all the faces of the diagonal pyramid of the normal group. For example, in a given case the faces for the plus and minus forms are;

Plus (above) $ll32, i5ll2, lSl2$; (below) $12T3, TT23, 2liS$, Minus (above) $!2l3, 1122, 2ll2$; (below) $Sll3, l5l5, 1133$.

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HEXAGONAL 8T8TElf.

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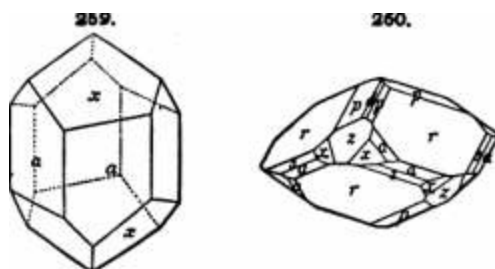
The rhombohedron of the third order has the general symbol (hkl) , and the symbol G of the six like faces of the plus right-handed form $(3l3l)$ are: Above $2l3l, 3211, lS21$; below $1351, 3l3l, 3511$.

There are three other complementary rhombohedra, distinguished respectively as l left-handed $\{aig\}$, minus right-handed $(l32l)$, and minus left-handed $(l!i, 1l)$. These four sets of faces make up the twenty-four like faces of the dihexagonal pyramid of Fig. 201. In Fig. 259 (drawn in the reverse position) the minus right-handed form $(l3S2)$ is shown.

The complementary right and left hexagonal prisms of the third order embrace all the faces of the dihexagonal prism (Fig. 197). The faces in a given case $(2l30)$ have the symbols:

Right $2130, 1350, 3210, Sl30, l320, 32l0$, Left $1230, 53l0, 3120, 1230, 2310, 3l20$.

148. The remaining forms are geometrically like those of the rhombohedral group, viz.: Base c (0001) ; unit prism wt (1010) ; diagonal prism a $(llS0)$; rhombohedra of the first order, sa $(l0ll)$ and $(0lll)$, etc.



160. The form of this group a species phenacite, with which the n

) best illustrated by the crystals of the nns right rhombohedron x (1322) is not

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infrequently the predominating form, as in Fig. 359. The true position of this rhombohedron is shown by its relation to the prisms m (lOlU) and a (llSO). A more complex crystal of the same species* is given in Fig. 260, and a basal projection (in normal position) is shown in Fig. 261. The faces present are: rhombohedrons, first order, plus, r (loll). minus, z (Olll), d (o112); second order, right, p (US3), left, p, (2!l3), oAmi) ; third order, plus right, s (2l3l), minus left, x, (1232), minus right, x (1333),

In order to make clearer the relation of the faces of the different types of forms under this group. Fig. 262 is added. Here the zones of the plus and minus rhombohedrons of the first order are indicated, also the plus and minus ectante corresponding to each.

The following scheme may also be helpful in connection with Fig. 262. It shows the distribution of the faces of the four rhombohedrons of the third order (-J- r, -\ -1, - r, - I) relatively to the faces of the unit hexagonal prism (lOlO).

Phbmacitb Ttpk.

6. TRAPEZOHEDRAL GKROUP (

QUARTZ TYPE.

\ "'

151. Symmetry.—This group includes, among minerals, the species quartz and cinnabar. The forms have no plane of symmetry and no center of symmetry; the vertical axis is, however, an axis of ³ trigonal symmetry, and there are also three

horizontal axes of binary symmetry, coincident in direction with the crystallographic axes; ci. Fig. 263.

158. Typical Forms. — The characteristic form of the group is the triangular trapezohedron shown in Fig. 264. This is the general form corresponding to the symbol (hkil), the faces being distributed as indicated in the accompanying spherical projection (Fig. 263). There are four such trapezohedrons, two plus, called respectively right-handed (Fig. 264) and left-handed (Fig. 265), and two similar minus forms, also right- and left-handed (see the scheme given in Art. 164). It is obvious that the two forms of Figs. 264, 265 are enantiomorphous, and circular polarization is a striking character of the species belonging to the group as elsewhere discussed.

The symbols of the six faces belonging to each of these will be evident on

* Down with the zone of minus rhombohedrons is drawn to better show the modifying facet. Fig. 200 is similar. ■

■•/ V

,ab,GoOgIc

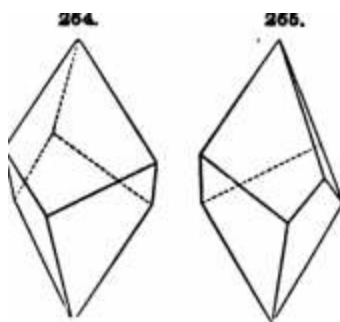
comparing Figs. 263 and 264, p. 70. The complementary forms (r and T) of a given symbol include the twelve faces of a plus rhombohedron, while the faces of all four include the twenty-four faces of the dodecahedral pyramid.

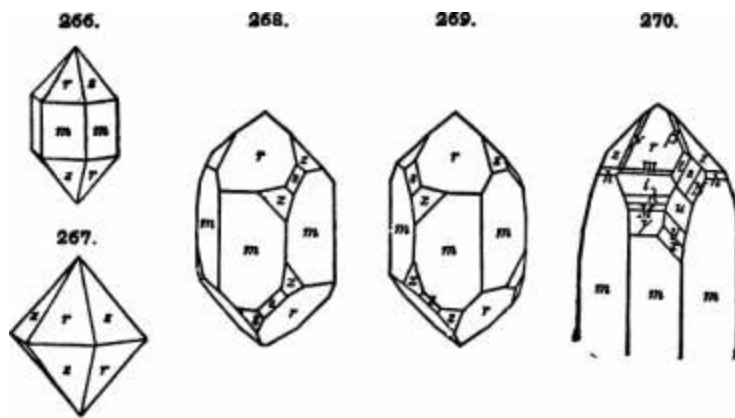
Corresponding to these trapezohedrons there are two ditrigonal prisms, respectively right- and left-handed, $aa(2130)$ and (ZWi) .

The remaining characteristic forms are the right- and left-handed trigonal prism $s(113u)$ and $a(2110)$; also the right- and left-handed trigonal pyramid, $as(11SS)$ and (2112) . They include respectively the faces of the hexagonal prism of the second order $(11SO)$ and those of the corresponding pyramid (1152) ; these are shown in Figs. 196 and 200.

153. Other Form*.—The other forms of the group are geometrically like those of the normal group. They are the base $c(0001)$, the hexagonal unit prism $m(110)$, and the plus and minus rhombohedrons $a(1011)$ and (0111) . These cannot be distinguished geometrically from the normal forms.

Illustration.—The forms of this group are best shown in the species quartz. As already remarked (p. 75), simple crystals often appear to be of normal hexagonal symmetry, the rhombohedrons $r(1011)$ and $x(1011)$ being equally developed (Fig. 267, 367). In many cases, however, a difference in molecular character between them can be observed, and more commonly one rhombohedron, $r(1011)$, predominates; the distinction can always be made out by etching. Some crystals, like Fig. 268, show as modifying faces the right trigonal pyramid $s(1131)$, with a right plus trapezohedron, $a(5151)$. Such crystals are called right-handed and rotate the plane of polarization of light transmitted in the direction of the vertical axis to the right. A crystal, like





FtgR. 366-370, QuarlB.

Fig. 269, with the left trigonal pyramid $s(21)$ and one or more left trapezo-faedrons, as $x(011)$, is ciuled left-handed, and aa regards light has the opposite

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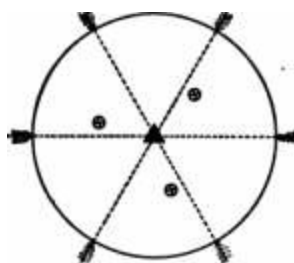
character to that of Fig. 368. Fig. $'^{\wedge}TO$ shows a more complex right-banded cryBtal with ser[^]ral plus and minus rhombohedrons, several [^]ua rigot trapezo-bedrons and the minus left trapezohedron, N.

The following scheme shows the distribuciou of t[^]e faces of the four trape-Eoliedrons ($- \{ -r, -j- I, - r, - I$) relatively to the faces of the unit hexagonal prism (1010) ; it is to be compared with the correspondiug scheme, given iu Art. 160, of crystals of the phenacite type. In the case of the minus forms some authors prefer to make the faces $ai5I, I23I$, etc., right, and $3I3I, i33I$, etc., left.

QoARTZ Type.

1B5. Other Oronpt.—The next group (33) has one plane of synimetry—that of the horizontal axes, and one axis of trigonal symmetry—the vertical axis. There is no center of symmetry. Its characteristic forms are the three typtat of trigonal prisms and the three corresponding types of trigonal pyramide. Ct Fig. 271.

271. 273.



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\ / X # • • *

group (34) has no plane of symmetry and no center of symmetry, ical axis is an axis of trigonal symmetry. The forms are all henii-e prisms trigonal prisms, and the pyramids hemimorphic trigonal r<i 1?:™ or-o Neither of these groups is known to be represented

The last o-. but the vertical morphic, the pyramids. Cf. Fig among crystals.

Hatheu ATicAL Relations ot ths Hbzaoonai. SystBif.

ISE. OholM of Axil.—The position of tlie vertical cryBtinlographic osIh Is flxed in all the groups of ihie system since it colnciilea with the axis of hexagonal symmetry Iu the hexagonal dlvlBiun and that of trleoual symmetry in the rhombohedral division. The three lateral axes are also fixed in alreclion except in the normal group and the subordinate hemimorphic ^roiip of the hexagonal division ; Iu these there Is a choice of two poelilona according to which of the two sets of vertlcnl plnnes of symmetry Is taken as the axial set.

157. Azlal and Angular Elamsnti.—The hiIhI element Is the length of Ibe vertical axis, i. Id terms of a lateral ails, a; in other words, the axial lailo ot a:i. A single measured

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HBXAGONAL STSTEM. maltej mil; be taken as the

Tbeau^uUrelemeiitiausiiially taken as tbeaagk between the biisee (0001) and the unit p;niniul (101), thai is, 0001 a lull.

The KlfttiuD between this aogle aod the axis t is given by the formula

$\tan(0001 \text{ A } 1011) X \wedge V'S = h$. Tue vertical axis is also easllj obtained from the unli diagonal pynmid, dnce

lau (0001 A ll2a) = i. These lelatlons become geeral by writing tbum as follows:

lau (0001 A hOhH X ^1^3 = jXi; tan (0001 A hh-2hti = ^xi.

In general It is easy to obtiia niy reiiuireil angle l'einifeii the poles of two faces on the spberlcat projection eillier by the iiee of tlie InngeiU (or ci>liui^eQt) relation, or by tba BulutiMi of split-ncal IriaoL'les, or by die appUcat ou of liolli metlids. Iu practice most of the triangles iiseil in culculntiOD nre risljt-augU-rl.

168. Taageat and CoUngant B«IaUon* —The tangent relation holds good in any zod« from e (0001) to a lace iu the pristnullc zone. For example:

.n(OQQl A a-a-27lt) _Z>i

tan(0001 A lOTi) I' tan (0001 A 1132)

& siroplifed form; for example, for a

cot(10iOAAKO) = ?^/jj

cot (Il5o A hklO) = ■^^f'8.

The sum of the angles (1010 a UiO) and (llSO A hkm) Is equal to 80*.

Purllier, the last equations can be written In a more general form, applying to any
pjTamicI (fikU) in a zone, first between lOLO and a face In the zone 0001 to OUO. where
tbe angle between lOLO and this face Is known; or again, for tbe same pyramid, in a zone
between 1120 and a face In the zone 0001 to 1010, the angle between 11^ and this face
being given. For example (ef. Fig. 209, p. 70). If the first-men^oned zoneia lOiOAAiiOlll
and ilie second is 1120' AiI:U-1011. then

cot (lOLO A AMI) = cot (lOiO A Olll). — 't -*.

Also simllBrly for other zones,

cot (lOTO A hm = cot {lOiO A

= cot(lOiO A 3(fil).^^,etc

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CRY8TA LLOGRAPHT.

uiii(1oil A Olll) = sliiCf'i, wberoUuiEs^ ud la general

Uo i (MM A OAU) = an H^, where Un S. = j^

^) For a Hagonal pi/ramid, u (llSz),

9 Bfc i(l]S2 A I3l2) = aln £, and Un f = &.

(8) For a rtonioAadnm

rin l(10!l A ilOl) = sin af/}. where a = (0001 A lOll);

Id general

aln i (AOAt A AAOO = du a.V), where a, = (0001 A MJU).

160. Zonal Belatious.—The zonal equations, described by Arts. 43, 44. apply here as in other systems, only it is to be noted that one of the indices referred to the lateral axes, preferably the third, is to be dropped in the calculations and only the other three remain. The faces (Am), (pqr) lie on

XXX

where $u = jh - jl$, $v = (p - Ai)$, $w = hq - kp$.

For example (Fig. 216), the face u lies in the zone mt , 1010 and also in the zone $m'y$. For the first zone the values obtained are $u = 0$, $v = 1$, $w = 1$; hence for any face in this zone the relation $k = t$ holds true, and for the second zone $u = 1$, $v = 0$, $w = -8$, or $ft = 2i$. Therefore the symbol of the given face is given either by the scheme

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311

or from the two equations $k = t$ and $A = 21$. The face m has, therefore, the symbol $219l$, where further $i = -(h + k)$.

161. Formulas.—The following formulas are sometimes useful:

(1) The distances (see Fig. 200) of the pole of any face (Aitnf) from the poles of the faces (lolo), (olio), (iloo), and (wol) are given by the following equations:

$$\cos PA = \cos (lko) \cos (loio) - \sin l \sin k \cos 2h$$

$$\frac{t}{ac} + \frac{i}{ik} + \frac{w}{w} + \frac{z}{z} = \cos (ftu) \cos (olio) = \cos^2 \theta + \cos^2 \phi$$

Dni,t,z.ctvGoogIc

CM PO 3 CM (AM) (0001)

HBZAOOHAL SYSTBH.

'yw + i}(h} + *• + AA) '

_ ZU + W{hq + p* + 3ftp + a*j>

$$\cos^2 [W + W > (A' + *• + hk) \cos [W + 4i'(p' + g* + pj)] \cdot$$

of the normal angles for the several forms in the attached. They are as follows: (a) Hemigonal Pyramid (AOU). Fig. 190:

$$\cos X \{ \text{term} \ln \langle l \rangle - \langle jr \rangle^{\wedge} \langle jjip \rangle - ; oo. Z (Iwal) = -g^{\wedge} r^{\wedge} p^{\wedge} j^{\wedge} ip.$$

^ JHagotuiipyramid{hh-Sh-l). Fig. 900;

<>.T (termini)-^^^; oo. Z (b<.l) = ^^ . (M<):

oo. X(.e< Fig. Ml) - 8f* + 4i'(A' + ^ + AA) -

v, » «... «• + 2i'l3A' + 2AA - A<) oo. T (see Fig. »1) = gji + 4J'(A-+ A* + AA) '

,^ ,. 4i'(A* + A* + AA) - 8i»

W^Zd^-l) = 8/+4i'(A' + A' + AA) -

(<) /MbinvonaJ Pritm (AAiO), Fig. 1S7:

_ , , „ A* 4- A" + 4AA ,, ,. SA* + aAA - *

<-^<"" - 2tA'VAA*) <><T«Hagonal)= g^^.^.^^^ .

(A AlMMMbibvn (1oll)i

IU* — SA'i* CO. X (terminal) ■ g^-j-jpjj-.

oo.x («e Fig. 942) - Bf+^-ty + y + M) -

■B-, » n^n. 8''' + 2*'(3A' 4- 2At - A*) Oo.T(..eng.M8)- 8i. + 4^.(. + y + a^; .

<>iZ(bua]) = «qr4j.(ft. + *> + AA) •

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CBY6TALLooBAPHX.

IBS. AaglM.—The angles for aome commonl}' occurring priama are given in Uu follow tug (able:

m(10iO. i)

BISO *4 8* 6r

4l50 H 10 58t

8140 ^1 18 54

S2To t-i 16 o

SlSO f-l IB 6}

G4do H ^ IS)

IBS. Th* &hamboh*dnl 8jit«ta ^ KUlai.—The following projection (Fig. 2TS) It added

lo order Co show the relatlua oF llie forma in the lieiagonal Hysiem aa referred by

373, Miller to three equal oblique iixea psrallel

lo the faces of the fuudameental rhombo-

hedron. The forms are ns followa :

TL e faces having IL>- iudlces 100, 010, 001 are those of ibe (plus) fundamental rhomhohedrou. while ibe face 111 la the base The facea 331, 133, 312 are ihoa« of the mJDUB fuudameental rhombohedroii; with 100, 010, 001 they form the unit hexFiKooal pymmid.

The facea of lIe hexagoual unit prfam, m (lOTO). h»ve the eymboU 211, 112, isi. §11, ii2, 121. Those of Llie secomi, or diagonal, hexagonal prism, a (1120), have the aymbols 101, Oil, 110, 101, Oil,

lio.

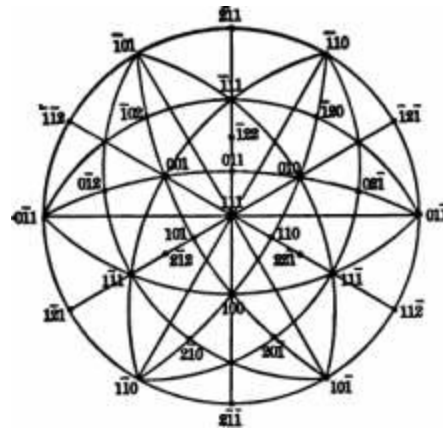
The diheragnnnl pynimid embraces. like the simple hexngouat pyramid, two forma (AW) and {^g): of iliese ihe symbol (hkl) hence beloujcb lo tlie plus scalenolie-droii, and [ffg] lo the mlitius. Id ibis as in other cases it la Irue thnt $t = 2li + Vi - l. / = 2A - * + 2i, ff - A J - 24 + 2i$. For example, lIe facea 210, 301, 021, 130, i02, "Ois (Kg. 978) belong Ju the Rhombohedral DiviBioo of thia system to the Bcalenohedroii (2191). The complemeulary minus scalenobedron woidd have lIe faces 6S4, etc. The tweny-four faces of these two forms taken together would embrace all the faces of tbe dibezagonal pyramid of the Hexagonal Division (2131). Gf. Fig. 20B. p. 70, and Fig. 853, p. 79, with Fie. 373 given here.

Similarly the dlhexagoijal priam Includea tbe ais faces ot tbe form (AM), and the remaining six of the form (e/o|.

It is seen at once that the Indices fdveD above are those of the [aometric system, where the cube cori-esponds to arhombobedron of 90'; the projection of Fig. 110, p. 41, is brought into relation with the above if an octahedral axis is placed vertical.

The iuconveulence of having tbe faces of tbe same form (a.g., the diheiagonal prism nr pynmid ot beryl) represented by two seta of indices Is obvious, and this method, inlniducefl by HilliT. is now seldom employed. This objection, however, disappears if tbe axes iini) indices described are used for rhombohedritt forms oaly. that is, for forms

belonging to the groups which are characterized by a vertical axis of trigonal symmetry. This is the method adopted by Orosi (1806). It is believed by the author, however, that the mutual relations of all the groups or both divisions of the hexagonal system among themselves (as also to the groups of the tetragonal system), both morphological and physical, are best brought out by keeping throughout the same axes, namely, those of Fig. 198, An. I U.



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IV. ORTHORHOMBIC SYSTEM.

164. The Orthorhombic System includes all the forms which are referred to three unequal axes at right angles to each other.

Of these axes the shorter lateral axis, or brachy-axis,* is represented by the letter a, the longer lateral axis, or macro-axis, by h, and the vertical axis by c (cf. fig. 2T5). In the statement of the axial ratio c is uniformly taken as the unit.

1. KOENIG GROUP (25). BAKITE TYPE.

165. Symmetry.—The forms of the normal group of the orthorhombic system are characterized by three unlike planes of

symmetry, at right angles to each other, and therefore, coincident with their intersection-lines, there are three axes of binary symmetry, which directions are also those of the crystallographic axes. These axes are consequently fixed in position by the symmetry, but any one of them may be made the vertical axis. -*

The symmetry of the group is exhibited in the accompanying spherical projection. Fig. 1T4. This should be compared with Fig. 69 (p. 33) and Fig. 148 (p. 53), representing the symmetry of the normal groups of the isometric and tetragonal systems respectively. It will be seen that while normal isometric crystals are developed alike in the three axial directions, those of the tetragonal type have a like development only in the direction of

the two lateral axes, and those of the orthorhombic type are unlike in the three axial directions. Compare also Figs. 70 (p. 34), 149 (p. 54), and 276 (p. 90).

166. Forms.—The various forms possible in this group are as follows:

Miller. Naumau.

3. Base or opinacoid (001) $a > b > c$ \cdot $\{hk0\}$ or o, c

(Unit prism (110) $d > e > f$ \cdot $\{hk0\}$ or o, c

4. \ Macroprisms. \cdot $\{hk0\}$ \cdot $\{hk0\}$ or $t-n$, as (210) $t-s$ (Brachypri8m8..(AtO) $h < k < n$ \cdot $\{hk0\}$ or $t-n$, as (120) $t-5$

5. Macrodomes (h0l) $t-i$ \cdot $\{hk0\}$ or $m-;$, as (201) $3 I$

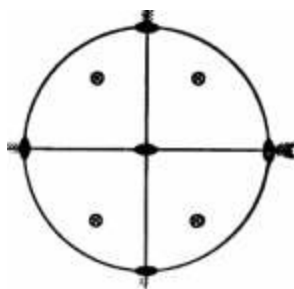
6. Brachydomes {0kl) $a > d > h > m$ \cdot $\{hk0\}$ or $m-l$, as (021) $2-i$

(Unit pyramids {hhl) \cdot $\{hk0\}$ or m , as (221) 2

(111) \cdot $\{hk0\}$ or l

Maon)pyramid8(Sii) $A > i > d > h > m$ \cdot $\{hk0\}$ or $m-H$, as (211) $2-2$

Brftchypyramids(Aii) $7^* < A > d > h > m$ \cdot $\{hk0\}$ or m \cdot $\{hk0\}$ or m , as (121) $2-^$



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Id f-Deral, IIS defined a

I form wbose (hcm are parallel to two of il pluQe; u prim is one whosf fiiCiB an: pHrahbl tii tLe verlicul uxin. Iiul iilercucl llie iwo lateral axes; udome* is one whuBc faces niv jmrallal 10 t>ue uf ilie la'eml Bies, but latersect lu the vertical axis. A dome is suiiihtiueBCHlledii Aortcontu^ pritm: a iiyroniid is a form whose facea mtei ul) tilt iLitv «xl-b.

These lenns are used in the aboTe seuKe Dot ously In Uie oriliui'']oiuib:c system, but al-o

111 ibe tuuDOcliuic iLud triclinic syaiema ; In iLe last encl funii rctisist ui iwu pliines oo.y.

167. *Pinacoid*.[^]—The macropinacoid includes two faces, each of which is parallel both to the macro-axis *i* and to the vertical axis *d*; their symbols are respectively 100 and 100 . This form is uniformly designated by the letter *a*, and is conventionally and briefly called the *a-pinacoid*.

The *irachypinacoid* includes two faces, each of which is parallel both to the brachy-axis *a* and to the vertical axis *i*; they have the symbols 010 and 010 . This form is designated by the letter *b*; it is called the *b-pinacoid*.

The base or basal pinacoid includes the two faces parallel to the plane of the lateral axes, and having the symbols 001 and 001 . This form is designated by the letter *c*; it is called the *c-pinacoid*.

Each one of these three pinacoids is an open-form, but together they make the so-called diametral prism, shown in Fig. 275, a solid which is the analogue of the cube of the isometric system. Geometrically it cannot be distinguished from the cube, but it differs in having the symmetry unlike in the three axial directions; practically this may be shown by the unlike physical character of the faces, *a*, *b*, *c*, for example as to luster, striations, etc.; or, again, by the cleavage. Further, it is proved at once by optical properties. This diametral prism, as just stated, has three pairs of unlike faces. It has three kinds of edges, four in each set, parallel respectively to the axes *a*, *i*, and *d*; it has, further, eight similar solid angles. In Fig. 275 the dimensions are arbitrarily made to correspond to the relative lengths of the axes, but the student will understand, that a crystal of this shape gives no suggestions as to these values.

188, *Prism*.—The *prism* proper includes those forms whose faces are parallel to the vertical axis, while they intersect both the lateral axes; their general symbol is, therefore, $\{hk0\}$. These all belong to one type of rhombic prism, in which the interfacial angles corresponding to the two unlike vertical edges have different values.

The unit prism, (110) , is that form whose faces intersect the lateral axes in lengths having a ratio corresponding to the accepted axial ratio of $c:h$ for the

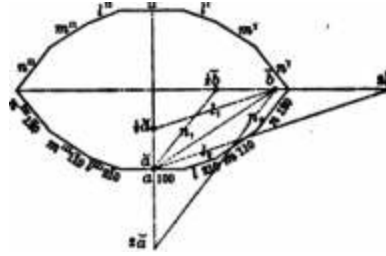
various species; in other words, the angle of this unit prism fixes the relative lengths of the lateral axes. This form is shown in combination with the basal

* From the Latin *domus*, because it resembles the root of a house; cf. Flg. STV, 280.

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pinacoid in Fig. 276; it is uniformly designated by the letter *m*. The two faces of the unit prism have the symbols 110 , 110 , 110 , 110 .

The macTOpTisms lie between the macropinacoid, a (100), and the unit prism m (110), and consequently for them the ratio of h to k is greater than 1:1; in other words, the ratio of the intercepts on the axes i and k is greater than that for the unit prism. Common forms have the symbols (410), (310), (210), (320), (430), etc., given in order from 100 toward 110; cf. the spherical projection of Fig. 3u3, The face / of Fig. 277 is the macroprism (310); for this form the axial intercepts (see the basal projection, Fig. 278) are in the ratio of $\sqrt{3} : 1$, or $1 : \sqrt{3}$; a similar relation holds for the other forms (410), etc.



The bracht/prisms lie between the unit prism and the brachypinacoid b (010), and consequently for them the ratio of the first two indices is less than 1:1, or the ratio of the intercepts on B, & is less than that of the unit prism. Common forms are (340), (230), (120), (130), given in order from 110 toward 010. For the form « (120), shown in Figs. 377, 278, the axial intercepts are in the ratio of $1/\sqrt{3} : 1$, or $2d:l$. Other examples of these prisms are given later (see Figs. 296-299).

All NanmniD's Bjmbola the Dumber n, the multiple of the lateral tula, is always made greater [biiD unity. Hence while the uHcruprbnj, t, of Fig. S!77 has the symbol $i:ib:-at.$ or $btiedy$ oo P5 (or $i-2$), the brnchypriaiii is irrillen & i ; 6 ; a)4, or-el's (or *-g), inslenl of the eqiiivHltHt form a: \h: aai. In i)tbi.-r wiirds, wiiL lUu macruuinrms (and macTopyramida) ike value of the bi-acby axis U made equal to unity, while witL tUe brachiy* prisms (and brtieliypyramids) the macro-aiaa is taken as Ibe unit.

169. MacTodomei, Bnohydomet.—The tnacrodomes are forms whose faces are parallel to the macro-axis, J, while they intersect the vertical axis d & the lateral axis a; hence the general symbol is (kOl). The angle of the unit macrodome, (101), fixes the ratio of the axes (i:d. This form is shown in Fig. 279 combined (since it is an open form) with the brachypinacoid.

In the macrodome zone between the base c (001) and the macropinacoid a (100) there may be a large number of macrodomes having the symbols, taken in the order named, (103), (10i), (303), (101), (302), (201), (301), etc Cf. Figs. 298 and 302 described later.

The brachydomes are forms whose faces are parallel to the brachy-axis, S, while they intersect the other axes i and h\ their general symbol is [OkI). The angle of the unit brachydome, (0il), which is shown with a (100) in Fig. SHO, determines the ratio of the axes h:i.

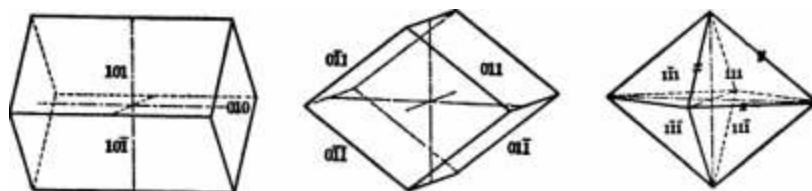
The brachydome zone between c(001) and 4(010) includes the forms (013), (012), (023),

(Oil), (032), (021), (031), etc. Cf. Figs. 298 and 302.

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Both sets of domes are often epokea of as horizontal prigma. The propriety of tliis expresBion is obvious, since th^ are in fact prisms in geometrical form; further, the choice of position for the axes which makes them domes, iast^d of prisms ic the narrower sense, is more or less arbitrary, as already explained elsewherefl.



170. Pyramids. —The pyramids in this system all belong to one type, the double rhovbicpyramid, bounded by eight faces, each a scalene triangle. This form has three kinds of edges, X, Y, Z (Fig. 281; cf. also Fig. 290), each set with a different interfacial angle; two of these angles suffice to determine the axial ratio. The symbol for this, the general form for the system, is [hkl].

The pyramids fall into three groups corresponding respectively to the three prisma just described, namely, unit pyramids, macro pyramids, and bracby pyramids.

The unit pyramids are characterized by tlie fact that their intercepts on the lateral ases have the same ratio as those of the nnit prism; that is, the assumed axial ratio {a:h} for the given species. For them, therefore, the general symbol becomes {hhl}.

For different nnit pyramids on crystals of the same species the vertical axes may have diiferent lengths bearing usually some simple numerical ratio to each other (and always commensurate), and these form a zone of faces lying between the base c (001) and tlie unit prism m (110). This zone, for example as shown in the basal projection of a sulphur crystal given in Fig. 302, includes the forms V^ (119), m (117), / (115), o (114). 5 (H3), y (112), ju (1U).

Cf. als

I Fig. 66, p. 30, of the same species, and the spherical projection, """""" I of this zone /* — k, and the

Fig. 303. In the symbol of all of the

lengths of the vertical axes are hence, in the example given, \, \, -J, \, j, 1 of the vertical axis i of the pyramid ;i (Fig. 290), which in this species ii taken as the unit pyramid. The

axial ratio for snliur is given on p. 22.

The macropyrnids and hrnchypyrniiil" are related to each other and to the unit pyramids, as were the mncroprisms and brachyprisms to themselves and to the unit prism. Further, each vertical zone of niacropyramids (or brachypvramids), having a common ratio for tiie lateral axes (or of A: it in the symtiol), belongs to a particular macroprism (or brachyprism) characterized by the same ratio. Thus the macropyramids (314), (213), (212), (421), etc, all belong in a common vertical zone between the base (001) and the prism (210). Similarly the brachypyramids (123), (122), (121), (241), etc., fall in a common vertical zone between (001) and (120). Cf. Fig. 299, where « and o are the brachy pyramids (134), (131), falling in the same vertical zone as the brachyprism (f(130)). See also the basal projection. Fig. 303, and the cpfaenal projection, Fig. 303, both of sulphur, noting the relation of the

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macropyramid (315) to the macroprism (310) and tlie brachypynunids (135),

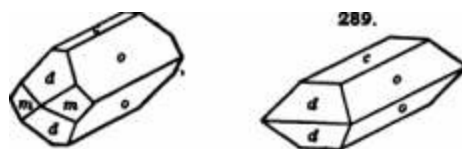
(133), (131) to the brachypriBm (130).

171. Illastiationi.~The following figures of bai-ite (282-289) give excellent illustrstiona of cryBtala of a typical orthorhombic Bpeciea, aud show also bow the habit of one and the same species may vary. The axial ratio for this species is gjea on p. 96. Here d is the macrodome (103) aud o the unit brachydome (911); m is, as always, the anit prism (110). Figs. 2SS-2S5 and

383. 3B3. flS4. 28B.

d

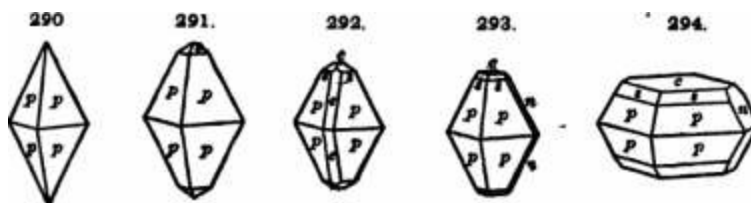
^^



Bariie Cryslalg.

28T are described as tabular \ c; Fig. 286 is priamatic in habit in the direction of the macro axis (5), and 288,289 prismatic in that of the brachy-axis (di). Figs. 290-394 of native sulphur show a aeries of crystals of pyramidt^ habit with tbe unit domes e (101), w

(Oil), and the unit pyramids p (111), «(II3). Note that e and n truncate respectively the terminal Ridges of the fundamental pyramid p . In general it should be remembered that a macro-dome truncating the edge of a pyramid must have the same ratio of $A: I$; thns, (\blacksquare ;ui) truncates the edge of (221), etc. Similariy of the brachydomes: (021) truncates the edge of (221), etc. Cf. Figs. 303 and 303.



Sulphur Cry stale.

Again, Fig. 295, of staurolite, shows the pinacoids ft (010), c (001), the nnit prism m (110), and the unit macrodome r (101).

Figs. 396-298 are prismatic cryBtals of topaz. Here m is the unit prism (110); I and n are the brachypriama (120), (140); d and p are the macroomes (301) and (401); l' and y are the brachydomes (021) and (041); i , u , and o are the nnit pyramids (223), (111), (221).

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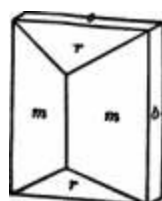
CBY 8TA LLOOaA PHT.

In Fi^{\wedge} . 399, of iolite, s and r are the unit pyramids (113), (111); d is the brachypnsm (130), and u , o are the correspond!og brachypyramids (134), (131). Fig. 300, of brookite, simulates a tetragonal crystal since the prismatic $ftDg$ le is not very far from 90° ; here $t = (113)$. In Fig. 301 of the same species, e is the brachypyramid (133); this crystal closely resembles an hexagonal pyramid with ite axis placed horizontal since the angles me (110 A 122) and ee' (123 A 132) are approximately equal.

396.

387.

298.



Staurolite.

299.

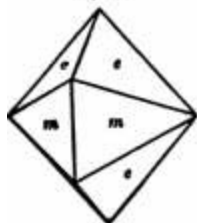
f^

$k^{\wedge\wedge}$

$\wedge\wedge\wedge f^{\wedge\epsilon^{\wedge}}$



301.



Brouklte.

Brook iie.

172. Fr[^]jectioiu.— Basal and spherical projections of a typical ortho-302. rhombic species nave already been given in Figs.

63 and 65 op pp. 27, 28. The subject ia so important, however, that others are given here (Figs. 302, 303) for the species eilphiir, cf. Figs. 290-294, also Fig. 66, p. 30. In Fig. 303 besides the pinacoids a (100), d (010), c(O0l), the positions of the prisms* (310), w(110), (1301 are shown; the macrodornes m (103), e(10l) and the brachydomes v (013), n (Oil); the remarkable zone of unit pyramids i//{n5), co (117), t (115), o(114), s (113), y (112), J (111); finally the macropyramid /S (315) and the braehvpyramids z (135) and x (133). Both projections exhibit clearly the symmetry

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characteristic of the groap; the promiuent zones, already spokeo of, sbonld 1[^]80 be noted.

2. HEMIMORPHIC GROUP (36). CALAMINE TYPE.

173. Symmetry and Typical Fomu.—The forms of the nrthorliombic-hemi-morpki[^] rotip are characterized by two unlike planes of symmetry and one aais of binfiry symmetry, the line in which they [^]304.

intersect; there ia no cenber of Byrametry. The " ,[^]

forms are therefore hemimorphic, as defined in Art. 29. For example, if, as is usually the case, x^m is

the vertical axis is made the axis of symmetry, the two planes of symmetry are parallel to the pinacoid, (100) and (010) . The prisms are then,

geometrically like those of the normal prisms, as Mf

are also the macropinacoid and brachypinacoid; but the two basal planes become two independent forms, (001) and $f(K)T$. There are also two macro-domes, (101) and $(10\bar{1})$, or in general (101) and $(10\bar{1})$

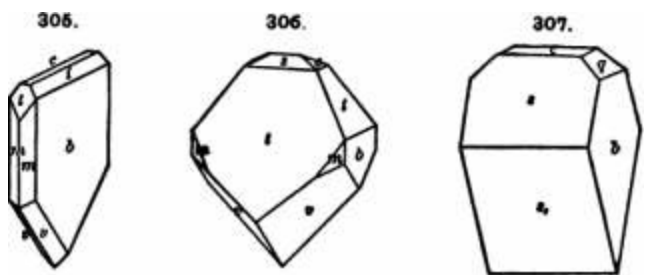
and similarly two sets, for a given symbol) of brachydomes and pyramids.

The general symmetry of the group is shown in the spherical projection of Fig. 304. Further, Figs. 305, 306, of calamine, and 307, of strotite, represent

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typical crystals of this group. In Figs. 305, 306 the forms present are I (301) , a (101) , i (031) , e (011) , v (111) ; in Fig. 307 they are a (101) , s (111) , g (011) .



3. SPHENOIDAL GROUP (27). EPSOMITE TYPE.

174. Symmetry and Typical Form.—The forms of the remaining group (of the hexagonal, the orthorhombic-sphenoidal group, are characterized by three unlike rectangular axes of binary symmetry, but they have no plane and no center of symmetry (Fig. 308). The general form hkl here has four faces only, and the corresponding solid is a rhombic sphenoid, analo.

dal group, are characterized by three unlike rectangular axes of binary symmetry, but they have no plane and no center of symmetry (Fig. 308). The general form hkl here has four faces only, and the corresponding solid is a rhombic sphenoid, analo.

h^3

$\hat{=} \hat{k} /$

gouBtothe sphenoid of thcteCragonal \

system. The complementary plus \

and minus sphenoids are enantio- \ O

tnorphous. Fig. 309 represents a \ ^

typical crystal, of epsomite, with the * ^

plus sphenoid, £ (111). Other crystals

of this species often show both plus and minus complementary forma, bnt

usually unequally developed.

Hatbemtical Relations or thb Orthobhoubic Stbtbu.

171. CholM of AXM.— Ab explaiiieri tri Art. IU, tliu three cryHtnllograpliic axes are Bied as regFirilBdlrcciiiiii in nil oi ^liiorlionible crysiaU. but anyone of tLera maybe mucie Ibe TerlUal axia. i; and nt Ihe two lateral iiJtes. wbicli is Ihe longer (S) nnil wbirli Ibe shorter (,) eajiLint be (ieleriiiiin.—(l until it is decided which faces lo iiflsume i* the fundameulBl. or unit, pyramid, prism, or domes.

The chuiice is generally so made, in a Riven case, as to best bring out Ihe relation of the cryalala of the species in baud to others allied lo them in form or In chemical comnosition or Iq both respects; or, so na lo make the cleavage parallel lo the fundamentHl form- or' u augijesled by the common habit of the crysials, or other cousideralioua

17*. Axial and Anpilar BlMuanU—Tbe aaiat eltmenU are given by the ratio of the lengths of the three axes Id terms of the macro-axis, I. as unity. For example, with barite the axial ratio ia

d:J:i = 0-81520:1 :1-81306.

100 A 110 = 39° 11' 18", 001 A 101 = 68' 10' 86", 001 a Oil = 88' «" «\

;;CoogIC

Two of these angles obviously determiuc the tbinl angle as well as the aztal ratio. The degree of accuracy to be attempted iu the slHltiment oF the hiililI ratio depeoda upon the ctisracleT of the fuudmneiitul measure men Is from which this ratio has been deducetl, Tbei-e fs no good reawu for glviug the values of d bdiL^ to mauy decimal placeB if the probable error of the measurements amounls 10 mauy mii'ieB. lu the above case the

measurements (by Uelinhacker) are supposed to be accurate within a few seconds. It is, however, to have the angular elements correct, they, within 10". So that the calculated angles obtained from the measurements will not vary from those derived direct from the measured angles by more than 30" 1» 1'.

$$\tan(\angle(100 \text{ A } 010) = (\dots, \tau(001 \text{ A } 011)) =$$

$$D(001 \text{ A } 101) =$$

These equations serve to give either the axes from the angular elements, or the angular elements from the axes. It will be noted that the axes are not needed for simple purposes of calculation, but it is still important to have them, for example to use in comparing the morphological relations of allied species.

In practice it is easy to pass from the measured angles, assumed as the basis of calculation (or deduced from the observations by the method of least squares), to the angular elements, or from either to any other angles by the application of the tangent principle (Art. 47) to the pinacoidal zone, and by the solution of the right-angled spherical triangles given on the sphere of projection.

Thus any face hkl (see p. 55) lies in the three pinacoidal zones, 100 and 010 and $h0l$, 001 and hkl . For example, the position of the face, 821 , is fixed if the positions of two of the poles, 101 , 021 , 820 , are known. These last are given, respectively, by the equations

$$\tan(\angle(001 \text{ A } 101)) = 8 \times \tan(\angle(001 \text{ A } 101)). \quad \tan(\angle(001 \text{ A } 021)) = 8 \times \tan(\angle(001 \text{ A } 021)), \quad \tan(\angle(100 \text{ A } 820)) = 1 \times \tan(\angle(100 \text{ A } 110)).$$

178. Sample — Fig. 810 represents a crystal of sillite from Japan, with the faces $j)(kl)$, $r(343)$, $7(858)$, etc. On the following measured angles were taken as (fundamental:

$$\angle(100 \text{ A } 858) = 55^\circ 1' 0''$$

Hence, the angles $\angle(858 \text{ A } 010) = 40' 10''$ and $\angle(858 \text{ A } 058) = 27^\circ 80' 1''$ are known without calculation. The right-angled spherical triangle $01001S$ yields the angle $(010 \text{ A } 053)$ and hence $(001 \text{ A } 053)$; also the angle at 010 , which is equal to $(001 \text{ A } 101)$. But $\tan\{001 \text{ A } 010\} = \frac{1}{8} \times \tan(001 \text{ A } 058)$, and $\tan(001 \text{ A } 010) = \frac{1}{8}$. Also since $\tan(001 \text{ A } 101) = \frac{1}{8}$, the axial ratio U thus known, and two of the angular elements.

The third angular element $(001 \text{ A } 101)$ can be calculated independently, for the angle at 001 in the triangle 001058 is equal to $(010 \text{ A } 050)$ and $\tan(\angle(010 \text{ A } 850)) \times (\dots) = (\angle(010 \text{ A } 110))$, the complement of $(100 \text{ A } 110)$.

Then since $\tan(\angle(100 \text{ A } 110)) = a$, this can be used to check the value of d already obtained. The further use of the tangent principle with the occasional solution of a right-angled spherical triangle will serve to give any desired angle from either the fundamental angles direct, or

from the angular elements.

Again, the symbol of any unknown face can be readily calculated if two measured angles of tolerable accuracy are at hand. For example, for the face m , suppose the measured angles to be

$\angle(010 \text{ A } hkt) = 30^\circ 15'$, $\angle(hkl \text{ A } ftii) = 51^\circ 32'$. The solution of the triangle bat gives the angle $\angle(010 \text{ A } O'U) = 16^\circ 55' 20''$, and

$$\tan(\angle(001 \text{ A } Oii)) = \tan 78^\circ 84' 00''$$

$$\frac{\sin 16^\circ 55' 20''}{\sin 78^\circ 84' 00''} = \frac{\sin 83^\circ 8' 00''}{r}$$

$$r(\angle(001 \text{ A } Oil)) = \tan 46^\circ 80' 00''$$

" r

* The student in this as in every similar case should draw a spherical projection (not necessarily accurately constructed) to show. If only approximately, the relative position of the faces present.

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But the $\angle(010 \text{ A } k)$ must be measured and the number derived from \sin with $10:8$.

Again, the angle $\angle(001 \text{ A } iO)$ may now be calculated from the same triangle and the value $59^\circ 55'$ obtained. From this the \angle of A to f U derived since

$$\tan(\angle(001 \text{ A } ftOf)) = \tan 59^\circ 55' 00'' \cdot \frac{\sin 16^\circ 55' 20''}{\sin 78^\circ 84' 00''} \sim \tan 48^\circ 00' \sim r$$

This ratio U nearly equal to $6:8$, and the two values thus obtained give the symbol $D'10'8$. If, however, from the triangle $001'OAr$, the angle at 001 is calculated, the value $36^\circ 42'$ is obtained, which is also the angle $\angle(010 \text{ A } hko)$. From this the ratio $ft: t$ is deduced, since

$$\tan(\angle(010 \text{ A } 110)) = \tan 45^\circ 18' 00'' = k$$

$$\tan 1010 \text{ A } / \gg \ll \sim \tan 43^\circ 43' 00'' = k$$

The value of r is hence closely equal to 2 ; this combined with that A is obtained

$\frac{1}{r} = \frac{1}{2}$ gives the same symbol $6'10'8$.

The symbol being more than usually complex for fairly accurate measurements. Hence the symbol obtained is can best be judged by comparing the measured angle with those calculated from the symbol. For example. In the given case the calculated angle for $\sin iU$ are $\angle(010 \text{ A } 6-10-8) = 80^\circ 16'$, $\angle(B'108 \text{ A } fi'10'8) = 51^\circ 51'$.

The correctness of the value deduced is further established if it is found that the given face falls into prominent zones.

It will be understood further that the zonal Truiltons, explained on pp. 50, 80, play an important part in all calculations. For example, in Fig 310, if the symbol of r were unknown, it could be obtained from a single angle (as bi), slope for this zone $h = 1$.

170. Formula.—Although it is not often necessary to employ formulas in calculations, a few are added here for sake of completeness.

(I) For the distance between the pole of any face P (hkl) and the poles a, b, c, \dots have in general:

$$\cos Pa = \cos \{AW A_{100}\} = \dots$$

(•(AW A 010):

$$\cos V H - i'a'c' + Co^*$$

of faces (UQ and (pgf)t

$$\cos tpe' + Agi't' + tr a^*$$

$$\cos V \Delta AV''T''We' + i'a'J[pV + \dots + fa^*]$$

V. MONOCLINIC SYSTEM.

180. The Monoclinic System includes all the forms which are referred to three unequal axes, having one of their axial inclinations oblique.

The axes are designated as follows: the inclined or clinodiagonal axis is i ; the orthodiagonal axis is S , the vertical axis is d . The acute angle between the axes i and S is represented by the letter β ; the angles between a and i and S and i are right angles. See Fig. 312. Crystals are usually drawn with the axis d vertical and the axis i directed to the front and inclined downward.

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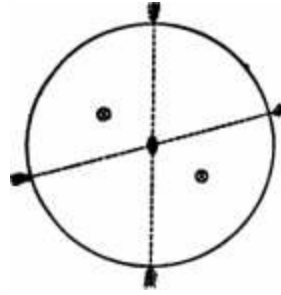
MONOCLINIC SYSTEM.

1. NORMAL GROUP (28). GYPSUM TYPE.

181. By symmetry in the normal group of the monoclinic system there is one plane of symmetry and one axis of binary symmetry normal to it. The plane of symmetry is always the plane of the axes a and i , and the axis of symmetry coincides with the axis S , normal to this plane. The position of one axis (β) and α that of the plane of the other two

axes (d and i) is thus fixed by the symmetry; but the latter axes may occupy different positions in this plane. Fig. 311 shows the typical spherical projection, projected on the plane of symmetry. Fig. 337 is the projection of an actual crystal of epidote; here, aa IB UBual, the plane of projection is normal to the prismatic zone.

182. FormB.—The Tarions-formB* belonging to this group, with their symbols, are given in the following table. As more particularly explained later, an orthodome includes two faces only, and a pyramid four only.



1. Orthopinacoid or

n-pinacoid

2, Clinopinacoid or]

A-pinacoid I

Miller. ..(100)

(010) 3. Base or c-pinacoid (001)

(Unit prism (HO)

4- j Orthoprisms (kkO) $h > k$

' Clinoprisms (AAO) $h < k$

5. Orthodomes....

6. Clinodomes

[Unit pyramids

((AOO

• \ (hOl)

...{OkI)

(ihk!)

\ (hhl)

J Orthopyramids. | [J*^*»t

Clinopyramids. jj*^J)A<i

•xi i : oio '. i

ni ;}> ita6

A : (tit -.mi d:'x,'b:.mi vid -.o -.mb

&'o : —md

d : «S : 'Vti

7tti -.h : -mi

Naumao. oo PS or i-l, a

oo P3) or i-i, b

OP or o, c

_ oo P_or /, m

« Pn or i-«, as (210) i-iS

oo PA ori'-n, as (120) t-2

—mPS or—fft-i,as(101) — l-i

mPm or m-i, as (101) l-j

mPco or m-l, as (011) l-i

— mP or -m, aa (111) — 1

mP, as (111) 1

—mPh or -m-n, as (211)-2-i5

«iPn or m-w, as (Sll) 2-S —mPh or -m-h, as (13I)-2-2 mPn or m-h, as (121) 2-3

The Nauinaon aymboli) jfiven nboTO are nnBlog-'iis to those of iLe ortliorhomhlc

Byslem. The long mark ;m|>loyed U lo be understwxl to be conveatlODal only and tu referring to the ortho-nxls. b. Il does Dot Imply tli:it iliU axis Is longer tban the clino-axis, d, ttiungh this U commonly the case. The Inclined mark rerers to the locllned axis, d. Willi aoDie authora iheae mnрка paw through tlie P, lusieail of being writieu over the letter (or Dumber) following.

- On the geeetal u»e of the terma pliiacold, rriamB, domes, pyramids, tee pp. 8d, 90,

..Coogic

CRTSTALLOSRAPHT.

1S3. PinacoidB.—The piuacoids are the orthopinacoid, clinopinacoid, and the baeal plane.

The orthopinacoid, (100), includes the two faces parallel to the plaue of the ortho-axis, t and the vertical axis A They have the aymbole 100 and 100. This form is designated by the letter a, Binoe it is situated at the extremity of the (i axis; it is hence conveniently called the a-piniicoid.

The clinopinacoid, (0101, includes the two faces parallel to the plane of symmetry; that is, the plane of the clino-axis a and to the axis t. i'hey have the symbols 010 and 010. The clinopinacoid is designated by the letter b, and is called the b-pinacotd.

The base or basal pinacoid, (001), includes the two terminal faces, abuve and below, parallel to the plane of the lateral axes a, t; they have the symbols 001 and 001. The base is designated by the letter c, and is often (»lled the opinacoid. It is obviously inclined to the orthopinacoid, and the normal an^le between the two faces (100 A 001) is the acute axial angle ff.

The diametral prism, formed by these three pinacoids, taken together. Fig. 312, is the analogue of the cube in the isometric eyatem. It is bounded by three sets of unlike faces; it has four similar vertical edges; also four lateral similar edges parallel to the axis a, but the remaining edges, parallel to the axis h, are only similar two-and-two. Of its eight solid angles there are two sets of four each; the two above in front are similar to those below behind, and the two below in front to those behind abovs

184. Friams.—The prisma are all of one type, the oblique rhombic prism, hey include the unit prism, (110), designated by the letter m, shown ia Fig. 313; also the orthoprisms, {hkO) where h >k, lying between «(100)

and »i(110), and the clinoprisms, {hhQ) where Ji<k, lying between m (110) and b (010). The orthoprisma and cHnoprisma correspond respectively to the macroprisms and brachyprisms of the ortborhombic system, and the explanation on p, 91 will hence make their relation clear. Common cases of these prisms are shown in the figures given later.

186. Orthodomei.—The four faces parallel to the ortho-axis, and meeting

the other two axes, fall into two sets of two each, having the general symbols (AOf) and (AO[^]). These forms are called orthodomes, they are strictly hemi-nrtbodomea. For example, the unit orthodome (101) has the faces 101 and 101̄; they would replace the two obtuse edges between a (100) and c (001) in Fig. 31a. The other unit orthodome (101̄) has the faces 101̄ and 101, and they would replace the acute edges between a and c. These two independent forms are shown together, with b (010), in Fig. 314.

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UOKOCLINIC STSTEU. 101

Similarly the faces 201, 301 belong to the form (201), and 201, 201̄ to the independent but complementary form (S01).

In the Symbols of Naumann's BDD, the lines between the base and the front ortho-axis above, hence corresponding to the obtuse edge between them, are designated by the minus sign ($-1\bar{i} = 10\bar{1}$, etc.), while those between the base and the ortho-pinacoid behind are designated plus* ($+1\bar{i} = i0\bar{1}$); the $-j$ -sign, however, is usually omitted. The two sets of hemi-pyramids (see beyond) are distinguished.

186. Clinodomei. — The clinodomes are the forms whose faces are parallel to the inclined axis, a, while intersecting the other two axes. Their general symbol is hence (Qkl) and they lie between the base (001) and the clinopinacoid (010). Each form has four faces; thus for the unit clinodome these are the symbols. Oil, Oil, Oil, Oil. The form n (031) in Fig. 321 is a clinodome.

187. Pyramids.—The pyramids in the monoclinic system are all hemi-pyramids, embracing four faces only in each form, corresponding to the general symbol (hkl). This obviously follows from the symmetry; it is shown, for example, in the fact already stated that the solid angles of the diametral prism (Fig. 312, see above), which are replaced by these pyramids, fall into two sets of four each. Thus any general symbol, as (321), includes the two independent forms (321) and (31̄1) with the faces

321, 351,

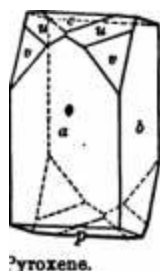
521,

351.

and 321,

321,

The pyramids ma^{\wedge} be unit pyramids, (hhl), orthopyramids, (kkl) when $A > A$, or clinopyramids, {hM) when $h < k$. These correspond respectively to the three prisms already named. They are analogous also to the unit pyramids, macropyramids and brachypyramids of the orthorhombic system, and the explanation given on pp. 91, 92 should serve to make their relations clear. But it must be remembered that each general symbol embraces two forms, (hhl) and {hkl) with four faces each, as above explained.

4[^]

^^?K

"Mii[^]

188. Distributions.—Figs. 315-318 of pyroxene {[^]:h:i= 1.092:1.0589, /3[^] 74° = ac) show typical monoclinic forms. Fig. 315 shows the diametral

•This choice of signs by Naumann was unfortunate, being contrary to ordinary usage; It is, however, too generally accepted to admit of being reversed. He was led to adopt it because the internal angle of the upper front edge between 001 and 100 is obtuse and hence the cosine {«.g. in the general cosine formula for the angle between two faces) is negative.

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prism. Of the other forms, m is the unit prism (110); p (101) is an ortho-dome; u (111), V (221), 3 (111) are unit pyramids; for other figures see p. 387. Again, Figs. 319-321 represent common crystals of orthoclase (d:i:i = 0.659:1.055, /J = 64° = ac). Here « (130) is a clinoprism; f(101) and y (301) are orthodomes; n(031) is a clinodome; o(111) a unit pyramid. Since (Fig. 319) c and x happen to make nearly equal angles with the vertical edge of the prism m, the combination often stimulates an orthorhombic crystal.

Si

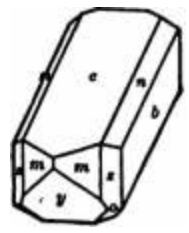
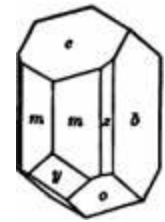
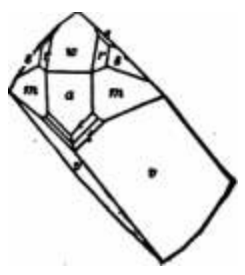


Fig. 333 shows a monoclinic crystal, epidote, prismatic in the direction of the ortho-axis; the forms are r (101) and n (111). Fig. 334 of gypsum is flattened | b; it shows the unit pyramid / (111) with the unit prism vi (110). Fig. 334 of monazite is prismatic in habit by extension of the pyramid v (111). It shows also the orthodome w(101); the clinodome «(011); the pyramiae r (111), «(121), a (311), t (311).

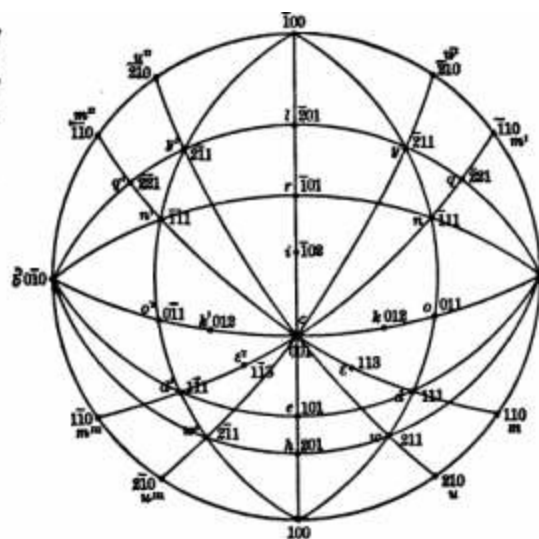
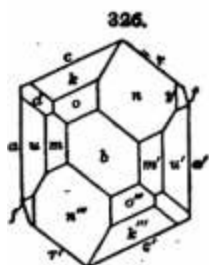
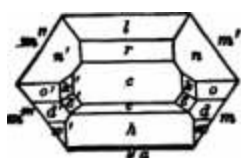
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189. FmJMtioiu. —Fig. 325 shows a projection of a crystal of epidota (cf. Fig. 853, p. 43S) on a plane normal to the prismatic zone, and Fig. 326 one of a similar crystal on a plane parallel to b (010); both should be carefully studied, as also the spherical projection of the same species. Fig, 327. The* symbols of the prominent faces are given in Fig. 327.

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MONOCLINIO ST8TBH.

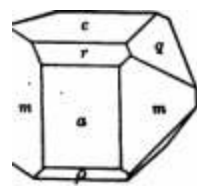


2. HEMIMORPHIC GROUP (29).

190. The monodinic-hemimorphic group is characterized by a single axis of binary Bymmetry, the CTystallographic axis %, but it has do plane of syni-metry. It is illuatrated by the spherical projection (Fig. 338); rIbo by Fig. 329, a commoQ form of tartaric acid; sugar crystals also belong here.

336.

/'
\
J



Tarlaric Acid. Forms: f (101), fl(101), 5(011),

The hemimorphio character is distinctly shown in the distribution of the clinodomee and pyramids; corresponding to this the artificial salts belonging here often exhibit marked pyro-electricaj phenomena.

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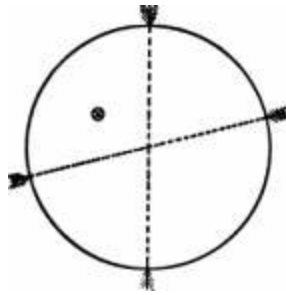
3. CLINOHEDRAL GROUP (30j. CLINOHEDRITE TYPE.

191. The monoclinic-elinokedrai group (the domatic oiasa of Grotb) is characterized by a single plane of symmetrj, parallel to the clino-piiiacoid, b (010), but it has no axis o(Bynimetry (cf. Fig. 330). In this group, ttierfore, the forms parallel to the S-axis, viz., c

(001), a (100) and the orthoiloines, are represented by a single face only. The other forms hare each two faces, bot it is to be noted that, with the single exception of the clino-pinacoid b (010), the faces of a given form are never parallel to another. The name given to the group is based on this fact.

Several artificial salts belong hero in their crystallization, but the only ktiowti representative among minerals is the rare silicate, clinohedrite (Il.CaZnSiO₂),* a complex crystal of which is shown in two positions in Figs. 331, 331«. As seen in those figures, the crystals of the group have a hemi-morphic aspect with respect to their development in the direction of the vortical axis, although they cannot properly be called hemimorphic since this is not an axis of symmetry. The forms shown in Figs. 331, 331a are iis follows: pinacoid, b (010); prisms, m (110), m, (!lo), h (320), n (120), I (130); orthodomes, fl (101), e, (101); pyramids, p (111), p, (III), q (III); r(53l), g, (551), t (?71), « (531), o (I31), X (i3l), y (i2l).

330. 831. 331a.



It is to be noted that crystals of the common species pyroxene {also of legirite and titanite) occasionally show this habit in the distribution of their faces, but it is not certain that this may not be accidental.f

Mathematical Rklatioss ov the Mosoclinic Sybtkii.

188. CholMof Aim.— Il is rept-nled here (Art, 181). thai tlie flued position of the plane of symmetry eHtablishes the direclion of the axis b Biid tlie plane of llie axes li aيتد i. The liKter axes, huwever, may hAve varying positions in tlii« plane ucLonling as lo which faces are rakeii as the pioacoids a auJ e, and which the unit pyi-amid. priam. or domes.

19S. AziU Md Anplu ElBmanti.—Tlie niaal eUtnenU ai-e llie UngtAt of ihe axes d and ^ In terms of the udU axis b. That is, the axial ratio, with also Uie acute angle of tnclluation of Ibe axes i and h, called /9, Thus for ortboclaBe the axial elements are : d ; S :<! = 0B685: 1 : 0-5554 /S = 68* 56|'.

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The aogukr eleroent are uaunlly taken as the angle (100 A 001) which Is equal to the aogle B: also the angles between \he three plnacolds 101), 010. Oul. reBpectiTely, aod the udU prism 110, the unit orthodome (191 or 101) md the unit ciluodoDie Oil. Thus again

for clinoclase, the angular elements are :

$$\alpha_{100} = 68^\circ 56' . \alpha_{100} = 80^\circ 86' .$$

$$\alpha_{100} = 50^\circ \text{ ISf}, \alpha_{001} = 36^\circ 31' .$$

Isl. The mathematical relations connecting axial and angular elements are as follows:

$$\alpha_{100}, (\alpha_{100} \text{ A } 110) \text{ , , , } (\alpha_{100} \text{ MIO}) = \langle \langle , * \langle ; \rangle \rangle ; (\langle \langle$$

AafS

$$\alpha_{100} \text{ I. } (\alpha_{100} \text{ A Oil}) \text{ , , , } (\alpha_{001} \text{ A Oil}) = \langle \dots \langle , ! ; (S)$$

Blu P

$$\alpha = d.H - \langle \rangle \text{ l7, . } u(UO \text{ lA } lOI) \text{ " } \bullet \text{ " } \langle \dots \text{ " } \text{ " } \text{ " } \text{ " } = a + i, 0, 11 - 1$$

$$\text{ , } \rangle . m (\alpha_{001} \text{ A } 101) \text{ , , , } \alpha \text{ . . } \text{ , , , } \text{ _ } \langle \langle l^\circ \rangle \text{ ' I } \text{ " } \text{ " } \text{ " } \text{ " }$$

These relations may be made more general by writing in the several cases—

iii(1) JAO for lO and -Idford; in (2) OW for Oil and *ifori;

In (8) AO for lOO and -ktoxt.

Also

$\alpha_{100} \sin(\alpha_{001} \text{ A } 101) \sim \sin(\alpha_{001} \text{ A } lOl) \sim \sin(100 \text{ A } lOi) \sim \sin(l''00 \text{ A } lOl)$ and more generally

$\alpha_{100} \sin(\alpha_{001} \text{ A } A OQ) \sim \sin(\alpha_{001} \text{ A } ^\circ Of) \sim \sin(100 \text{ A } m) \sim \sin(100 \text{ A } AOi)$ Note also that $\alpha_{100} \sin \alpha = (i \text{ and } -\tan Z = h$

where α is the angle (Fig. 82?) between the zones (001,100} and (001,110); also C the angle between (100, 001) and (100, Oil).

All the above relations are important and should be thoroughly understood.

19fi. The problems which usually arise have as their object either the deducing of the axial elements, the angle α and the values of l and h in terms of α ($=1$). from three measured angles, or the finding of any required interfacial angles from these elements or from the fundamental angles.

The simple relations of the preceding article connect the angular and axial elements, and beyond this all ordinary problems can be solved* either by the solution of spherical

triangles on the sphere of projection, or by the use of the cotangent relation.

It is to be noted, in the first place, that all great circles on the sphere of projection {Fig. 337} from 010 cut the zone circle 100, 001, 100 at right angles, but those from 100 cut the zone circles 010, 001, 010 obliquely, as also those from 001 cutting the zone circle 100, 010, 100.

196. Tangent and Cotangent Relations.—The simple tangent relation holds good for all zones to any pole on the zone circle 100, 001, 100; in other words, for the prisms, and zones of pyramids in which the ratio of $h : l$ is constant (from 001 to 100 or 100 to 001). Thus it is still true, as in the orthorhombic system, that the tangent of the angles

* The general formulas from which it is possible to calculate directly the angles between any face and the plane, or the angle between any two faces whatever, are so complex as to be of little value.

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of the prism 310, 110, 180 from 100 are in the ratio of $1 : 2^8$, or, more generally,

tant

tan(100 A 101) = cot(100 A 001) = h

tan(100 A 110) = cot(100 A 010) = h

Also for the clinodomes the tangents of the angles of 012, 011, 021 from 001 have the ratio of 1:1, etc. A similar relation holds for the angles of the faces of the prism. In the zone mentioned, as 131, 111, 312, etc.

For zones other than those mentioned in the preceding article, or from 100 to a clinodome, or from 001 to a prism, the more general cotangent formula given in Art. 47 must be employed. This relation is simplified for certain cases.

For any zone from 001, as the zone 001, 100, or 001, 110, or 001, 180, etc; if two faces are known, viz., the angles between 001 and these two faces in the given zone which fall (1) in the zone 010, 101, and (3) in the prismatic zone 010, 100; then the angle between 001 and any other face in the given zone can be calculated. Thus.

Let 001 A 101 = PQ and 001 A 100 = PR, or " 001 A 101 = PQ " 001 A 100 = PR, or " 001 A 110 = PQ " 001 A 100 = PR, etc.

Then for these, or any similar zone (as 201, or 221, or 421, etc.,

} Too. to 110, to SLO, etc., the expression has Um

$PQ = 001 A iOI$. $PR = 001 A iOO$. $PB = 001 A AOI$, or $001 A in$, etc.. $001 A UO$, etc., $001 a$
'M. etc

If, however, 100 is the starting point, and

$100 A 101 = PQ$, $100 A 001 = PR$, or $100 A 111 = PQ$. $100 A Oil = PR$. etc, then the relation becomes

$\cot P8 - \cot P R_ A \cot PQ - \cot PR''$ (■)

VI. TRICLINIC SYSTEM.

197. The Triclinic System includes all the forms which are referred to three unequal axes with all their intersections oblique.

The axes are here designated as in the orthorhombic system, the letters used for the lateral axes i, B (or d, b), having a short or long mark over them to indicate which is the shorter and which the longer axis. In the majority of cases, d is the brachy-axis and h the macro-axis. But this is not invariably true; thus with rhodonite the ratio of $d : S = 10TS : 1$. The vertical axis is always designated by i. The angle between the axes h and i is called a, that between a and i is o, and that between a and b is y (Fig. 333).

It is to be noted that there is no necessary relation between the values of a, >3 , and y, any one may be greater or less than 90° ; this is determined by the choice of the fundamental forms.

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-"-r

TRICLINIC SYSTEM. 107

1. NOEMAL GROUP (31). AXINITE TYPE.

188. Symmetry.—The normal group of the triclinic system is characterized by a center of symmetry, the point of intersection of the three axes, but there is no plane and no axis of symmetry.* This system

is shown in the accompanying spherical projection

(Fig. 332). ,---'---.,

199. Forms.—Each form of the group includes v^{\wedge}

' two faces, pitrallel to one another and symmetrical y with reference to the center of symmetry. This ia / true as well of the form with the general symbol / (

{hkl) as of one of the special forms, as, for example, { the R-pinacoid (100), \ .-"

Hence, as shown in the following table, the four *'\, prismatic faces 110, llo, llo, llo include two forma, \ namely, 110, llo, and 110, llo. The same ie true '-.,_

of the domes. Further, any eight corresponding pyramidal faces, as, for example. 111, 111, til, ill. Hi, III, III, III belong to four dietinct forms, namely, 111, III; TU, 111; Hi, III; ill. III, and similarly in general.

The TftriouB types of forma are given in the following table:

Hitler. Nnumana.

""piKd""} -ITM) •«-«-' «P[^].r«,«

Base or c-pinacold (001) 'x, tf: •n I: i OP or o, c

,... .^ j(110).!'» Hzl-.o[^]i <BP'oTr,m

Macroprisms) ({kkO) 3:nl -.coi oo /"« or I'-w'

'* > ^) t (hhO) a-.-nh-.'[^]i -KThoT 'i-n

Brachyprisma) j (AAO) n[^] : 5 : oo rf oo Pn or i-»

'* < *)] {/ikO) ni-.h-.'xi ccPnori-n

Macrodom,, 1 jS " i'[^]i"[^], .»'P'« or W

((hOl) —d:<x>h:mi m[^]P.-n or ,m.V

Brachydomes iLin » i i , t.

{(Okt) a:>d: — h:mi 'mPas or tn-l

{{hhl) a:b:mc mP' or «t'

(hhl) ~a:h:mc m[^]P or ,m

(hhl) -a-.-btmc mP, or m,

(hhl) a:-b:mc m'Por'm

{{hkl) S mh: m6 mP'n or w-w'

(hki) -£:nh:mi mJ'H or ,v>-n

(hkl) -S:-nh:m6 mP/i or m-n,

{Ji'kl) (i : -ni:mi m^Pn or '«i-5

* On llie »tber metbed nf viewing tlie symmelry here described, see An. 17. p. 10. \ In tlie dbove table It is awiimml tbat lIe ftxial raltii iaA:t:k. If h-tieTe a-l^:i. tbfr n:int<'>. Iiracby- aud macro- would be exchANGED, ami also tbe long nod short miirk!i in tbe NftkiTnann eymbols. The use of acceats Co distiuguisb priams, domes, and pyramids accord-iug to their poattton Is to be noted.

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CBT8IAU.oQBAPHT.

Brachypjr&midB)

1 (*«)

■ 1 (Mi)

I. ("o

m£ : S : rru!

: - b imi

- himi m'J'

mP'n or m-fi' tn^Fn or ,»»-«

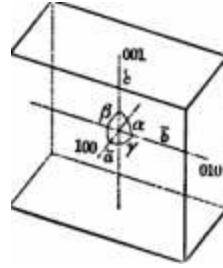
200. The explanations given under the two preceding sjatemB make it unnecessary to discusa in detail the various forma individually, except as 333. illustrated in the case of crystals belonging to cenain

typical triclinic species.

It may be mentioned, however, that Fig. 333 shows the diametral prism, which is bounded by three sets o(unlike faces, the pinacoids a, b, and c. This is the analogue of the cube of the isometric system, but here the like faces, edges, and solid angles include only a given face, edge aud angle, and that opposite to it.

201. Illntrationi. —A typical triclinic crystal is shown in Fig. 334 of ainite, already introduced. Here a (H;o) is the macropinacoid; m (110) and .¥(110) the two unit prisms; s {iO\} a macrodome, and x (111) and r (ill) two unit pyramids. The axial ratio is as

follows: $1 : 0.48$, $a = 82^\circ 54'$, $\beta = 91^\circ 52'$, $\gamma = 131^\circ 32'$. Figs. 335, 336 show two crystals of rhodonite, a species which is allied to pyroxene, and which approximates to it in angle and habit. Here the faces

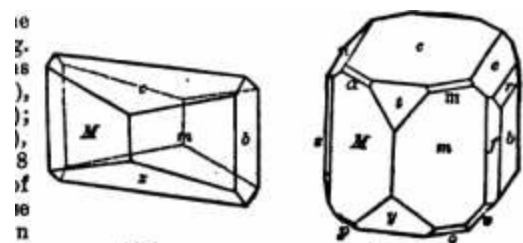


$= 0.49$;



are: Pinacoids a (100), b (010), c (001); prisms m (110), M (110); pyramids y (221), t (321), x (331), r (111).

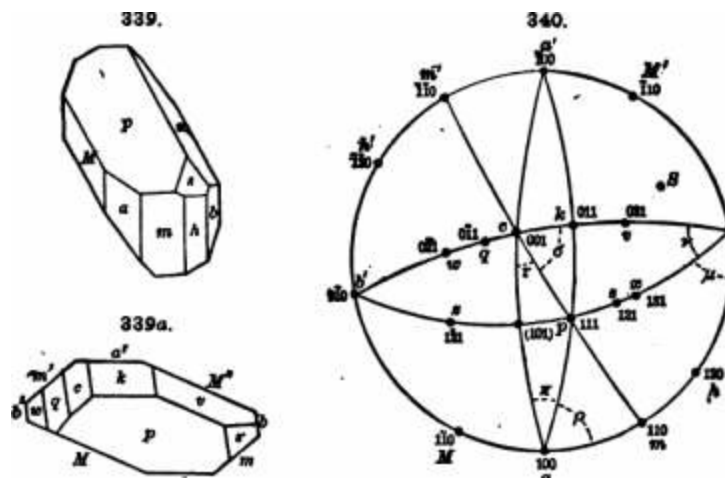
Further illustrations are given by Fig. 337 of albite and Fig. 338 of anorthite. The symbols of the faces, besides the pinacoids and the unit prisms, are as follows: Fig. 337, X (101); Fig. 338, prisms z (130), z (130); domes (307), y (201), e (021), r (061), M (051) pyramids m' (111), a (111), p (511). In Fig. 338 of anorthite the similarity of the crystal to one of orthoclase is evident on slight examination (cf. Figs. 319, 320), and careful study with the measurement of angles shows that the correspondence is very



TBIOLINIC BTBTEU.

close. Hence in this case the choice of the fundamental planes is readily made.

Fig. 339 represents a crystal (artificial) of blue vitriol, the mineral cobalt-



canthite; Fig. 339it gives a projection on a zone normal to the prisms, and Fig. 340 a spherical projection. The last figure also shows the STmhols of the different faces-

2. ASYMMETRIC GROUP (32).

202. Besides the normal group of the triclinic system there is another possible group, possessing symmetry neither with respect to a plane, axis nor center; in it a given form has oite face only. Tliis group, the asymmetric class of Groth, finds eicamples among a niimbpi' of artificial salts. One of these is calcium thiosiiiiphate (CaS,o,. 6H o); as yet no mineral species is known to be included here. This Is the most general of all the thirty-two types of forms classified according to their symmetry and comes first, therefore, if the groups are iiranged in order according to tbe degreeof symmetry characterizing them. This group is one of those whose crystals may show circular polarization. This ia true of eleven of the groups which have been described in the preceding pages.

Mathematical Rblationb op tbc TatcLtNio Stbtbh.

SOS. Olioim of AzM.—It U obvjouE, from wliat)ins beu said m to lbe symmetry of this Bystf^m. lljHt ang three fnces uf a trldinlc ciysial may be chostu as lhc piDocolds. or lhc faces wliicli lit the posiiion of lhc aiiiiil pliiiiies ntid tlie directions of tlie aiea; moreover, there U n like llberly tn tbe cboice of tbe unit prisma, domes or pyramids which further flz the lengths of the axes.

,ab,GoOgIc

J10 CETSTALIOGBAPHT.

Wbeo Uie C178IBI ia baod 1b ftliled Id form or compoeltlon to other Epecler sfUecIM of the stinio or (lifiereut ByHtisma, this fact bhiiplifles lhc problem unci makes tbe uculcu o1 .he fuudmealBl Tciraia eauf. 1'lils Is well iltuetraltl, as already uoted, by llie trlcliuic feldBiurB (e.g.. ulbiie iio.l aiiortliite. Figa. 837, 8B8j which are uear in Bugle to tbe allied mohocloic spedeB urtbucloM. KbodunUe (Figs. iiSS, iiStt), ihe iiiiicloic member of lbe pyroieuo gmu)], ia iuuuiber good example.

lu oHiii- caatB, wLere 110 aucb relationship eials, and wbre varied habit makes difleteni urifnlaiions plausible, Ibere is but little lo guide tlje choice. This ia illuslrated iJi lUe CB'jt! of axiuite ^Flg. »84), wbero at least tec dlsiinct posiiioDS have been assumed by diltereut auLiiora.

SM. Askl and Angular Elamanti.—Tbe aaini eUvuntt of a tridlnic crystnl are: (1) the axial t-atio, wliicb expresses tbe leugtbs of lbt axes a ood i in tenns of the tbiid axis. 6 ; And (2) The auglea between tbe axes a. ff, y {Fig. 838). There are here fl»e quaDtitles to be determined whicb obviously require the meaarenient of Ave ludepeudeul augles betweea The Isces.

The anguiav elemenli are usuallT taken as the anc «(iditioD, those between each piuauid aud the unit u pinacoldsi that is,

Ob. 100 A 010, ac. 100 A 001, be. 010 A 001; also

am 100 A 110. 001 A 101, 001 A Oil;

or, iostead, any one or all of these,

oJf. 100 A liO, 001 A iOl, 001 A Oil.

o1 these six angles taken, one Is delennined when the others are known.

SOS. The mHthematcal relations existing between tbe axial angles and axfa1 ratio, on th« one band, and tbe angles between the facet on the other, admit of being dnwii out with great completeness, but they are neces-iarly complex and in (irrend have litle practical value. In Tnct. most of tbe problems likely 10 arise ciuu be solved by means of ibe tilangles of tbe spherical projectioD. together wiib the cotangent foininia conuecMng four planes in the Mine zone (Art. 47. p. 311; tbiis will often lie laborious and may require some iogenully, but in general Involves no serious diffclulty.]n connection 'niib tbe u^e of the cotangent formubi. it is lo be noted IbKt in cerinbi commonly occurring cases its form la much 'almpllfled; some of these have already been eipbnned under ibe mouocllulc system <Art. 106). The formulas giv^ n tber ure of course equally applicnble here.

206. Tbe Brat problem may be to Qud lbe axiul elements from measured angles. Since these e'cments include five unknown quautilies. viz.. the three nxinl nnglesit, fl, y and tbe lengths of the axea a and i iu terms of b. five measured anglea are required, us ulieudy «taled.

These anglea, by use of three or more spherical trinngles. will serve to sive tbe angles (see Fig. 840) Jt, (>. fi, V, z, o-(or the coiTesponding angles n", o', etc., in tbe adjacent quadmuts). The ratio of tbe slues of each pair of Uiene angles Hies the rntlos of the corresponding axes (see'.below). The full axial ratio mity be obtained fmni any two pairs and the ibira ratio Mrres as a check upon tbe results given by the other two.

The simple formulas required are ;

$$\frac{\sin r}{\sin r'} = \frac{T \sin \alpha}{\sin \beta} \quad \frac{\sin \alpha}{\sin \beta} = \frac{\sin \gamma}{\sin \delta} \quad \frac{\sin \gamma}{\sin \delta} = \frac{\sin \epsilon}{\sin \zeta} \quad \frac{\sin \epsilon}{\sin \zeta} = \frac{\sin \eta}{\sin \theta} \quad \frac{\sin \eta}{\sin \theta} = \frac{\sin \iota}{\sin \kappa} \quad \frac{\sin \iota}{\sin \kappa} = \frac{\sin \lambda}{\sin \mu}$$

If the corresponding angles for the general case are given (not those of the unit zones), the relations are similar. That is, if for the face UI the corresponding angles be represented by $r, \alpha, \beta, \gamma, \delta, \epsilon, \zeta, \eta, \theta, \iota, \kappa, \lambda, \mu$ (where $\alpha, \beta, \gamma, \delta, \epsilon, \zeta, \eta, \theta, \iota, \kappa, \lambda, \mu$ are the angles between the zone circles 100, 001 and 100. 010 respectively and the zone circle 001. $\alpha, \beta, \gamma, \delta, \epsilon, \zeta, \eta, \theta, \iota, \kappa, \lambda, \mu$ (and similarly for $r, \alpha, \beta, \gamma, \delta, \epsilon, \zeta, \eta, \theta, \iota, \kappa, \lambda, \mu$ in the adjacent quadrant, also similarly $r, \alpha, \beta, \gamma, \delta, \epsilon, \zeta, \eta, \theta, \iota, \kappa, \lambda, \mu$. etc.), these relations may be expressed in the general form

$$\frac{\sin r}{\sin r'} = \frac{\sin \alpha}{\sin \alpha'}$$

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For the face 821 the formula become

It is also to be noted that

$$a = 180^\circ - \alpha, \quad \beta = 180^\circ - B, \quad \gamma = 180^\circ - C,$$

where α, β, γ are the angles in the planar triangle 100-010-001 at the poles respectively. That is,

$$A = 180^\circ - \alpha, \quad B = 180^\circ - \beta, \quad C = 180^\circ - \gamma;$$

$$B = 180^\circ - \beta, \quad C = 180^\circ - \gamma;$$

$$C = 180^\circ - \gamma, \quad A = 180^\circ - \alpha;$$

$$180^\circ - \alpha, \quad A = 180^\circ - \alpha, \quad B = 180^\circ - \beta, \quad C = 180^\circ - \gamma.$$

When given, by measurement or calculation, (the angles between the faces 100, 010, 001), and the angles α, β, γ are calculated and their supplements are the axial angles a, β, γ respectively.

Still another series of equations are those below, which give the relations of the angles $\alpha, \beta, \gamma, \delta, \epsilon, \zeta, \eta, \theta, \iota, \kappa, \lambda, \mu$, to the axes and axial angles. By means of them, with the sine formulas given above, the angular elements (and other angles) can be calculated from the axial elements.

These equations apply when $\alpha + \beta, \beta + \gamma, \gamma + \delta, \delta + \epsilon, \epsilon + \zeta, \zeta + \eta, \eta + \theta, \theta + \iota, \iota + \kappa, \kappa + \lambda, \lambda + \mu$ is less than 90° ; if their sum is greater than 90° the denominator is negative.

The following equations are also often useful; they give the relations between the

angles a , f , y , and the angles j , v , etc., already defined.

$$\sin 3 \text{ Bin } /> \text{rin fl}' \sim^* \sin (p - p')'$$

$$\sin (M - \text{Sstn rf})$$

$$a + M + p = f - i \cdot f_i + r = y + r + iT = 180^*.$$

The calculation from the angular elements or from the fundamental measured angles, either (1) of the position of any face whose symbol is given, or (2) of the symbol of an unknown face for which measured angles are at hand, requires no further explanation. The following formula is all that is needed in a single zone, and the solution of spherical triangles on the projection (with the use of the sine formulas) will suffice in addition in all ordinary cases.

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OBTSTALLOOBAPET.

MEASUREMENT OF THE ANGLES OF CRYSTALS.

208. Band-Goniometers.—The interfacial angles of crystals are measured by means of instruments which are called goniometers.

The simplest form is the hand-goniometer, represented in Fig. 343. It consists of an arc, graduated to half-degree or finer, and two movable arms. In the instrument figured, one of the arms, ab , has the motion forward

343.

and backward by means of slits gh , ik ; the other arm, cd , has also a similar slit, and in addition it turns around the center of the arc as an axis, the faces whose inclination is to be measured are applied between the arms ao , co , and the latter adjusted so that they and the surfaces are in close contact; further, the arms must be exactly at right angles to the intersection-edge. This adjustment must be made with care, and when the instrument is held up to the light none must pass through between the arm and the surface. The number of degrees read off on the arc between k and the left edge of d (this edge being in the line of the center, o , of the arc) is the angle required. The motion to and fro by means of the slits is for the sake of convenience in measuring small or embedded crystals. In a better form of the instrument the arms are wholly separated from the arc; and the arc is a delicately graduated circle to which the arms are adjusted after the measurement.

The hand-goniometer is useful in the case of large crystals and those whose faces are not well polished; the measurements with it, however, are seldom accurate within a quarter of a degree. In the finest specimens of crystals, where the faces are smooth and lustrous,

results far more accurate may be obtained by means of a different instrument, called the reflecting goniometer.

209. Reflecting Goniometer.—This instrument, devised by Wollaston (1809), has been much improved in its various parts since his time by Mitscherlich and others. The principle on which it is constructed may be understood by

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METHOD OF MEASURING THE ANGLES OF CRYSTALS 113

reference to the following figure (Fig. 343], which represents the section of a crystal, whose angle, abc , between the faces ah , be , is required.

The eye at P , looking at the face of the crystal, be , observes a reflected image of m , in the direction of Pn . The crystal may now be so changed in its position that the same image is seen reflected by the next face and in the same direction, Pn . To effect this, the crystal must be turned around, until abd has the present direction of be . The angle dOc measures, therefore, the number of degrees through which the crystal must be turned; it may be measured by attaching the crystal to a graduated circle, which turns with the crystal. This angle is the supplement of the interior angle between the two faces, or in other words is the normal angle, or angle between the two poles (see Art, 41, p. 28). The reflecting goniometer hence gives directly the angle needed on the system of IVfiller here followed.

Many different forms of reflecting goniometer of simple type are in use. The accompanying figure (Fig. 34-1) will suffice to make clear the general character of the instrument, as well as to exhibit some of the refinements added for the sake of greater exactness of measurement.

The circle, C , is graduated), in this case, to twenty minutes, and by means of the vernier at v the readings may be made to minutes and half-minutes. The crystal is attached by means of wax to the little plate at k ; this may be removed for convenience, but in its final position it is, as here, at the extremity of the axis of the instrument. This axis is moved by means of the wheel, m ; the graduated circle is moved by the wheel, m . These motions are so arranged that the motion of n is independent, its axis being within the other, while on the other hand the revolution of m moves both the circle and the axis to which the crystal is attached. This arrangement is essential for convenience in the use of the instrument, as will be seen in the course of the following explanation. The screws, c , d , a - e for the adjustment of the crystal, and the slides, a , b , serve to center it.

The method of procedure is briefly as follows: The crystal is attached by means of suitable wax at k , and adjusted by the hand so that the direction of the combination-edge of the two faces to be measured coincides with the axis of the instrument; the wheel, n , is

turned until an object {e.g., a window-bar) reflected in one face is seen to coincide with another object not reflected {e.g., a chalk-line on the floor); the position of the graduated circle is observed, and then both crystal and circle revolved together by means of the wheel, m, till the same reflected object now seen in the second face again coincides with the fixed object (that is, the chalk-line); the angle through which the circle has been moved, as read off by means of the vernier, is the normal angle between the two faces.

In order to secure accuracy, several conditions must be fulfilled, of which the following are the most important:

1. The position of the eye of the observer must remain perfectly stationary.
- 2- The object reflected and that with which it is brought in coincidence should be at an equal distance from the instrument, and this distance should not be too small.
3. The crystal must be accurately adjusted; this condition is satisfied when the line seen reflected in the case of each face and that seen directly with which it is in coincidence are horizontal and parallel. It can be true only when the intersection-edge of the two faces measured is exactly in the direction

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' 114 CBTSTALLOOBAPHT.

of the axis of the instrument, and perpendicular to the plane of the circle. The adjustment is accomplished roughly by the hand and accurately by the screws c and t.

4. The crystal must be centered as nearly as possible, or, in other words, the same intersection-edge must coincide with a line drawn through the

8*4.

revolving axis. This condition will be seen to be distinct from the preceding, which required only that the two directions should be the same. The error arising when this condition is not satisfied diminishes as the object reflected is removed farther from the instrument, and becomes zero if the object is at an infinite distance. In the centering of the crystal the slides a and b are employed.

The first and second conditions are both satisfactorily fulfilled by the use

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of a telescope, as t (Fig. 344), with slight magnifying power. This is arranged for parallel light, and provided with a hair-cross in its focus. It admits also of some adjustments, as seen in the figure, but when used it must be directed exactly toward the axis of the

goniometer. This telescope has also a little magnifying-glass (g) attached to it, which allows of the crystal itself being seen when mounted at k. This latter is used for the crystal, and then dipped aside, when some desired object which has not been selected must be seen in the field of the telescope as reflected, first by the one face and then by the other as the wheel h is revolved. When the final adjustments have been made so that in each case the object coincides with the center of the hair-cross of the telescope, and when further the edge to be measured has been centered, the crystal is ready for measurement.

This telescope, obviously, can be used only when the face is smooth and large enough to give distinct and brilliant reflections.* In many cases sufficient accuracy is obtained without it by the use of a window-bar and a white chalk-line on the floor below for the two objects; the instrument in this case is placed at the opposite end of the room, with its axis parallel to the window; the eye is brought very close to the crystal and held motionless during the measurement.

The best instruments are provided with two telescopes. The second stands opposite the telescope, i (see figure ^, the centers of both telescopes being in the same plane perpendicular to the axis of the instrument. This second telescope has also a hair-cross in the focus, and this when illuminated by a brilliant gas-burner (the rest of the instrument being protected from the light by a screen) will be reflected in the successive faces of the crystal. The reflected cross is brought in coincidence with the cross in the first telescope, first for one and then for the other face. As the lines are delicate, and as exact coincidence can take place only after perfect adjustment, it is evident that a high degree of accuracy is possible. Still more than before, however, are well-polished crystals required, so that in many cases the use of the ordinary double telescopes is impossible. Very often the hair-cross of the second telescope may be advantageously replaced by a bright hue or cross, the light shining through a cross cut in tin-foil (Schraiff), or as given by the analogous Weisky signal. This light-signal is visible in the first telescope even when the planes are extremely minute, or, on the other hand, somewhat rough and uneven; even if the image is not perfectly distinct, it may be sufficiently so to admit of fairly good measurements (e.g., within two or three minutes).

210. Horizontal Goniometer.—A form of reflecting goniometer well adapted for accurate measurements is shown in Fig. 34f. It is made on the Babinet type, with a horizontal graduated circle; the instruments of the Mitscherlich type, just described, having a vertical circle. The horizontal circle has many advantages, especially when it is desired to measure the angles of large crystals

* When planes are rough and destitute of luster the angles can often best be obtained by use of a candle-flame, the diffuse reflection of which in the given face indicates the place of more distinct images. For embedded crystals, and often in other cases, measurements may be advantageously made from impressions in some material, like sealing-wax. Angles thus obtained ought to be accurate within one degree, or even less, and suffice

for many purpose*. It is sometimes of aivniitpp to altacb to the planes to be monsiired. wliion quite rough, fraicment^ of ihtn glnw, from whicli reflections can be obtained; this must, however, be clone with cnre. to avoid coDsidrralile error. OccasI'mally dusting the surface with graplille makes a " sidmme'-' mensiiri'ini.'iit with the caudfe-flame possible, or, again, cuVL-rlog It wtlfa a thin film of gum arable.

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or those which are attached to the rock. Thiii particular form of instrameot hfere figured* is made hy R, Fuess in Berlin.

The instrnmeDt stands on a tripod with leveliug screws. The central axis, o, bae within it a hollow axis, b, with which turns the plate, d, carrying the verniers and also the observing telescope, the upright support of which is shown at B. Within A is a second hollow axis, e, which carries the graduated circle,/, above, and which is turned by theecrew-head, ^; the tangent screw, or, serves as a fine adjustment for the observing telescope, B, the screw, c, being for this purpose raised so as to bind b and e together. The tangent screw, p, is a fine adjustment for the graduated circle. Again, within e is the third axis, h, turned by the screw-head, t, and within h is the central rod, which carries the support for the crystal, with the adjusting and centering contrivances mentioned below. This rod can be raised or lowered by the screw, k.

So as to bring the crystal to the proper height—that is, up to the axis of the telescope; when this has been accomplished, the clamp at p, turned by a set-key, binds s to the axis, h. The movement of h can take place independently of g, but after the crystal is ready for measurement these two axes are bound together by the seI screw, /. The signal telescope is supported at C, firmly attached to one of the legs of the tripod. The crystal is mounted on the plate, w, with wax, the plate is clamped by the screw, v. The centering apparatus consists of two slides at right angles to each other (one of these is shown in the figure) and the screw, a, which works it; the end of the other corresponding screw is seen at a'. The ailjusling arrangement consists of two cylindrical sections, one of them, r, shown in the figure, the other is at r'; the cylinders have a common center.

* The figutv hem used la from the catalogue of Fueu.

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The circle is gradnated to degrees and quarter degi^eB, and the vernier gives the readinzs to 30", but by estimate they can be obtained to 10". The eignals provided are four in number, each in its own tube, to be inserted benind toe collimator lens; these are: (1) the ordinary telescope with the hair-cross, to be used in the case of the most perfect faces; (3) the commonly nsed signal, proposed by Websky, canBisting of two small opaque circles, whose distance apart can be adjusted by a screw between them; the light passing between these circles enters the tube in a form resembling a double concave lena; also (3) an

adjustable slit; and, finally, (4) a tube with a single round opening, very small. There are four observing telescopes of different angular breadth of field and magnifying power, and hence suitable for faces varying in size and in degree of polish.

The methods to be employed, both in making the preliminary adjustments required by every instrument before it can be used and in the actual measurement of the angle θ of crystals, have been described by Websky with a fullness and clearness which leave nothing to be desired, and reference must be made to his memoir,

211. Theodolite-Goniometer.—A form of goniometer having many practical advantages has two independent circles; it can be used in a manner analogous to that of the ordinary theodolite, as will appear below. Instruments of this type have been devised independently by Fedorow, Czapaki, and Goldschmidt. In addition to the usual graduated horizontal circle of Fig. 346, and the two accompanying telescopes, a second graduated circle is added which revolves in a plane at right angles to the first; to the latter the crystal to be measured is attached, with the addition of suitable adjusting and centering contrivances.

By this instrument, instead of the interfacial angles being measured directly, the position of each face is determined independently of others by the measurement of its angular coordinates. These coordinates are the angles (θ and ϕ of Goldschmidt) measured, respectively, in the vertical and horizontal circles from an assumed pole and meridian, which are fixed, in most cases, by the symmetry of the crystal. In practice the crystal is usually so mounted that the prismatic zone is perpendicular to the vertical circle and a pinacoid in this zone is the zero point. For example, with an orthorhombic crystal, for the face 111, the angle θ is equal to $010 \wedge 110$ and ϕ to $001 \wedge 111$ for the given species. Goldschmidt has shown that this instrument is directly applicable to the system of indices and methods of calculation and projection adopted by him, which admit of the deducing of the elements and symbols of a given crystal with a minimum of labor and calculation. Fedorow has also shown that this instrument, with the addition of the appliances devised by him, can be most conveniently used in the crystallographic and optical study of crystals.

•See Websky. *Zs. Kryst.*, 3, 341. 1879; 4, 15. 1880; also *Lfgblsch. Bericht Ober die Verhandlungen des Berliner Oewerbeausstellungskomitees in Jahre 1879*, pp. 88-83

Fedorow, Universal or Theodolite Goniometer. *Zs. Kryst.*, 21, 574, 1898; 23, 29, 1893; *Monatsh. f. Chem. u. Physik*, 1. 1898; Goldschmidt. *Zb. Kryst.*, 21, 310, 1893; 34, 810. 1895; 36, 821, 538. 1896. On the method of Goldschmidt. *we Psluche, Am. J. Be.* 3, 275. 1897. A simplified form of the theodolite-goniometer is described by Stobber, *Ze. Ervnt.*, 39, 3R, 1897.

The following is a list of the works on this subject for his work, *Krystallographische*

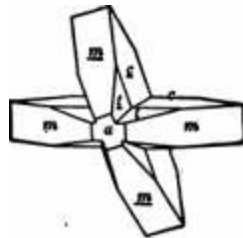
Winkelmann (1899 pp., Binn. ISHT). This gives the angles required by his system for
Miller species. See also Zs. Kryst., 39, 1898.

;Cooper

CRYSTALLOGRAPHY.

COMPOUND OR TWIN CRYSTALS.

218. Twin Crystals.—Twin crystals are those in which one or more parts regularly arranged
are in reverse position with reference to the other parts. They often appear
externally to consist of two or more crystals symmetrically united, and sometimes have
the form of a cross or star. They also exhibit the composition in the reversed arrangement
of part of the faces.



Thomson;

Crystallography.

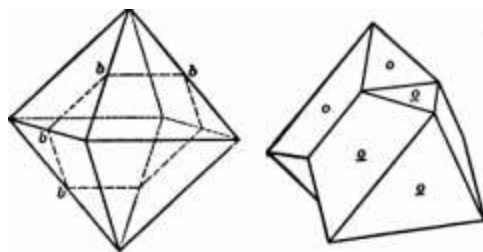
340.

in the strike of the surface, and in re-entering angles; in certain cases the compound
structure can only be entirely detected by an examination in polarized light. The above
figures (Figs. 346-348) are examples of typical kinds of twin crystals, and many others are
given on the pages following.

To illustrate the relation of the parts in a twin crystal, Figs. 349, 350 are given. Fig. 349
shows a regular

octahedron divided into two halves by a plane parallel to an octahedral face. If now the
lower half be supposed to be revolved 180° about an axis normal to this plane, the
twinned octahedron of Fig. 350 results. This is a common type of twin in the isometric
system, and the method here employed to describe the position of the parts of the crystal
to one another is applicable to nearly all twins. It is important to understand that
crystals, or parts of crystals, so grouped as to occupy parallel positions with reference to
each other—that is, those whose similar faces are parallel—are not called twins; the term
is applied only where the crystals or parts of them are united in their reversed position in
accordance with some deducible mathematical law. Thus Fig. 351, which represents a
cluster of partial crystals of analcite, is said to be a case of parallel grouping simply (see

Art. 231); but Fig. 369 illustrates twinning, and this is true of Fig. 378 also. Since though in these cases the axes remain parallel the similar faces (and planes of symmetry) are reversed in position.



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214. Twinning-axis. —The relative position of the parts of a twinned crystal can be best described as just explained, by reference to that line or axis called the twinning-axis, a revolution of 180° about

which would serve to bring the twinned part parallel to the other, or in other words, which would cause one of the parallel parts to take a twinned position relatively to the other.

The twinning-axis is always a possible crystalline line—that is, either a crystallographic axis or the normal to some possible face on the crystal, usually one of the common fundamental forms.

It is not to be supposed that ordinary twins have actually been formed by such a revolution of the parts of crystals, for all twins (except those of secondary

origin, see Art. 221) are the result of regular molecular

growth or enlargement, like that of the simple crystal. This reference to a revolution, and an axis of revolution, is only a convenient means of describing the forms.

In certain rare cases, particularly of certain pseudo-hexagonal species, a revolution of 60° or 120° about a normal to the base has been assumed to explain the complex group observed,

215. Twinning-plane. —The plane normal to the axis of revolution is called the twinning-plane. The axis and plane of twinning bear the same relation to both individuals in their reversed position; consequently, in the majority of cases the twinned crystals are symmetrical with reference to the twinning-plane.

The twinning-plane is, with rare exceptions, parallel to a possible occurring face on the given species, and usually one of the more frequent fundamental forms. The exceptions occur only in the triclinic and monoclinic systems, where the twinning-axis is sometimes one of the oblique crystallographic axes, and then the plane of twinning normal to it is

obviously not necessarily a crystallographic plane; this is conspicuously true in albite.

216. Composition-plane,—The plane by which the reversed crystals are united is the composition-plane. This and the twinning-plane very commonly coincide; this is true of the simple example given above (Fig. 350), where the plane about which the revolution may be conceived to take place (normal to the twinning-axis) and the plane by which the individuals are united are identical. When not coinciding, the two planes are generally at right angles to each other—that is, the composition-plane is parallel to the axis of revolution. Examples of this are given beyond. Still again, where the crystals are not regularly developed, and where they interpenetrate, the contact surface may be interrupted, or may be exceedingly irregular. In such cases the axis and plane of twinning have, as always, a definite position, but the composition-plane loses its significance.

Thus in quartz twine the interpenetrating parts have often no rectilinear boundary, but mitigate in the most irregular manner throughout the mass, showing this composite irregularity by abrupt variations in the character of the surfaces. This irregular internal structure, found in many quartz crystals, even the common kinds, is well brought out by means of polarized light; also by etching with hydrofluoric acid.

The composition-plane has sometimes a more definite signification than the twinning-plane. This is due to the fact that in many cases, whereas the former

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120 CRYSTALLOGRAPHY,

is fixed, the twinning-axis (and twinning-plane) may be exchanged for another line (and plane) at right angles to each, respectively; since a revolution about the second axis will also satisfy the conditions of producing the

required form. An example of this is furnished by Fig. 352, of orthoclase; the composition-plane is here fixed—namely, parallel to the crystal face, b (010). But the axis of revolution may be either (1) parallel to this face and normal to a (100), which is then commonly the twinning-plane, though the axis does not coincide with the crystallographic axis; or (2) the twinning-axis may be taken as coinciding with the vertical axis, and then the twinning-plane normal to it is not a crystallographic face. In other simpler cases, also, the same principle holds good, generally in consequence of the possible mutual interchange of the planes of twinning and composition. In most cases the true twinning-plane is evident, since it is parallel to some face on the crystal of simple mathematical ratio.

317. An interesting example of the possible choice between two twinning-axes at right angles to each other is furnished by the specimen in Fig. 400 above, a prismatic

twla . from FanDin Oo., Ga, The measured angle for U was $70^{\circ} 30'$. The twinning-axis deduced from this may be normal to the face (230), which would then be the twinning-plane. Or, instead of this axis, its complementary axis at right angles to it may be taken, which would equally well produce the observed form. Now in this species it happens that the faces, 180 and 230 (over 100), are almost exactly at right angles with each other, and, according to the latter supposition, 180 becomes the twinning-plane, and the axis of revolution is normal to it. Hence, either 230 or 180 may be (be twinning-planes, either supposition agreeing closely with the measured angle (which could not be obtained with great accuracy). The former method of twinning (tw. pi. 230) conforms to the other twins observed on the species, and hence it may be accepted. What is true in this case, however, is not always true, for it will seldom happen that of the two complementary axes each is so nearly normal to a face of the crystal. In most cases one of the two axes conforms to the law in being a normal to a possible face, and the other does not, and hence there is no doubt as to which face

Another interesting case is that furnished by columbite. The common twins of the species are similar to Fig. 347, p. 118, and have $\{021\}$ as the twinning-plane; but twins also occur like Fig. 404, p. 128, where the twinning-plane is $\{028\}$. The two faces, 031 and 023, are nearly at right angles to each other, but the measured angles are in this case sufficiently exact to prove that the two kinds cannot be referred to one and the same law.

218. Contact- and Penetration-twins.—In contact-twins, when normally formed, the two halves are simply connate, being united to each other by the composition-plane; they are illustrated by Figs. 347, 350, etc. In actually occurring crystals the two parts are seldom symmetrical, as demanded by theory, but one may preponderate to a greater or less extent over the other; in some cases only a small portion of the second individual in the reversed position may exist. Very great irregularities are observed in nature in this respect. Moreover, the re-entering angles are often obliterated by the abnormal developments of one or other of the parts, and often only an indistinct line on some of the faces marks the division between the two individuals.

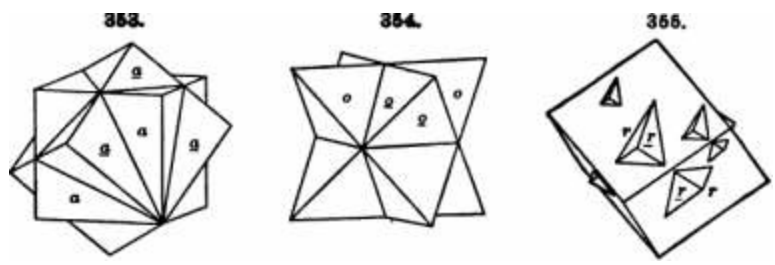
Penetration-twins are those in which two or more complete crystals interpenetrate, as if they were crossing through each other, formally, the crystals have a common center, which is the center of the axial system for both; practically, however, as in contact-twins, great irregularities occur.

Examples of twins of this second kind are given in the annexed figures. Fig. 353 of fluorite. Fig. 354 of tetrahedrite, and Fig. 355 of chabazite. Other examples occur in the pages following, as, for instance, of the species staurolite (Figs. 408-411), the crystals of which sometimes occur in nature with almost the perfect symmetry demanded by theory. It is obvious that the distinction

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COUPOL'ND Oil TWIN CRYSTALS.

between contact- and penetration-twins is not of great importance, and the line cannot always be clearly drawn between them.

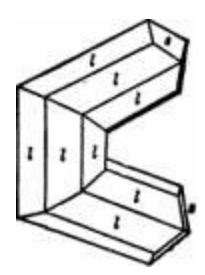


Fluorite.

Tetrahedrite.

Chabazite.

■IB. PangndB and MaUganis Tvtni.-



Butte.

—The distinction of pangndB and metagealc twins is not always so obvious.

The distinction is important.

In ordinary twins, the compound structure had its beginning in the

> nucleal compound molecule, or was compound in its very origin; and whatever
mechanisms in the result, these are only irregularities in the development from such a
nucleus. But in others, the crystal is at first simple; and afterwards, through some
change in the position of the material applied to its increase, receives new
layers, or a continuation in a reversed position. This mode of twinning is metagenetic, or
a result subsequent to the origin of the crystal; while the ordinary mode is paragenetic.
One form of it is illustrated in Fig. 556. The middle portion had attained a length of half
an inch or more, and then became geniculated alternately at either extremity. These
geniculations are often repeated in rutile, and the ends of the crystal are thus joined
to one another, and occasionally produce nearly regular prismatic forms.

This metagenetic twinning is sometimes presented by the successive layers of deposition
in crystals, as in some quartz crystals, especially smoky, the inseparable layers,
exceedingly thin, being of opposite signs. In a similar manner, crystals of the triclinic
feldspars, albite, etc. are often made up of thin plates parallel to b (010), by oscillatory
cumposition, and the face c (001), accordingly, is finely striated parallel to the a , b .

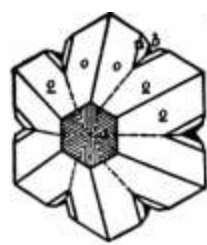
220. Repeated Twinning, Polysynthetic and Symmetrical.—In the preceding paragraph
one case of repeated twinning has been mentioned, that of the feldspars; it is a case of
parallel repetition or parallel grouping in reversed position of successive crystalline
lamellae. This kind of twinning is often called polysynthetic twinning, the lamellae in
many cases being extremely thin, and giving rise to a series of parallel lines (striations)
on a crystal face or a surface of cleavage. The triclinic feldspars show in many cases
polysynthetic twinning and not infrequently on both c (001) and b (010), cf. p. 130. It is
also observed with magnetite (Fig. 456), pyroxene, barite, etc.

Another kind of repeated twinning is illustrated by Figs. 357-363, where the successively
reversed individuals are not parallel. In these cases the axes may, however, lie in a zone,
as the prismatic twins of aragonite, or they may be inclined to each other, as in Fig. 359 of
staurolite. In all such cases the repetition of the twinning tends to produce circular forms,
when the angle between the two axial systems is an aliquot part of 360° (approximately).
Thus six-rayed twinned crystals, consisting of three individuals (hence called trillings),
occur with chrysoberyl (Fig. 357), or cerussite (Fig. 358), or staurolite

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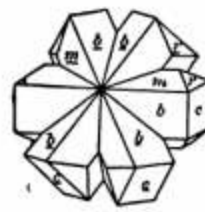
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(Fig. 359), since three times the angle of twinning in each case is not far from 360° . Again,
five-fold twins, or fivefolds, occur in the octahedrons of gold a



Chrysoberyl.

360.

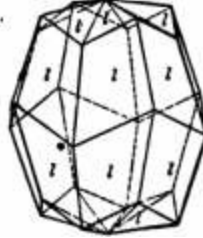


Cerussite.

361.

Staurolite.

362.



Spinel.

Rutile.

Phillipsite.

spinel (Fig. 360), since $5 \times 70^\circ 32' = 360^\circ$ (approx.). Eight-fold twins, or eightlings, of rutile (Figs. 361, 357) occur, since the angle of the axes in twinned position goes approximately eight times in 360° .

Repeated twinning of the symmetrical type often serves to give the compound crystal an apparent symmetry of higher grade than that of the simple individual, and the result is often spoken of as a pseudo-symmetry (Art. 20), cf. Fig. 397 of aragonite, which represents a pseudo-hexagonal crystal. Fig. 363 of phillipsite (cf. Figs. 423--424) is an interesting case, since it shows how a multiple twin of a monoclinic crystal may simulate an isometric crystal (dodecahedron).

Compound crystals in which twinning exists in accordance with two laws at once are not of common occurrence; an excellent example is afforded by staurolite, Fig. 411. They have also been observed on albite, orthoclase, and in other cases.

221. Secondary Twinning.—When there is reason to believe that the twinning has been produced subsequently to the original formation of the crystal, or crystalline mass, as, for example, by pressure, it is said to be secondary. Thus the calcite grains of a crystalline limestone often show such secondary twinning lamellae. The same are occasionally observed ($\{c, 00i\}$) in pyroxene crystals. Further, the polygenetic twinning of the triclinic feldspars is often

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EXAMPLES OF METHODS OF TWINKING.

secundarily in origin. It is also explained in a crystallographic treatise, in calcite (1

in origin. This subject is further discussed on a later page, where it is also explained that in certain cases twinning may be produced artificially (B Art. 261).

EXAMPLES OF IMPORTANT METHODS OF TWINNING.

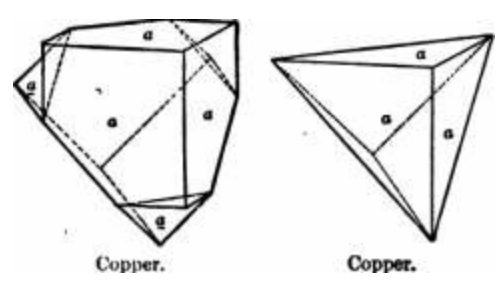
222. *Trigonal System*.—With few exceptions the twins of the normal group of this system are of one kind, the twinning-axis an octahedral axis, and

the twinning-plane consequently parallel to an octahedral face; in most cases,

also, the latter coincides with the composition-plane. This kind as applied to the simple octahedron; it is especially spinel group of minerals, and is hence called in general

Fig. 350, p. 112,* shows a form commonly associated with the epitaxial-twin. Fig.

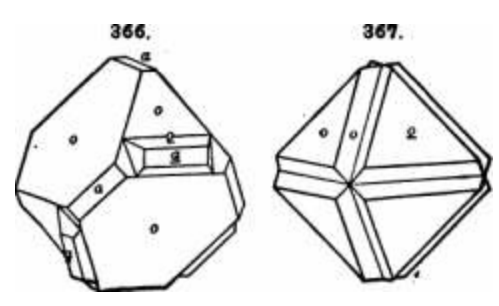
363 is a similar more complex form; Fig. 364 shows a cube twinned by this



Copper

method, and Fig. 365 represents the same form but shortened in the direction of the octahedral axis, and hence having the anomalous aspect of a triangular pyramid. All these cases are contact-twins.

Penetration-twins, following the same law, are also common. A simple case of fluorite is shown in Fig. 353, p. 121; Fig. 366 shows one of galena;



Galena

Fig. 367 is a repeated octahedral twin of barytine, and Fig. 368 a dodecahedral of sodalite.

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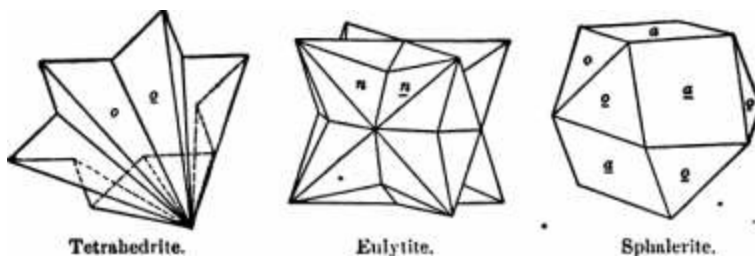
223. In the pyritohedral group of the isoiuetric system penetration-twins 369. of the type shown in Fig. 369 are common (this form of

pyrite is ofteu called tiie iron crosK). Here the cubic axis U the twimiiug-axie, and obviously such a twin is Impossible in the Dormal group.

Figs. 370 and 371 show analogous forms with parallel ^ aies for crystals belonging to the tetraliedral group. 'Tlie peculiar development of Fig. 370 of tetrahedrite is to be noted. Fig. ;iT2 is a twin of the ordinary spinel type of another tetrahedral species, sphalerite; with it, complex forms with repeated twinning are not uncommona Pyrite. and sometimes polysynthetic twin lamellffl are noted.

224. Tetragonal System.—The most common method is that where the twinning-plane is parallel to a face of the pyramid, e (lOlJ. It is especially characteristic of the species of the rntile group—viz., rutile and cassiterite;

370. 311. 372.



Tetrahedrlie.

Euly

SpbuU'ril

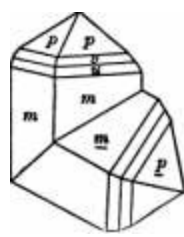
also similarly the allied species zircon. This is illustrated in Fig. 373, and

again iu Fig, 374. Fig. 375 shows a repeated twin of rntile, the twinning

^according to this law; the vertical axes of the snccessive six indiviituals lie

in a plane, and an inclosed circle is the result. Another repeated twin of rutile

373. 374. 37B.



Rutile.

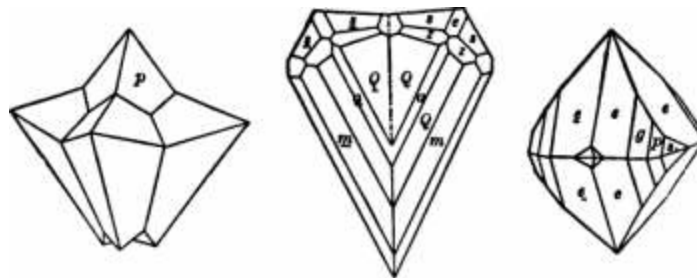
Cas[^]iturite. Zircon.

according to the same law is shown in Fig. 361; here the successive vertical axes form a zigzag line; Fig. 376 shows an analogous twin of hausmannite. Another kind of twinning, twinning-plane parallel to a face of the pyramid (301), is shown in Fig. 377.

225. In the pyramidal group of the same system twins of the type of Fig. 378 are not rare. Here the vertical axis, *d*, is the twinning-axis; such a crystal may simulate one of the normal group.

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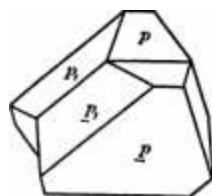
EXAMPLES OF IMPORTANT METHODS OF TWINNING. 376. 377. 378.



Hausmannite. Rutile. Sellaite.

In chalcopyrite, of the sphenoidal group, twinning with a face of the unit prism, $\{111\}$, as the twinning-plane is common (Fig. 379). As the angles differ but a small fraction of a degree from those of a regular octahedron, such twins often resemble closely spinel-twins. In Fig. 380 a repeated twin of this type has a pseudo-rhombic aspect. In Fig. 381 the twinning-plane is $e(101)$. Other rarer kinds of twinning have also been noted.

379. 380. 381. ■



Chalcopyrite.

226, Hexagonal System.—In the hexagonal division of this system twine

382. are rare. An example is furnished by pyritohite, Fig.

382, where the twinning-plane is the pyramid (1011), the

vertical axes of the individual crystals being nearly at

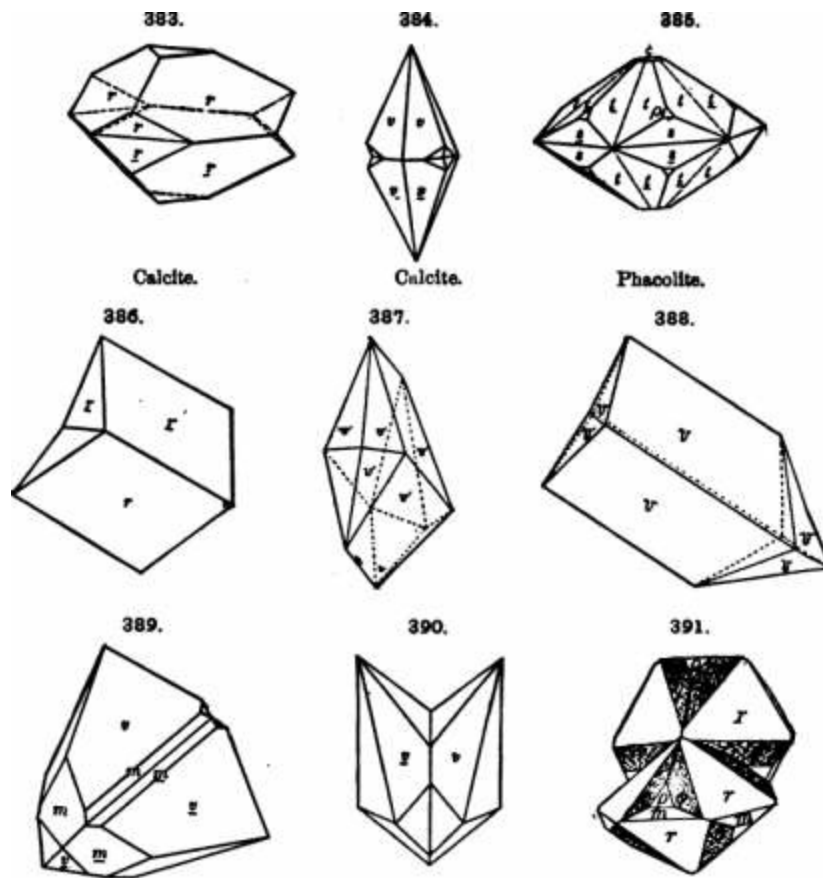
right angles to each other (since $OUOl A 1011 = 45^\circ 8'$).

(Apparent cruciform twins of apatite, of the pyramidal group, have been noted in some rare cases. Here the diagonal pyramid $s(1131)$ was the plane which seemed to be the twinning-plane. These cases need confirmation.)

237. In the species belonging to the trigonal or rhombohedral division, twins are common. Thus the twinning-axis may be the vertical axis, as in the contact-twins of Figs. 383 and 384, or the penetration-twins of Figs. 348, 385. Or the twinning-plane may be the obtuse rhombohedron $e(0113)$, as in Fig. 386, the vertical axes crossing at angles of 137° and 52° ; these forms are often curiously distorted, as in Figs. 33T, 388. Again, the twinning-plane may be $r(1011)$, as in Fig. 389, the vertical axes nearly at right angles (90°); or (0531) , as in Fig. 390, the axes inclined 53° and 120° . In Fig. 391 of gmelinite the twinning-plane is the rhombohedron (30.12) , which corresponds in angle with the common fundamental form of the allied species chabazite.

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Figs. 386-390, *Calcite* Pfg. 391, *Gmelinite*. In the trapezohedral group, the species quartz shows several methods of twinning. In Fig. 392 the twinning-plane is the diagonal pyramid $\{112\}$, the axes crossing at angles of 84° and 954° . In Fig. 393 the twinning-planes are $\{101\}$ and $\{011\}$, the axes hence parallel, the individuals both right- or both left-handed but unsymmetrical. The resulting forms, as in Fig. 393, are mostly penetration-twins, and the parts are often very irregularly united, as shown by dull areas (z) on the plus rhombohedral face (r); otherwise these twins are recognized by pyro-electrical phenomena. In Fig. 394 the twinning-plane is a $\{110\}$ —the Brazil law—the individuals respectively right- and left-handed and the twin symmetrical with reference to an a -face; these are usually irregular penetration-twins; in these twins r and r , also z and z , coincide. These twins often show, in converging polarized light, the phenomenon of Airy's spirals. It may be added that pseudo-twins of quartz are common—that is, groups of crystals which nearly

;*CoogIC*

EXAMPLES OF IMPROPER TWINNING OF CaCO_3 .

conform to some more or less complex twinning law, but where the twinning is nevertheless only accidental. See also the remarks in Art. 232 about Fig. 433.

393. 393. 394. 3»B.

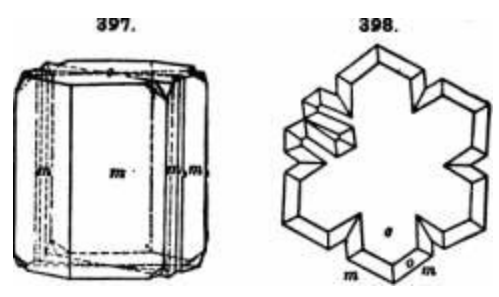


Fig. 893-894, Qtinrrz. Fig. 393. Pheoaclte.

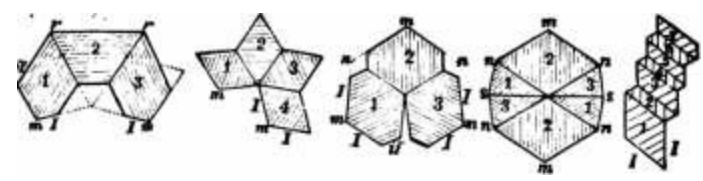
In Fig. 395 of phenacite (cf. p. 80 et seq.) the vertical axis is obviously the twinning-axis.

228. Orthorhombic System.—In the orthorhombic system the commonest method of twinning is that where the twinning-plane is a face of a prism of

$\frac{1}{2}N$
 $\frac{1}{2}E$



Figs. 397-398, Aragoite. 399.



60°, or nearly 60'. This is well shown with the species of the aragonite group. In accordance with the principle stated in Art. 220, the twinning after this law is often repeated, and thus forms with pseudo hexagonal symmetry result. Fig. 396 shows a simple twin of aragoite; Figs. 397, 398 repeated forms exhibiting the irregularities on the faces due to the fact that the prismatic angle is not exactly 60°. Fig. 399. a-e, show further some of the methods of composition which have been noted; in e the twinning is poly-synthetic.

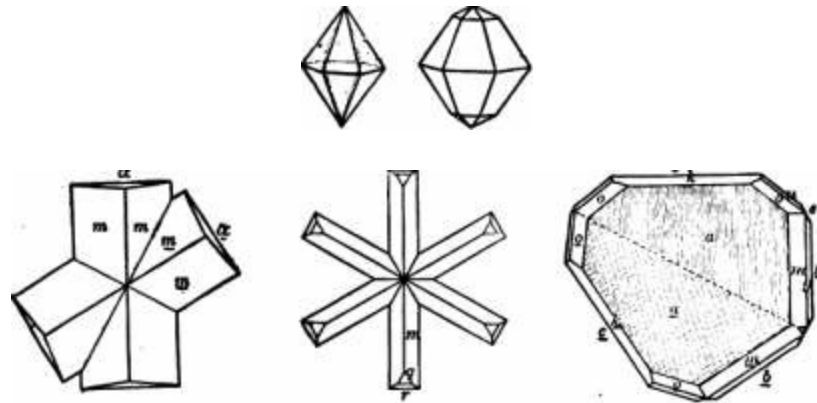
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With witherite (and bromilite), apparent hexagonal pyramids (Figs. 400, 401) are common, but the true complex twinning is revealed in polarized light, as noted later.

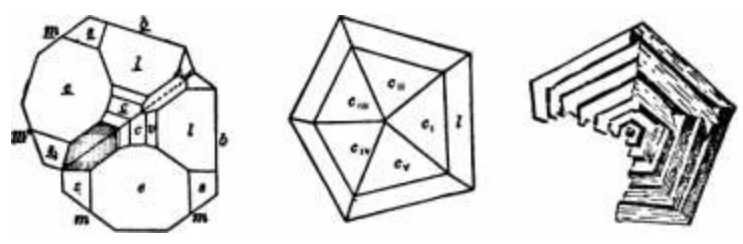
Twinning of the same type, but where a plane of 60° is twinning-plane, is common with

arsenopyrite (tw. pi. 6 (101), as shown in Figs. 40-3, 403; also Fig. 404 of columbite, but compare Fig. 347 and remarcia iii Art 217, Another example is given in Fig. 357 of alexandrite (cbrysoberyl). Chrysolite, manganite, hrrnite, are other species with which this kind of ' twinning is common.

403. 401.



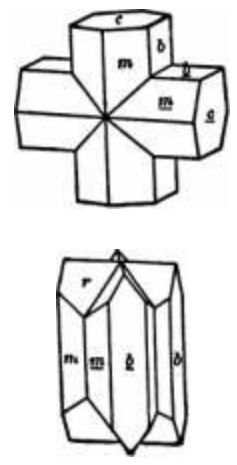
Another common method of twinning is that where the twinning is parallel to a face of a prism of about 70° , ^ shown in Fig 405 With this method symmetrical twinnings not infrequently occur (Figs. 406, 407).

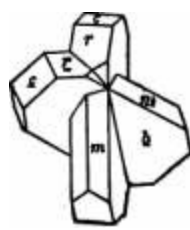


Harcasite. Mnrcaalte. AraennpTnta

The species stanrolite illustrates three kinds of twinning. In Fig 408 the

40S. 409. 410.





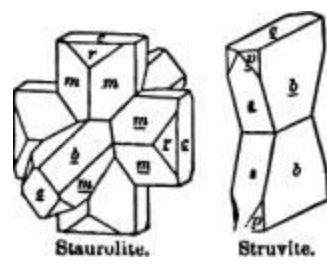
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twinning-plane m (032), and since $(001 \wedge 032) = 45^\circ 41'$, the crystals cross nearly $^\wedge$, iio at right angles. In Fig. 409 the twinning-

plane is the prism (330). In Fig. 410 it is the pyramid (5;3"2); the crystals lie crossing at angles of about $60''$, stellate trillings occur (see Fig. 359), and indeed more complex formt. In Fig. 411 there is twinning according to both (0152) and (232).

In the heniimorphic group, twins of the type shown in Fig. 413, with c its the twin-ning-plaie, are to be noted.

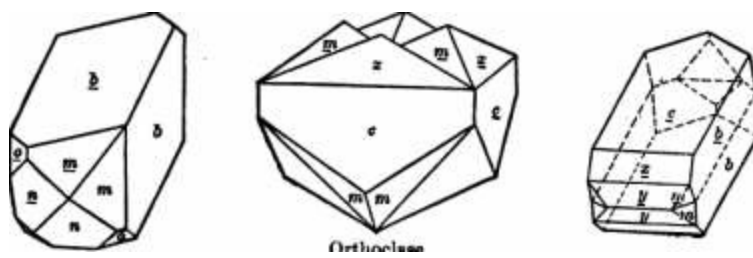
22B. Honooluua System.—In the mono-clinic system, twins with the vertical axis as twinning-axis are common; this is illustrated by Fig. 413 of angite (pyroiene). Fig. 414 of gypsum, and Fig. 415 of ortho-elase (see also Fig. 352, p. 120). With the latter species these twins are called VarUbad twint (oecause common in the trachyte of Carlsbad, Bohemia); they may be contact-twins (Fig. 352), or irregular penetration-twins (Fig. 415). In Fig. 352 it is to be noted that c and x fait nearly in the same plane.



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Augile. Gypsum. Orthoclaae.

In Fig. 416, also of orthoclaae, the twinning-plane is the clinodome (021), and since $(001 \wedge 021) = 44^\circ 56'$, this method of twinning yields nearly square prisms. These twins are called Breno twina (from a prominent locality at Baveno, Italy); they are often repeated (Fig. 417). In Fig. 418 a



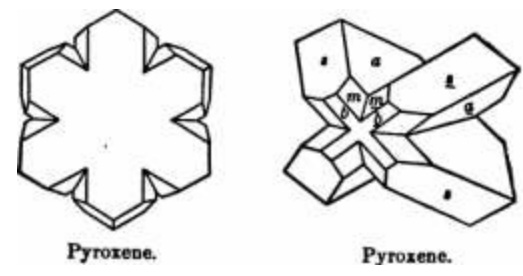
Orthoclase

Manebach twin is shown; here the twinning-plane is c (001). Other rarer types of twinning have been noted with orthoclase. Polysynthetic twinning with c (001) as twinning-plane is common with pyroxene (cf. Fig. 430, p. 131).

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Twins of the aragonite-chrysoberyl type, are not uncommon with monoclinic species, having it prominent 60° prism (or dome), as in Fig. 419. Stellate twins after this law are common with chondrodite and clinohumite. An analogous twin of pyroxene is shown in Fig. 430; here the pyramid (122) is the twinning-plane, and since $(010 \wedge 23) = 59^\circ 31'$, the crystals cross at angles of nearly 60°; further, the orthopyramids fall nearly in a common zone, since $(100 \wedge 33) = 90^\circ 9'$, In Fig. 421 the twinning-plane is the orthodome

419. 420. 42t.



Pyroxene

(101). Phillipsite and harmotome exhibit multiple twinning, and the crystals often show pseudo-symmetry. Fig. 422 shows a cruciform fourling with c (001) as twinning-plane, the twinning shown by the striations on the side face. This is compounded in Fig. 423 with twinning-plane (010), making nearly square prisms, and this further repeated with m (110) as twinning-plane yields the form in Fig. 424, or even Fig. 362, p. 122, resembling an isometric dodecahedron, each face showing a fourfold striation.

Feldspar

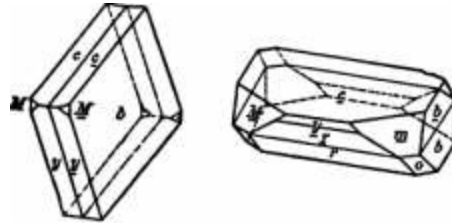
230. Triclinic System.—The most interesting twins of the triclinic system are those shown by the feldspars. Twinning with b (010) as the twinning-plane is very common, especially polysynthetic twinning yielding thin parallel lamellae, shown by the striations on the face c (or the corresponding cleavage-surface), and also clearly revealed in polarized light. This is known as the albite law (Fig. 425). Another important method (Fig. 426) is that of the pericline law; the twinning-axis is the crystallographic axis J . Here

the twins are united by a section (rhombic section) shown in the figure and further explained under the feldspars. Polysynthetic twinning after this law is common, and hence a cleavage-mass may show two acts of striation, one on the surface parallel to c (001) and the other on that parallel to ft (010). The

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angle made by these last striations with the edge 001/010 is characteristic of the particular triclinic species, as noted later.

426. 43S. 427.



Labradorite. Albite. Albiti.'

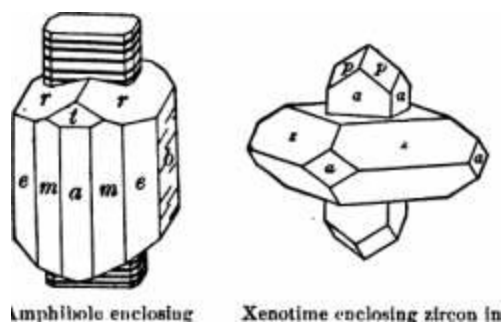
Twins of albite of other rarer types also occur, and further twins similar to the Carlsbad, Baveno, and Manebach twins of orthoclase. Fig. 437 shows twinning according to both the albite and Carlsbad types.

REGULAR GROUPING OF CRYSTALS. 231. Parallel Grouping. —Connected with the subject of twin crystals is that of the parallel position of associated crystals of the same species, or of different species.

Crystals of the same species occurring together are very commonly in parallel position. In this way large crystals, as of

calcite, quartz, fluorite, are sometimes built up of smaller individuals grouped together with corresponding faces parallel. This parallel grouping is often seen in crystals as they lie on the supporting rock. On glancing the eye over a surface covered with crystals a reflection from one face will often be accompanied by reflections from the corresponding face in each of the other crystals, showing that the crystals are throughout similar in their positions.

With many species, complex crystalline forms result from the growth of parallel partial crystals in the direction of the crystallographic axes, or axes of symmetry.



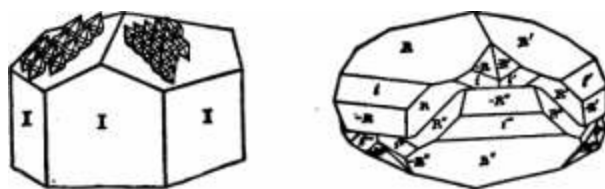
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Such dendritic forms, resembling branching vegetation, often of great delicacy, are seen with gold, copper, argentite, and other species, especially those of the isometric system. This is shown in Fig. 16 (ideal), and again in Fig. 49, where the twinned and flattened cubes (cf. Fig. 365, p. 123) are grouped in directions corresponding to the diagonals of an octahedral face which is the twinning-plane.

232. Parallel Grouping of Unlike Species.—Crystals of different species often show the same tendency to parallelism in mutual position. This is true most frequently of species which are more or less closely similar in form and composition. Crystals of albite, implanted on a surface of orthoclase, are sometimes an example of this; crystals of amphibole and pyroxene (Fig. 430), of zircon and xenotime (Fig. 431), of various kinds of mica, are also at times observed associated in parallel position.

The same relation of position also occasionally occurs where there is no connection in composition, as the crystals of rutile on tabular crystals of hematite, the vertical axes of the former coinciding with the lateral axes of the latter. Crystals of calcite have been observed whose rhombohedral faces



had a series of quartz crystals upon them, all in parallel position (Fig. 433); sometimes three such quartz crystals, one on each rhombohedral face, entirely envelop the calcite, and unite with re-entering angles to form pseudo-twins (rather trillings) of quartz after calcite. A similar occurrence from Specimen Mountain, in the Yellowstone Park, is shown in Fig. 433.

IRREGULARITIES OF CRYSTALS.

233. The laws of crystallization, when unmodified by extrinsic causes, should produce forms of exact geometrical symmetry, the angles being not only equal, but also the

homologous faces of crystals and the dimensions in the directions of like axes. This symmetry is, however, so uncommon that it can hardly be considered other than an ideal perfection. The various possible kinds of symmetry, and the relation of this ideal geometrical symmetry to the actual crystallographic symmetry, have been discussed in Arts. 14 and 18 et seq. Crystals are very generally distorted, and often the fundamental forms are so completely disfigured that an intimate familiarity with the possible irregularities is required in order to unravel their complexities. Even the angles may occasionally vary rather widely.

The irregularities of crystals may be treated of under several heads: 1, Variations of form and dimensions; 2, Imperfections of surface; 3, Variations of angles; 4, Internal imperfections and impurities.

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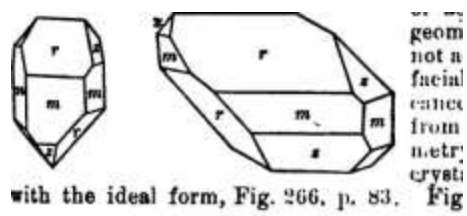
1. VARIATIONS IN THE FORMS AND DIMENSIONS OF CRYSTALS.

234. Distortion in General—The variation in the forms of crystals, or, in other words, their distortion, may be irregular in character, certain faces being larger and others smaller than in the ideal geometrical solid. On the other hand, it may be systematic, giving to the distorted form the symmetry of a group or system different from that to which it actually belongs. The former case is the common rule, but the latter is the more interesting.

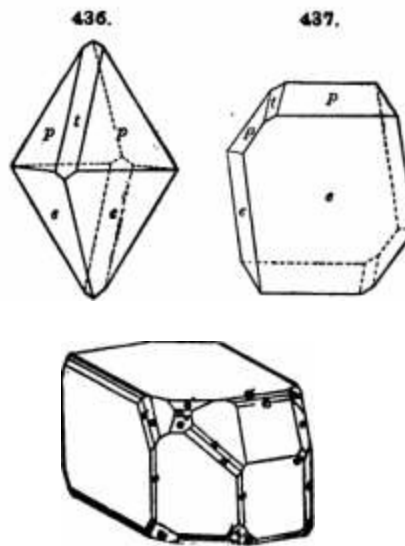
235. Irregular Distortion.—As stated above and on p. 11, all crystals show to a greater or less extent an irregular or accidental variation from the ideal geometrical form. This distortion, if not accompanied by change in the interfacial angles, has no particular significance, and does not involve any deviation from the laws of crystallographic symmetry. Figs. 434, 435 show distorted crystals of quartz; they may be compared

Fig. 436 is an ideal and Fig. 437 an

43B.



with the ideal form, Fig. 266, p. 83. Fig



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actual crystal of lazuhite. So, too, Fig. 438 is a distorted crystal of apatite, to be compared with Fig. 216, p. 73.

The correct identification of the forms of a crystal is rendered much more difficult because of the prevailing distortion, especially when it results in the destruction of certain faces by the enlargement of others. In describing the distorted crystalline forms it must be remembered that while the appearance of the crystals may be entirely altered, the interfacial angles remain the same; moreover, like crystals are physically alike—that is, alike in degree of luster, in striations, and so on. Thus the prismatic faces of quartz show almost always characteristic horizontal striations.

In addition to the variations in form which have just been described, still greater irregularities are due to the fact that, in many cases, crystals in nature are attached either to other crystals or to some rock surface, and in consequence of this are only partially developed. Thus quartz crystals are generally attached by one extremity of the prism, and hence have only one set of pyramidal faces; perfectly formed crystals, having the double pyramid complete, are rare,

236. Symmetrical Distortion.—The most interesting examples of the symmetrical distortion of crystalline forms are found among crystals of the isometric system. An elongation in the direction of one cubic axis may give

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the appearance of tetragonal symmetry, or that in the direction of two cubic axes of orthorhombic symmetry; while in the direction of an octahedral axis a lengthening or shortening gives rise to forms of apparent rhombohedral

Bjmmetry, Such cases are common with native gold, silver, and copper.

A cube lengthened or shortened along one axis becomes a rhombic prism, and if varied in the direction of two axes is changed to a rectangular prism. Cubes of pyrite.

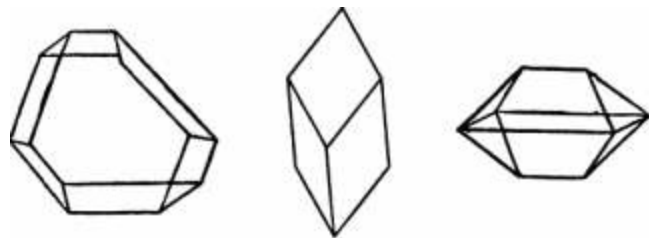
Siderite, fluorite, etc., are often thus distorted. It is very unusual to find a cubic crystal which is a true symmetrical cube. In some species the cube or octahedron (or other isometric form) is lengthened into a capillary crystal or needle, as happens in cuprite and pyrite.

An octahedron which is not parallel to a face that is, in the direction of a trigonal interaxis—is reduced to a tabular crystal resembling a rhombohedron crystal with basal plane (fig. 439). If lengthened in the same direction, to the obliteration of the terminal octahedral faces, it becomes an acute rhombohedron (cf. Fig. 440),

43».

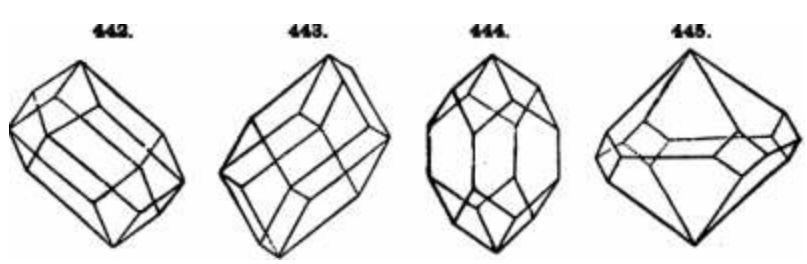
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When an octahedron is extended in the direction of a line between two opposite edges, or that of a rhombic interaxis, it has the general form of a rectangular octahedron; and still farther extended, as in Fig. 441, it is changed to a rhombic prism with dihedral summits (spinel, fluorite, magnetite). The figure represents this prism lying on its acute edge.

The octahedron lengthened in the direction of a diagonal between the obtuse solid



angles—that is, that of a trigonal interaxis—becomes a rhombic prism with three-sided summits, as in fig. 442. If shortened in the same direction, it becomes a short prism of the

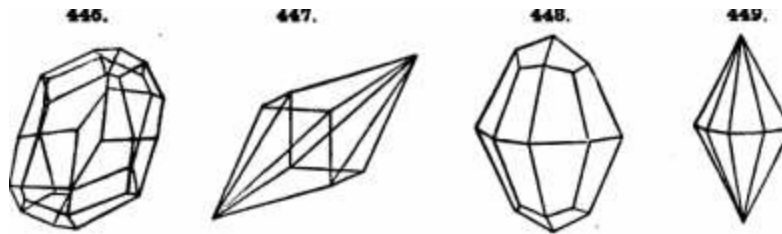
same kind (Fig. 448). Both resemble rhombohedral forms and are common in garnet (compare Fig. 234, p. 76, of calcite). When lengthened in the direction of one of the cubic axes, the dodecahedron becomes a square prism with pyramidal summits (Fig. 444). and shortened along the same axis it is reduced to a square octahedron, with truncated basal angles (Fig. 446).

The trapezohedron elongated in the direction of an octahedral (trigonal) axis assumes a rhombohedral (trigonal) symmetry. The resulting forms referred to the usual hexagonal axes for both (311) and (811) are as follows:

Fig. 445, 446, 447, 448

IRREGULARITIES OF CRYSTALS.

For $\langle 211 \rangle$ the resulting form is that of Fig. 446, and if still further lengthened, to the obliteration of some of the faces, it becomes a tetrahedron (Fig. 447). This has been observed in fluorite. Only two faces are here present out of the twelve. If the



elongation of this trapezohedron (311) takes place along a cubic axis, it becomes a double six-sided pyramid with four-sided summits (Fig. 446); or if these planes are further iterated by a further extension, it becomes a complete eight-sided double pyramid (Fig. 448).

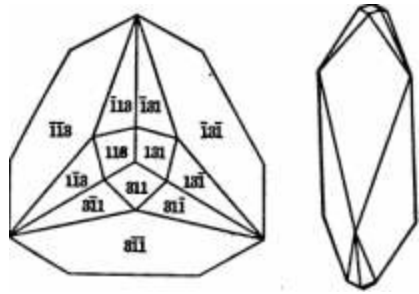
The accompanying figures illustrate the rhombohedral distortion of the trapezohedron (311), cf. Fig. B3, p. 8. Fig. 450 shows the faces as projected on a plane normal to an

Fig. 450.

Fig. 451.

Fig. 452.

Fig. 453.

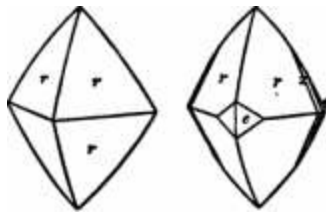


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The native gold from Ihu White H ,

Uregod, aoinetimes consists of n slender !<ti'iiig of such rUomtxihedral crystals, Hud not infrqieilly there are minor bruncliea in the direction of two or more of the other ociahBiiral axes (Figs. 452. ira). The Irisoctahedron Is mrp ns n prominent form, hut ft Piirloiaa example of Its diatort'on in given in Figs. 454. 455 of pyrite from French Creeli. Ptnn. The form, apparently letragonal (or orlliorhombie), shows only eight fjices of the trisoctihedron » '{3S3|, and these are strongly rounded; faces of ihe pyriio-bedroD e (210) also appear as a subordlonte fotiu. Hlmllarly tbe tetr&bezahedion and bexociahedrou may show distortion of the same kind. Further examples are to be found in Ibe other lyiums.



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2. IMPERFECTIONS OF THE SURFACES OF CRYSTALS.

237. Striatiftiu Dae to Oscillatory Gombinationi.—The parallel lines or furrows on the Burfaces of crystals are called slrUa or striatione, and ench surfaces are said to be striated.

Eiicli little ridge on a striated surface is enclosed by two narrow planes more or less regular. These planes often correspond in position to dinereDt faces of the crystal, and these ridges have been formed by a continued oscillation in the operation of the causes that give rise, when acting uninter-riiptetlly, to enlarged faces. By this means, the surfaces of a crystal are marked in pinullel lines, with a succession of narrow planes meeting at an angle and cuneituting the ridges referred to.

This combination of different planes in the formation of a surface has been termed

oscillatory combination. The horizontal striations on prismatic crystals of quartz are examples of this combination, in which the oscillation has taken place between the prismatic and rhombohedral faces. Thus crystals of quartz are often tapered to a point, without the usual pyramidal terminations.

Other examples are the striations on the cubic faces of pyrite parallel to the intersections of the cube with the faces of the pyritohedron; also the striations on magnetite due to the oscillation between the octahedron and dodecahedron. Prisms of tourmaline are very commonly bounded vertically by three convex surfaces, owing to an oscillatory combination of the faces in the prismatic zone.

23B. Striations Due to Repeated Twinning.—The striations of the basal plane of albite and other triclinic feldspars, also of the rhombohedral surfaces

of some calcite, have been explained in Art. 230 as due to polysynthetic twinning. This is illustrated by

Fig. 45(i) of magnetite from Port Henry, N. V.

(Kemp.)

239. Markings from Erosion and Other Causes.—The faces of crystals are not uncommonly uneven, or have the crystalline structure developed as a consequence of etching by some chemical agent. Crystals of galena are often thus uneven, and crystals of lead sulphate (anglesite) or lead carbonate (cerussite) are sometimes present as evidence with regard to the cause. Crystals of numerous other species, even of Mugwort, corundum, spinel, quartz, etc., sometimes show the

same result of partial change over the surface—often the incipient stage in a process tending to a final removal of the whole crystal. Interesting investigations have been made by various authors on the action of solvents on different minerals, the actual structure of the crystals being developed in this way. This method of etching is fully discussed, with illustrations, in another place (Art. 265).

The markings on the surfaces of crystals are not, however, always to be ascribed to etching. In most cases such depressions, as well as the minute elevations upon the faces having the form of low pyramids (so-called vicinal prominences), are a part of the original molecular growth of the crystal, and often serve to show the successive stages in its history. They may lie Imperfections arising from an interrupted or disturbed development of the form, the perfectly smooth and even crystalline faces being the result of completed

action free from disturbing causes. Examples of the markings referred to occur on the crystals of most minerals, and conspicuously so on the rhombohedral faces of quartz.

Faces of crystals are often marked with angular elevations more or less distinct, which are due to oscillatory combination. Octahedrons of Quartz are common which have for each face a surface of minute cubes, proceeding from an oscillation between the cube and octahedron. Sometimes an examination of such a crystal shows that though the form is apparently octahedral, there are no octahedral faces present at all. Other similar cases could be mentioned,

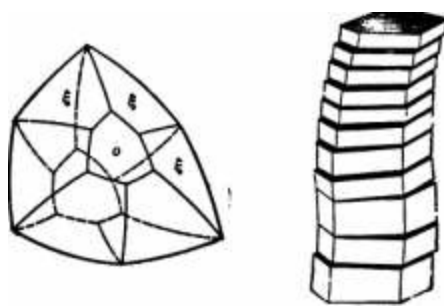
Whatever their cause, these minute markings are often of great importance as revealing the true molecular symmetry of the crystal. For it follows from the symmetry of crystallization that like faces must be physically alike—that is, in regard to their surface character; it thus often happens that on all the crystals of a species from a given locality, or perhaps from all localities, the same planes are etched or roughened alike. There is much uniformity on the faces of quartz crystals in this respect.

240. Curved Surfaces may result from (a) oscillatory combination; or (b) some independent molecular condition producing curvatures in the laminae of the crystal; or (c) from a mechanical cause.

Curved surfaces of the first kind have been already mentioned (Art. 287). A singular curvature of this nature is seen in Fig. 457, of calcite; in the lower

part traces of a scalenohedral form are apparent which was in oscillatory combination with the prismatic form. Curvatures of the second kind sometimes have all the faces convex. This is the case in crystals of diamond (Fig. 458), some of which are almost spheres. The mode of curvature, in which all the faces are equally convex, is less common than that in which a convex surface is opposite and parallel to a corresponding concave surface. Rhombohedrons of dolomite and siderite are usually, thus curved. The feathery curves of frost on windows and the flagging-stones of pavements in winter are other examples. The alabaster rosettes from the Mammoth Cave, Kentucky, are similar. Stibnite crystals sometimes show very remarkable curved and twisted forms.





A third kind of curvature is of mechanical origin. Sometimes crystals appear as if they had been broken transversely into many pieces, a slight;

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displacement of which has given a curved form to the prism. This is common in tourmaline and beryl. The beryls of Monroe, Conn., often present these interrupted curvatures, as represented in Fig. 459.

Crystals not infrequently occur with a deep pyramidal depression occupying the place of each plane, as is often observed in common salt, alum, and sulphur. This is due in part to their rapid growth.

3. VARIATIONS IN THE ANGLES OF CRYSTALS.

241. The greater part of the distortions described in Arts. 236, 236 occasion no change in the interfacial angles of crystals. But those imperfections that produce convex, curved, or striated faces necessarily cause such variations. Furthermore, circumstances of heat or pressure under which the crystals were formed may sometimes have resulted not only in distortion of form, but also some variation in angle. The presence of impurities at the time of crystallization may also have a like effect.

Still more important is the change in the angles of completed crystals which is caused by subsequent pressure on the matrix in which they were formed, as, for example, the change which may take place during the more or less complete metamorphism of the enclosing rock.

The change of composition resulting in pseudomorphous crystals {see Art. 293} is generally accompanied by an irregular change of angle, so that the pseudomorphs of a species vary much in angle.

In general it is safe to affirm that, with the exception of the irregularities arising from imperfections in the process of crystallization, or from the subsequent changes alluded to, variations in angles are rare, and the constancy of angle alluded to in Art. 11 is the universal law.

In cases where a greater or less variation in angle is observed in the crystals of the same species from different localities, the cause for this can usually be found in a difference of

chemical composition. In the case of isomorphous compounds it is well known that an exchange of corresponding chemically equivalent elements may take place without a change of form, though usually accompanied with a slight variation in the fundamental angles.

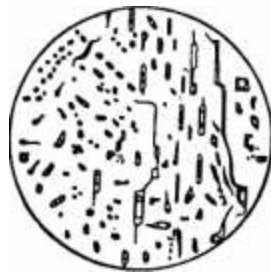
The effect of heat upon the form of crystals is alluded to in Art. 41B.

i. INTERNAL IMPERFECTIONS AND INCLUSIONS.

242. The transparency of crystals is often destroyed by disturbed crystallisation; by impurities taken up from the solution during the process of crystallization; or, again, by the presence of foreign matter resulting from partial chemical alteration. The general name, inclusion, is given to any foreign body enclosed within the crystal, whatever its origin. These inclusions are extremely common; they may be gaseous, liquid, or solid; visible to the unaided eye or requiring the use of the microscope.

Rapid crystallization is a common explanation of inclusions. This is illustrated by quartz crystals containing large cavities full or nearly full of water (in the latter case, these showing a movable bubble); or, they may contain sand or iron oxide in large amount. In the case of calcite, crystallization from a liquid largely charged with a foreign material, as quartz sand, may result in the formation of crystals in which the impurity makes up as much

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INTERNAL IMPERFECTIONS OF CRYSTALS. 139

as two-thirds of the whole mass; this is seen in the Fontainebleau limestone, and similarly in that from other localities.

243. Liquid and Gas Inclusion. —Attention was early called by Brewster to the presence of fluids in calcite in certain minerals, as quartz, topaz, beryl, cleveite, etc. In later years this subject has been thoroughly studied by

Sorby, Zirkel, Vogelsang, Fischer, Eosenbach,

and others. The nature of the liquid can often be determined, as by its refractive power, or by special physical test {e.g., determination of the critical point in the case of CO₂), or by chemical examination. In the majority of cases the observed liquid is simply water; but it may be the salt solution in which the crystal was formed, and not infrequently, especially in the case of quartz, liquid carbon dioxide (CO₂), as first proved by Vogelsang. These liquid inclusions are marked as such, in many cases, by the presence in the cavity of a movable bubble of gas. Occasionally cavities contain two liquids, as water and liquid carbon dioxide, the latter then inclosing a bubble of the same substance as gas (cf. Fig. 460). Interesting experiments can be made with sections showing such inclusions (cf. literature, p. 141). The mixture of gases yielded by smoky quartz, meteoric iron, and other substances, on the application of heat, has been analyzed by Wright.

In some cases the cavities appear to be empty; if they then have a regular form determined by the crystallization of the species, they are often called negative inclusions. Such cavities are commonly of secondary origin, as remarked on a later page.

244. Solid Inclusions. —The solid inclusions are almost infinite in their variety. Sometimes they are large and distinct, and can be referred to known mineral species, as the scales of goethite or hematite, to which the peculiar character of aventurine feldspar is due. Magnetite is a very common impurity in many minerals, appearing, for example, in the primary mica; quartz is also often mechanically mixed, as in staurolite and kyanite. On the other hand, quartz crystals very commonly inclose foreign material,

such as chlorite, tourmaline, rutile, hematite, asbestos, and many other minerals. (Cf. also Arts. 245, 246.)

An excellent example of the inclusion of one mineral by
I & another is seen in the figures of tourmaline envelop-
ment of a crystal of tourmaline; a few
cross-sections of it at the points indicated (a, b, c) are given by
Figs. 463, 464, 465. The latter shows the feldspar in-
clusion in the tourmaline, the tourmaline being
merely a shell. Similar specimens from the same locality
(Port Henry, Essex Co., N. Y.) show that there is no necessary
connection between the position of the tourmaline and that of
the feldspar.

Similar occurrences are those of trapezoidal crystals of garnet, where the latter is merely
inclusion, enclosing calcite, or sometimes epidote or quartz (Fig. 475).

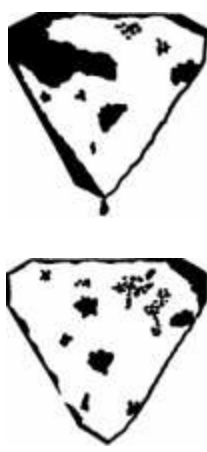


• E. H. Williams, Am. J. Sci. 11. 878. 1876.

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452.



The inclusions may consist of a heterogeneous mass of material; as the granitic matter seen in orthoclase crystals in a porphyritic granite; or the feldspar, quartz, etc., sometimes included in large coarse crystals of beryl or spodumene, occurring in granite veins.

245. Microlites, Crystallites—The microscopic crystals observed as inclusions may sometimes be referred to known species, but more generally their true nature is doubtful. The term microlite, proposed by Vogelsang, is often used to designate the minute included crystals; they are generally of needlelike form, but sometimes quite irregular, and often very remarkable in their arrangement and groupings; some of them are exhibited in Fig. 470 and Fig. 471, as explained below. Where the minute individuals belong to known species they are called, for example, feldspar microlite, etc.

Crystallites is an analogous term used by Vogelsang to cover those minute forms which have not the regular exterior form of crystals, but may be considered as intermediate between amorphous matter and true crystals. Some of

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469.



the forms are shown in Figs. 465-469; they are often observed in glassy volcanic rocks, and also in furnaces. A series of names has been given to varieties of crystallites, such as globulites, margarites, etc. Trichite and belonite are names introduced by Zirkel; the former name is derived from Greek, hair; trichites, like that in Fig. 469, are common in obsidian.

The microscopic inclusions may also be of an irregular glassy nature; this kind is often observed in crystals which have formed from a molten mass, as lava or the slag of an iron furnace.

246. Symmetrically Arranged Inclusions.—In general, while the solid inclusions sometimes occur quite irregularly in the crystals, they are more generally arranged with some evident reference to the symmetry of the form, or external faces of the crystals. Examples of this are shown in the following

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IRREGULARITIES OF CRYSTALS.

Fig. 470 exhibits a crystal of aegite, enclosing magnetite, feldspar and nepheline, microcline, etc. Fig. 471 shows a crystal of leucite, a species



whose crystals very commonly enclose foreign matter. Fig. 475 of a crystal of garnet, containing quartz.

Another striking example is afforded by andalusite (Fig. 473), in which the enclosed carbonaceous impurities are of considerable extent and remarkably arranged, so as to yield symmetrical figures of various forms. Stanrolite occasionally shows analogous carbonaceous impurities (Fig. 474).

symmetrically distributed.

The magnetite common as an inclusion in muscovite, as alluded to above, is always symmetrically disposed, usually = parallel to the directions of the percussion-figure (Fig. 477, p. 149). The asterism of phlogopite is explained by the presence of symmetrically arranged inclusions (cf. Art. 342).

Fig. 474 shows an interesting case of symmetrically arranged inclusions due to chemical alteration. The original mineral, beryl, from Brantford, Ohio, has been altered in substance, apparently homogeneous to the eye, and found under the microscope to exhibit the structure shown in Fig. 474. Chemical analysis proves it to be based on the original and the enclosed hexagonal mineral to be a silicate (LiAlSiO₄) called eucryptite. It has not yet been identified except in this case.

LITERATURE.

Some of the most important works on the subject of microscopic inclusions are referred to here; for a fuller list of papers reference may be made to the work of

RoseDbu3ch(1878, 1892); also that of Zirkel and others mentioned on p. 4.

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(Preisachrin.)

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CRYSTALLINE AGGREGATES.

247. The greater part of the specimen or mass of minerals that occur may be described

as aggregations of imperfect crystals. Many specimens whose structure appears to the eye quite homogeneous, and destitute internally of distinct crystallization, can be shown to be composed of crystalline grains. Under the above head, consequently, are included, all the remaining varieties of structure among minerals.

The individuals composing imperfectly crystallized individuals may be:

1. Columns, or fibers, in which case the structure is columnar or fibrous.
2. Hemilaminar, producing a lamellar structure.
3. Oranular, constituting granular structure.

248. Columnar and Fibrous Structure.—A mineral possesses a columnar structure when it is made up of slender columns, as some amphibole. When the individuals are flattened like a knife-blade, as in cyanite, the structure is said to be bladed.

The structure again is called fibrous when the mineral is made up of fibers, as in asbestos, also the antinapar variety of gypsum. The fibers may or may not be separate. There are many gradations between coarse columnar and fine fibrous structures. Fibrous minerals have often a silky luster.

The following are properly varieties of columnar or fibrous structure:

Reticulated: when the fibers or columns cross in various directions and produce an appearance having some resemblance to a net.

Star-shaped: when they radiate from a center in all directions and produce star-like forms. Ex. stilbite, wavellite.

Radial, divergent: when the crystals radiate from a center without producing stellar forms. Ex. quartz, stibnite.

248. Lamellar Structure,—The structure of a mineral is lamellar when it consists of plates or leaves. The laminae may be curved or straight, and thus give rise to the curved lamellar and straight lamellar structure. Ex. wollastonite (tabular spar), some varieties of gypsum, talc, etc. If the plates are approximately parallel about a common center the structure is said to be concentric. When the laminae are thin and separable, the structure is said to be foliated or micaceous. Mica is a striking example, and the term *micaceous* is often used to describe this kind of structure.

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260. Granular Structure. —The particles in a granular structure differ much in size. When coarse, the mineral is described as coarse-granular; when fine, fine-granular; and if not

distinguishable by the naked eye, the structure is termed impalpable. Examples of the &t'st may be observed in granular crystalline limestone, sometimes called saccharoidal; of the second, in some varieties of hematite; of the last, in some kinds of sphalerite.

The above terms are indefinite, but from necessity, as there is every degree of fineness of structure among mineral species, from perfectly impalpable, through all possible shades, to the coarsest granular. The term *micro-crystalline* has been used for varieties in which the grains are distinct, and *amorphous* for those in which they are not discernible, although an indistinct crystalline structure can be proved by the microscope.

Granular minerals, when easily crumbled in the fingers, are said to be friable.

291. Imitative Shapes. — The following are important terms used in describing the imitative forms of massive minerals.

Reniform: kidney-shaped. The structure may be radiating or concentric. Ex. hematite.

Botryoidal: consisting of a group of rounded prominences. The name is derived from the Greek (*botrys*, a bunch of grapes, Ex. limonite, chalcedony, prehnite.

Mammillary: resembling the botryoidal, but composed of larger prominences. Ex. malachite.

Globular: spherical or nearly so; the globules may consist of radiating fibers or concentric coats. When attached, as they usually are, to the surface of a rock, they are described as implanted globules.

Stippled: in tuberoso forms, or having irregular protuberances over the surface.

Amygdaloidal: almond-shaped, applied often to a rock (as diabase) containing almond-shaped or sub-globular nodules.

Coralloid: like coral, or consisting of interlaced fleshy branchings of a white color, as in the variety of aragonite called *coralloid*.

Dendritic: branching tree-like, as in crystallized gold. The term *dendritic* is used for similar forms even when not crystalline, as in the dendrites of manganese oxide, which form on surfaces of limestone or are inclosed in " moss-agates."

Moss-like: like moss in form or appearance.

Filiform or Capillary: very slender and long, like a thread or hair; consists ordinarily of a succession of minute crystals. Ex. millerite.

Acicular: slender and rigid, like a needle. Ex. stibnite.

Reticulated: net-like. See Art, 248.

Drusy: closely covered with minute implanted crystals. Ex. quartz.

Stalactitic: when the mineral occurs in pendent columns, cylinders, or elongated cones. Stalactites are produced by the percolation of water, holding mineral matter in solution, through the rocky roofs of caverns. The evaporation of the water produces a deposit of the mineral matter, and gradually forms a long pendent cylinder or cone. The internal structure may be imperfectly crystalline and granular, or may consist of fibers radiating from the central column, or there may be a broad cross-cleavage. The most familiar example of stalactites is afforded by calcite. Chalcedony, gibbsite, limonite, and some other species, also present stalactitic forms.

The term amorphous is used when a mineral has not only no crystalline

structure.

Amorphous.

form or imitative shape, but does not polarize the light even in its minute particles, and thus appears to be destitute wholly of a crystalline structure internally, as roost opat. Such a structure is also called colloid or jelly-like, from the Greek *kolle* (see p. 6), for glue. The word amorphous is from a privative, and *morphē*, shape.

262, Pseudomorphous Crystals.—Every mineral species has, when distinctly crystallized, a definite and characteristic form. Occasionally, however, crystals are found that have the form, both as to angles and general habit, of a certain species, and yet differ from it entirely in chemical composition. Moreover, it is often noted in such cases that, though in outward form complete crystals, in internal structure they are granular, or waxy, and have no regular cleavage. Even if they are crystalline in structure the optical characters do not conform to those required by the symmetry of the faces.

Such crystals are called pseudomorphs, and their existence is explained by the assumption, often admitting of direct proof, that the original mineral has been changed into the new compound; or it has disappeared through some agency, and its place been taken by another chemical compound to which the form does not belong. In all these cases the new substance is said to be a pseudomorph after the original mineral.

Common illustrations of pseudomorphous crystals are afforded by malachite in the form of cuprite, limonite in the form of pyrite, barite in the form of

quartz, etc. This subject is further discussed in the chapter on Chemical Mineralogy.

PAET II. PHYSICAL MINERALOGY.

263. The physical characters of minerals fall under the following

heads:

I. Characters depending upon Cohesion and Elasticity—viz., cleavage, fracture, tenacity, hardness, elasticity, etc.

II. Specific Gravity, or the Density compared with that of water.

III. Characters depending upon Light—viz., color, luster, degree of transparency, special optical properties, etc.

IV. Characters depending upon Heat—viz., heat-conductivity, change of form and of optical characters with change of temperature, fusibility, etc.

V. Characters depending upon Electricity and Magnetism.

VI. Characters depending upon the action of the senses—viz., taste, odor, feel.

264. General Relation of Physical Characters to Molecular Structure.—It has been stated on pp. 5, 6 that the geometrical form of a crystallized mineral is the external evidence of the internal molecular structure. A full knowledge in regard to this structure, however, can only be obtained by the study of the various physical characters included in the classes enumerated above.

Of these characters, the specific gravity merely gives indication of the atomic mass of the elements present, and further, of the state of molecular aggregation. The first of these points is illustrated by the high specific gravity of compounds of lead; the second, by the distinction observed, for example, between carbon in the form of the diamond, with a specific gravity of 3.5, and the same chemical substance as the mineral graphite, with a specific gravity of only 2.

All the other characters (except the relatively unimportant ones of Class VI) in general vary according to the direction in the crystal; in other words, they have a definite orientation. For all of them it is true that directions which are crystallographically identical have like physical characters.

In regard to the converse proposition—viz., that in all directions crystallographically dissimilar there may be a variation in the physical characters, an important distinction is to be made. This proposition holds true for all crystals, so far as the characters of Class I are concerned; that is, those

depending upon the cohesion and elasticity, as shown in the cleavage, hardness, the planes of molecular gliding, the etching figures, etc. It is also true in the case of pyro-electricity and piezo-electricity.

It does not apply in the same way with respect to the characters which influence the propagation of light (and radiant heat), the change of volume with change of temperature; further, electric radiation, magnetic induction, etc.

Thus, although it will be shown that the optical characters of crystals are in agreement in general with the symmetry of their form, they do not show all the variations in this symmetry. It is true, for example, that all directions are optically similar in a crystal belonging to any group under the hexagonal system; but this is obviously not true of its molecular cohesion, as may be shown by the cleavage. Again, all directions in a tetragonal crystal at right angles to the vertical axis are optically similar; but this again is not true of the cohesion. These points are further elucidated under the description of the special characters of each group.

I. CHARACTERS DEPENDING UPON COHESION AND ELASTICITY.

255. Cohesion, Elasticity.—The name cohesion is given to the force of attraction existing between the molecules of one and the same body, in consequence of which they offer resistance to any influence tending to separate them, as in the breaking of a solid body or the scratching of its surface.

Elasticity is the force which tends to restore the molecules of a body back into their original position, from which they have been disturbed, as when a body has suffered change of shape or of volume under pressure.

The varying degrees of cohesion and elasticity for crystals of different minerals, or for different directions in the same crystal, are shown in the prominent characters: cleavage, fracture, tenacity, hardness; also in the gliding-planes, percussion-figures or pressure-figures, and the etching-figures.

258. Cleavage.—Cleavage is the tendency of a crystallized mineral to break in certain definite directions, yielding more or less smooth surfaces. It obviously indicates a minimum value of cohesion in the direction of easy fracture—that is, normal to the cleavage-plane itself. The cleavage parallel to the cubic faces of a crystal of galena is a familiar illustration. An amorphous body (p. 6) necessarily can show no cleavage.

As stated in Art. 31, the consideration of the molecular structure of crystals shows that a cleavage-plane must be a direction in which the molecules are closely aggregated together; while normal to this the distance between successive layers of molecules must be relatively large, and hence this last is the direction of easy separation. It further

follows that cleavage can exist only parallel to some possible face of a crystal, and, further, that this must be one of the common fundamental forms. Hence in cases where the choice in the position of the axes is more or less arbitrary the presence of cleavage is properly regarded as showing which planes should be made fundamental. Still Main, cleavage is the same in all directions in a crystal which are crystallographically identical.

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Cleavage is defined, (1) according to its direction, as cubic, octahedral, rhombohedral, basal, prismatic, etc. Also, (2) according to the ease with which it is obtained, and the smoothness of the surface yielded. It is said to be perfect or eminent when it is obtained with great ease, affording smooth, lustrous surfaces, as in mica, topaz, calcite. Inferior degrees of cleavage are spoken of as distinct, indistinct or imperfect, interrupted, in traces, difficult. These terms are sufficiently intelligible without further explanation. It may be noticed that the cleavage of a species is sometimes better developed in some of its varieties than in others.

U7. *GI* BTfg* 1q tha *DI* Sraent iTtHW.—(1) In the *isometric* system, cleavage is cubic, as in *fluorite* and *calcite*; this is the common one, as illustrated by *fluorite* and *calcite*. It is also of the octahedral—that is, parallel to the octahedral faces, in *fluorite* and *calcite*. Less frequently it is dodecahedral, or parallel to the faces of the rhombic dodecahedron, as in *sphalerite*.

In the *tetrahedral* system, cleavage is tetrahedral, or parallel to the faces of the tetrahedron, as in *zinc blende*; also prismatic, or parallel to one (or two) of the square prisms, as in *quartz*; less frequently it is pyramidal, or parallel to the faces of the square pyramid, as in *beryl*.

In the *trigonal* system, cleavage is usually either basal, as in *beryl*, or prismatic, parallel to one of the six-sided prisms, as in *nepherite*; pyramidal cleavage, as in *pyrite*, is rare and imperfect.

In the *rhombic* division, besides the basal and prismatic cleavages, rhombic cleavage, parallel to the faces of a rhombohedron, is also common, as in *calcite* and the allied species.

In the *orthorhombic* system, cleavage parallel to one or more of the rhombic faces is common. Thus in *topaz*, and in all three rhombic directions with *anhydrite*. Prismatic cleavage is also common, as in *barite*; in this case the arbitrary position of the crystal may make this cleavage parallel to a "horizontal" or "dome."

In the *monoclinic* system, rhombic cleavage, parallel to the rhombic faces, is common, as in *orthoclase*, *gypsum*, *hemimorphite* and *enclase*; also basal, as in the

in: cnsaid orihocliise. or parallel to the nrlbopiiiiatoid; also prMtnad'c. as with amphibolc. Leiw frequently cleivage is parallel to a heiiii-pyramiil, 04 wilh gypsnm,

III the TBiCLisic BVSTBM, il la uatial iind proper lo so select the fundamental foi-m as to niske the cleiiviige liireclious correspond with the plnacoids.

IGS. iu some cases clenvase which is ordinarily not observed maybe developed by a shsip blow or by snililen change of temperature. Thus, qimrlz is ii=ually con»ip!ciLonsly free from cleavage, but a ■41111112 crystrl heated and phingcd Inio cold water ofien showa pUnes of separaliou* parallel to both the + and - rhnmboliedrons and to the priftn ns well. Similarly, the prismatic cleavage of pyroxene la observed with treat distinctness in iiiiId sections, made by grinding, while not so readily noted in laree cryiials

When the cleavage is parallel lo a closed form—that is. when it Is cubic orlnhedra) .I'Klecnhcdral, or rhonibolKdral (also pyramidal in the letragoual. hexagonal and onhol rhombic syateiiml—a solid resembling a cr\ -Bial may often he broken out from a slnelo crystnlline individual, nod al! the fraLrmenta have the same angles. Il is in geiieimi eiisv to distinguish such a clenvnge form, as a cleavage ocUhedron ot fluorite. from a ini'a irys'al by the splintery diaracter of the faces of the fonner.

2J9. Cleavage and iMtw.-Tbc face of a crystiil parallel to which there is perfect rlravBL-e often .tliows a pearly luster (see p, tB91, due to the partfil separation of the crysmI into parallel plates. Tins is illustrated by th« Im.sal plane of apophyllite, iho diiiopiuacind of sidbite and heulim dita. An iridescent play of colors is also often seen M with calcite, when ihe separation has been sufflcenl lo produce ihe prismatic colors by interference.

260. GUdiny-planefc—Closely related to the cleavage directions in their connection with the cohesion of the molecnlea of a crystal are the glidihg-

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US PHYSICAL UINSEALOOT.

planet,* or directions parallel to which a elipping of the molecnles may take plitce under the applicatiou of mechaolical force, as bj preaaure.

This may have the reaiik of aimplly producing a auparatioti into layers in the given direction, or, ou the other hand, and more commonly, there may be

4<jg a revolution of the molecula into a new twinning-

poaition, bo that secondary twinning-lameUte are

) formed.

Thus, if a crystal of halite, or rock salt, be subjected to gradual pressure in the direction of a dodecahedral face, a plane of separation is developed normal to this and hence in the direction of another face of the same form. There are six such directions of molecular slipping and separation in a crystal of this substance. Certain kinds of mica of the biotite class often show pseudo-crystalline faces, which are undoubtedly secondary in origin—that is, have been developed

by pressure exerted subsequently to the growth of the crystal (cf. Fig. 475).

In general, the base, (001) . Normal to the plane of perfect cleavage, is a glide plane. This is a slip of the molecules without their separation may be made in the direction of perfect cleavage $\{100\}$. A single crystal supported near the ends and pressed downward by a vertical edge is readily bent, or kinked, in this direction without the parts beyond the support being affected.

261. Secondary Twinning.— The other case mentioned in the preceding article, where molecular slipping is accompanied by a half-revolution (180°) of the molecules into a new twinning-position (see p. 118 et seq), is well illustrated by calcite. Pressure upon a cleavage-fragment may result in the formation of a number of thin lamellae in twinning-position to the parent mass, the twinning-plane being the obtuse negative rhombohedron, (0112) . Secondary twinning-lamellae similar to these are often observed in natural cleavage-masses of calcite, and particularly in the grains of crystalline limestone, as observed in thin sections under the microscope.

Secondary twinning-lamellae may also be produced (and are often noted in nature) in the case of the triclinic feldspars, pyroxene, barite, etc. A secondary lamellar structure in quartz has been observed by Judd, in which the lamellae consisted of right-handed and left-handed portions.

By the proper manipulation a conical twin may be artificially produced by pressure. *Tim. feldspar* in prismatic form. say 6-8 mm. in length

and 3-6 mm. in breadth. be pinned with the obtuse edge on a firm support, and pressed by the blade of an ordinary table-knife on the other edge (at n, Fig. 476). the result is that the portion of the crystal lying between a and b is reversed in position, as if twinned parallel in the horizontal plane (0112) . If this is done, the twinning surface, if perfectly smooth, and the re-entrant angle corresponds exactly with that required by theory.

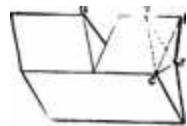
262. Parting. —The secondary twinning-planes described

are often directions of an easy separation—conveniently called parting—which may be mistaken for cleavage. The basal parting; of pyroxene is a common example of such pseudo-cleavage; it was long mistaken for cleavage. The basal and rhombohedral $(10\bar{1})$

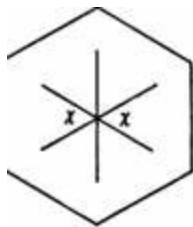
and the less distinct prismatic (1150) parting

" From the German, *Uitfiihen*.

f The lamellar stmctiie of a musivo mineral, wllbout Iwinnlng, may also be the cauae of a fracture which cat) be mistaken for cleavage.



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of cornndum; the octahedral parting of magnetite (cf. Fig. 456, p. 136), are other examples.

An important distinction between cleavage and parting is this: parting can exist only in certain definite planes—that is, on the surface of a twinning-jamella—while the cleavage maj take place in ajig plane having the given direction.

863. PeronMlon-flgUTBB.—Immediately connected with the gliding-planes ■re the figures—ctilled percussion-figures* —produced upon a crystal section by a blow or preasure with a suitable point. In sucli cases, the method described serves to develop more or lees well-defined cracks whose orientation varies with the crystallographic direction of the surface. 477.

Thus upon the cubic face of a crystal of halite a four-njed, star-shaped figQre is produced with arms parallel to the diagonals—that is, parallel to the dodecahedral r-faces. On an octahedral face a three-rayed star is obtained.

The percussion-figures in the case of the micas have beeu often investigated, and, as remarked later, they form n Dieaus of fixing the true orientation of a cleavage-plate having no crystalline outlines. The figure (Fig. 477) is ^ here a six-rayed star one of whose branches la parallel to the clinopinacoid (6), the others approximately parallel to the intersection edges of the prism (m) and base (c).t

Pressure upon a mica plate produces a less distinct Bix-r»ed star, diagonfd to that just named; this is called a pressKve-jigure (Germ. Dntckfiache).

set. flotartioD-pluN.— In the cum of many cryslala, ft fi posatble lo prove tbe ezlBence

ofcen lu dlrectioDB, orstnicturs-planes, in whlcL clie m leal at: I Ion takes place moat
reail I ly— for example, when a crystal ia undeTKraiLt preiaure. These directinne of
ciemlcal weakue» have been csllcd iolmUva--piaiu9. Ttmy often mantfeet lliemsclvea by
the preatnce of a muliltuda of oiieoted caviuea of cryaianlne ouillne (Go-cal1ed oegatlve
crystals) In tbe given illrectioD.

Tbe«e Bolutlon-planes in certain caves, aa ahown by Judd, are the same aa tbe dlrectiona
of aecondnry lamellar twlnniue, as la Illustrated by calclte. Connected with thla Is the
tchMtriaation (tee Art. MS) observed In certain mluemls lo rocks (as dlalJoge, BChlller-
■par).

269. Xtohing'^vni. —Intimately connected with tbe general subjects here considered, of
cohesion in relation to crystals, are the figures produced by etching on cryetalUne faces;
these are often called etchtng-figures.X This method of investigation, developed
particularly by Baumhauer, is of high importance as revealing tbe molecular structure of
the crystal faces under examination, and therefore tbe symmetry of the crystal itself.

The etching is performed mostly by solvents, as by water in some cases, more generally
the ordinnrv mineral acids, or caustic alkalies, also by steam at : liieh pressure and
hydrofluoric acid; the last is especially powerfal in its action, and is used frcc|uently with
the silicates. The figures produced are in

'From llie OennsD, Bchlagflgitren. The percussion-flKures are best nbtlued it the ayslal
plate under Investigation be supported upon a liard cushion and a blow be struck with a
li^t hammer upon a steel rod the slightly rounded point of which Is held flrmly igahiat tte
surface.

tCf. Walker, Am. J. Be , 2, 5, ISH. and o. Priedel. Bull. Soc. Mb., 19, IS. 1866. Walker found
the angle oppoatte ((010) (x In Fig. 477) to be 08* to H* for muscovile, GO*

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the majoritj of cases angular depressioiiiB, such as low triaugular or quadri-

lateral pyramids, whose out!in

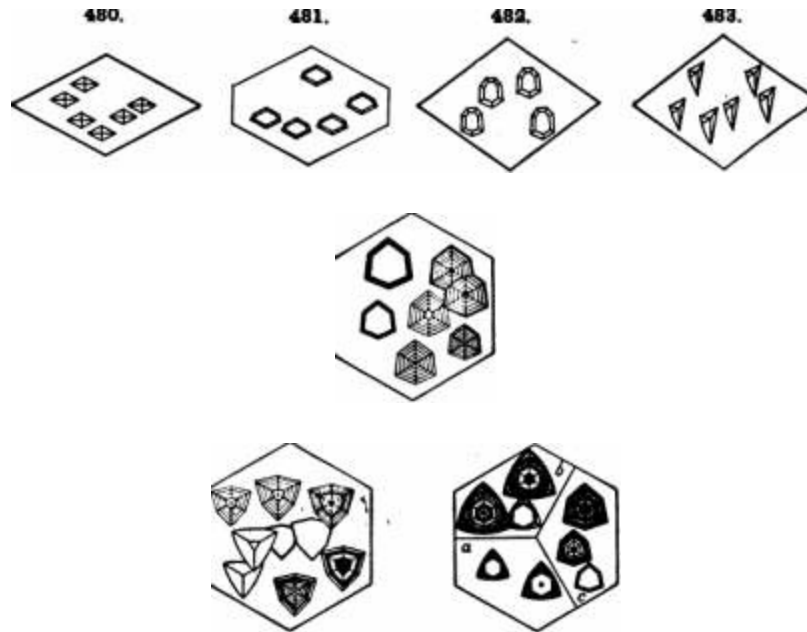
run parallel to some of the crystalliL edges, ill some cases the planes produced can he
referred to occurriug crystallo-graphic faces. They nppcar alike on similar (aces of
crystals, ai)d hence serve to , distiiigiiieh different forms, perhaps in

appearance identical, as the two sets of faces in the ordiuary donhle pyramid of ! quartz;
so, too, they reveal the compound

twinning-Structure common on some crystals, as Quartz and aragonite. Further, their form in general corresponds to the symmetry of the group to which the given Quartz. High Quartz. Irregular crystal belongs. They thus reveal the

liuudeJ crystal hmiert rryeial. trapezohedral symmetry of quartz and the

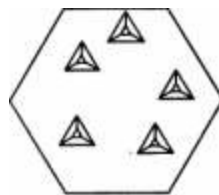
difference between a right-handed and left-handed crystal (Figs. 478, 479); the distinction between calcite and dolomite (Figs. 483, 484); the distinctive character of apatite, pyromorphite, etc.; the hemimorphic symmetry of calamine and nepheline (cf. Fig. 220, p. 73), etc.; they also prove by their form the monoclinic crystallization of muscovite and other micas (Fig. 481).



The shape of the etching-figures may vary with the same crystal with the nature of the solvent employed, though their symmetry remains constant. For example, Fig. 484 shows the figures obtained with aqua regia by the action of sulphuric acid. Fig. 485 by the same diluted, and Fig. 486 by hydrochloric acid of different degrees of concentration.

Of the same nature as the etching-figures artificially produced, in their

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Related to the symmetry of the crystal, are the markings often observed on

the natural faces of crystals. 4B^ 4gg

These are sometimes acc-

ooidary, caused by a natural
etching process, but are more
often an irregularity in the
crystalline development of
the crystal. The inverted
triangular depressions often
seen on the octahedral faces
of diamond crystals are an
example. Fig. 487 shows
lateral depressions, rhombo-

hedral in character, observed on corundum crystals from Montana (Pratt).

Fig. 488 shows a twin crystal of fluorite with natural etching-figures (Pirsson);
these are minute pyramidal depressions whose sides are parallel to the faces
of the trapezohedron (311).

266. Corrosion Forms.—If the etching process spoken of in the preceding article—whether
natural or artificial—is continued, the result may be to destroy the original crystalline
surface and to substitute for it perhaps a multitude of minute elevations, more or less
distinct; or, further, new faces may be developed, the crystallographic position of which
can often be determined, though the symbols may be complex. This is illustrated by Fig.
489 of beryl; here x is the berylloid (36-34-60-5). The mere loss of water in some cases
produces certain corrosive forms (see Pape, literature).

Penfield subjected a sphere of quartz (from a simple rhombohedral individual) to the
prolonged action of hydrofluoric acid. It was found that it was attacked rapidly at the
extremities of the vertical axis, but barely at the horizontal extremities of the base. Figs. 489,
490, 491 show the form remaining. After the sphere had been etched for seven weeks;

489. 490. 491.



Fig. 400 is a basal view; Fig. 491, a front view; the circle shows the original form of the sphere, the dotted hexagon the position of the axes. Compare also the results of Heyer on calcite (see literature. p. 155).

267, Fracture. —The term fracture is used to define the form or kind of surface obtained by breaking in a direction other than that of cleavage in crystallized minerals, and in any direction in massive minerals. When the

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PHYSICAL MINERALOGY.

Fracture is highly perfect in several directions, as the rhombohedral cleavage of calcite, fracture is often readily obtainable.

Fracture is defined as:

(a) Conchoidal; when a mineral breaks with curved concavities, more or less deep. It is so called from the resemblance of the concavity to the valve of a shell, from *concha*, a shell. This is well illustrated by obsidian, also by flint. If the resulting forms are small, the fracture is said to be small-conchoidal; if only partially distinct, it is subconchoidal.

(b) Even; when the surface of fracture, though rough with numerous small elevations and depressions, still approximates to a plane surface.

(c) Uneven; when the surface is rough and entirely irregular; this is true of most minerals.

(d) Hackly; when the elevations are sharp or jagged; broken iron. Other terms also employed are earthy, splintery, etc.

268. Hardness.—The hardness of a mineral is measured by the resistance which a smooth surface offers to abrasion. The degree of hardness is determined by observing the comparative ease or difficulty with which one mineral is scratched by another, or by a file or knife.

In minerals there are all grades of hardness, from that of talc, impressible by the finger-nail, to that of the diamond. To give precision to the use of this character, a scale of hardness was introduced by Mohs.* It is as follows:

1. Talc. 6. Orthoclase.
2. Oypium. 7. Quartz.
3. Calcite. 8. Topaz.
4. Fluorite, 9. Sapphire.
5. Apatite. 10. Diamond.

Crystals in varieties with smooth surfaces should be taken so far as possible.

If the mineral under examination is scratched by the knife-blade as easily as calcite its hardness is said to be 3; if less easily than calcite and more so than fluorite its hardness is 3.5. In the latter case the mineral in question would be scratched by fluorite but would itself scratch calcite. It need hardly be added that great accuracy is not attainable by the above methods, though, indeed, for purposes of the determination of minerals exactness is quite unnecessary.

It should be noted that minerals of grade 1 have a greasy feel to the hand; those of grade 2 are easily scratched by the finger-nail; those of grade 3 are rather readily cut, as by a knife; of grade 4, scratched rather easily by the knife; grade 5, scratched with some difficulty; grade 6, barely scratched by a knife, but distinctly by a file—moreover, they also scratch ordinary glass. Minerals as hard as quartz (H. = 7), or harder, scratch glass readily and are little touched by a file; the few species belonging here are enumerated in Appendix B; they include all the gems.

269. Sclerometer.—Accurate determinations of the hardness of minerals can be made in various ways, one of the best being by use of an instrument called a sclerometer. The mineral is placed on a movable carriage, with the

* The interval between 3 and 8, and 5 and 6, in the scale of Hobs, being a little greater than between the other numbers, Breithaupt proposed a scale of twelve minerals; but the scale of Hobs is now universally accepted.

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Surface to be experimented upon horizontal; this is brought in contact with a steel point (or diamond point), fixed on a support above; the weight is then determined which is just sufficient to move the carriage and produce a scratch on the surface of the mineral.

By means of such an instrument the hardness of the different faces of a given crystal has been determined in a variety of cases. It has been found that different faces of a crystal {e.g., cyanite) differ in hardness, and the same face may differ as it is scratched in different directions. In general, differences in hardness are noted only with crystals which

show distinct cleavage; the iijirdeBt face is that which is intersected by the plane of most complete cleavage. Further, of a single face, which is intersected by cleavage-planes, the direction perpendicular to the cleavage-direction is the softer, those {nruUel to it the harder.

Tlii" Euliject has been investigated by Kiiier (p. 1S5}, who has given the form of the esru of bardnett for lUe dlfferuul fiices iif many ciystaU. These cqi ves ai-e obtained as fo lows ^ the least weight reqiiiii'ed to Benitcli n crystBltine Biirrnre Id different directions, fur eacli lO" Or 16'. from 0° to 180°, is delermim-ii wiith the sclerometer; these directions ire laid nff as radii f rojii a center, niid the length of each Is irtnila propnrllonal to the weight nxml by eKperiment—that is. lo the harrtvesa thus deieimintd; the line couueotiug the Mirfniilies of these rndii is the curve of Iianlnesa Tor the given face.

The Following table gives the resulis ohla-tied ■ (see tileratiire) in comparing the hard* tiesauf llie miDt'Dils of tlie scale from corundum, No. 9, taken as IDOO, to gypsum, No. 3. PfaS useil [be muthoil of boring with a stHndard point, the hardness lieing determied by ilie number of roiatlous: Kosfwal used h stiudaril powder lo grind the surface. Jnggar emplu/ed his microBclerouielier, the method being esscDllally a modification of that of Pfiiff. By means of this instrument he is able to test the hardness of the minerals present ina tijin section under the microscope. Measurements of absolute baf dneM have also been made by Auerbach.

VttB. 1884. Roslml, IBOI. Jaggar. IStT.

». Corundum 1000 1000 1000

8. Topaz 459 188 152

» 7. Quana 264 149 40 ,

«. Orthoclae 161 38'7 2S

S. Apatite 88-5 6-30 1-38

4 Ffuorito 878 470 -76

8. Calcito 15-8 3-68 -26

a. Gypsum 13 08 84 04

tro. BaUtion et Haidnea to OhrailDal Cttmpadtinu—Some general facts of importance «u be Slated t in regard to the conoectiou butweea lbe bnr dneas of a mineral and its chemical composition.

1. Compounds of the heavy metals, as silver, copper, mercuri^-, lead, etc., are «^, Their

hardness seldom exceeding 3½ to 8.

Among the compounds of the common metals, the sulphides (arsenides) and oxides of iron (also of nickel and cobalt) are relatively hard (e.g., for pyrite H. = 6 to 8; for hematite H. = 5, etc.); here belong also columbite, iron niobate, tantalite, iron tantalate, wolframite, iron tungstate.

2. The sulphides* are mostly relatively soft (except as noted in 1). Also most of the carbonates, sulphates, and phosphates.

3. Hydrated salts are relatively soft. This is most distinctly shown among the silicates—e.g., compare the feldspars and zeolites.

4. The conspicuously hard minerals are found chiefly among the oxides and aluminates; many of them are compounds containing aluminium—e.g., corundum, diaspore, chrysoberyl, and many aluminosilicates. Outside of these the borate, boracite, is hard (H. = 7); also iridosmine.

On the relation of hardness to specific gravity, see Art. MO.

* The minerals are here given as tabulated by Jsggar. — See further in Appendix B.

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871. *Prüfung des Buffens*.—Seven things should be regarded in the trials of hardness:

(1) If [the mineral is] slightly allured, it is often the case with corundum, garnet, etc., the surface may be readily scratched with the nail itself; a trial with the edge of the nail will often give a correct result in such a case.

(2) A mineral with a granular surface often appears to be scratched when the grains have been only torn apart.

(3) A relatively soft mineral may leave a faint white ridge on a surface, as of glass, which may be mistaken for a scratch if carelessly observed.

(4) A crystal of quartz, is often slightly scratched by the edge of another of the same species and like bodies.

(5) The scratch should be made in such a way as to disfigure the specimen as little as possible.

272. *Teigigkeit*.—Minerals may be either brittle, sectile, malleable, or flexible.

(a) Brittle; when parts of a mineral separate in powder or grains on attempting to cut it, as calcite.

(5) Sectile; when pieces may be cut off with a knife without falling to powder, but still the mineral pulverizes under a hammer. This character is intermediate between brittle and malleable, as gypsum.

(c) Malleable; when slices may be cut off, and these slices flattened out under a hammer; native gold, native silver.

(d) Flexible; when the mineral will bend without breaking, and remain bent after the bending force is removed, as talc.

The tenacity of a substance is properly a consequence of its elasticity.

273. Elasticity—The elasticity of a solid body expresses at once the resistance which it makes to a change in shape or volume, and also its tendency to return to its original shape when the deforming force ceases to act. If the limit of elasticity is not passed, the change in molecular position is proportional to the force acting, and the former shape or volume is exactly resumed; if this limit is exceeded, the deformation becomes permanent, a new position of molecular equilibrium having been assumed; this is shown in the phenomena of gliding-planes and secondary twinning, already discussed. The magnitude of the elasticity of a given substance is measured by the coefficient of elasticity, or, better, the coefficient of restitution. This is illustrated by the relation, for example, between the elongation of a bar of unit section to the force acting to produce this effect; similarly of the bending or twisting of a bar. The subject was early investigated acoustically by Savart; in recent years, Voigt and others have made accurate measures of the elasticity of many substances and of the crystals of the same substance in different directions. The elasticity of an amorphous body is the same in all directions, but it changes in value with change of crystallographic direction in all crystals.

The distinction between elastic and inelastic is often made between the species of the mica group and allied minerals. Muscovite, for example, is described as "highly elastic," while phlogopite is much less so. In this case it is not true in the physical sense that muscovite has a high value for the coefficient of elasticity; its peculiarity lies rather in the fact that its elasticity is displayed through unusually wide limits.

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II. SPECIFIC GRAVITY OR RELATIVE DENSITY.

274. Definition of Speoiflo Gravity.—The specific eraTity of a mineral is the ratio of its deosity • to that of water at 4° C. (39-2° I.). This relative detisitj may be learned in auy case by comparing the ratio of the weight of a certain volume of the given substance to that of an equal volume of water; heuce ilio specific gravity ia often defined as: the weigki of the body divided by Ike n-eig'.l of ail equal volume af water.

The statement that the specific gravity of graphite is 2, of corundum 4, of galena 7'5, etc., means that the densities of the minerals named are 2, 4, and 7'5, etc., times that of water; in other words, as familiarly expressed, any volume of them, a cubic inch for exitmple, weighs 2 times, 4 times, ?' & times, etc., as much as a like volume, a cubic iicli, of water.

Strictly speaking, since the density of water varies with its expansion or contraction under change of tempenanre, the comparison should be made with water at a fixed temperature, niunely 4" C (39'2° F.), at which it has its maai-miim density. If made at a Jiiglier temperature, a suitable correction should be introduced by calculation. Practically, however, since a high degree of accuracy is not often called for, and, indeed, in many

cases is impracticable to attain in consequence of the nature of the material at hand, in the ordinary work of obtaining the specific gravity of minerals the temperature at which the observation is made can safely be neglected. Common variations of temperature would seldom affect the value of the specific gravity to the extent of one unit in the third decimal place.

For the same reason, it is not necessary to take into consideration the fact that the observed weight of a fragment of a mineral is less than its true weight by the weight of air displaced.

Where the nature of the investigation calls for an accurate determination of the specific gravity (e.g., to four decimal places), no one of the precautions in regard to the purity of material, exactness of weight-measurement, temperature, etc., can be neglected. The accurate values spoken of are needed in the consideration of such problems as the specific volume, the relation of molecular volume to specific gravity, and many others. Some of these have been discussed by Schroder, Hunt, also (for salts) by Playfair and Joule, etc.

275. Determination of the Specific Gravity by the Balance.—The direct comparison by weight of a certain volume of the given mineral with an equal volume of water is not often practicable. By making use, however, of a familiar principle in hydrostatics, viz., that a solid immersed in water, in consequence of the buoyancy of the latter, loses in weight an amount which is equal to the weight of an equal volume of the water (that is, the volume it displaces)—the determination of the specific gravity becomes a very simple process.

The weight of the solid in the air (w) is first determined in the usual manner; then the weight in water is found (w'); the difference between these weights—that is, the loss by immersion ($w - w'$)—is the weight of a volume of

* The density of a body is strictly that of one unit volume. Thus if a cubic centimeter of water (at its maximum density, 4° C. or 39.2° F.) is taken as the unit of mass, the density of any body—as gold—is given by the number of grams of mass (at 15° C.) in a cubic centimeter; in this case the same number, 19, gives the relative density or specific gravity. If, however, a pound is taken as the unit of mass, and the cubic foot as the unit of volume, the mass of a cubic foot of water is about 62.5 lbs., that of gold about 1200 lbs., and the specific gravity is the ratio of the second to the first, or about 19.

† Cf. *Philosophy of Berkeley* in *Proc. Roy. Soc.*, 11, p. 180B.

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water equal to that of the solid; finally, the quotient of the first weight (w) by that of the equal volume of water as determined ($w - w'$) is the specific gravity (G). Hence,

A common method of obtaining the specific gravity of a firm fragment of a mineral is as follows: First weigh the specimen accurately on a good chemical balance. Then suspend it from one pan of the balance by a horse-hair, silk thread, or, better still, by a fine platinum wire, in a glass of water conveniently placed beneath, and take the weight again with the same care; then use the results as above directed. The platinum wire may be wound around the specimen, or where the latter is small it may be made at one end into a little spiral support,

276. The Jolly Balance.—Instead of using an ordinary balance and determining the actual weight, the spiral balance of Jolly, shown in Fig. 412, may be conveniently employed; this is also suitable when the mineral is in the form of small grains. The mineral is first placed on the upper pan, and the amount that the spring is stretched noted by the scale number (V), opposite to which the index line comes to rest (the eye sees the reflection in the mirror and thus avoids error by parallax). If from V , be subtracted the number n , expressing the amount to which the scale is stretched by the weight of spring and pans alone, the difference will be proportional to the weight of the mineral. Next, the mineral is placed in the lower pan, d , immersed in the water, and again the corresponding scale number, A^* , read. The difference between these readings ($V - JV$) is a number proportional to the loss of weight in water. The specific gravity is then

It is obviously necessary to have the wires supporting the lower pan immersed to the same depth in the case of each of the three determinations. If care is taken the specific gravity can be obtained accurate to two decimal places.

277. Pycnometer.—If the mineral is in the form of grains or small fragments, the specific gravity may be obtained by use of the pycnometer. This is a small bottle (Fig. 493) having a stopper which fits tightly and ends in a tube with a very fine opening. The bottle is filled with distilled water, the stopper inserted, and the overflowing water carefully removed with a soft cloth. It is now weighed, and also the mineral whose density is to be determined. The stopper is then removed and the mineral in powder or in small fragments inserted with care, so as not to introduce air-bubbles.* The water which overflows on replacing the stopper is the amount of water displaced by the mineral. The weight of the pycnometer with the enclosed mineral is determined, and the weight of the water lost is obviously the difference between this last weight and that of the bottle and mineral together, as first determined.

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The specific gravity of the mineral is equal to its weight alone divided by the weight of the equal volume of water thus determined. Where this method is followed with sufficient care, especially avoiding any change of temperature in the water, the results may be highly accurate.

If the mineral forms a porous mass, it may be first reduced to powder, but 493. It is to be noted that it has been shown by Rose that when-

ical precipitates have uniformly & higher density than belongs to the same substance in a less finely divided state. This increase of density also characterizes, though to a less extent, a mineral in a fine state of mechanical subdivision. It is ex-

posed by the condensation of the water on the surface of the powder. 278. Use of Liquids of High Density.—It is often found convenient both in the determination of the specific gravity and in the mechanical separation of fragments of different

specific gravities (e.g., to obtain pure material for analysis, or again in the study of rocks) to use a liquid of high density—that is, a so-called heavy liquid. One of these is the solution of mercuric iodide in potassium iodide, called the Stoufflet solution or Thoulet solution. When made with care* it has a maximum density of nearly 3.3, which by dilution may be lowered at will.

A second solution, often employed, is the potassium nitrate of cadmium, having a maximum density of 3.6. This again may be lowered at will by dilution, observing certain necessary precautions. Still a third solution of much practical value is that proposed by Brauns, methyl iodide, which has a specific gravity of 3.35. A number of other solutions, more or less practical, have also been suggested (see papers referred to in the literature on p. 180, which also give the necessary directions for the use of the liquids). When one of these liquids is to be used for the determination of the specific gravity of fragments of a certain mineral it must be diluted until the fragments just float and the specific gravity then obtained, most conveniently by the Westphal balance (Art. 279).

When, on the other hand, the liquid is to be used for the separation of the fragments of two or more minerals mixed together, the material is first reduced to the proper degree of fineness, the dust and smallest fragments being sifted out, then it is introduced into the solution and this diluted until one constituent after another sinks and is removed. For the convenient application of this method a suitable tube is used and certain precautions must be observed; compare the papers noted in the literature (p. 180), especially one by Penfield.

278. Westphal's Balance.—The Westphal balance is conveniently used to determine the specific gravity of a mineral when a heavy solution is employed (278). It consists essentially of a graduated arm, upon which the weights

of specific gravity of a mineral when a heavy solution is employed (Art. 278). It consists essentially of a graduated arm, upon which the weights in the form of lead, are placed. These should be so adjusted that the balance in free equilibrium

he given liquid while the Index at llic end poluts lo the zeri of the scale ami shows ' the arm is horizontal (cf. Itoaenbscli. Mikr. Phys. Hiu., p. 24G). The grndnatlou usually allows of the spcciBc gravity being rend off diieclly without ca'culntloQ.

SBO. Balation of Dsntltr to Hardniu, GhBiiii«ftl Compoaitlim. (to.—The densily, or specific gravity, of a solid depends, flnt, upon the aature of the cheniicitl substances wjiicli it con-Xnlofi. and, second, upon the stale of molecular ag!;regiliuD.

Thus, as iiu illuBtratioQ of the first polal, all liiid couipouuds have a high deusiiy (O. = about S), since lead U a heavy melal, or, chemically expressed, has a high atomic weigliC (206'4) Sltnllarly, bftrium sulphate, barite, lias a specitic gravity of 4 5, while for

t See the directions by Goldschmltdt, refereuce od p. 160.

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csldum aulphaU or anhydrite the value Is OD]y 3'95 (at. weight for barium 187, for caldum about 40).

On the other hand, wblle aluminium la a metal of low denally (o. = 2'S and atomic weight = 27), fia oxide, corundum, has a remarkably Aigh deoBlty (Q. = 4) aud Is also verr luudtH. =9). Again, carbon (at. weight = 12) bass b!gh density In the diamond (o.=8'5) aDdlowlnBTaphlle(o. = 2): also, the Brit U hard (H. = 10). iLe second soft (H. = 1'6). In these and •unilar cases the high deuslty ilxnifles great molecular aggregaiioD, and hence It la natural that It should be accompaQted by great faardnesa aad reautauce lo the attack of aclda.

Aa bearing upon iLIs point, It Is to be noted that the density of many Bubsauce Is altered by fusion. Again, the same mineral in different states of molecuHr aggregation may dlOer (but only slightly) In density. Furthermore, minerals havlDK the same chemical composition have sometimes dISerent densities, corresponding to llie different cryslallioe forms in which lthey appear. Thus in (be case of calcium carbonate (CaCOi), calcite his G. = 3-7, aragonlte lias G. = i-0.

Ml. ATirage Bpsstfls SraTltiss.—It is to be noted chat among minerals of unuetaluc LDBTER the aMraga specific gravity ranges from 26 to 8. Here belong quartz {2'M), calcille (2-7), the feldspars (3'6-2'7G). muscovlte (28). A speclQc gravity of 2'S or less Is loa, and Is chanicterUtlc of soft minerals, and often those whilcb are bydruus {e.g., gypsum, O. =: 2'8). The common species fluorite, tourmaline, apatite, vesuvisuUe, ampblbcie. prrozene, and epldote lie juSl above the limit given, uamely, 'iO lo 3'G. A specific gnivlly o13'5 or above Is relatively hig/i, and belongs to bard minerals (ns corundum, see An. StO), or lo those conlaEoIug a heavy metnl, as compounds of Btronlium, barium, also iron, tungsten, copper, silver, lead, mercury, etc.

With minerals of metallic lustkb, thie iiverage is about 5 (here belong pvriie, beinmiie,

etc.), while if below 4 it is relatively low (2, 3); if 7 or above, relatively high (as galena, 7.6).

Tables of minerals arranged according to their specific gravity are given in Appendix B.

2.2. Constant of Specific Gravity.— The specific gravity of a mineral species is a character of fundamental importance, and is usually constant for different specimens of the same species, if pure, free from cavities, solid inclusions, etc., and if essentially constant in composition. In the case of many species, however, a primary or less variable exists in the chemical composition, and this at once causes a variation in specific gravity. The different kinds of garnet illustrate this point: also the various minerals intermediate between the iron (and manganese) and the niobate, varying from $D = 3.7$ or above to

US. Practical Geology.— It should be noted that the determination of the specific gravity is possible unless the fragment taken is pure and is free from impurities, internal and external, and not porous. Care must be taken to exclude air-bubbles, and it will often be found well to moisten the surface of the specimen before inserting it in the water, and sometimes boiling (or the use of the vacuum-pump) is necessary to free it from air. If it absorbs water this latter process must be allowed to go on until the substance is fully saturated. No accurate determinations can be made unless the changes of temperature are rigorously excluded and the actual temperature noted.

In a mechanical mixture of two constituents in known proportions, when the specific gravity of the whole and of one are known, that of the other can be readily obtained. This method is often important in the study of rocks.

It is to be noted that the method may be soon learned to detect a difference of specific gravity, if like volumes are taken. even in a small fragment—this is the difference between calcite and barite, even a slight difference between a small diamond and a quartz crystal, can be detected.

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ni. CHABACTEES DEPENDING UPON LIGHT.

GENERAL PRINCIPLES OF OPTICS.

SB4. Before considering the optical caractere of minerals in general, and mriTe particulurlj thioee thiat belong to thie crj'Btals of the different syetems, it is desirable to review brieSy some of the more important principles of optics upon which the phenomena in qiieation depend.

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285. Sature of Light—The propagation of light from a luminous body, ae the sun or a caTidie-flame, is believed to be accompliahed by a very rapid wave-motion* in the medium culled the Inminiferous ether, which, it is assumed, pervades all space as well us all material bodies.

286. The Etiier—The assumption of the medium called the ether is necessary, since without this it is impossible to explalin the transmission of light throtigli K|i:ice where no ordinary medium (as the air) is present. Furtlicrmore, as thie velocity of light even within solid media, though less than that in a vacuum or in air, is still enormously rapid, it is inconceivable that it should be propagated by the molecules of the body; hence the

assnmp-tiou, otherwise verified, that the ether pervades all material bodies. The properties of the ether, however, are modified by the molecular structure of the given body, as is proved by the fact that the velocity of light varies with the chemical nature of the substance, and also in certain cases with the direc-

* It is generally accepted that light is an electromagnetic phenomenon and that the nature of the periodic motion in the ether by which light is propagated is that it is resolved in the transmission of electric waves produced, for example, by a very rapid oscillatory electric discharge between two spark-knobs. In fact these electric waves have been shown to travel with the same velocity as light-waves, and to exhibit like phenomena of reflection, refraction, polarization, etc.; hence they are believed to differ from light-waves only in their much greater length. For the purposes of the present work, however, light-waves are treated of as if a mechanical phenomenon, but all assumption of vibrations of the "elasticity of the ether" in crystals as an explanation of the observed variation of light-velocity is avoided.

A reference is made to an article by Clerk Maxwell in the *Encyclopaedia Britannica* for a discussion of the general properties of the luminiferous ether.

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tion in the uncrystallized medium as corresponding to its particular molecular structure.

287, Wave-motion in General—A familiar example of wave-motion is given by the series of concentric waves which on a surface of smooth water

radiate out from a center of disturbance, as the point where a pebble has been dropped in. These surface-waves are propagated by a motion of the water-particles which is transverse to the direction in which the waves themselves travel; this motion is given from each particle to the next adjoining, and so on. Thus the particles of water at any one spot oscillate up and down,* while the wave moves on as a circular ridge of water of constantly increasing diameter, but of diminishing height. The ridge is followed by a valley, indeed both together properly constitute a wave in the physical sense. This compound wave is followed by another wave and another, until the original impulse has exhausted itself.

Another familiar kind of wave-motion is illustrated by the sound-waves which in the free air travel outward from a sonorous body in the form of concentric spheres. Here the actual motion of the layers of air is forward and back—that is, in the direction of propagation of the sound—and the effect of the transfer of this impulse from one layer to the next is to give rise alternately to a condensed and rarefied shell of air, which together constitute a sound-wave and which expand in spherical waves of constantly decreasing intensity (since the mass of air set in motion continually increases). Sound-waves, as of

the voice, may be several feet in length, and they travel at a rate of 1110 feet per second at ordinary temperatures.

28S. It is important to understand that in both the cases mentioned, as in every case of free wave-motion, each point on a given wave may be considered as a center of disturbance from which a system of new waves tend to go out, these individual wave-systems ordinarily destroy each other except so far as the onward progression of the wave as a whole is concerned. This is further discussed and illustrated in its application to light-waves (Art. 292 and Figs. 495, 496).

In general, therefore, a given wave is to be considered as the resultant of all these minor wave-systems. If, however, a wave encounters an obstacle in its path, as a narrow opening (i.e., one narrow in comparison with the length of the wave) or a sharp edge, then the fact just mentioned explains how the waves seem to bend about the obstacles, since new waves start from them as centers. This principle has an important application in the case of lightwaves, explaining the phenomena of diffraction (Art. 308).

SBS. But no other case of wave motion may be mentioned, since it is particularly helpful in giving the correct impression of light-phenomena. If a long rope, attached to a wall at one end, be moved up and down, will give rise to a half-wave—on one side a crest, on the other a trough—which will travel quickly to the other end. If it be reflected back with a reversal in its position; that is, if it went forward as a crest-like wave, it will return as a trough. If, instead, the wave is reflected (the end of the rope) like one that is started, the two will meet and pass in the middle, but here for a brief interval the rope is at rest, since it feels two equal and opposite impulses. This will be seen later to be a case of the simple interference of two like waves opposed in phase.

Again, a double motion of the hand, up and down, will produce a complete wave, with a crest and a trough, as the result, and this again is reflected back as in the simpler case. Still again, if a series of fingers move rhythmically and so timed that each finger is at an even part of the whole rope, the two systems of equal and opposite waves* (moving in the two directions) will interfere and a system of so-called stationary waves will

* Strictly speaking, the path of each particle approximates closely to a circle.

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be the result, the rope seeming to vibrate in segments to and fro about the position of equilibrium.

Finally, if the end of the rope be made to describe a small circle at a rapid, uniform, rhythmic rate, a system of stationary waves will again result, but now the vibrations of the string will be equally in circles about the central line. This last case will be seen to

roughly indicate the kind of transverse vibrations by which the waves of circularly polarized light are propagated, while the former case represents the vibrations of waves of what is called plane-polarized light.

All these cases of waves obtained with a rope deserve to be carefully examined and studied by experiment, for the sake of the assistance they give to an understanding of the complex phenomena of light-waves.

290. Wave-length, Amplitude, etc.—In the cases mentioned, as in all kinds of simple wave-motion, the length of a wave is the distance from any one particle of the medium to the next which is moving in the same direction with the same velocity, or, technically expressed, which is in the same phase. The amplitude of the wave is the excursion to or from its position of equilibrium made by each particle in succession. Further, the wave-system travels onward the distance of one wave-length in the time that a given particle makes a complete excursion to and from.

291. Light-wave.—The propagation of ether-waves involves the same fundamental principles as the familiar forms of wave-motion just considered. Here the motion of the medium is transverse to the direction of propagation, and this motion may be regarded as communicated from one set of particles to the next and so on, the ether-waves traveling as concentric spherical waves (in an isotropic medium) outward in all directions from the luminous point.

The nature of the vibrations will be better understood from Fig. 494. If AB represents the direction of propagation of the light, each particle of ether must vibrate at right angles to this as a line of equilibrium. The vibration of the first particle induces a similar movement in the adjacent particle; this is communicated to the next, and so on. The particles vibrate successively; from the line AB to a distance corresponding to $\frac{1}{2}\lambda$, they return to the position of equilibrium, then return to b and pass on to b' , and so on. Thus at a given instant there are particles occupying all positions, from that of the $\frac{1}{4}\lambda$.

extreme distance $\frac{1}{2}\lambda$, or $\frac{1}{2}\lambda$, from the line of equilibrium to that on this line. In this way the wave moves forward, while the motion of the particles is only transverse. The distance between any particle and the next which is in a like position—i.e., of like phase, as $\frac{1}{2}\lambda$ and $\frac{3}{2}\lambda$ —is the wave-length; and the time required for this completed movement is the time of vibration, or vibration-period. The intensity of the light varies with the amplitude of the vibration, and the color, as explained in a later article, depends upon the length of the waves; the length of the violet waves is about one-half the length of the red waves.

In ordinary light the transverse vibrations are to be thought of as taking place in all planes about the line of propagation. In the above figure, vibrations in one plane only are represented; light which is thus one-sided or has only one direction of transverse vibration is said to be plane-polarized.

Light-waves have a very minute length, only 0.000023 of an inch for the yellow sodium flame, and they travel with enormous velocity, 186,000 miles

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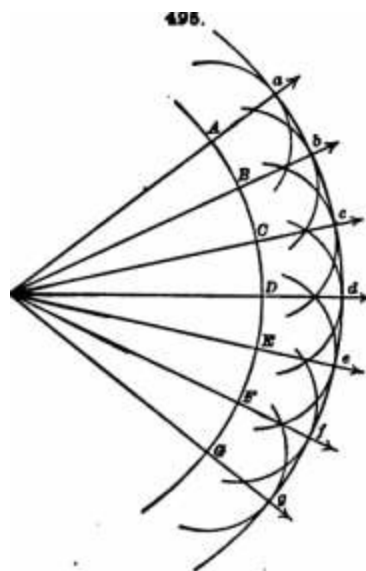
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per second in a vacuum; the light passes from the sun to the earth in about eight minutes. The time of one oscillation, or time of one oscillation, is consequently extremely brief; it is given by dividing the distance traveled by light in one second by the number of waves included.

292. Wave-front.—In an isotropic medium, as air, water, or glass—that is, one in which light is propagated in all directions about the luminous point with the same velocity—the waves are spherical in form. The wave-front is the continuous surface, in this case spherical, which includes all particles which commence their vibration at the same moment of time. Obviously the curvature of the wave-front diminishes as the distance of the source of light increases, and when the light comes from an indefinitely great distance (as the sun) the wave-front becomes sensibly a plane surface. Such waves are usually called plane waves. These cases are illustrated by Figs. 495 and 496. In Fig. 495 the luminous point is supposed to be at O , and the medium being isotropic, it is obvious that the wave-front, as $JBC^{\wedge} \dots O$, is spherical. It is also made clear by this figure how, as briefly stated in Art. 288, the resultant of all the individual impulses which go out from the successive points, as a, b, c , etc., as centers, form a new wave-front, $dbc \dots g$, concentric with $ABC \dots G$. In Fig. 496 the luminous body is supposed to be at a great distance, so that the wave-front $AB \dots F$ is a plane surface. Here also the individual impulses from A, B , etc., unite to form the wave-front $ab \dots f$ parallel to $AB \dots F$.

293. light-ray.—The study of light-phenomena is, in certain cases.



facilitated By the conception of a 496.

t-ray, a line drawn from the Inminous

Point to the wave-front, and whose direction is taken so as to represent that of the wave itself. In Fig. 49. OA, OB, etc., are diverging light-rays, and in Fig. 496 OA, OB, etc., are parallel light-rays. In both these cases, where the medium is assumed to be isotropic, the light-ray is normal to the wave-front. This is equivalent to saying that the light-wave moves onward in a direction normal to the wave-front.

It must be understood that the "light-ray" has no real existence and is to be taken only as a convenient method of representing the direction of motion

of the light-waves under varying conditions. Thus when by appropriate means

by the use of

{e.g., the use of lenses) the direction of the wave-front is altered—for example, if from being a plane surface it is made sharply convex—then the light-rays, at first parallel, are said to be made to diverge. Again, if the convex wave-front is made plane, the diverging light-rays are then said to be made parallel.

294. Wave-length. Color. White light.—Notwithstanding the very small length of the waves of light, they can be measured with great precision. The visual part of the waves going out from a brilliantly incandescent body, as the glowing carbons of an electric arc-light, may be shown to consist of waves of widely varying lengths. They include red waves whose length is about seven-eighths of an inch and waves whose length constantly diminishes without break, through the orange, yellow, green, and blue to the violet, whose minimum length is about half of that of the red. The length of each group of these waves determines the sensation of color which the eye perceives. This color is strictly monochromatic only when it corresponds to one definite wave-length; this is nearly true of the bright-yellow sodium line, though strictly speaking this consists of two sets of waves of slightly different lengths.

The effect of "white light" is obtained if all the waves from the red to the violet come together to the eye simultaneously; for this reason a piece of platinum at a temperature of 1500° C. appears "white hot."

The radiation from the sources named, either the sun, the electric carbons, or the glowing platinum, includes also longer waves which do not affect the eye, but which, like the light-waves, produce the effect of sensible heat when received upon an absorbing surface, as one of lampblack. There are also, particularly in the radiation from the sun, waves shorter than the violet which also do not affect the eye. The former are called infra-red, the latter ultraviolet waves.

299. Complementary Color.—The sensation of white light mentioned above is also obtained when to a given color—that is, light-waves of given wavelength—is combined a certain other so-called complementary color. Thus certain shades of pink and green combined, as by the rapid rotation of a card on which the colors form segments, produce the effect of white. Blue and yellow of certain shades are also complementary. For every shade of color in the spectrum there is another one complementary to it in the sense here defined. The most perfect illustration of complementary colors is given by the examination of sections of crystals in polarized light, as later explained.

296. Reflection.—When light-waves come to the boundary which separates one medium from another, as a surface of water, or glass in air, they are, in general, in part reflected or returned back into the first medium.

The reflection of light-waves is illustrated by Figs. 497 and 498. In Fig. 497, MM is the reflecting surface—here a plane surface—and the light-waves have a plane wave-front {Abcde); in other words, the light-rays (OA, Of, etc.) are parallel. It is obvious that the wave-front meets the surface first at A and successively from point to point to E. Each of these points is to be regarded as the center of a new wave-system which unimpeded would be propagated onward in a given time distances equal to the lines Aa' Bb', etc. Hence the common tangent to the circular arcs drawn with these radii from A, B, etc., represents the direction of the new or reflected wave-front. But geometrically the angle eAE is equal to fEA, or the incident and reflected wave-fronts make equal angles with the reflecting surface. If XA is a normal at A, the angle OAV called the angle of incidence—is equal to 2fAF, the angle of reflection. Hence the familiar law:

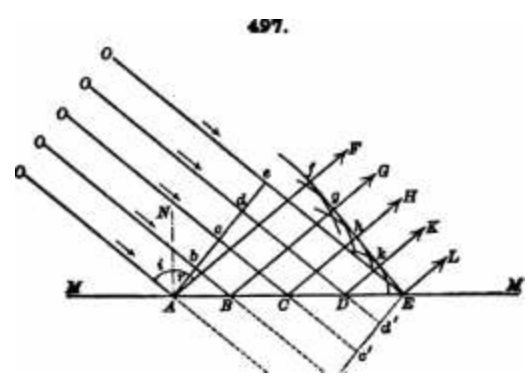
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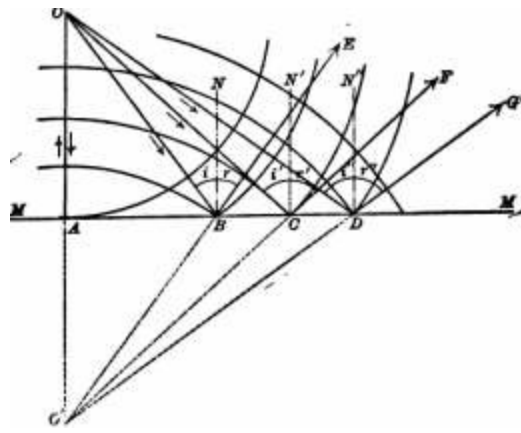
The angle of incidence is equal to the angle of reflection, the incident and reflected rays lie in the same plane with the normal to the reflecting surface.

Angle of reflection.

The rays both lie in the same plane



In Fig. 49S, where the Incident point is at O, the waves go out from it will meet the plane mirror MM at the point A and successively at points.



M B, C, D, etc., farther away to the right (and left) of A. Here also it is easy to show that all the new impulses, which have their centers at A, B, C, etc., must together give rise to a series of reflected waves whose center is at O', at a distance equally distant from MM measured on a normal to the surface {OA - OA'}.

Now the lines OA, OB, etc., which are perpendicular to the wave-front, represent certain incident light-rays, and the eye placed in the direction BE,

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OF, etc., will see the luminous point as if at O'. It follows from the construction of the figure and can be proved by experiment that if BN, ON, etc. are normals to the mirror the angle of incidence, OBN, OCN', etc., are equal to the angles of reflection, NBE, N'BF, etc., respectively. Hence the above law applies to this case also.

If the reflecting surface is not plane, but, for example, a concave surface, as that of a spherical or parabolic mirror, there is a change in the curvature of the wave-front after reflection, but the same law still holds true.

The proportion of the reflected light decreases with the smoothness of the surface and also as the angle of incidence diminishes. The intensity of the reflected light is a maximum for a given surface when the angle of incidence is 90°. (Fig. 498).

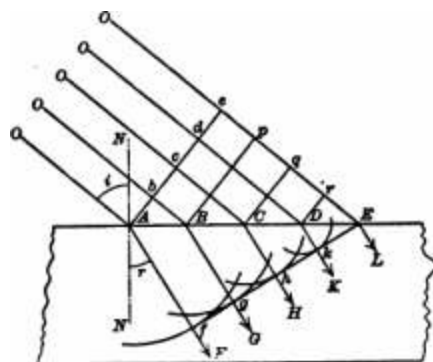
If the surface is not perfectly polished, diffuse reflection will take place, and (here will be) the reflected light. It is the diffusely reflected light which makes the reflected surface visible; if the surface of a mirror were absolutely smooth the eye would see the reflected body itself, not the surface itself. Optically speaking, the surface is to be considered smooth if the distance between the scratches upon it is considerably less (say one-fourth) than the wave-length of light.

297. Refraction.—When a system of light-waves of the same wave-length passes from one medium into another there is, in general, an increase or decrease in the velocity of

the light,

and this results in the phenomenon of refraction—that is, a change of direction at the bounding surface. The principles applicable here can be most easily shown in the case of light-waves with a plane wave-front, as shown in Fig. 499—that is, where the light-rays OA, OB, etc., are parallel. Suppose, for example, that a light-wave, part of whose wave-front is Abcde, passes from air obliquely into glass, in which its velocity is about two-thirds as great, and suppose the surface of the glass to be plane. The points A, B, etc., will be successively centers of disturbance which will be propagated in a given time, not to distances equal to bE (from A in the line o^), topE^, etc., but only two-thirds of these distances. Circles drawn from the points A, B, C, etc., with radii equal to these diminished values (two-thirds of eE.pE, etc.), will have a common tangent in the plane fghkE, and this will be then the new wave-front in the second medium. Here it is seen that there is a change of direction in the wave-front, or otherwise stated, in the light-ray, the magnitude of which depends on the ratio between the light-velocities in the two media, and, as discussed later, also upon the wave-length of the light. The light-ray is here said to be broken or refracted, and for a medium like glass, optically denser than air {i.e., with a lower value of the light-velocity), the refraction is toward the perpendicular. In the opposite case—in an optically rarer medium—the refraction is away from the perpendicular, the angle of refraction is larger than that of incidence Art, 303).

398. Refractive Index.—It is obvious from the figure that whatever the direction of the wave-front—that is, of the light-ray—relatively to the given surface, the ratio of tE to Af, which determines the direction of the new



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wave-front (t.e., the direction of a refracted ray, AF) is constant. This ratio is equal to — where V is the value of the light-velocity for the first medium

(Here air) and v for the second (as glass). If this constant ratio be represented by n , we may write:

$$\frac{\sin i}{\sin r} = \frac{v_1}{v_2} = n$$

Here i ($\angle OAN$) is the angle of incidence and r ($\angle FAN$) the angle of refraction; thus, in its last form.

we have the familiar relation usually expressed as follows:

The sine of the angle of incidence bears a constant ratio to the sine of the angle of refraction.

It is also true that the incident and refracted rays lie in a common plane with the normal to the surface.

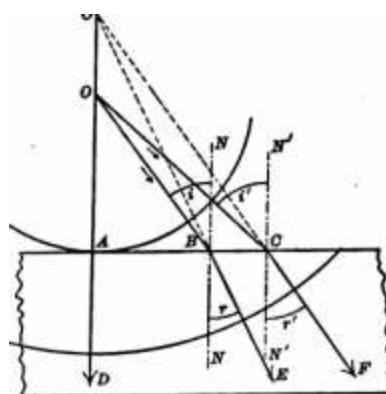
The above relation holds true for any wave-system of given wave-length in passing from one medium into another, whatever the wave-front or shape of the bounding surface. In Fig. 500 the luminous point is at O , and it can be readily shown that the new wave-front propagated in the second medium (of greater optical density) has a flattened curvature and corresponding to this a

center at O' where $r = \dots$ Here

the incident rays OB, OC , are refracted at B and C , the corresponding refracted rays being BE and BF . For, in this case also the relation holds good,

$$\frac{\sin i}{\sin i'} \cdot n = \dots; \text{ etc. } \frac{\sin r}{\sin r'}$$

This constant ratio for light of a given wave-length passing from one medium to another, expressed here by n , is called the index of refraction or refractive index. In the examples given for air and



the refractive index, or n , for this kind of glass.

The relation between wave-length and refractive index is spoken of in Art. 309.

If the bounding surface is not plane but curved, as in lenses, there is a change in the curvature of the wave-front in the second medium, but the

so long as the medium is isotropic.

* See S. P. Thompson (Light Visible and Invisible. 1897) who develops the formulae etc., on the basis of light-waves instead of light-rays.

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PHYSICAL OPTICS.

299. Relation of Refractive Index to Light-velocity.—The discussion of the preceding article shows that if n is the refractive index of a given substance for waves of a certain length, referred to air, V the velocity in air and v the velocity in the given medium, then

For two media whose indices are n_1 and n_2 , respectively, it consequently follows that

Therefore, The indices of refraction of two given media for a certain wavelength are inversely proportional to their relative light-velocities.

Since light-waves are propagated by a transverse motion in the ether which pervades the given body, and is as it were weighted down by its molecules, it is obvious that the velocity of the light-wave itself is measured by the rate of this transverse motion in the ether; hence for waves of the same length traveling through media of different refractive power, this latter velocity of transverse vibration is inversely proportional to the refractive indices.

300. Principal Refractive Indices.—The refractive index has, as stated, a constant value for every substance referred, as is usual, to air (or it may be to a vacuum). In regard to solid media, it is evident from Art. 298 and will be further explained later that those which are isotropic, viz., amorphous substances and crystals of the isometric system, can have but a single value of this index. Crystals of the tetragonal and hexagonal (and rhombohedral) systems have, as later explained, two principal refractive indices, c and v , corresponding to the velocities of light-propagation in certain definite directions in them. Further, all orthorhombic, monoclinic, and triclinic crystals have similarly three principal indices, a , b , γ . In the latter cases of so-called anisotropic media, the mean refractive index is taken, namely, as the arithmetical mean

301. Examples of Refractive Indices.—The following table includes the values of n for a variety of substances, for sodium light. For minerals other than those of the isometric system the average value (as defined in the preceding article) is given here.

The refractive index for air referred to the ether of a vacuum is 1.000292 for a wavelength equal to that of yellow sodium light ($\lambda = 0.000589$ cm.).

SOS. I^mUo BafnwtlTa ?ow*r.—The relation between the refractive Index and the chemical composition of a given substance is expressed by what has been called the Gladstone law,* namely.

Here n is refractive Index (for anisotropic substances, the mean index), and d is the density. The value of the constant is called the specific refractive power. The product of the specific refractive power into the molecular weight gives the refractive equivalent. Thus for quartz, n = 1.5, d = 2.65, therefore the value of the specific refractive power is 0.2, and the refractive equivalent U equal to this number multiplied into the molecular weight (80) or 16 (H = 1 X 16). Similarly the value obtained for CaO is 0.17, and for MgO 0.17.

In the case of a complex molecule, it is assumed that the sum of the refractive equivalents of the parts of the molecule divided by the sum of the corresponding molecular weights is equal to the specific refractive power of this given compound. Thus for grossular garnet whose formula may be written 3CaO . 3SiO₂, the above relation gives

U = 0.17.

Further, -

$$U = \frac{3 \times 16 + 3 \times 18}{3 \times 80 + 3 \times 60} = 0.17$$

U = 0.17, and n = 1.5796; experiment gives n = 1.5747.

303. Total Reflection. Critical Angle.—In regard to the principle stated

in Art. 298 and expressed by the equation $\sin r = n_2 \sin i$, two points are to be

noted. First, if the angle $i = 0^\circ$, then $\sin i = 0$, and obviously also $r = 0$; in other words, when the ray of light (as OA, Fig. 500) coincides with the perpendicular, no change of direction takes place, the ray proceeds onward (AD) into the second medium without deviation, but with a change of velocity.

Again, if the angle $i = 90^\circ$, then $\sin i = 1$, and the equation above

becomes $n_2 = n_1 \sin r$

1

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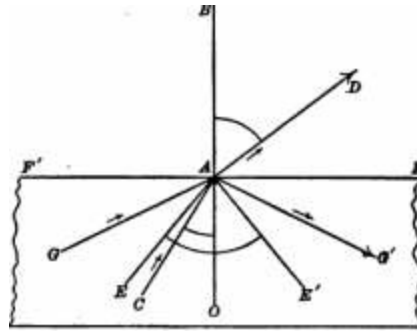
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n has a fixed value for every substance, it is obvious that there will also be a corresponding value of the angle i - for the case mentioned. From the above table it is seen that

for water, $\sin r$

1

and $r = 48^\circ 1$



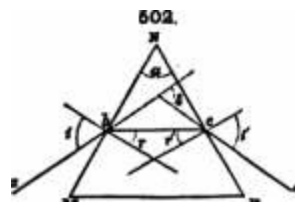
31'; for crown glass, $\sin t = \wedge$

and $r = 38^\circ 27'$; for diamond, \sin

$= \wedge;'$ - and $r = 24^\circ 25'$.

In Fig. 501 the ray OA in the glass is refracted on passing into the air in the direction AD, but if the angle EA O = $38' 27'$, the ray EA will graze the surface or take the direction AF. Any ray, OA, for which the angle GA O is greater than $38^\circ 27'$ will not emerge at all, but suffer total reflection, being returned in the direction AG'. The surface of glass illuminated from beneath in the direction last named has a brilliant, almost metallic luster. This is the

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appearance also, under like conditions, of the surface of a transparent solid— for example, of a glass prism or a cut gem. The value of r thus found is called the critical angle; the smaller this angle the greater the apparent brilliancy of the given substance.

The relative refractive power of a given substance compared with that of the enclosing medium {e.g., Canada balsam with $n = 1.539$) determines on the principle of total reflection whether the surface appears rough with dark cracks ("high relief") as in garnet

and zircon, or smooth and even ("low relief") as in quartz.

30C Determination of the Refractive Index.—B7 means of a prism, as MNP in Fig. 503, it is possible to determine the value of n , or refractive index of a given substance. The angle of the prism MNP, a , is, in each case, measured in the same manner as the angle between two faces of a crystal, and then the minimum amount of deviation (J) of a monochromatic ray of light, *fig.* yellow sodium light, passing from a slit through the prism is also determined. The amount of deviation of a ray in passing through the prism varies with its position; but when the prism is so placed that the ray makes equal angles with the sides of the prism, that is, with the normals ($i = i'$, Fig. 503), when entering and emerging, this deviation has its minimum value.

If δ = the minimum deviation of the ray, and a = the angle of the prism, then

$$\delta = 2i - a$$

The application of this method is given in a later article. Several other methods are also explained—for example, one depending upon total reflection.

305. Dispersion.—Thus far the change in direction which light suffers in reflection and refraction has alone been considered. It is further true that the amount of refraction differs for the waves of different length, that is, the different colors of which ordinary white light is composed, being greater for blue than for red. In consequence of this fact, if ordinary light be passed through a prism, as in Fig. 502, it will not only be refracted, but it will also suffer dispersion or be separated into its component colors, thus forming the prismatic spectrum.

This variation for the different colors depends directly upon their wavelengths; the red waves are longer, their transverse vibrations are slower, and it may be shown to follow from this that they suffer less change of velocity on entering the new medium than the violet waves, which are shorter and whose velocity of transverse vibration is greater. Hence the refractive index for a given substance is greater for blue than for red light. The following are values of the refractive indices for diamond determined by Schrauf:

2.40845 red (lithium flame).

2.41733 yellow (sodium flame).

2.42549 green (thallium flame).

306. Spectroscopie.—The instrument most simply used for the analysis of the light by dispersion is familiar to all as the spectroscope.* In it the light

* A. de Arago has shown that the direct spectroscopic examination of many minerals

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from the given source, received through a narrow slit at the end of one tube, is made to fall as a plane-wave that is, as a "pencil of parallel rays" upon one surface of a prism at the center, and the spectrum produced is viewed through a suitable telescope at the end of a second tube.

If the light from an incandescent solid—which is "white hot" (Art. 294)—is viewed through the spectroscope, the complete band of colors of the spectrum is seen from the red through the orange, yellow, green, blue, to the violet. If, however, the light from an incandescent vapor is examined, it is found to give a spectrum consisting of bright lines (or bands) only, and these in a definite position characteristic of it—as the yellow line {double line} of sodium vapor; the more complex series of lines and bands, red, yellow, and green, characteristic of barium; the multitude of bright lines due to iron vapor (in the intensely hot electric arc), and so on,

307. Absorption. — Of the light incident upon the surface of a new medium, not only is part reflected (Art. 296) and part transmitted and refracted (Art. 297), but, in general, part is also absorbed at the surface and part also during the transmission. Physically expressed, absorption in this case means the transformation of the ether-waves into sensible heat, that is, into the motion of the molecules of the body itself.

The color of a body gives an evidence of this absorption. Thus a sheet of red glass appears red to the eye by transmitted light, because in the transmission of the light-waves through it, it absorbs all except those which together produce the effect of red. For the same reason a piece of jasper appears red by reflected light, because it absorbs part of the light-waves at the surface, or, in other words, it reflects only those which together give the effect of this particular shade of red.

Absorption in general is selective absorption; that is, a given body absorbs particular parts of the total radiation, or, more definitely, waves of a definite wave length only. Thus, if transparent pieces of glass of different colors are held in succession in the path of the white light which is passing into the spectroscope, the spectrum viewed will be that due to the selective absorption of the substance in question. A layer of blood absorbs certain parts of the light so that its spectrum consists of a series of absorption bands. Certain rare substances, as the salts of didymium, etc., have the property of selective absorption in a high degree. In consequence of this, a section of a mineral containing them often gives a characteristic absorption spectrum.

The dark lines of the solar spectrum, of which the so-called Fraunhofer lines are the most prominent, are due to the selective absorption exerted by the solar atmosphere upon the waves emitted by the much hotter incandescent mass of the sun.

308. Diffraction.—When monochromatic light is made to pass through a narrow slit, or by

the sharp edge of an opaque body, it suffers diffraction, and there arise, as may be observed upon an appropriately placed screen, a series of dark and light bands, growing fainter on the outer limits. Their presence is explained (see Arts. 312, 313) as due to the interference, or mutual reaction, of the adjoining systems of waves of light, that is, the initial light-waves, and further, those which have their origin at the edge or sides of the slit in question. It is essential that the opening in the slit should be small as compared with the wave-length of the light. If ordinary light is employed,

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the phenomena are the same, and for the same causes, except that the bands are successive colored spectra.

Diffraction spectra, explained on the principles alluded to, are obtained from gratings of fine parallel lines (say 15,000 or 20,000 to the inch) ruled with great regularity upon a glass surface or upon a surface of speculum metal. The glass grating is used with the most brilliant light: the Rowland grating of the latter kind has a concave surface. Each

grating gives a number of spectra, of the first, second, third order, &c. These spectra are the most beautiful, as compared with those given by prisms, and are in fact the result of the diffraction of light: the Rowland grating of the latter kind has a concave surface. Each

309. Double Refraction.—In the discussion of Art. 297, applying to isotropic media, it was shown that light-waves passing from one medium into another, which is also isotropic, suffer simply a change in wave-front in consequence of their change in velocity. In anisotropic media, however, which include all crystals but those of the isometric system, there are, in general, two wave-systems propagated with different velocities and only in certain limited cases is it true that the light-ray is normal to the wave-front. This subject cannot be adequately explained until the optical properties of these media are fully discussed, but it must be alluded to here since it serves to explain the familiar fact that, while with glass, for example, there is only one refracted ray, many other substances give two refracted rays, or, in other words, show double refraction.

The most familiar example of this property is furnished by the mineral calcite, also called on account of this property "doubly-refracting spar." If a thin (Fig. 50S) be a cleavage piece of calcite, and a ray of light meet it at S, it will, in passing through, be divided into two rays,

one bright and one dark. For this reason, a dark spot or a line seen through a piece of calcite ordinarily appears double. As implied above and also in Art. 300 the same property is enjoyed by all crystallized minerals, except those of the isometric system. The wide separation of the two refracted rays by calcite, which makes the phenomenon so striking, is a consequence

of the large difference in the values of its indices of refraction, in other words, as technically expressed, it is due to the presence of its double refraction, or its birefringence.

310. When the incident light is perpendicular to the surface of the doubly-refracting substance, there is. In the more commonly occurring cases, no change of direction in the transmission; but even then it is usually still true that the incident ray is divided into two rays, which, though they may travel in the same path, yet have different velocities, so that one falls behind the other. Further, as later explained, each is in general plane-polarized. For each of these rays, it is true that for waves of the same length the rate of transverse vibration, and hence the velocity of the ray itself, is inversely proportional to the respective refractive index.

311. Interference of Waves in General.—The subject of the interference of light-waves, alluded to in Art. 308, requires detailed discussion. It is one of great importance, since it serves to explain many noble and beautiful phenomena in the optical study of crystals, for example, the axial interference fringes shown on the plate forming the frontispiece.

Referring again to the water-waves spoken of in Art. 287, it is easily understood that when two wave-systems, going out, for example, from two centers of disturbance near one another, come together, if at a given point



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they meet in the same phase (as crest to crest), the result is to give the particle in question double amplitude of motion. On the other hand, if at any point the two wave-systems come together in opposite phases, that is, half a wave-length apart, the crest of one corresponding to the trough of the other, they interfere and the amplitude of motion is zero. Under certain conditions, therefore, two sets of waves may unite to form waves of double amplitude; on the other hand, they may mutually interfere and destroy each other. Obviously an indefinite number of intermediate cases lie between these extremes. What is true of the waves mentioned is true also of Bound-waves and of wave-motion in general. A very simple case of interference was spoken of in connection with the discussion of the waves carried by a long rope (Art. 289).

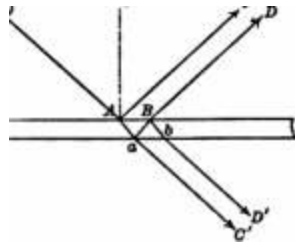
312. Interference of Light-waves.—Interference phenomena can be most satisfactorily studied in the case of light-waves. The extreme cases are as follows: If two waves of like length and intensity, and propagated in the same direction, meet in the same phase, they unite to form a wave of double intensity {double amplitude). If, however, the waves differ in phase by half a wavelength, or an odd multiple of this, they interfere and extinguish

each other. For other relations of phase they are also said to interfere, forming a new resultant wave, differing in amplitude from each of the component waves. In these cases monochromatic light-waves were assumed (that is, those of like length). If ordinary white light is used, the waves in the case of interference will overlap, and their interference will be indicated by the appearance of the colors of the spectrum.

313. Illustration of Interference.—A simple illustration is afforded by the bright colors of very thin films or plates, as a film of oil on water, a soap-bubble, and like cases. To understand these, it is only necessary to remember that the incident light-waves are reflected in 50%

part from the upper and in part from the lower surface of the film or plate. Hence if the thickness is very small, these two reflected wave-systems, when they come together (represented in Fig. 504 by the two rays AC, CD) will differ from one another in phase, and interfering give rise (in ordinary light) to the colored phenomena spoken of. It is to be noted that the phenomena of interference by reflection are somewhat complicated by the fact that there is a reversal of phase (that is, a loss of half a wave-length) at the surface which separates the medium of greater optical density from the rarer one. Hence the actual relation in phase of the two reflected rays, as AC, CD (supposing them of the same wave-length) is that determined by the retardation due to the greater length of path traversed by CD, together with the loss of a half wavelength due to the reversal of phase spoken of. As shown in the figure, there are also two transmitted waves which also interfere in like manner.

A plano-convex lens of long curvature, resting on a plane glass surface (Fig. 505), and hence separated from it, except at the center, by a film of air of varying thickness, gives by reflected monochromatic light a dark center and about this a series of light and dark rings, called Newton's rings. The dark center is due to the interference of the incident and reflected waves, the latter half a wave-length behind the former. The light rings correspond



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to the distances where the two sets of reflected waves meet in the same

position, phase, that is (noting the explanation

above) where the retardation of those

waves is $\frac{1}{2}\lambda$ having the longer path is a half wave-

length or an odd multiple of this ($\frac{1}{2}\lambda$).

(See Fig. 505, etc.). Similarly the dark rings fall between these and correspond to the points where the two waves meet in opposite phase, the retardation being a wave-length or an even multiple of this. The rings are closer together with blue than with red because of their smaller wave-length. In each of the cases described the ring is properly the intersection on the plane surface of the cone of rays of like retardation.

In ordinary white light there can be no dark rings because of the difference of length of the component waves; on the contrary, the overlapping of these waves produces a series of colored rings, each showing the successive colors of the spectrum. The series of colors are distinguished as of the first, second, third, etc., order; for a given color, as red, may be repeated a number of times as the waves overlap. After a certain number of waves have overlapped in this way, white light ("of a higher grade") results.

Similarly in the case of the thin plate in white light, a certain thickness and consequent retardation produces a superposition of the waves which yields, for example, a shade of red; a greater thickness (and retardation) a shade of the second order, etc. If the plate is not very thin, simply white is reflected from it.

The interference of light is a subject of the highest importance in the history of science, and the subject cannot be

furtLer discussed in tllila pllice.

Other cases of the compoaiton of two systems of light-waves will be considered after some remarka on polnrized light.

314. Polarisation and Polarized Light.—Ordinary light is propagated by transverse vibrations of the ether which take place alike in all planes about the line of propagation. A ray of ordinary light is, therefore, alike or aym-melrical in all directions about this line; it may be most simpW thought of as being propagated by two equal sets of transverse vibrations taken in any two planes at right angles to each other.

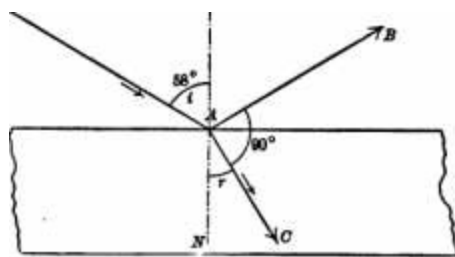
Plane-polarized light, on the other hand, as stated briefly in Art. 291, is propagated by ether-vibrations which take place in one plane OJih/. The change by which ordinary light is changed into a polarized light ia called polarisnfinn, and the plane at right angles to the plane of transverse vibration is called the plane of polarization.*

Polarization may be accomplished (1) by reflection and by single refraction, and (2) by double refraction.

316. Polarisation by Reflection and Single SefrHotion.—In general, light which has suffered reflection from a surface like that of polished glass is more or less completely polarized; that is, the reflected waves are propagated by vibrations to a large extent limited to a single plane, viz, (as assumed) the plane normal to the plane of incidence, which last is hence the plane of polarization. "Furthermore, in this case, the light transmitted and refracted by the reflecting medium is also in like manner partially polarized; that is, the

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of incidence. For a given angle of incidence, varyine for each substance, bnt aach toat the

reflected and refracted rays [AB and AC, Fig. 506) make an angle of 90° with each other, this polarization is a maximum. For this case it is hence true, if we require that this angle of polarization be such that

$\tan i' = n$. This law, established by Brewster, may be stated as follows:

The angle of polarization is that angle whose tangent is the index of refraction of the reflecting substance. For crown glass this angle is about 58° (see Fig. 506). If light suffers repeated reflections from a series of thin glass plates, the polarization is more complete, though its intensity is weakened. Metallic surfaces polarize the light very slightly.

If the polarized light-waves fall upon a second similar reflecting surface at the same angle, they will be reflected again unchanged, on the condition that the two planes of incidence (and hence the two planes of polarization) of the two mirrors coincide.

If, however, these planes are at right angles to each

other, the light polarized by the first mirror will be

extinguished by the second. As the polarization is

in no position absolutely complete, the light is not

completely arrested, but only reduced to a minimum in the second position.

This case is illustrated by Fig. 507. Here the incident ray AH is reflected by the first mirror in the direction AC and polarized in a plane normal to the plane of incidence—the angle ABH being equal to the angle of polarization. If now the second mirror occupy either of the positions op or o'j', the planes of incidence (and of polarization) of both mirrors coincide and the light-ray, BC, is, therefore, reflected a second time in the direction of oD, or oD. If, however, the second mirror be revolved about a vertical axis the reflected light becomes gradually weaker and is sensibly extinguished when the two planes of incidence are at right angles to each other.

316. Polarization by Double Reflection.—When light in passing through a crystalline medium is doubly refracted (Art. 308) or divided into two sets of waves, it is always true that both are completely polarized and in planes at right angles to each other. This subject can only be satisfactorily explained after a full discussion of the properties of anisotropic crystalline media, but it may be alluded to here since this principle gives the most satisfactory method of obtaining polarized light. For this end it is necessary that one of the two wave-systems should be extinguished, so that that due to a single set of vibrations only is transmitted. This is accomplished by natural absorption in the case of tourmaline plates and by artificial means in the Nicol prisms of calcite.



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PHYSICAL OPTICS.

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317. Polarized Light by Absorption.—If from a crystal of tourmaline, which is suitably transparent, two sections be obtained, each cut parallel to the vertical axis, it will be found that these, when placed together with the direction of their axes coinciding, allow the light to pass through. If, however, one section is revolved upon the other, less and less of the light is transmitted, until, when their axes are at right angles (90°) to each other, the light is (almost perfectly) extinguished. As the revolution is continued, more and more light is obtained through the sections, and after a revolution of 180° , the axes being again parallel, the appearance is as at first. A further revolution (270°) brings the axes again at right angles to each other, when the light is a second time extinguished, and so on around.

The explanation of this phenomenon, so far as it can be given here, is analogous to that employed for the case of polarization by reflection. Each plate doubly refracts the light; but one of the two sets of waves is absorbed, and only that set whose vibrations are parallel to the vertical axis are transmitted. If now the two plates are placed in the same position, *abcd*, and e/h (Fig. 508), the light passes through both in succession. If, however, the one is turned upon the other, only that portion of the light can pass through which vibrates still in the direction *ac*. This portion is determined by the resolution of the existing vibrations in accordance with the principle of the parallelogram of forces. Consequently, when the sections stand at right angles to each other (Fig. 509) the amount of transmitted light is nearly zero, that is, the light is extinguished. Instead of tourmaline, an artificial salt, the sulphate of iodoquinine (herapathite) is sometimes employed, but it has little practical value.

318. Polarized Light by Nicol's Prism.—The most satisfactory method of obtaining polarized light is by means of a prism of transparent calcite (Iceland spar). Fig. 510 shows the principle involved in the prism early constructed by Nicol, which transmits one only ray.

of the two refracted rays, that represented by the line *bde* (the extraordinary ray, as later defined). The other ray, *be*, suffers total reflection at the surface where the two sections are united together by Canada balsam and is then absorbed by the black surface of the sides. Here the vertical faces are natural cleavage-faces; the face *PJ'* is ground on so as to make an angle of 68° with the obtuse vertical angle; the prism *BO* formed is cut

diagonally across (////), and then the parts, cemented together. This form of prism, as well as other somewhat different in form but accomplishing the same end with the use of less material, is ordinarily called a Nicol prism, or briefly a nicol. The section of the ordinary nicol of Fig. 510 is



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rhomb-shaped (Fig. 511); the plane of polarization, PP, passes through the acute angles of the rhomb, and the vibration-plane, here as usual represented by a double-headed arrow, passes through the obtuse edges. The other prisms alluded to may have a rectangular cross-section. The vibration-plane can be readily determined in any case by examining with it the light reflected from some suitable surface (e.g., of a wood table). Twice in a revolution of the prism through 180° about its axis this light will be weakened; when this is true, the vibration-plane of the prism must be perpendicular to that of the partially polarized reflected light, that is, it must be vertical, since the latter is taken as horizontal.

319. Polarizing. Polarizer. Analyzer.—The combination of two nicols, or other polarizing contrivances, for the examination of a substance in polarized light is called, in general, a polariscope; the common forms are described later. In any polariscope the prism, or other contrivance, which polarizes the light from the outside source is called the polarizer; the other is the analyzer. If the prisms have their vibration-planes at right angles to each other, they are said to be crossed and the incident light polarized by the polarizer is then extinguished by the analyzer; briefly, it is said to suffer extinction.

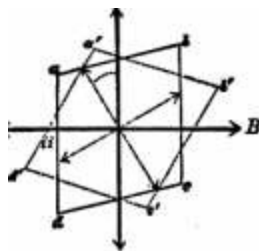
320. Interference of Plane-polarized Waves. Interference-color. — The simplest case of the interference (Art. 312) of polarized light to consider is that where the two light-waves, or, more simply expressed, two rays, are polarized in the same plane. They may then interfere to extinguish each other, or they may give rise to beautiful color-effects.

Suppose, for example, that in a polarization-microscope (Art. 328) parallel light passes upwards through the lower nicol, whose vibration-plane is shown in the cross-section of Fig. 512 by the arrow AA;

this light is polarized in a single plane. Now let

this polarized light pass through a thin cleavage-plate of selenite; it will in general be separated into two rays, each polarized in planes at right angles to the other, having a definite position peculiar to this substance. Thus, in Fig. 512, if $abcd$ represents the selenite plate, its vibration-planes have the directions of the dotted arrows, the two rays corresponding to them travel through the section with unequal velocity, and on emerging one is slightly retarded as compared with the other. Now let these light-rays pass through a second nicol, with its vibration-plane at right

angles to that of the first nicol, that is, in the direction of the arrow BB . Then each of the two sets of vibrations (represented by the dotted arrows) will have a component in the direction of BB , and these will emerge now polarized in the same plane, and hence capable of interfering, for light-rays can only thus completely interfere when their vibrations are in a common plane. Further, an amount of light corresponding to the other components (in the direction AA) will be extinguished. One of these emergent rays is, as stated, slightly retarded as compared with the other. The amount of this retardation obviously varies with the strength of the double refraction (in this case $y - a$), and also with the thickness of the section taken. The interference-color of the section, supposing ordinary light to be used, depends upon these two conditions, and may be calculated for a given substance. Thus a plate of selenite of a thickness of 0.066 mm. will give a red (of first order), and if thinner.



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a yellow or gray. As the thickness increases, the colors (now of the second order) pass through successive shades of blue, green, yellow, orange, and if the plate is of sufficient thickness a second red and so on (see, further, Arts. 369 and 382). A mineral of very strong double refraction, as calcite, shows only the white of the higher order unless extremely thin.

If the section had happened to have the position of $a'b'e'd'$ (Fig. 513), its vibration-planes would have coincided with those of the two nicols, and the light, after passing through the first nicol and the section, would have been propagated by vibrations in the direction AA only, and hence have been completely extinguished by the second nicol. The plate would then have appeared

The successive interference-colors of the first order pass from an indigo-gray through blue-gray to lilac, yellow, and red; then follow indigo, blue, green, yellow, orange, and red of the second order; likewise a similar but paler series of colors of the third order, and

finally the very yellow; aliadta of green and red of the first order. Beyond this the colors are not very distinct; white of a higher order finally results from the interference.

An excellent colored plate allowing these colored bands to be given by Levy and Lacrola (*Les Mineraux des Koches*, 1888). It is so arranged as to give the thickness of the section of a given mineral (all important species present in rocks being included) which will yield any one of the different shades of color mentioned. The use to which such a plate may be put will be the practical determination of the birefringence of a given mineral will be referred

321. Complementary Colors in Polarized Light.—If in the examination of the selenite plate, as just described (Art. 320), one of the nicols had been rotated 90° , or, in other words, if the vibration-planes of the two nicols had been made parallel, then it is obvious that interference would also have taken place between the emerging rays, but the color resulting in each case would have been exactly the complementary tint to that obtained at first when the nicols were crossed. The section in the position 'at right angles' between parallel nicols obviously would appear white.

322. In the preceding articles the two interfering light-rays, after emerging from the second nicol, were assumed to be polarized in the same plane; for them the resulting phenomena as indicated are comparatively simple. If, however, two plane-polarized rays propagated in the same direction have their vibration-directions at right angles to each other, and if they differ one-quarter

of a wave-length (λ) in phase (assuming monochromatic light), then it may easily be shown that the composition of these two systems results in a ray of circularly polarized light, briefly expressed, this is a ray which looked at

end-on would seem to be propagated by ether-vibrations taking place in circles about the line of transmission. From the side, the onward motion would be like that of a screw, and either right-handed or left-handed.

If, again, two light-rays meet as above described, with a difference of phase differing from $\lambda/2$ (but not equal to an even multiple of λ), then the resulting composition gives rise to elliptically polarized light, that is, a light-ray propagated by ether-motions taking place in ellipses.

The above results are obtained most simply by passing plane-polarized light through a doubly refracting medium of the proper thickness (e.g., a mica plate) which is placed with its vibration-plane inclined 45° to that of the polarizer. If the thickness is such as to give a difference in phase of $\lambda/2$ or an odd multiple of this, the light which emerges is circularly polarized. If the phase differs

from $\lambda/2$ (but is not equal to $\lambda/2$ or λ), the emergent light is elliptically polarized. In the following article: also the explanation of the

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The following table from Klein • gives the relation between the retardation δ to W . ($A = \text{wave-length}$) For a section of h doubly refractive substance, the interference-color it yields, and the state of the transmitted light as regards polarization. The section U supposed to be observed in parallel sodium light with crossed nicols; further, the vibration-direction corresponding to the greater refractive index in the section to right be indicated.

\B from left in front

for Na light. δ in μ

*A

iK

iX k

|A

|A δ

»i VA δ

LBTeader-grsj Qmyisli-blue Clearer-gray Pal* straw-yellow Btlgt yellow Brownish-yellow

Red

Indigo Azure-blue

- Brighter green Yellow Orange

Reddish-orange Dark violet-red

Kind of Polarization.

Elliptic, TighMiBoded, Circular, ""

Elliptic,

- Plane-polarized. Elliptic, left-handed.

Elliptic. Plane-polarized. Elliptic, right-handed. Circular, ""

Elliptic, Plaiie-polariud. Eliipic, lefl-banded. Circular, " Elliplio, Plant-polariud.

323. CryataU OiTin; Circular FoUiizatio. — In the case of certain doubly refracting crystallized media (as quartz), and also of certain solutions (as of sugar), it can be shown that the light is propagated by two sets of ether-vibnitious which tal(e place, not in definite transverse planes—as in pliiie-polarized light—but in circles; that is, each ray is circularly polarized, one being ri^ht-handod, the other left-handed. Further, of these rays, one will uniformly gain with reference to the other. The result is, that if a ray of plane-polarized light fall upon such a medium (assuming the simplest case, as of a section of quartz cut normal to the axis), it is found that the two rays circularly polarized within unite on emerging to a plane-polarized ray, but the plane of polarization has suffered an angular change or rotation, which may be either to the right (to one looking in the direction of the ray), when the substance is said to be right-handed, or to the left, when it is called left-handed.

This phenomenon is theoretically possible with all crystals of a given system belonging to any of the groups of lower symmetry than the normal group which show a prismatic development of the faces; or, more simply, those in which the corresponding right and left (or + and —) typical forms are enantiomorphous (pp. 50, 82), as noted in the chapter on crystallography. In mineralogy, this subject is most important with the common species quartz, of the rhombohedral-trapezohedral group, and a further discussion of it is postponed to a later page (Art. 366).

• Ber. Ak, Berlin, 231, 1863.

Of the thirty-two possible groups among crystals, the following eleven may be characterized by circular polarization : Group 4, p. 60 ; 5, p. 61 ; U and 13, p. 68 ; 17, p. 78, S^p. 82 ; 23 and 24, p. 84 ; 37, p. 96 ; S9, p. 103 ; S3, p. 109.

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PHYSICAL MINERALOGY.

OPTICAL INSTRUMENTS AND METHODS.

324. Determination of Refractive Index. Refractometer. — For the determination of the refractive indices of crystallized minerals various methods are employed. The most accurate results, when suitable material is at hand, may be obtained by the ordinary refractometer. This requires the observation of the angle of minimum deviation (d) of a light-ray on passing through a prism of the given material, having a known angle (n), and with its edge cut in the proper direction. The measurements of d and n can be made with an ordinary refractometer or with the horizontal goniometer described in Art. 210. For the latter instrument, the collimator is made stationary, being fastened to a leg of the tripod support, but the observing telescope with the verniers moves freely. Further, for this

object the graduated circle is clamped, and the screw attachments connected with the axis--carrying the support, and the vernier circle and observing telescope are loosened. Light from a monochromatic source passes through an appropriate slit and an image of this is thrown by the collimator upon the prism. With a doubly refracting substance two images are yielded and the angle of minimum deviation must be measured for each; the proper direction for the edge of the prism in this case is discussed later. In cases where the highest degree of accuracy is desired sunlight is employed and the angle of deviation measured for the prominent Fraunhofer lines (p. 171). When a and S are known the formula in Art. 304 is used.

325. Total Reflectometer. —The principle of total reflection (Art. 303) may also be made use of to determine the refractive index. No prism is required, but only a small fragment having a single polished surface; this may have any direction with an isometric crystal, but in other cases must have a definite orientation, as described later. The arrangements required (as developed by F. Kohlrausch) are, in their simplest form, a wide-mouthed bottle filled with a liquid of high refractive power, as carbon disulphide ($n_D = 1.6442$ at 20°C) or α -bromonaphthalene ($n_D = 1.6626$ at 20°C). The top is formed by a fixed graduated circle, and a vertical rod, with a vernier attached, passes through the plate and carries the crystal section on its extremity, immersed in the liquid. The angle through which the crystal surface lying in the axis is turned is thus measured by the vernier on the stationary graduated circle. The front of the bottle is made of a piece of plate glass, and through this passes the horizontal observing telescope, arranged for parallel light. The rest of the surface of the bottle is covered with tissue-paper, through which the diffuse illumination from say a sodium flame has access; the rear of the bottle is suitably darkened. When now the observer looks through the telescope, at the same time turning the axis carrying the crystal section, he will finally see, if the source of illumination is in a proper oblique direction, a sharp line marking the limit of the total reflection. The angle is measured off on the graduated circle, when this line coincides with one of the spider lines of the telescope. Now the crystal is turned in the opposite direction, and the angle again read off. Half the observed angle ($2a$) is the angle of total reflection; if n_1 is the refractive index of the carbon disulphide, then the required refractive index, n , is equal to

$$n = n_1 \sin a.$$

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Under favorable conditions the results are accurate to four decimal places. This method is limited, obviously, to substances whose refractive index is less than that of the liquid medium with which the bottle is filled.

Different forms of total reflectometers * have been devised by Soret, Polfrich, Czapski, and others.

326. The method of obtaining the refractive index of a transparent medium, first

described by Duke de Uhaulnes (1767), has been shown by Sorby to allow, under suitable conditions, of determinations of considerable accuracy. This method consists essentially in observing the distance $\{d\}$ which the focal distance of the objective is changed when a plane-parallel plate of known thickness (t) is introduced perpendicular to the axis of the microscope between the objective and the focal point, here

t

Sorby made use of a glass micrometer, upon which two systems of lines perpendicular to each other were ruled. A micrometer-screw in the microscope makes it possible to measure the distance through which the tube is raised and lowered down to 0.01 mm.; consequently both t and d can be obtained with a high degree of accuracy.

327. Tonn[^]luie Tongs.—A very simple form of polariscope for converging light is shown in Fig. 513; it is convenient in use, but of limited application. Here the polarizer and analyzer are two tourmaline plates such as were described in Art. 317. They are mounted in pieces of cork and held in a kind of wire pincers. The object to be examined is placed between them and supported there by the spring in the wire. In use they are held close to the eye, and in this position the crystal section is viewed in converging polarized light, with the result of showing (under proper conditions) the axial interference-figures (Arts. 360 and 387).

B13.

328. Polariiscope. Conoscope. Fitauroscope.—The common forms of polariscope § employing Nicol prisms are shown in Figs. 514 and 515, | Fig. 514 represents the instrument arranged for converging light, which is often called a conoscope.

The essential parts are the mirror S , reflecting the light, which after passing through the lens e is polarized by the prism p . It is then rendered strongly converging by the system of lenses nn , before passing through the

•See Groth, Phys. Kryst., 1891 pp. 654-679; also Das Heflectometer, etc., von Dr. C. Pflücker. Leipzig, 1890.

IMN. Hag., 3. 1, 101, 1878.

Cr. Rosenbuech, Micr. Phys. Hie., p. 84, 1893, who mentions particular methods applicable to mineral thin sections.

^8 See Brewster, Groth, Phys. Kryst. (also FogK. Add., 144, 34, 1871).

These figures, also Figs. 516, 517, 544, are taken from the catalogue of R. Tschermak, Steg-Us. Berlin.

FHTSICAL MINERALOGY.

section under examination placed on a plate at k. The upper plate can be revolved through any angle desired, measured on its circumference. The upper tube 614. 616.

contains the converging system oo, the lens t, and the analyzing prism q. The arrangements for lowering or raising the tubes need no explanation, nor indeed the special devices for setting the vibration-planes of the nicols at right angles to each other.

The accompanying tube (Fig. 515) shows the arrangement (or observations in parallel light, the converging lenses having been removed. In this form it

is especially used for stereoscopic measurement, as later explained. In some forms of polariscope of the above type the place of the analyzer is taken by a pair of black glass mirrors set at the proper polarizing angle.

329. Polarization-Microscope.—The investigation of the form and optical properties of minerals when in microscopic form has been much facilitated by the use of microscopes specially adapted for this purpose. First arranged with reference to the special study of minerals as seen in thin sections of rocks, they have now been so elaborated as largely to take the place of the older optical instruments. They not only allow of the determination of the optical properties of minerals with greater facility, but are applicable to many cases where the crystals in hand are far too small for other means.

A highly serviceable microscope, for general use, is that described by Rosenbusch in 1876 and later much improved. A sectional view of one form is shown in Fig. 516, and a later and improved pattern is given in Fig. 617. The essential arrangements of Fig. 516 are as follows: The tube carrying the eyepiece and objective has a fine adjustment-screw at ff; the coarse adjustment is accomplished by the hand. The screw-head g is graduated and turns about a fixed index attached to the tube [^]; by this means the distance through which the tube is raised or lowered can be measured to 0.001 mm. The polarizing prism is placed below the stage at r, in a support with a graduated circle, so that the position of its vibration-plane can be fixed. The analyzing prism is contained in a cap, ss, which is placed over the eyepiece; this may be revolved at pleasure, its edge being graduated. When both prisms are set at the zero mark, their vibration-planes are crossed (J.); when either is turned 90°, the planes are parallel (H). The stage is made to rotate about the vertical axis, but otherwise (in this simple form) is fixed; its edge is graduated, so that the angle through which it is turned can be measured to 4°. Three adjustment-screws, of which one is shown at n, n, make it possible to bring the axis of the object-glass in coincidence with axis of rotation of the stage (see, further, the detailed drawing at title

side).

The instrument here described may be used in the first place as an ordinary microscope with magnifying power adapted to the apertural case in hand. In the second place, with polarizing prisms and the usual arrangement of lenses, it serves for determining the planes of vibration (like the selenite of Art. 328); also for observing the interference-colors of doubly refracting sections and so on. Finally, with eyepiece removed and special condensing lenses placed beneath the object on the stage (as more fully described later), it may be used, like the conoscopic microscope, for observing axial interference-figures, etc.

330. A later and improved form of microscope shown in Fig. 517 is essentially like that of Fig. 516, but has various refinements for accurate work. Thus, a screw is added for the coarse adjustment; another screw to raise and depress the lower nicol; a mechanical stage, etc. A more essential improvement is the insertion of the upper nicol in a support, which can be pushed in or out at will between the eyepiece and objective. The upper nicol above the eyepiece is, however, also needed in certain cases, for example with the Bertrand ocular, described later.

The microscope which has been briefly described is, as stated, especially applicable to the study of the form, optical properties, and mutual relations of minerals as they are found in thin sections of rocks; it has therefore become an important adjunct to geological research. It can also be used to great advantage for the study of small independent crystals

•See Smithsonian, Mikr. Phys., 117-130, 1892; also Groth, Phys. Kryst., 788-766, 1896.

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PHYSICAL MINERALOGY.

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in crystalline sections or in thin sections. The more important points to which (the attention of the student should be directed, more particularly in the case of minerals in sections of rocks, are: (1) crystalline form, as shown in the outline; (2) direction of cleavage; (3) optical character

Index: (4) Uniaxial or biaxial character, i.e., dichroism or pleochroism; (5) Isotropic or anisotropic character, and if the latter, the direction of the plane of light-vibration; (6) position of the axial plane and nature of the axial interference-figures; (7) the length of the double refraction; (8) Inclusions, solid, liquid, or gaseous.

The explanation in regard to the special optical points mentioned is deferred to later pages.

PHYSIOAL HIITERALOOT.

GENEEAL OPTICAL CHARACTERS OF MINERALS.

331. There are certain characteristice belonging to all minerals alike^ ciyatallized and non-crystallized, in tlieir relation to light. These are:

1. DiAPHAKEiTT: depending on the relative quantity of light transmitted.
2. Color: depending on the kind of light reflected or transmitted, as determined by the selective absorption.
3. Ldsteb: depending on the power and manner of reflecting light.

1. DXAPHANEITT.

332. Decreet of Trwuparanoy.—The amount of light trananiitted by a solid varies in intensity, or, in other words, more or less light may be absorbed in the passage through the given substance (see Art. 807). The amount of absorption is a minimum in a trsLsparent solid, as ice, while it is greatest in one which is opaque, as iron. The following terms are adopted to express the different degrees in the power of transmitting light:

Transparent: when the outline of an object seen through the mineral is perfectly distinct.

Suitransparent, or setni-lraitsparent: when objects are seen, but the outlines are not distinct.

Translucent: when light ia transmitted, but objects are not seen.

Siibtranslucent: when merely the edges transmit light or are translucent.

When no light is transmitted, even on the thin edges of small splinters, the mineral is said to be opnqve. This is properly only a relative term, since no substance fails to transmit some light, if made sufficiently thin. Magnetite is translucent in the Pennshury mica. Even gold maj be beaten out so thin as to be translucent, in which case it transmits a, greenish light.

The property of diaphaneity occnra in the mineral kingdom, in everr degree from nearly perfect opacity to transparency, and many minerals present, in their numerous varieties, nearly all the different shades.

2. COLOE.

333. ITature of Color.—As briefly explained in Art. 294, the sensation of

color depends alone upon the length of the waves of light which meet the eye. If they are all of the same length. If the light consists of various wavelengths, it is to the combined effect of these that the sensation of color is due.

Further, since the light ordinarily employed is essentially white light, that is, consists of all the wave-lengths corresponding to the successive colors of the spectrum, the color of a body depends upon the selective absorption (see Art. 307) which it exerts upon the light transmitted or reflected by it. A yellow mineral, for instance, absorbs all the waves of the spectrum with the exception of those which together give the sensation of yellow. In general the color which the eye perceives is the result of the mixture of those waves which are not absorbed.

All minerals may be divided into two classes: (1) those whose color belongs to the finest particles mechanically made; and (2) those whose color in the state of fine powder is different from what it is in the mass.

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To the first class belong the metallic and many minerals having a metallic luster; for instance, the powder of the black magnetic oxide of iron, magnetite, is black; that of benidite, which, though often black on the surface, is red by transmitted light, is red, and so on.

To the second class belong the silicates, and in fact the large part of all minerals having an unmetallic luster. With them the color is often quite essential, being generally due to small admixtures of some metallic oxide, to some carbon compound, or to some foreign substance in a finely divided state. With most of these, the fine powder is either white or light-colored; for example, the streak (Art. 334) of black, green, red, and blue tourmaline varies little from white.

334. Streak.—The color of the powder of a mineral as obtained by scratching the surface of the mineral with a knife or file, or still better, if not too hard, by rubbing it on an unpolished porcelain surface, is called the streak. It is obvious from the distinctions stated above that the streak is often a very important quality in distinguishing minerals. This is especially true with minerals of the first class mentioned above, that is, those with metallic luster, as defined in Art. 338.

335. Dichroism; Pleochroism.—The selective absorption, to which the color of a mineral is due, more especially by transmitted light, varies according to the molecular structure of the crystals. It is hence one of the special optical characters depending upon the crystallization, which are discussed later. Here belong dichroism or pleochroism, the property of exhibiting different colors in different directions by transmitted light. This subject is explained further in Arts. 366 and 363.

freeu, with more yellDw ; greeo dLdlage.—8. Piitaehio-grten: yellowish green, with Biiiiiie rown ; epldote,—9. Atparagiugreea: pale green, wiili mucji yellow; aspunigus siune (upeilte).—10. Bladdth green: serpeutlDe.^-II. OUm green: dark grecD, with much tiruwu uLd yellow; chrysolite.—13. Oil-grten: the color of ollve-oii; beryl, pilchstoie.—18. SUkin-green: light green, much iiiiicllDiiig to yellow ; urauite.

F. Yellow. 1. Salphiiryellow: Guliiljur.—3. Sraa-gallou: pale yellow; lopiiK.—8. Wax yelltm; grayish yellow wjlh some brown ; blende, opal.—4. Boneg-yeilote: yellow, with some red and brown ; cnlcite.—Et. L'noit-yellua: sulphur, orplmeul.—o. Oeher-f/eUow: yelloT, wiilh browu ; yellow ocher—7. Wiae-yeilovi: topaz and lluorite.—8. Cream-yeUtna: some varieties of litbomarge.—o. Oraiiige-yeUovi; orpimeut.

G. Red. 1. Aurora-red: reil. with iiiiiicb yellow; some realgar.—2. Hyaeinth'Ted: red. with yellow and some brown ; liyaciulb giinot.—3. Biidc-rtd: polyliialile, some jasper.—4, 8cartet-red: bright red, with n tinge of yellow; ciDUftbnr.~5. Blood-red: dark red. with some yellow; pyrope.— i. Kiei/i-red: ffldspnr—7, Oarmine-red: pure red ; ruby sapphire. — 8. fioie-red: rose qunrtz —8. CHmiion-rtd: ruby.—IU. Peacliblouom-red: red, with white null gray : lepidoliti;—11. Oolun^iu red : deep red, with some blue ; garnet.—13. C/terry-red.' iliirk red, with some blue and bniwu ; spinel, some jasper.—13. Brownith-red: jasper. Hmoiiiite,

H. Brown. 1. Beddith brouin: garnet, zircon.—2. Clotebroien: brown, with red and some blue; azioite.—3. Hairbroan; wood opal.—4. Broceoli-brovin: brown, with blue. red, and gray ; zii-con. —5. Chetlnut-br-mcn: pure brown.—8. Yeliowieh brown; Jasp^T.—T. Pinchbeek-brown: jellowiali-bniwa. wiilh a metallic or metallic-pearly luster: sev.Tal varieties of talc, bronziti'.—8. Wood-broaiH: color o(old wood nearly rotten ; somt- .lin-ci-meJis of usbestus.—9. Liver-brnren: brown, with some gray and green ; jasper.—10. Black-uhbrown: bituminous coal, brown coal.

337. Nature of Lmter.—The luster of minorals varies with the nature of their surfaces. A variation in the quantity of light refiected produces different degrees of inteDsity.of luster; a variation in the nature of the reflecting surface produces different kinds of luster.

338. Eindi of Laater.—The kinds of luster recognized are aa follows:

1. Metallic: the luster of the metals, as of gold, copper, iron, tin.

In genera!, a mineral is not said to have metallic luster unless it is opaque in the mineraiological sense, that is, it transmits no light on the edges of thin splinters. Some minerals have varieties with metallic and others with nnmetal-lic hister; this is true of hematite.

Imperfect metallic hister is expressed by the term swJ-me/«2Wc, as illustrated by

colitmbite, wolframite. Other kinds of luster are described briefly as

UNMETALLIC.

2. Adamantine: the luster of the diamond. When also sub-metallic, it is termed metallic-adamantine, as cerussite, pyrargyrite.

Adamantine luster belongs to substances of high refractive index. This may be connected with their relatively great density (and hardness), as with the diamond, also corundum, etc.; or because they contain heavy molecules, thus most compounds of lead, not metallic in luster, have a high refractive index and an adamantine luster.

3. Vitreous: the luster of broken glass. An imperfectly vitreous luster is termed sub-vitreous. The vitreous and sub-vitreous lusters are the most common in the mineral kingdom. Quartz possesses the former in an eminent degree; calcite, often the latter.

4. Resinous: luster of the yellow resins, as opal, and some yellow varieties of sphalerite.

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CHARACTERS DEPENDING UPON LIGHT. §9

5. Greasy; luster of oily glass. This is near resinous luster, but is often quite distinct, as efflorescence.

6. Pearly: like pearl, as talc, brucite, bitumens etc. When united with sub-metallic, as in hypersthene, the term metallic-pearly is used.

Pearly luster belongs to the light reflected from a pile of thin glass-plates; similarly it is exhibited by minerals, which, having a perfect cleavage, may be practically separated into thin plates, as on the basal plane of apophyllite, it is also shown for a like reason by foliated minerals, as talc and brucite.

7. Silky: like silk; it is the result of a fibrous structure. Ex. fibrous calcite, fibrous gypsum.

The different degrees and kinds of luster are often exhibited differently by unlike faces of the same crystal, but always similarly by like faces. The lateral faces of a right square prism may thus differ from a terminal, and in the right rectangular prism the lateral faces also may differ from one another. For example, the basal plane of apophyllite has a pearly luster wanting in the prismatic faces, they having a vitreous luster.

As shown by the following, the vitreous, adamantine, and metallic luster belong to faces of the same crystal. In the first, the refractive index of the mineral is 1.818; in the second, 1.825; in the third, about 2.0. The true difference between the

light: fii gei]i;riil. ihe luster ia prodiicud by Ihe util'iu nt two slmulKtneoua tui|ii'easioDS niqde upon . the eye. If Ihe light ruUeclcd from u metullic Biirfiice iie exidniDid by a Di<-ol prism (or the dichro!^ei>pv of HaiJbiger, AK. S66). It will be rmmd that boib rnys, thai vibriliug lu the pUiue of incidence iinil tbiil wbiiae vibmtioitR are normal lo It, are alike, eiicb having the I'oliir of tbe materiil, only differing a little in bnliiriDcy; on Ihe conirary. of the liglil rt^HecIeil by n vilruoua kdIi.i lance, tbose ruys wliose vil>iilioiis nre at right iuglt-s lo the plane or incidence are more or lesa polarized, atid nre colorleia. if bile those wiose vibmtiona are in ihls plane, having peieirated snmewbtit into the medium iinii siiSired »onie absorption, ahow tie color of the a'lbaieuce llselC A pliile of red glass Thus examined will alioW a colorless and a ri'd image. Adamantine liisiui' occupies a position between Ihe olliera.

339. Degrees of Luster.—The degress of intensity of luster are denominated OS follows;

1. Splendent: reflecting with brilliancy and giving well-defined images, as hematite, cassiCerite.
2. Shining: producing an image by reflection, but not one well-defined, as celfstite.
3. GUsfenin/j: affording a general reflection from tbe surface, but no image, as talc, chalcopyrite.
4. Glimmering: affording imperfect reflection, and apparently from points over the surface, as flint, chalcedony.

A mineral is said to be dull when there is a total absence of luster, as chalk, the ocbers, kaolin.

340. Flay of Colors. Opalescence Irideseenoe.—The term play of colors is used to describe the appearance ot several prismatic colors in rapid sticcession on turning the mineral. This property belongs in perfection to the diamond, in which it is due to its high dispersive power. It is also observed in precious opal, where it is explained on tlie principle of interference; in this case it is most brilliant by candle-light.

The expression change of colors is used when each particular color appears to pervade a larger space than in tbe play of colors and the succession produced by turning the mineral is less rapid. This is shown in lahradonte, as explained under that species,

Ojjialeacence ia a milky or pearly reflection from the interior of a specimen. Observed in some opal, and in cat e-eye.

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Iridescence means the exhibition of prismatic colors in the interior or on the Bui'face of a

mineral. The phenomena of the play of colors, iridescence, etc., are sometimes to be explained by the presence of minute foreign crystals, in parallel positions; more generally, however, they are caused by the presence of fine cleavage-lamellae, in the light reflected from which interference takes place, analogous to the well-known Newton's rings (see Art. 313).

341. Tarnish.—A metallic surface is tarnished when its color differs from that obtained by fracture, as is the case with specimens of bornite. A surface possesses the steel tarnish when it presents the superficial blue color of tempered steel, as columbite. The tarnish is iridescent when it exhibits fine colors, as is common with the hematite of Elba. These tarnish and iridescent colors of minerals are owing to a thin surface or film, proceeding from different sources, either from a change in the surface of the mineral or from foreign incrustation; hydrated iron oxide, distally formed from pyrite, is one of the most common sources of it, and produces the colors on anthracite and hematite.

313. Asterism.—This name is given to the peculiar star-like rays of light observed in certain directions in some minerals. This is seen by reflected light in the form of a six-rayed star in sapphire, and is also well shown by transmitted light (as of a small flame) with the phlogopite mica from South Burgess, Canada. In the former case it is explained by the presence of thin twinning-lamellae symmetrically arranged. In the other case it is due to the presence of minute included crystals, also symmetrically arranged, which are probably rutile or tourmaline in most cases. Crystalline faces which have been artificially etched also sometimes exhibit asterism. The peculiar light-figures sometimes observed in reflected light on the faces of crystals, either natural or etched, are of similar nature.

343. Schillerization.—The general term schiller (from the German) is applied to the peculiar luster, sometimes nearly metallic, observed in definite directions in certain minerals, as conspicuously in schiller-spar (an altered variety of bronzite), also in diallage, hypersthene, sunstone, and others. It is explained by the reflection either from minute included plates in parallel position or from the surfaces of minute cavities (negative crystals) having a common orientation. In many cases it is due to alteration which has developed these bodies (or the cavities) in the direction of solution-planes (see Art. 264). The process by which it has been produced is then called schillerization.

344. Fluorescence.—The emission of light from within a substance while it is being exposed to direct radiation, or in certain cases to an electrical discharge in a vacuum tube, is called fluorescence. It is best exhibited by fluorite, from which the phenomenon gained its name. Thus, if a beam of white light be passed through a cube of colorless fluorite a delicate violet color is called out in its path. This effect is chiefly due to the action of the ultra-violet rays, and is connected with a change of refrangibility in the transmitted light.

The electrical discharge from the negative pole of a vacuum tube calls out a brilliant fluorescence not only with the diamond, the ruby, and many gems, but also with calcite and other minerals. Such substances may continue to emit light, or phosphoresce, after

the discharge ceases,

345. Phosphorescence.—The continued emission of light by a substance (not incandescent) produced especially after heating, exposure to light or to an electrical discharge, is (if λ) phosphorescence.

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Fluorite becomes highly phosphorescent after being heated to about 150° U. Different varieties give off light of different colors; the chlorophane variety, an emerald-green light; others purple, blue, and reddish tints. This phosphorescence may be observed in a dark place by subjecting the pulverized mineral to a heat below redness. It is even called out by a sharp blow with a hammer. Some varieties of white limestone or marble, after slight heating, emit a yellow light; so also tremolite, danburite, and other species.

By the application of heat minerals lose their phosphorescent properties. But on passing electricity through the calcined mineral a more or less vivid light is produced at the time of the discharge, and subsequently the specimen when heated will often emit light as before. The light is usually of the same color as previous to calcination, but occasionally is quite different. It is in general less intense than that of the unaltered mineral, but is much increased by a repetition of the electric discharges, and in some varieties of fluorite it may be nearly or quite restored to its former brilliancy. It has also been found that some varieties of fluorite, and some specimens of diamond, calcite, and apatite, which are not naturally phosphorescent, may be rendered so by means of electricity. Electricity will also increase the natural intensity of the phosphorescent light.

Exposure to the light of the sun produces very apparent phosphorescence with many diamonds, but some specimens seem to be destitute of this power. This property is most striking after exposure to the blue rays of the spectrum, while in the red rays it is rapidly lost.

LITKRAVOR.

Ashentm,

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SPECIAL OPTICAL CHARACTERS BELONGING TO CRYSTALS OF THE DIFFERENT SYSTEMS.

346. All crystallized minerals may be grouped into three grand classes, which are dietinguiehed by their physical properties, as well as their geometrical form. These three classes are as follows:

A. Isometric class, embracing crystals of the isometric system, which are referred to three equal rectangular aiaa.

B. Isodiametric class, embracing crystals of the tetragonal and hexagonal systems, referred to two, or three, equal lateral axes and a third, or fourth, axis nnequal to them at right angles to their plane. Crystals of this class have a fixed principal axis of crystallographio symmetry.

C. Anisoiitric class, embracing the crystals of the orthorhombic, mono-clinic, and triclinic systems, referred to three unequal axes. Crystals of thia class are without a fixed axis of cryatallographic symmetry,

347. I[^]tropio Crystals.—Of the three classes, the laoMETRic class includes all crystals which, with respect to light and related phenomena involving the ether, are isotropic; that is, those which have like optical properties in all directions. Specifically, a light-wave is propagated in them with the same Telocity in all directions, and its wave-tront is therefore a sphere. Hence, also, the sphere may be regarded as representing the optical structure of an isometric crystal. The geometrical property of the sphere that every cross-section is a circle corresponds to the optical property of the isotropic medium in which the velocity of light-propagation is the same in every direction, for this being true, the

medium must have like properties of the ether in any plane normal to such a line.

It must be repeated here, however, that such a crystal is not isotropic with reference to those characters which depend directly upon the molecular structure alone, as cohesion and elasticity, (See Art, 264.)

Further, amorphous bodies, as glass and opal, which are destitute of any oriented molecular structure—that is, those in which all directions are sensibly the same—are also isotropic, and not only with reference to light, but also as regards their strictly molecular properties,

348. Anisotropic Crystals; Uniaxial and Biaxial—Crystals of the isodia-METRIC and ANISOMETRIC CLASSES, on the Other hand, are in distinction anisotropic. Their optical properties are in general unlike in different directions, or, more particularly, the velocity with which light is propagated varies with the direction.

Further, in crystals of the isodiametric Class that variable property of the light-ether upon which the velocity of propagation depends remains constant for all directions which are normal, or, again, for all those equally inclined to the vertical crystallographic axis. In the direction of this axis there is no double refraction; it is hence called the optic axis, and the crystals of this class are said to be uniaxial. The optical structure of uniaxial crystals can be represented by a spheroid, that is, an ellipsoid of revolution whose axis of revolution is the optic axis, or axis of crystallographic symmetry. The direction and properties of this optic axis will be seen to correspond to the geometrical property of the spheroid, a section of which normal to this axis is always a

Crystals of the third or anisometric class have more complex optical relations requiring special explanation, but in general it may be stated that in them there are always two directions analogous in character to the single optic axis spoken of above, hence these crystals are said to be optically biaxial. Further,

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it will be shown that their optical structure may be represented geometrically by an ellipsoid with three unequal rectangular axes. Every ellipsoid has two directions in which it can be cut yielding cross-sections which are circles; the optic axes spoken of will be seen later to be normal to these planes after the analogy of uniaxial crystals.

In crystals of the orthorhombic system the axes of the ellipsoid coincide in direction with the crystallographic axes. In crystals of the monoclinic system one of these ellipsoidal axes coincides with the axis of crystallographic symmetry, the other two lie in the plane of symmetry. In crystals of the triclinic system there is no necessary relation between the position of the ellipsoidal axes and those assumed to describe the crystallographic form.

All of these points require detailed discussion, but the above statements will partially

serve to bring out the intimate connection between the molecular structure exhibited in the geometrical form and the optical characters depending upon the properties belonging to the light-ether within the crystal.

A. IsoMETBio Crystals. 349. It has been stated that crystals of the isometric system are optically isotropic, and hence light travels with the same velocity in every direction in them. Light can, therefore, suffer only single refraction in passing into an isotropic medium; or, in other words, there can be but one value of the refractive index for a given wave-length. If this be represented by n , while V is the velocity of light in air and v that in the given medium, then

The wave-front for light-waves propagated from any point within such an isotropic medium is a sphere, and, as already stated, this geometrical figure may be taken as representing the optical structure of an isometric crystal.

This statement holds true of all the groups of isometric crystals. In other words, a crystal of maximum symmetry, as fluorite, and one having the restricted symmetry characteristic of the tetrahedral or pyritohedral divisions, have alike the same isotropic character. Two of the groups, however, namely the plagioclinal and the tetartohedral groups, differ in this particular: that crystals belonging to them may exhibit what has already been defined (Art. 333) as circular polarization.

350. Behavior of Sections of Isometric Crystals in Polarized Light.—In consequence of their isotropic character, isometric crystals exhibit no special

phenomena in polarized light. Sections of transparent isometric crystals may be always recognized as such by the fact that they behave as an amorphous substance in polarized light. In other words, a section on the stage of the

polarization-microscope, when the nicols are crossed, appears dark, and a rotation of the section in any plane produces no change in appearance. Similarly, it appears light in any position when placed between parallel nicols. Some anomalies are mentioned on a later page (Art. 411).

The single refractive index may be determined by means of a prism cut with its edge in any direction whatever.

B. Uniaxial Crystals, General Optical Relations. 351, The crystallographic and optical relations of crystals belonging to crystals of the tetragonal and hexagonal systems have already been briefly

discussed by GoOgle-

Generalized (Art. 348); it now remains to develop their optical characters more fully. This can be done most simply by making frequent use of the familiar conception of a

light-ray to represent the character and motion of the lightwave.

352. Optic Axis. Ordinary and Extraordinary Ray.—The study of a crystal belonging to this class shows, in the first place, that light-rays which pass in the direction of the vertical axis suffer no double refraction. This direction is that called the optic axis. Since the rays spoken of are propagated by vibrations at right angles to the vertical axis, that is, in the plane of the lateral crystallographic axes, this observed fact proves that for such rays there is but one value of the refractive index, and, further, that all the lateral directions must be identical so far as those properties of the ether are concerned upon which the velocity of light depends.

On the other hand, light which passes through the crystal in any other direction than that of the vertical axis suffers double refraction; in other words, it is separated into two rays, which are propagated with different velocities. This is true (see Art. 310) even when the rays follow the same path, as in the case of perpendicular incidence upon the given face.

Both of these rays are completely polarized, and their planes at right angles to each other.

It is found, further, that for one of these two rays, namely, that propagated by vibrations normal to the vertical axis, there is a constant value of the refractive index, whatever its direction; moreover, this value follows the usual law as to the constant ratio between the sines of the angles of incidence and refraction (Art. 293). It is hence called the ordinary ray, and the corresponding refractive index is uniformly represented by the letter μ_o .

For the other ray, on the other hand, it is found that the refractive index varies, and in general it does not obey the sine law. It is hence called the extraordinary ray. Further, if the direction of propagation changes progressively from that nearly coinciding with the vertical axis to that in the lateral plane normal to it, it is found that the value of the refractive index of the extraordinary ray deviates more and more widely from the constant value for the ordinary ray, and this difference becomes a maximum when the former is propagated in a lateral plane normal to the vertical axis, that is, by transverse vibrations in the direction of this axis. This last value of the refractive index is represented by the letter μ_e . These two indices, μ_o and μ_e , are called the principal indices of a uniaxial crystal. A principal section of a uniaxial crystal is a section passing through the vertical axis.

353. Positive and Negative Crystals.—Uniaxial crystals are divided into two classes. Those in which the refractive index of the extraordinary ray, μ_e , is greater than that of the ordinary ray, μ_o , are called positive. This is illustrated by quartz for which (for yellow sodium light):

$$\mu_o = 1.544. \quad \mu_e = 1.553.$$

On the other hand, if μ_o is greater than μ_e , the crystal is said to be negative. Calcite is an

example, for which (for sodium light):

$$w = 1.658. \quad e = 1.486.$$

Other examples are given later (Art 866V

364. Wave-anrfaoe.—Remembering that the velocity of light-propagation is always inversely proportional to the corresponding refractive index, it is obvious that the velocity of the ordinary ray for all directions in a uniaxial

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crystal must be the same, being uniformly proportional to $\frac{1}{n}$. In other

words, the wave-front of the ordinary ray must be a sphere.

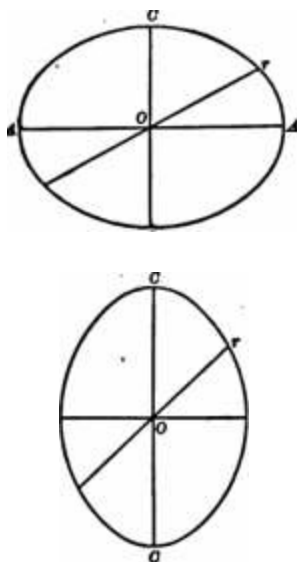
For the extraordinary ray, however, the Velocity varies with the direction,

being proportional to $\frac{1}{n}$ in a lateral direction and becoming sensible equal to $\frac{1}{n}$ when nearly coincident with the direction of the vertical axis. The

law of the varying change of velocity between these values, $\frac{1}{n}$ and $\frac{1}{n}$ is given

by an ellipse whose axes (OO, OA, Figs. 518,519) are respectively proportional to the above values.

618. SIS.

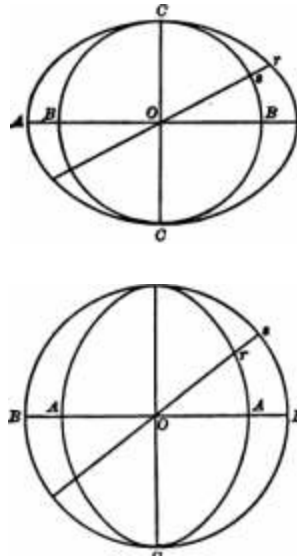


$$OO : OA = \frac{1}{n} : \frac{1}{n} = B : w.$$

This law, suggested by Huygens, has since been verified by accurate experiments by several observers for typical substances, as calcite and soda niter; hence it is accepted without question as a law of nature. The wave-front of the extraordinary ray is then a

spheroid, or an ellipsoid of revolution whose axis coincides with the vertical crystallographic axis, that is, the optical axis. In the direction of the vertical axis it is obvious that the two wave-fronts coincide.

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Negative crystal, $a > e$. Positive crystal, $e > a$.

Figures 520 and 521 represent vertical sections of the combined wave-fronts for both rays. Figure 520 pivots about the vertical axis for a negative crystal like calcite.

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$a > e$); Fig. 521 that of a positive crystal like quartz ($e > a$). Fig. 522 is

an attempt to show the relations of the two

wave-fronts of a negative crystal in perspective.*

The constant value of the velocity of the ordinary

ray, whatever its direction in this plane, is

expressed by the radius of the circle ($= OC$). On the other hand, the velocity of the extraordinary ray in the lateral direction is given

by OA , while in a direction as Oer , Fig. 520

(Ors, Fig. 521), it is expressed by the length of

this line, becoming more and more nearly equal to OC as its direction

approaches that of the vertical axis.

365. Indioatrlx.—It will now be understood what was meant by the statement in Art. 348 that the optical structure of a uniaxial crystal may be represented by an ellipsoid of revolution, and it is further obvious that the ratio between the axes of this ellipsoid must be as already given:

$$OC : OA = \dots \text{ or } e : \dots$$

$O \} \in$

It has been shown by Fletcher f that this ellipsoid, called by him the iodicatrix, may be taken to repreaeut the optical characters of both rays without reference to the wave-surface, since it can be proved geometrically % that for a given direction, as Or, the velocity of the extraordinary ray is expressed not only by Or bnt also by the inverse of the normal upon it from the point R (determined

by the tangents to the ellipse), that is, by \dots ;

also this normal fixes the plane of polarization which ia perpendicular to UN. Further, the velocity of the ordinary ray, having the same direction (cf. p. 195), is expressed by the inverse of the second normal upon the same line, that

is, T—T, since this normal is always in the equatorial plane, the section of which

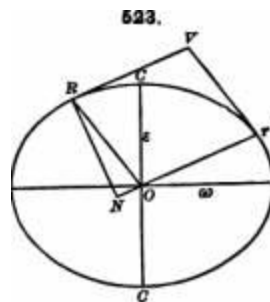
is a circle. Fig. 523 shows the form of the indicatrix for a negative crystal

* Figs. 622 and 536 Are taken from Mllller-Pouillet's L«btbucli der PbyBik. fThe Opllcnl Jodlcilrlz &d<1 lbe TraDsmisaloQ of Light \a Crytala, by h. Plelcher, LoDdoD. 18S3.

{This follows, from the property of lbe ellipse Id geDeral, since the paiiallelogara

$$OnVr = OA \cdot OC. \text{ llMit li. } RN \cdot Or = OA \cdot OC \text{ and } Or = \dots \text{ Or} = \dots$$

In other words, the velocity of the eztraordlDarr ray (Si) varies tovenely as \dots Bhnllarly, tb U repreteDted by Ort, that is, Id the iodicatrix by Ml rt. $n/i \ o^{\circ}$ Constants



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like calcite ($e > t$); that for a positive cr78ta1, like quartz ($e > t$) would be a prolate

Bpueroid.

ue. EzamplM of FcdtlTi and H*K>tiT« or7(taIi.~The following listii give promloeDt
pomtive aud Degfttlve uniaxial crystals, with the values of Itie refr&clive iLdlces, m nud e,
for each, correspoudiue lo yellow •oUium light." Tlie differeuce bttwten lUesc, m— e or f
—ao, fe also Kiveo: llifs measures the blrefrlugence or ilrtTigtA of Iht: double rvfnictlon.

It Tiiay b« remarked that Id tome ipecies bnib -4- Rnd — varielii's have been observed.
CertAio cryalals of upopbyllite ftre pwiliTe for out: eud of the ipectrum and Degnllve for
the oiber, aud cuDiequeully for some colnr between the two ixlremea <1 has no double
refreciiOD. Tbe Mme It true fnr tome other species (t.g.. cbibazile) of weak double
lefntloD. It if to be noted also that while eudlalyte Is positive, the related eucoUte Is
negative.

Neoativb Cbibtai.

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Beryl l-fi8M 15821 0-0078

Nephellie IMia 15876 00040

Apatite. 1-QlQl 1-6417 00084

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Rulile 3-6158 2'9029 02871

Cassllerlte 1 oBS6 30934 00068

Zircon 1-9818 18831 00618

Phenacle.. 16540 1-6697 00157

Brudte 1-5500 1-S705 00205

Qunrtz 1-5448 15588 000»1

Examination of Uniaxial Crystals in Polarized Light.

357. Section Normal to the Axis in Parallel Polarized Light.—Suppose a

section of a uniaxial crystal to be cut perpendicular to the vertical axis. It has already been shown that light passing through the crystal in this direction suffers no double refraction; consequently, such a section examined in parallel polarized light in the instrument called an orthoscope (Fig. J15), or in the polarization-microscope (Figs. 516, 517), behaves as a section of an isometric crystal, or of an anisotropic substance, if the nicols are crossed it appears dark, and remains so when revolved.

358. Section Parallel to the Axis.—A section cut parallel to the vertical axis, as already explained, has two directions of light-vibration, one parallel to this axis and the other at right angles to it. A ray of ordinary light falling upon such a section at right angles is divided into the two rays, ordinary and extraordinary, which, however, in this special case of perpendicular incidence travel on in the same path through the crystal, but one of them retard relative to the other. In parallel polarized light between crossed nicols each section will appear dark if the directions of its two vibration-planes coincide with the vibration-planes of the nicols. Thus in Fig. 554, A being

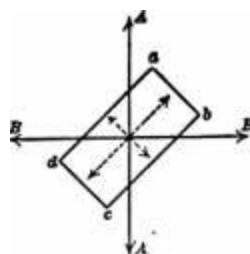
For authorities, see Dana's System, 1892. For calcite and aragonite the values of μ_o and μ_e are given.

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the vibration-plane of the lower nicol (polarizer) and BB of the upper nicol (analyzer), the light that has passed through the polarizer has its vibrations limited to the plane AA; these, therefore, pass through the section abed, but they are retarded or extinguished by the second nicol. The same will be true if this section is turned at right angles to the first position, that is, into the position a'b'c'd', represented by the dotted lines.

If the section stand obliquely, as abed in Fig. 535, it will appear light to the eye (and usually colored). For the vibrations parallel to AA that have



passed through the polarizer have upon resolution a component in the direction of each of the vibration-planes of the section. Again, each of these components can be resolved along the direction of the vibration-plane of the upper nicol, BB. Therefore, two rays will emerge from the analyzer, both having the same vibration-plane, but one more or less retarded with reference to the other, the amount of retardation increasing with the birefringence and the thickness of the section. In general, therefore, these rays will interfere, and if the thickness of the section is small (and not too great) it will appear colored in white light and, supposing the thickness uniform, of the same color throughout.

Any section whatever of a uniaxial crystal appears dark between crossed nicols if its principal section (Art. 352) coincides with the vibration-plane of either nicol.

359. Color of a Section in Parallel Polarized Light. Birefringence.—The interference-color of a section under examination depends (Art. 320) upon its thickness and upon the birefringence; this birefringence has a maximum value, equal to $\frac{1}{2}(\mu - \nu)$, if the section is cut parallel to the optic axis (i.e. \perp to the axis).

The following table* gives the thickness (in millimeters) of sections of a few uniaxial crystals which yield red of the first order:

Birefringence

(to \perp to the axis). Millimeters.

Quartz 0.28 to 0.0019

Calcite 0.1 to 0.0032

Zircon 0.062 to 0.0089

Tourmaline 0.023 to 0.0340

Quartz 0.009 to 0.00612

Kephelite 0.004 to 0.01377

Leucite 0.001 to 0.05510

* See further. Rosenbusch (Hilber Pliya. Moll. 1892, p. 166). from whom these are taken. Compare also remarks made in Art. 350.

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Again, as an example, it may be noted that quartz ($\mu - \nu = 0.062$), a thickness of 0.0019 mm gives red of the first order; of 0.017 mm red of the second order; of 0.026 mm red of the third order.

of tilt: third urdr.

The method ordinarily used to determine the birefringence of a section (not $\pm i$) of a uniaxial crystal, as shown to fix the relative value of the two principal refractive indices, are the same as those employed for biaxial crystals. and the discussion of them is postponed to a later part (Art. 51)

360. Uniaxial Interference-figure.—If an axial section, that is, one cut normal to the vertical axis of suitable thickness, be viewed in converging polarized light in a polariscope, e.g., the conoscope (Art. 827, Fig. 514), or the tourmaline tongs (Fig. 513), or again in the microscope* arranged for the purpose, it no longer appears dark. On the contrary, a beautiful phenomenon is observed: a symmetrical black cross when the nicols or tourmaline plates are crossed—with a series of concentric rings, dark and light in monochromatic light, but in white light showing the prismatic colors in succession in each ring. This is represented without the colors in Fig. 536, and with the colors in Fig. 1 of the plate forming the frontispiece to this volume.

This cross becomes white when the nicols or tourmalines are in a parallel position, and each band of color in white light changes to its complementary tint (cf. Fig. 521). These interference-figures, seen in this form only in a plate cut perpendicular to the vertical axis, mark the uniaxial character of the crystal.

The explanation of this phenomenon, so far as it can be given in a brief statement, is as follows: All the rays of light perpendicular to the plane of the section, that is, those whose vibrations coincide sensibly with the vibration-planes of either of the crossed nicols, must necessarily be extinguished. This 636. 637.



gives rise to the black cross in the center, with its arms in the direction of the

planes mentioned. Obviously this cross will be darkest along its central axis, while it fades out on the sides. All other rays passing through the given plate obviously are doubly refracted, and after passing through the second nicol, thus being referred to the same plane of polarization, they interfere, and give rise

•After the section is in position on the stage, and properly focused, the eyepiece is removed and a condensing lens inserted over the lower nicol. It is important to use a relatively high-power objective. It is also possible to see axial figures without removing the eyepiece by using a converging glass above the latter. Cf. Klein. Jb. Min., Bell.-Bd., 3. 4(), 1855; also Berirand, Bull. Soc. Min., 1. 23. 90, 1876; 3, 97. 1860.

Uniaxial crystals which produce circular polarization exhibit an axial interference figure (Fig. 3 of plate referred to above) which differs somewhat from that described, M. Dotted in Art 366. Some anomalies are mentioned later. (Art. 411.)

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In series of concentric rings, light and dark in monochromatic light, but in ordinary light showing the successive colors of the spectrum. The phenomenon is closely analogous to that of the Newton's rings mentioned in Art. 313. A cone of converging rays passes through the crystal and, having traversed the second nicol, each is divided into two rays with common vibration-planes, but one of them (the ordinary ray in positive crystals) slightly retarded with reference to the other. When the amount of retardation is equal to a wave-length (supposing monochromatic light to be employed) the effect of the interference is to destroy the light and the plane section of the cone, or circle, appears dark. Other dark rings are seen at distances which correspond to a retardation of $\frac{1}{2}$, $\frac{3}{2}$, etc., of a wave-length. If, however, the retardation amounts to a whole wave-length or any multiple of this, the two rays unite to strengthen each other and give rise to a light ring. In ordinary white light is employed, the relations are similar but the retardation cones overlap because of the different values of the refractive indices (i.e., velocities) for the different wave-lengths, and the series of colored circles is the result.

The distance of each successive ring from the center obviously depends upon the birefringence, or the difference between the refractive indices for the ordinary and extraordinary ray, and also upon the thickness of the plate. The stronger the double refraction and the thicker the plate, the smaller the angle of the light-cone which will give a certain amount of retardation, or, in other words, the nearer the circles will be to the center. Further, for the same section the circles will be nearer for blue light than for red, because of their shorter wave-length. When the thickness of the plate is considerable, only the black brushes are distinctly seen.

361, Determination of the Refractive Indices.—A single prism suffices for the measurement of the indices of refraction, m and e , with the refractometer. Further, its edge may be either parallel to the vertical axis or at right angles to this direction. Such a prism yields two images of the slit, one corresponding to the ordinary and the other to the extraordinary ray, and for each the angle of minimum deviation is to be determined, that is, the angle d in the general formula of Art. 804. Which of the two rays corresponds to the ordinary and which to the extraordinary ray can be easily distinguished by means of a nicol, the position of whose vibration-plane is known. This will extinguish that ray whose vibrations take place in a plane at right angles to its own vibration-plane.

362. Other Methods for Determining the Refractive Indices.—The method of total reflection (Arts. 303 and 325) may also be employed to determine the values of m and e . The section taken of a uniaxial crystal has its surface most conveniently parallel to the vertical axis. It is so placed that the direction normal to the optic axis is horizontal. The

light is here separated into two rays, having separate limiting surfaces, and with a Nicol prism it is easy to determine which of them corresponds to the vibrations parallel and perpendicular, respectively, to the optic axis.

Again, it is possible to obtain the refractive indices with considerable accuracy from measurements, in the plane of the axes, of the distances between the black rings in the interference-figures as seen in homogeneous light. The relation between these distances and the optical "axes of elasticity" was established by Neumann (Pogg. Ann., 33, 357, 1834). Bauer has also developed the same method as applied to uniaxial crystals, and employed it in the case of bnicite (Ber. Ak. Berlin, 1877, 704, and 1881, 958).

With the polarizing-microscope the most simple method is that of the

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Dake de Chaulnes, explained in Art. 326; this requires, however, that the two quantities measured should be determined with a considerable degree of accuracy, if the result is to be more than an approximation. (Cf. further, Bosenbusch, Mikr. Phys., p. 155 et seq., 1893.)

563. Determination of the Positive or Negative Character.—The most obvious way of determining the character of the double refraction ($e > o$ or $e < o$) is to measure the refractive indices in accordance with the principles explained in the preceding articles. It is not always possible, however, to obtain a prism suitable for this purpose, and in any case it is convenient to have a more simple method of accomplishing the result.

In the case of uniaxial crystals, the method which is practically the most simple is that suggested by Dove—the use of a cleavage plate of muscovite of such thickness that the two rays in passing through suffer a difference of phase which is equal to a quarter wavelength, or an odd multiple of this, it is often called a quarter-undulation plate (see Art. 322).

Suppose that the section of the crystal to be examined, cut perpendicular to the axis, is brought between the crossed nicols in the polariscope; the black cross and the concentric colored rings are of course visible. Let now, while the given section occupies this position, the mica plate be placed over it, with the plane of its optic axes (determined beforehand, and the direction marked by a line for convenience) making an angle of 45° with the vibration-planes of the nicols. The interference-figure is completely transformed. The colored rings are broken by two more or less

distinct hyperbolic brushes which pass +

through two black spots near the center, while the rings in the corresponding quadrants are pushed out from the center, and in the two remaining quadrants pushed in.

If now the line joining the two ilark spots Is at right angles to the axial plane of the mica (shown in the figures

by the arrow) the Crystal under experiment is opposite in refractive character to the mica, that is, positive (Fig. 529); if this line coincides with axial direction, the crystal is like the mica, negative (Fig. 530).

364. With the microscope the above method may also be employed. A mica plate, usually in the form of a narrow strip whose length is a multiple of the thickness of the plate, is placed between the slit of the microscope and the analyzer. However, the field of view is smaller than in the case of the black sro mil always distinctly observed; this is particularly true if the section be very thin or the thickness of low birefringence. In such cases a mica plate is conveniently employed. This is of such thickness as to give rise to the first order, and the direction of elongation usually corresponds to the axis (Art. 378). The plate is inserted in the tube with its axis inclined 45° to the vibration-planes of the nicols. This serves to increase the distance between the two interference colors; the section in two adjacent quadrants in and out of phase in the others; the effect being to show the rise or fall of the interference colors, as coincident with the micrometer scale (Art. 320). For example, two blue areas (second order) may be seen in two opposite quadrants and two white areas (first order) in the others. The blue areas here correspond in position to the black dots in Figs. 529 and 530; hence if the line joining them is transverse to that of the axis of the crystal plate the crystal is positive; if it coincides with it the mineral is negative.

365. Absorption Phenomena of Uniaxial Crystals. Dichroism.—In uniaxial crystals it has been seen that there are two distinct values for the velocity of light transmitted by them, according as the vibrations take place parallel or at right angles to the vertical axis. Similarly the crystal may exert different

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PHYSICAL MINERALOGY.

degrees of absorption upon the rays transmitted by vibrations in these two directions. For example, a transparent crystal of zircon looked through in the direction of the vertical axis appears of a pinkish-brown color, while in a lateral direction the color is asparagus-green. This is because the rays vibrating parallel to the axis are absorbed with the exception of those which together give the green color, and those vibrating laterally (ordinary) are absorbed except those which together appear pinkish-brown.

Again, all crystals of tourmaline in the direction of the vertical axis are optically, since the ordinary ray, vibrating normal to the axis, is absorbed, while light-colored varieties,

looked through laterally, are transparent, for the extraordinary ray, vibrating parallel to the axis, is not absorbed; the color differs in different varieties, in all uniaxial crystals; be dichroic or have two distinct axial colors.

The absorption-colors are most satisfactorily investigated by examining a section cut parallel to the vertical (optic) axis under the microscope provided with a single polarizing nicol. The light that passes through the section is then that corresponding to vibrations coinciding with the known vibration-plane of the nicol, and as the section is rotated 90° , both the two axial colors are observed in succession. References to some important papers on this subject are given on p. 219.

An instrument called a dichroite, contrived by Huyghens, is sometimes used for examining this property of crystals, a long rhombic prism or Iceland spar has a glass

prism of 15°

the figure, as follows

; fixed to each extremity-

placed in a metallic

cells at one end, and a square hole at the other. On looking

631.

through it, the appearance is double; one image belongs to the ordinary and the other to the extraordinary ray. When a pleochroic crystal is examined with it by transmitted light, on revolving it the two squares, at intervals of 90° in the revolution, have distinct colors, corresponding to the vibration-planes of the ordinary and extraordinary ray in succession. Hence the two images are situated side by side, a very slight difference of color is perceptible.

366. Circular Polarization. —The subject of elliptically polarized light and circular polarization has already been briefly alluded to in Art. 323. This

phenomenon is most distinctly observed among minerals in the case of crystals belonging to the rhombohedral-trapezohedral group, that is, quartz and cinnabar. It has been explained that a section of an ordinary uniaxial crystal cut normal to the vertical (optic) axis appears dark in parallel polarized light for every position between crossed nicols. If, however, a similar section of quartz, say 1 mm. in thickness, be examined under these conditions, it appears dark in monochromatic light only, and that not until the analyzer has been rotated so that its vibration-plane makes for sodium light an angle of 34° with that of the polarizer. In other words, this quartz section has rotated the plane of

polarization (t.e., the vibration-plane normal to it) some S_4° , and here either to the right or to the left, looking in the direction of the light. The amount of this rotation increases with the thickness of the section, and as the wavelength of the light diminishes (for red this angle of rotation for a section of 1 mm. is about 19° , for blue 33°). The direction of the rotation is to the right

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CHARACTERISTICS DEPENDING UPON LIGHT. SO_3

or left, as defined above—according as the crystal is crystallographically right* handed or left-handed (p. 83),

If the B_{\parallel} section of quartz (cut perpendicular to the axis) be viewed between crossed nicols in converging polarized light, it is found that the interference-figure differs from that of an ordinary uniaxial crystal. The central position of the black cross has disappeared, and instead the space within the inner ring is brilliantly colored.* Furthermore, when the analyzing nicol is revolved, this color changes from blue to yellow to red, and it is found that in some cases this change is produced by revolving the nicol to the right, and in other cases to the left; the first is true with right-handed crystals, and the second with left-handed. If sections of a right-handed and left-handed crystal are placed together in the polariscope, the center of the interference-figure is occupied with a four-rayed spiral curve, called, from the discoverer, Airy's spiral. Twins of quartz crystals are not uncommon, consisting of the combination of right- and left-handed individuals (according to the Brazil law) which show these spirals of Airy. With similar phenomena are observed. Twins of this species also not infrequently show Airy's spirals in the polariscope.

C. Biaxial Crystals.

General Optical Relations.

867, Principal Refractive Indices.—All crystals of the third or anisometric class, that is, those of the orthorhombic, monoclinic, and triclinic systems, are optically biaxial. In the directions of the optic axes there is a single value only for the light velocity, but in other directions a light-ray is separated into two rays propagated with different velocities; that is, it suffers double refraction.

The study of biaxial crystals shows that there are two directions within them at right angles to each other, corresponding to which, as vibration-axes, the refractive indices have respectively a maximum (a) and a minimum value (γ) for the given substance. Further, in a third direction at right angles to each of those just named, the refractive index has a certain intermediate value, related to the others by a simple mathematical law. These three rectangular directions, or ether-axes, are properly axes of vibration, and

the three corresponding refractive indices determine the rate of this transverse vibration and hence the velocity of the light-ray which corresponds to each of them. The

Velocities of the velocities are respectively proportional to $\frac{1}{a}$, $\frac{1}{b}$, $\frac{1}{c}$.

The indices a , b , c are called the principal refractive indices for the given substance. The mean refractive power is given by their arithmetical mean,

viz., $\frac{a+b+c}{3}$. Further, the difference between the greatest and least index,

$c - a$, measures the birefringence or strength of the double refraction.

368. Optical Structure of Biaxial Crystals.—It is found further that the optical structure of a biaxial crystal can be represented by an ellipsoid having as its axes the three lines mentioned in the preceding article which are at right angles to each other and proportional in length to the indices a , b , c . It is indeed inferred (following Fresnel) from the analogy of uniaxial crystals. The position of the axes named, or, in other words, the symmetry of

• Very thin sections of quartz, however, show (e.g., with the microscope) the dark cross or an ordinary uniaxial crystal.

of the difference between the primary and the secondary optical axes. See Art. 361.

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this ellipsoid, is such as to correspond to the general crystallographic symmetry of the crystal. For example, to repeat the statement already made (Art. 348), in the orthorhombic system the axes of this ellipsoid coincide in direction with the crystallographic axes. In the monoclinic system, one of them coincides with the axis of crystallographic symmetry (Z), the other two lie in the plane of symmetry, that is, in the plane of the crystallographic axes X and Y . In the triclinic system there is no necessary connection between the position of the ether-axes and the crystallographic axes.

369. Indicatrix.—It may be shown, as is done by Fletcher, that the ellipsoid mentioned, whose axes repre-

sent in magnitude the three principal refractive indices, a , b , c (where $a < b < c$), not only exhibits the character of the optical symmetry, but also serves to represent the direction, velocity and plane of polarization of a light-ray in any direction whatever, precisely analogous to the spheroid mentioned on p. 198. See Fig. 533 (from Groth), also Figs. 518, 519. That is, for the two rays having any direction, as OA in the plane of the axes AA' , VC (Fig. 534), the velocities

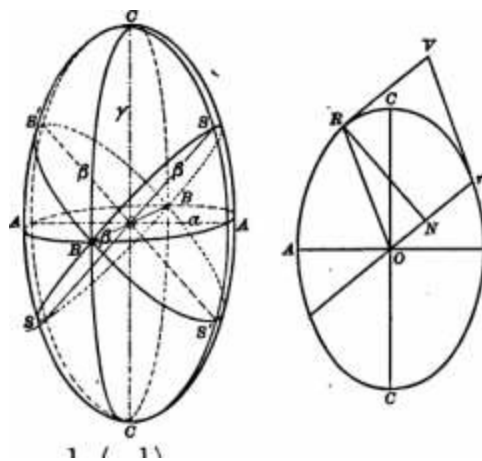
will be proportional to $\frac{1}{a}$, $\frac{1}{b}$

and $j^{\wedge} I = \wedge$) respectively, and the planes of polarization will be perpendicular to these lines. From the equation of this index-ellipsoid, called by Fletcher the indicatrix (see Art. 355), it is possible, as shown by the author named, to deduce by ordinary analytical methods the mathematical expression for the wave-surface, the position of the two sets of optic axes (later explained), etc.

One important relation appears at once from a first study of this ellipsoid. Obviously for two definite positions of transverse planes passing through the center (55, 5'5", Fig. 533), the positions depending upon the relative values of a and y the cross-sections will be circles each having a radius equal to the index (i , intermediate in value between a and y) similarly all sections parallel to these are also circles. Hence, light propagated in a direction normal to these planes, that is, by vibrations lying in them, will suffer no double refraction—and after the analogy of uniaxial crystals these directions are called optic axes; they are the primary optic axes mentioned in Art. 371.

370. Wave-surface.—Following out the analogy of uniaxial crystals, Fresnel deduced the now generally accepted "wave surface" for biaxial crystals. That it gives correctly the law of the varying refractive indices (that is, of varying light-velocity) in a biaxial crystal has been demonstrated by the agreement between the requirements of the theory and the results of experiment.

The form of the sections of this wave-surface with the three rectangular



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axial planes are easily deduced by very elementary considerations, though the analytical development is most satisfactorily derived from the equation of the indicatrix as shown by Fletcher.

First consider the section of the wave-surface for the transverse plane of the axes AB (Fig. 533). Light passing in the direction of the axis AA will be separated into two rays; for one of these the line of transverse vibration will correspond to the axis CC, and hence its rate,

or, in other words, the

tion of transverse vibration will be that of the axis BB , and its rate, that is, the velocity of the ray itself, will be proportional to v . Let these values be

represented in Fig. 535 by Oc' ($= v$) and Ob ($= v$). Again, in the direction

of the axis BB , there will be two rays whose vibration-directions are respectively parallel to the axes AA and CC , and their velocities in

the direction of BB proportional to v and v

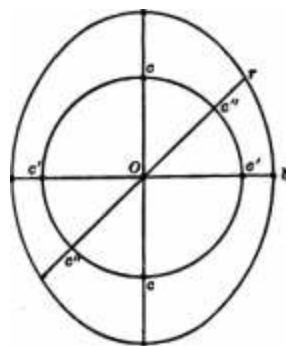
respectively. Let these be represented by Oa and Oc . For some other direction in the same plane, there will be two rays, one of whose vibration-directions corresponds to CC , and its

velocity to v , represented in Fig. 533 by Oc'' ,

while for the other there will be an intermediate

vibration-direction and a velocity between v and

v , and it can be shown (after the analogy of



uniaxial crystals and as proved before that this value is given by the line

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experiment)) Or in the ellipse whose major and minor

axes $\{Oa$ and $Oc\}$ are v and v . Hence the circle ec' represents the section of

the wave-surface for the rays in the given plane, whose vibration-direction

corresponds to the axis CC , and the velocity to the constant value v . While

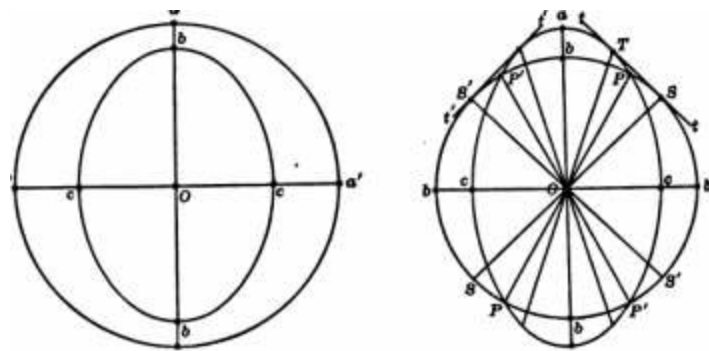
for other rays the vibration-directions change from AA to BB , and the velocity

from - to -3-

The ray propagated by vibrations in the direction of the axis CC, which has the constant velocity—, that is, the ray whose wave-front in this cross-section is a circle, is called the ordinary ray, since on refraction it remains in the plane of incidence. The other ray, whose velocity varies with the direction from — to 3, is called the extraordinary ray.

Again, take the plane of the axes BC (Fig. 533). Whether the direction of the light be that of B or of C, or any intermediate line in the same plane, there will be in each position one ray whose vibration-direction is that of the

axis A, and whose velocity is hence expressed by —; for it the section of the wave-surface will be a circle. For the other ray, if parallel to B, the vibration-direction will be that of the axis C, and its velocity is represented by -. If it is parallel to C, its vibration-direction is that of B, and its velocity is given by --=. As in the other case, intermediate values will be given by the ellipse S3«. 637.



having -. and — for its major and minor axes. The combined section of the wave-surface is shown in Fig. 536. Here also, the ray with the constant velocity - is called the ordinary ray; the other is the extraordinary ray.

For the third plane, that of the axes
 **'■ AC, one ray will always have as its
 vibration-dii'ection that of the axis B,
 and its velocity will heuce be expressed

by -3-. For the other ray, if parallel to A, the vibration-direction is that of C, and the velocity is expressed by -. If parallel to C, the vibration-directioQ is ' that of A and the

velocity -, and similarly for intermediate positions. The section of the wave-surface constructed from these values is given in Fig.

537. Here the circle (radius = r) cuts the ellipse at the points FP, P'P'.

The complete wave-surface, of which the three axial sections are given in

Figs. 535, 536, and 537, can be constructed, but it is not easy to form a

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complete knowledge of the form without having a model in hand. Some idea of it may be gathered from Fig. 538,

371. Primary and Secondary Optic Axes.—It has already been briefly stated (Art. 367) that there are two directions, namely, those normal to the circular cross-sections of the indicatrix (Fig. 533) in which the light

is propagated by transverse vibrations of like rate $I \sim J$. Hence in these directions in a crystal there is no double refraction within the crystal; nor is there when the ray emerges. These two directions bear so close an analogy to the optic axes of a uniaxial crystal that they are also called optic axes, and the crystals here considered are hence named biaxial. In Fig. 537, these optic axes have the direction $S'S'$, $S'S'$ normal to the tangent planes ti , ft' , and the direction of the external wave is given by the normal ST (Fig. 539).

Properly speaking the directions mentioned are those of the primary optic axes, for there are also two other somewhat analogous directions, PP , $P'P'$, of Fig. 537, called for sake of distinction the secondary optic axes. The properties of the latter directions are Kaa obvious from the following considerations.

In the section of the wave-surface shown in Fig. 537 (also enlarged, in Fig. 539), corresponding to the axial plane AC , it is

seen that the circle with radius Ob (

intersects the ellipse whose major and minor

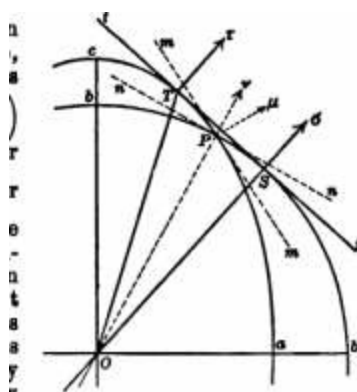
ueB are Oa

(= ^)

and Oc

(=1]

in the four



Joins P,P, P',P'. Corresponding to these irections the velocity of propagation is obviously the same for both mys. Hence within the crystal these rays travel together without doable refraction. Since, however, there is no common wave-front for these two rays (for the tangent for one ray is represented by vini and for the other by nn, Fig. 539) they do Buffer double refraction on emerging; in fact, two external light-waves are forined whose directions are given by the normals Pfi and Pv, These dii-ectione, PP, P'P', therefore have a relatively minor interest, and whenever, in tlie pages following, optic axes are spoken of, thoy are always the primary optic axes, that is, those having the directions SS, S'S" (Fig. 537), or OS, Fig. 539.* In practice, however, as remarked in the nest article, the angular variation between the two sets of axes is usually very small, perhaps 1° or leas.

ITS. Istwior ud Xxtarior OonlixJ BafrMtion.—The taDgent pinne to Ibe wavf-surface drawD normal [o the Hue OSthrough llie pnlnt iS'(Fig 5So) mity be abown to meet It In a iinsll circle on wbose circumference lie the points 8 and T. Tills circle la tlie bnne of tbe interior cone of raya 8o7', wboM remarknble properties will be brieflj btnted Ht. If a section of a biaxial crystal be cut wlib lla faces uormal in o8, tboae parallel njs. belongtue to a cylinder bawtce tilts circle aa Its bau. Incklent upon It from wltbout. will be propagated wllbiu ai tbe cone 9o7*. Conversely, rays from wl I bin correaponiliig In nnaitlon to rbe ■urface ot tbfs cone will emerKc parallel aod form a circular cylinder. This pbeuomeuon it c&IIRD intriar eonial rtfraction.

"Fletcher calls iliepTimaty axes MMrniaft, the secondary axe* MrottMi.

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On the other hand, if a section be cut with its face normal to OP, those rays having the direction of the surface of a cone formed by perpendiculars to the face and normal will be propagated without deviation parallel to OP, and emerging on the other surface form without a similar cone on the other side. This phenomenon is called *double refraction*.

In the various experiments given (685-58U) the relations are much exaggerated for the sake of clearness; in practice the relatively small difference between the indices of refraction μ_x and μ_y make this cone of small angular size, rarely over 2° . For example, with sulphur, which has very strong double refraction $\mu_x - \mu_y = 0.29$; compare the values given in Art. 59 and Art. 58, the values of μ_x, μ_y for yellow sodium light were measured by Schrauf as follows:

$$\mu_x = 1.6047, \mu_y = 1.30889, \mu_z = 1.24068.$$

373. *Axis of Elasticity*.—As intimated in Art. 368, Fresnel appears to have deduced the wave-surface of biaxial crystals, as it is here called following him, by a generalization from that accepted for the more simple uniaxial crystal. The explanation of the observed phenomena, attempted by him, was based upon the assumption that the varying velocity of light shown by the varying values of the refractive indices depended upon the variable elasticity of the ether within the crystal. Since, as stated on p. 160, it seems better not to insist upon this hypothesis and since, further, it is possible to describe all the phenomena without attempting to explain the properties of the ether upon which the ultimate values of the pulses depend which manifest themselves as light-waves, all mention of elasticity has been thus far avoided.

These "axes of elasticity" are of great convenience in describing the optical properties

of crystals. And it is hence necessary to make

frequent use of them. They are uniformly represented by the letters a, b, c , where $a > b > c$ as shown in Fig. 540, and where

further it is true that $a : b : c = v : v' : v''$. a, b, c being the three principal refractive indices (μ_x, μ_y, μ_z). The three axes as given in Fig. 540 then have the direction of the three ellipsoidal axes (Fig. 533); the maximum value of the elasticity, a , corresponds to the minimum value of the observed refractive index μ_x , and this in turn corresponds to the maximum velocity of a ray propagated perpendicular to the plane of the axes a and b by vibrations having a direction parallel to this axis a ; similarly for the axes b and c .

Whenever in this work the axes a, b, c , which may be simply called the ether-axes, are spoken of in describing the optical characters of crystals, it is to be understood that they have the directions indicated, corresponding respectively, as just explained, to the ellipsoidal axes; moreover, their relative magnitude is expressed as follows: $a > b > c$.

374. Bisectrices, or Uean-lineB.—Aa shown in Art. 371, the optic axes always lie in the plane of the axes a, y of the indicatrix (that is, of the ether-axes a and c); this is called the optic axial plane (or briefly, ax. pi.). The value of the optic axial angle is known when the values of the refractive indices, a, o, y, are given, as stated in the next article. That axis (cf. Fig. 539, also Figs. 54L, 543) which bisects the acute angle of the optic axes is called the acute bisectrix, or first mean-line, and that bisecting the obtuse angle is the obtuse bisectrix, or second mean-line.

The acute bisectrix is often represented by Bx^\wedge , the obtuse bisectrix by Bz^\wedge .

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,ab,GoOgIc

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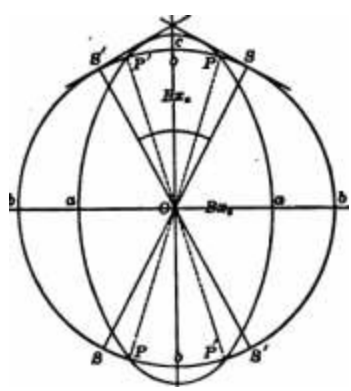
If the word bisectrix is used alone without special qualification it is always to be understood as referring to the acute bisectrix.

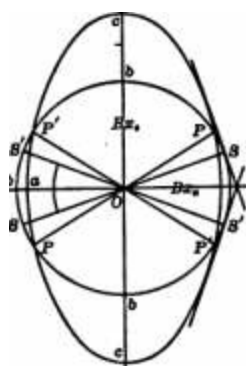
375. Relation of the Axial Angle to the Refractive Indices.—If in a given case the values of a, i, and v are known, the value of the interior optic axial angle (ω) can be calculated from them by the following formula:

I'

376. Positive and Negative Crystals.—Biaxial crystals are distinguished as

optically positive { -{- } } or optically negative { - } after a manner analogous to the usage with uniaxial crystals. Referring to Fig. 533 of the ellipsoid, and also to Fig. 539, it will be obvious at once that for certain relative values of the indices, a, fi, y, the interior optic axial angle must be 90° . In other words, in this





PoritWe CiTBtal, &■_ = t.

Negative Crystal, Bxa = a.

case the planes of the optic axes will be equally inclined to the two planes of the ether-axes. Such a case, however, is rare in practice, and when it occurs it is true for light of a certain color * (wave-length) only, and not for others.

■ Fbr danburite 2F = Be* U'for green (thallnm) and 90° 14'for blue (CuSOt).

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RoDghly expressed,* the optic axes will lie nearer to c than to a—that ia, c will be the bisectrix—when tne valne of the intermediate index, fi, is nearer that of a than to that of y, Snch a crystal, for which Bx^ = C, is called optically positive. It is obvious (cf. Fig. 541) that in this case, as the angle dimiuishes and becomes nearly equal to zero, the form of the ellipsoid then approaclies that of the prolate spheroid of the positive uniaxial crystal as its limit (Fig. 531, p. 195); this showB the appropriateness of the -|- sign here used.

On the other hand, the optic axes will lie nearer to a than to c—tli^t is, a will be the bisectrix—if the value of the mean index fi is nearerf to that of y than to that of a. Snch a crystal, for which £Xa = o, is called optically negative. It is seen that in this case (Fig. 542) the smaller the angle the more the ellipsoid approaches the oblate spheroid of the negative uniaxial crystal

(Fig. 520, p. 195). f he following

are a few examples of positive and negative biaxial crystals:

Positive (+). Negative (-).

Sulphur. Aragonite.

Enstatite. Hypersthene.

Topaz. Muscovite.

Barite. Orthoclaee.

Ghi[^]solite. Epjdote.

Albite. Axinite.

377. Dispersion of the Bisectrices.—In certain cases the ether-axes of Fig. 533 may have different positions in the crystal for different colors; that is, for different wave-lengths. This is true of the two axes which lie in the plane of symmetry of a monoclinic crystal, and of all the three axes in a triclinic crystal. This results in a phenomenon which is often called the dispersion of the bisectrices, and which, if pronounced, is always manifest in the axial interference-figure, as explained beyond.

378. Dispersion of the Optic Axes.—Further, since the three refractive indices may have different values for the different colors, and as the angle of the optic axes is determined by these three values (Art. 376), the axial angle may also vary in a corresponding manner.

This variation in the value of the axial angle for light of different wavelengths is called the dispersion of the optic axes, and the two extreme cases are distinguished by writing $p > v$ when the angle for the red rays (p) is

greater than for the blue (violet, v), and $p < v$ when the reverse is true. These cases are illustrated later.

Examination of Biaxial Crystals in Polarized Light.

579. Sections in Parallel Polarized Light. Extinction-angle.—A section of a biaxial crystal appears dark between crossed nicols when its vibration-plane coincides with the vibration-planes of the nicols. In any other position of the

• $\tan F = 45^\circ$ and $2\psi = 90^\circ$ for a value of ψ given by the equation

To compare Figs. 642 and B20, the linearly biaxial of the former should be placed in a vertical position; that is, the axis AA of Fig. 580 corresponds to that of Fig. M8.

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CHARACTERISTICS DEPENDING ON POLARIZED LIGHT.

section i^* appears light and may show the usual bright interference-color.

g^* Hence the use of the polariscope, or polarization-

j^* microscope, by the method of extinction,

gives a quick means of determining the position of these vibration-planes in a given case.

For example, in Fig. 543, repeated from p. 177, let the two larger rectangular arrows represent the vibration-directions for the two nicols, and between the two prisms suppose a section of a biaxial crystal, $abcd$, to be placed so that one edge of a known crystallographic plane coincides with the direction of one of these lines. The field of the microscope, dark before, since the prisms were crossed, is no longer so, and becomes dark again, as

explained, only when the crystal is revolved so that its vibration-directions (the smaller dotted arrows) coincide with those of the nicols, as is indicated by the maximum extinction of the light. The crystal has then the position $a'b'c'd'$. The angle (indicated in the figure) which it has been necessary to revolve the plate to obtain the effect described, is the angle which one of the vibration-directions in the given plate makes with the given crystallographic edge ad ; it is often called the extinction-angle.

When the vibration-planes of a crystal-section coincide in direction with the planes of its crystallographic axes, the extinction is said to be parallel; if not, it is called oblique or inclined. On the practical determination of the extinction-directions see Art. 380 and 389.

380. Determination of the Extinction-directions with the Microscope.- In the use of the microscope, in cases to which the method of the stauroscope is not applicable, instead of depending upon the somewhat uncertain estimate of the point of maximum tight-extinction, a convenient way is to employ a plate of quartz, which for a certain position of the analyzer gives the field a tint of color (a purplish pink), to slight changes in which the eye is very sensitive. When the section is revolved on the stage till it has precisely the same tint as the surrounding field, its vibration-planes are those of the cross-hairs in the eyepiece (supposing the lower nicol has the normal position), and from the graduation of the stage their position in the section can be at once determined.

Instead of the quartz, a plate of selenite (Art. 364) of such a thickness as to give the red of the first order is often employed, the nicol being crossed. When this plate is inserted as usual, the crystal-section can have the same color only on condition that its vibration-directions coincide with those of the nicols; hence their crystallographic orientation in the section is readily determined and with considerable accuracy.

A still more delicate method involves the use of the Bertrand ocular. This has four sectors of quartz cut in it: two of these, diagonally opposite, are from a right-handed crystal. The other pair from a left-handed crystal. When the diameters in which the sectors meet coincide with the vibration-directions of the two crossed nicols, the two pairs of quadrants have precisely the same color; any change of position, however, of the upper nicol causes them to assume tints complementary to each other.

Assuming now the nicols to be crossed and in the normal position, If a section of a doubly refracting mineral be introduced on the stage of the microscope, the quadrants In general take unlike tints and are brought to the same color as before only when by the revolution of the stage the vibration-directions of the section shall correspond precisely to those of the crossed nicols, that is. of the diameters of the ocular. The adjustment can be made in this case with great accuracy,

general case, for any section whatever, the relation between the vibration-planes and the ether-axis is highly complex.* A common special case is where the section is parallel to one of the ether-axes; this then fixes one vibration-plane, and the other will obviously be at right angles to it. A still more special case is that of a section parallel to the plane of two of the ether-axes; these axes then at once fix the directions of light-vibration and conversely; or, in other words, these directions being determined by observation, the position of the axes in the crystal section is known from them.

The practical application of the above relations depends upon the crystallographic orientation of the ether-axes, and is spoken of later under the different systems.

§2. Color of a Section in Parallel Polarized Light.—The interference-color of the section under examination depends, as before explained, upon its thickness and upon the birefringence; the latter varies with the orientation of the section, but is a maximum (equal to $\frac{1}{2}(\gamma - \alpha)$) if the section is cut parallel to the axes a and c. that is, parallel to the plane of the optic axes. In any case for a given thickness the interference-color will depend upon the difference between the refractive powers of the two vibration-directions. For sake of illustration, the following table is added (from Kosenbnsch) giving the thickness of section for a few biaxial crystals which yield red of the first order, with also their maximum birefringence $\{\gamma - \alpha\}$.

Birefringence Thickness in

$\gamma - \alpha$ millimeters.

Brookite 0.158 0.0034

Muscovite 0.042 0.01318

Epidote 0.037 0.01490

Augite 0.022 0.02505

Gypsum 0.010 0.05510

Orthoclase 0.008 0.0687

Zoisite 0.006 0.09183

Turlher, it may be noted that with a section of epidote (OLD), while the thickness of about 0.015 mm, gives red of the first order, one of 0.026 gives red of the second order, and of 0.013 red of the third order.

383. Determination of the Birefringence with the Microscope.—The value of the maximum birefringence ($\gamma - \alpha$) is obviously given at once when the refractive indices are known. It can be approximately estimated for a section of proper orientation and of measured thickness by the comparison of the interference-color with the table referred to in Art. 380, p. 178. This is on the assumption that the thickness is such as to yield a tint of readily recognized position in the interference-scale. To which order a given interference-color belongs can be readily determined by the compensation method involving the use of the quartz-wedge or the mica-wedge of Fedorow (Art. 384).

* The method of extinction with division of the orientation of the section under examination is chiefly included in the microscopic study of rock-crystals. It is more fully discussed in the *Mineralogische Atlas* of M. Levy and Lacroix. The subject has also been extensively treated in the case of the plagioclase feldspar by M. Levy in a work in two parts entitled "Etude sur la détermination des feldspaths dans les plaques minces" (Paris, 1894, 1896). Cf. also Fedorow, *Z. Krystallogr.*, 16, 37, 1896; 39, 604, 1898; also Viola, *Min. petr. Hiith.*, 16, 481, 1895.

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More accurate measurements of the birefringence can be made by other methods, as with the quartz-comparator of M. Levy.*

384. Determination of the Relative Birefringence.—The relative refractive power of the two vibration-directions in a thin section is readily determined with the microscope (in parallel polarized light) by the method of compensation. This is applicable to any section, whatever its orientation and whether uniaxial or biaxial. Practically, however, it is only employed when the section is parallel to the plane of two of the ether-axes; it then serves to determine the relative magnitude of these axes. If the position of either bisectrix in the section is known, it also serves to determine whether the crystal is optically positive or negative. In practice a mica-plate or selenite-plate may be used with very thin sections; or with thicker ones a quartz-wedge or the excellent mica-wedge of Fedorow. The section under examination must be placed so that its vibration-directions make an angle of 45° with those of the crossed nicols; the undulating mica or selenite-plate is then inserted. The change in the interference-colors is noted, and again after the stage has been revolved 90° . In the case where the effect of the compensating plate is to reverse the interference-color in the scale the retardation of the section is added to that of the plate; for this position the plate and section are then alike in optical character. In the case of a fall of color the plate diminishes the retardation due to the section; obviously for this position they are opposed in optical character.

For thicker sections placed in position as before with vibration-directions inclined 45° to those of the crossed nicols, the quartz-wedge or mica-wedge is employed. This is advanced across the field until a thickness is reached which serves to make the given section, previously colored, dark, that is, to bring it sensibly to extinction. The optical character of the crystal and section are now opposed, and in fact just balanced against each other. By observing the position of the section the relative value of the refractive power is given.

A crystal-section is said to have positive elongation if its direction of extension approximately coincides with the ether-axis c ; if with a the elongation is negative. The same terms are also used in general, according to the relative refractive power of the two directions.

365. Determination of the relative elongation.—The values of the three refractive indices, a , b , c , for biaxial crystals, may be determined from three prisms cut with their refracting edges parallel respectively to the three axes, a , b , and c , corresponding to the indices a , b , c , respectively. See Art. 304.

It is possible, however, to obtain the values of a , b , and c by the use of two prisms; in this case one of the prisms must be so made that its vertical edge is parallel to one axis, while the line bisecting its refracting angle at this edge is parallel to a second. In the case of such a prism the angle of minimum deviation is obtained for both rays, that having its vibrations parallel to the prism-edge, and that vibrating at right angles to this, that is, parallel to the line bisecting the prismatic angle.

Of the three indices the mean index, b , is one which it is most important to determine, since by means of it, in accordance with the formulas in Art. 389, the true value of the axial angle can be calculated from its apparent value

see M. Levy, Bull. Soc. Min., 6, 143, 1858; also Levy, *Microirois*, p. 74; Roemer, Mikr. Phys., p. 105.

This is a series of overlapping step-like and all placed with their axial planes (i.e., the axis c) in a common line coinciding with the direction of vibration. In the case of crossed nicols, the axis c makes an angle of 45° with the principal planes, it gives a series of interference colors whose position in the scale is given (Art. 330; see also Art. 330).

« obvious (Art. 330). See Pedersen, *Za. Krysl.*, 26, 848, 1885.

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in air. The prism to give the value of b should obviously have its refracting edge parallel to the axis c , that is, at right angles to the plane of the optic axes.

The other methods* alluded to on pp. 160, 181 may also be applied here. For the method by total reflection the sections should be cut normal to the acute bisectrix. This will give by actual observation the values of a and y , and if iE , the apparent axial angle in air, is known, then p , the mean index, can be calculated.

386. Interference-figures for a Section Normal to an Optic Axis.—A section cut perpendicular to either optic axis will show, in converging polarized light, a system of concentric rings analogous to the concentric circles of uniaxial crystals, Fig. 516, but more or less elliptical in shape. There is, moreover, no black cross, but a single black line, which revolves as the section is turned around on the stage.

387. Interference-figures for Section Normal to a Bisectrix.—If a section of a biaxial crystal, cut perpendicularly to the acute bisectrix, is viewed in the conoscope (p. 181), two types of characteristic interference-figures are observed, according to the position of the optic axial plane relative to the vibration-planes of the crossed nicols.

First, suppose that the plane of the axes coincides with the vibration-plane of one of the crossed nicols; an unsymmetrical black cross is then observed, and also a series of elliptical curves, surrounding the two centers and, finally uniting, forming a series of lemniscates. If monochromatic light is employed, the rings are alternately light and dark; in white light each ring shows the successive colors of the spectrum. If one of the nicols be revolved, the dark hyperbolic brushes gradually become white, and the colors of the rings take the complementary tints after a revolution of 90° .

The smaller the axial angle the nearer together are the oval centers and the more the interference-figure resembles the simple cross of a uniaxial crystal. On the other hand, when the axial angle is large the hyperbolas are far apart, and may even be so far apart as to be invisible in the field of the polariscope. When this is not the case a micrometer scale in the polariscope. Fig. 514, p. 188, allows of an approximate measurement of the axial angle, the value of each division of the scale being known.

Again, suppose that the plane of the axes makes an angle of 45° with the vibration-planes of the crossed nicols; two black hyperbolas are then noted, marking the position of the axes; further, there is a system of curves similar to those described before.

Both of these figures are well exhibited on the plate forming the frontispiece (Plates 1 and 3A). The one gradually changes into the other as the crystal-section is revolved in the horizontal plane, the nicol remaining stationary.

A section of a biaxial crystal cut perpendicular to the obtuse bisectrix exhibits the same figures under the same conditions in polarized light, when the angle is not too large. This is, however, generally the case, and in consequence the axes suffer total reflection (Art. 303) on the inner surface of the section and no axial figures are visible. This is sometimes the case also with a section cut normal to the acute bisectrix, when the angle is large.

(See, also,

• Becke baa friven a Bimpla metLod for deterrntninK approilmatel v tbe retalm Tefracliva power of 1 wo adjacent mineralB (e.g.. quartz nn<i acerlaln kind orpl«(rfocIa8e) !n Uiin aectlong. Bar Ak. Wlen, 103 (1), July. 1898. See also Viola. Min. petr, Mittli., 18, 150, 18M.

t On Ihe special phenomens of secllona of Wnxial cryBtsls cut normal to an opilc axis, in pai»llet and converging pobrized Hglit, bm Kalkowsky, Zb. Krysl., «, 486,1884.

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Art. 389.) The pecnliavitiea in the interfecece-figures dae to the dtapersioii of the optic axes and that due to the dleperHiou of the bisectricee, or both together, are alluded to later.

388. The explanation of the biaxial interference-figures—moat simply understood for the first case mentioned—is analogons to that for the analogous phenomenon of the uniaxial crystals (Art. 360). The arma of the black cross mark the directions in which the tight-rays, which are sensibly norma! to the section, are extinguished, since the vibration-planes of the uicols coincide with those of the section. The dark ellipses and lemniscate curves seen in mono-chromatic light are dae to the interference of the two raya produced by the double refraction of the section and referred back to a common vibration-plane by the polarizer. This interference takes place when the rett^rdation of one ray relatively to the other is equal to half a wave-length, $\frac{1}{2}\lambda$, or to fA , $\frac{1}{2}A$, etc. The intermediate light-spaces correspond to a similar retardation of a whole wave-length, A , or $3A$, $3A$, etc. When ordinary tight is employed there is complete extinction only in the direction sf the vibration-planes of the nicols, and the curves become colored rings showing the prismatic colors. The number of colored rings noted in the field of view increases, and their distance from the axial centers and from eacli other grows less as the thickness of the

plate is increased, and also aa the strength of the double refraction is greater. f the plate is very thick, only the black cross may be distinctly visible.

389. Heasurement of the Axial Angle.—The determination of the angle made by the optic axes is ordinarily accomplished by use of the instrument shown m Fig. 544. The section of the crystal, cut at right angles to the bisectrix, is held in the pincers at p, with the plane of the axes horizontal, and

making an angle of 45° with the vibration-plane of the nicola. There ia a cross-wire in the focus of the eyepiece, and ns the pincers holding the section are turned by the screw at the top (liere omitted) one of the axes, that is, one

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black hyperbola, is brought in coincidence with the vertical cross-wire, and, then, by a further revolution, the section

is turned through the angle which the section has been turned

from one axis to the second, as indicated at the vernier on the graduated circle above, is the apparent angle for the axes of the given crystal as seen in the air ($E = 2E'$, Fig. 545). It is only the apparent angle, for, on passing from the section of the crystal to the air, the true axial angle is more or less increased, according to the refractive power of the given crystal. The relation between real interior angle and the measured angle is given below.

If the axial angle is so large that the axes suffer total reflection, oil or some other liquid with higher refractive power is made use of, into which the axes pass when no longer visible in the air. In the instrument described a small receptacle holding the oil is brought between the tubes, as seen in the figure, and the pieces holding the section are immersed in this and the angle measured as before.

In the majority of cases it is only the acute axial angle that it is practicable to measure; but sometimes, especially when oil (or other liquid) is made use of, the obtuse angle can also be determined from a second section normal to the obtuse bisectrix.

If E' = the apparent semi-acute axial angle in air (Fig. 545), H' = " " " " in oil,

F = the real (or interior) semi-acute angle, V_a = " " " " semi-obtuse angle, n = refractive index for the oil or other medium, μ = the mean refractive index for the given crystallized substance, the following simple relations connect the various quantities mentioned:

$$\sin E = n \sin E'$$

These formulas give the true interior angle ($2F$) from the measured apparent angle in air ($2E'$) or in oil ($2H'$) when the mean refractive index (μ) is known.

Instead of the oil, carbon disulphide with a refractive index of 1.6442 for $X(\lambda)$ may be employed; or the solution of mercuric iodide in potassium iodide, whose refractive index (μ_o) is 1.76. The axial angle measured in the latter is usually represented by $2K$. Methyl iodide may also be used, since its refractive index is also high; for it $n = 1.7466$. The axial angle in this case is called ZM .

Another modified form of axial instrument (the Adams-Schaeffer) encloses the section in a sphere of glass which can be turned at the proper angle; for the axial angle is

— ' - —•.>-<- '----umeul consist in the fact that the field of view - ' - -

If any desired position

• Almond-oil which has been decolorized employed; its refractive index is about 1.46.

' exposure to the light to commonly

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Furthermore, the angle measured in air made, so that the axes are visible in light, its reflection, either in air or in oil.

890 AzklAngl* Haunred vltb tlie]U«roMop«.—The microscope, with eyepiece removed (ami ciiiideusiig ieiisiur lensei.) added above the lower Nicol, often serves to show (with a high-power objective) the axial interference-figures in very thin sections with only the interference-bands belonging to the ordinary (utic p. 119). Bertrand* has shown that the addition of a simple piece of apparatus the axial angle can be measured with fair accuracy. Further, a very convenient apparatus for this object has been devised by Kline. Again, by the careful measurement of the linear distance between the two isotherms the axial angle can be calculated as shown by Mallard.

391. Determination of the Positive or Negative Character of Biaxial Crystals.

—The question of the positive or negative character of a biaxial crystal is determined from the values of the indices of refraction, where these can be obtained. If C , the ether-axis corresponding to the index y , is the acute bisectrix, the crystal is optically positive; if a , the ether-axis corresponding to a , is the acute bisectrix, the crystal is optically negative; as explained in Art. 376 and illustrated by Figs. 541, 543, this relation follows from the values of the refractive indices.

There are, however, more simple methods of determining the character by experiment. The quarter-wave mica plate may be employed just as with uniaxial crystals, but its use is not very satisfactory excepting when the axial divergence is quite small. In this case it can be used to advantage, the plane of the axes of the crystal investigated being made to coincide with the vibration-plane of one of the nicols. With the microscope the selenite plate may be employed after a manner similar to that explained in Art. 364,

392. The more general method is the employment of a thin, wedge-shaped piece of quartz; this is so cut that one surface coincides with the direction of the vertical axis, and the other makes an angle of 4° to 6° with it. By this means a wedge of varying thickness is obtained. The section to be examined, cut normal to the acute bisectrix, is brought between the crossed nicols of the polariscope (Fig. 513), and with its axial plane making an angle of 45° with the polarization-plane of the Nicol prisms; that is, so that the black hyperbolas are visible. The quartz-wedge is now introduced slowly between the section examined and the analyzer, first, in a direction at right angles to the axial plane, that is, to the line joining the hyperbolas, of the plate investigated; and second, parallel to the axial

plane, that is, in the direction of the line joining the hyperbolas. In one direction or the other it will be seen, when the proper thickness of the quartz-wedge is reached, that the central rings appear to increase in diameter, at the same time advancing from the center to the extremities

The effect, in other words, is that which would have been produced by the thinning of the given section. If the phenomenon is observed in the first case when the axis of the quartz is parallel to the axial plane, that is, to the obtuse bisectrix, it shows that this bisectrix must have an opposite sign to the quartz, that is, the obtuse bisectrix is negative, and the acute bisectrix positive. If the mentioned change in the interference-figures takes place when the axis of the quartz is at right angles to the axial plane, then obviously the opposite must be true and the acute bisectrix is negative. This method of investigation

•Berlinil- Bull- Boc. Min.. 3. 97. 1880; see, also, Sachet, Ibid., 10, 186. 1887; Klein. Ber. Ak. Berlin. 61, 1895. Also references in p- 199.

\ Mnllird, Bull. 8nc. Mlu., 6, 77, 1883; This last method is explained by Rosenbusch, Uikr. Pliyi,, 194, 1893.

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can be applied even in cases where the axial angle is too large to appear in the air.

The same effects may be obtained by bringing an ordinary quartz section of greater or less thickness, cut normal to the axis, between the analyzer and the crystal examined, and then inclining it, first in the direction of the axial plane, and again at right angles to it.

The method of determining the optical character in thin sections in parallel polarized light is mentioned in Art. 384.

393. Absorption Phenomena of Biaxial Crystals. Pleochroism. Biaxial crystals, corresponding to the fact that they have three principal refractive indices $[a, \beta, \gamma]$ may show different degrees or kinds of absorption in different directions, usually assumed as those of the ether-axes, viz., a, b, c . The degree of absorption is designated as $a > b > c$, etc. Further, according to the kind of selective absorption, the crystal may be dichroic or trichroic, or better, in general, pleochroic;* in this case the colors corresponding to the vibrations parallel to the ether-axes are usually given. It has been shown, however, that the axes of absorption do not in all cases coincide with the ether-axes.

In order to investigate the absorptive properties of a biaxial crystal, sections must be obtained which are parallel to the several ether-axes, cf. 646. Fig. 546. In an orthorhombic crystal the faces are those of

the three pinacoids; in a monoclinic crystal one side coincides with the clinopinacoid, the others are to be determined for each species. The light transmitted by this solid, or by the

corresponding sections, is examined by means of a single nicol prism. Suppose, first, that the light transmitted (Fig. 546) in the direction of the vertical axis is to be examined. When the shorter diagonal of the nicol coincides with the direction of the axis l , the color observed belongs to that ray with vibrations parallel to this direction; when it coincides with the axis n , the color for the ray with vibrations parallel to a is observed. In the same way the nicol separates the different colored rays vibrating parallel to C and Q respectively, when the light passes through in the direction of b .

So also finally when the section is looked through in the direction of the axis o , the colors for the rays vibrating parallel to b and c , respectively, are obtained. It is evident that the examination in two of the directions named will give the three possible colors. All of these observations are readily made with the microscope provided with one nicol.

For epidote, according to Klelu, the colors for the three axial directions are: f Vibrations parallel in i . brown (absorbed).

$\} \text{ " " " } \ll \text{. yellow.}$

$a \text{ I y t b r a t l o D S } \text{ parallel in } I \wedge \text{ r p e n. } * \blacksquare 1 \blacksquare \text{ " } \blacksquare \blacksquare \text{ a. yellow.}$

$\wedge \text{ j T i b r a t l o D B } \text{ parallel to } c \text{ freeo.}$

$\backslash \text{ " " " h. brown 'absorbed).}$

The colors observed by the eye alone are the result of the double refraction of vibrations.

In which the stronger color predominates. Thus, in the above example, the plane normal to

(a is brown, b yellowish-green, c a green. In any other direction in the crystal the

apparent color is the result of a mixture of those corresponding to the three directions of vibrations in different proportions. When a section (normal to an optic axis) of a crystal characterized by a high degree of color-absorption is examined by the eye alone (or with the microscope) Early observations were made by Haidinger, see literature below.

Microscopic

(microscopic) in strongly converging light, it often shows the so-called optic figures* or polarization-brushes, somewhat resembling the ordinary axial interference-figures. This is true of andalusite, epidote, iolite, also tourmaline, etc. A cleavage section of epidote || c (001) held close to the eye and looked through to a bright sky shows the polarization-brushes, here brown on a green ground.

It is to be noted that certain strongly absorbing crystals (e.g., biotite) often show

spots where the color is particularly deep; such areas are sometimes called pleochroic
halm.

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1854. Brewster. C. & S. 38. 101, 18M; or Fogg. Ann., 9, 491, 1854. Trichemuk. Ber. Ak.
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Ephemer. Til. Mid., n. 1872.

Bertin. "Uvula," etc. Ann. Cli. Phys., 16, 896, 1878, and Zs. Kryst., 3, 449. 1879. Also
Berl. Mon. Uebers. Mith., 2, 67, 1879; Mallard, ibid., p. 73. Mon. Piedmont. Zs.
Kunstl., 4, 444, 1880. Pulfrich. Zs. Kryst., 6, 142, 158, 1881. Mallard Bull. Sue. Mid. 2. 73.
1879; 6, 45, 1888. Voigt. Jb. Min., 1, 1875; Wied. Ann., 31, 288, 1887. Ranuay.
J. Phys. Zs. Erysl., 13, 97, 1887. Uebers. Nachr. Oest. Quedlinburg, No. 8, 202, 1888. Bland.
Jb. Hin., Beil.-Bd., 11, 269, 1897.

Special Optical Characters of Orthorhombic Crystals.

894. Formation of the Ether-axes.—In the orthorhombic system, in accordance with the
symmetry of the crystallization, the three axes of the indicatrix, that is, the ether-axes a ,
 b , c , coincide with the three crystallographic axes, and the three unlike crystallographic
planes of symmetry correspond to the planes of symmetry of the ellipsoid. Further than
this, there is no immediate relation between the two sets of axes in respect to magnitude,
for the reason that, as has been stated, the choice of the crystallographic axes is arbitrary
so far as relative length and position are concerned, and hence made, in most cases,
without reference to the optical character.

Sections of an orthorhombic crystal parallel to a pinacoid plane (a , b , or c) appear dark
between crossed nicols, when the axial directions coincide with the vibration-planes of
the nicols. In intermediate positions a section will appear light (or colored if of the proper
thickness). Hence such a section when revolved on the microscope will appear dark four
times.

The same will be true of a section cut in the prismatic zone (a & the vertical axis) or in
the plane parallel either to the axis b or c .

395. Determination of the Plane of the Optic Axes.—The plane of the optic axes, that is, the
plane of the axes a and c , corresponding to the indices μ_x and μ_y , must be parallel to one of the three
pinacoids. In order to determine in which plane the axes lie, it is necessary to examine
sections parallel to these directions. One of these will in all ordinary cases show, in
converging polarized light, the interference-figure peculiar to biaxial crystals. It is evident
also that two of the sections named determine the character of the third, so that the plane
of the optic axes and the position of the acute bisectrix can be in practice generally told
from them.

The position of the optic axial plane is briefly indicated according to the pinacoid to which it is parallel: as, ax, plane H rt (100), etc. Further, the

position of the acute bisectrix is described according to the pinacoid to which it is normal, as Bx, SI", etc.

396. Siipersleii of the Axes.—From the section showing the axial figures, that is, normal to the acute bisectrix, the axial angle can be measured in the manner which has been described (Art. 389). If it is practicable to determine also the obtuse axial angle from a second section normal to the obtuse bisectrix, the true axial angle ($2V$) and also the mean refractive index (μ) can then be calculated.

There is further to be determined the dispersion of the axes (see Ait. 378). Whether the axial angle for blue rays is greater or less than for red ($V_{\text{blue}} > V_{\text{red}}$, or $\mu_{\text{blue}} < \mu_{\text{red}}$) can be often seen immediately from the axial interference-figure in the colored plate (frontispiece); it is obviously true, from Fig. 3n, as also Fig. 3b, that the angle for the blue ray is greater than that for the red ($V_{\text{blue}} > V_{\text{red}}$), and so in general. This same point V is also accurately determined, of course, by

" the measured angle for the two colors.

In all cases the same line will be the bisectrix of the axial angle for both blue and red rays, so that the position of the respective optic axes is symmetrical with reference to each bisectrix, in Fig. 547, the dispersion of the axes is illustrated, where $\mu_{\text{blue}} < \mu_{\text{red}}$, it is shown also that the lines H'B' and B'B' bisect the angles of both red (V_{red}) and blue (V_{blue}) rays. It also needs no further explanation that for a certain relation of the refractive indices of the different colors, the acute bisectrix of the axial angle for red rays may be the obtuse bisectrix for the angle for blue rays. This is true, for example, in the case of the species danburite, as already noted (p. 209).

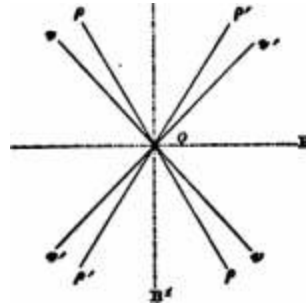
397. Refractive Indices, etc.—The determination of the refractive indices and the character (+ or -) of the acute bisectrix is made for orthorhombic crystals in the same way as for all biaxial crystals. It is merely to be mentioned that, since the ether-axes always coincide with the crystallographic axes, it sometimes happens that crystals, without artificial preparation, furnish, in their prismatic or dome series, prisms whose edges are parallel to these axes, and consequently at once suitable for the determination of the indices of refraction. This is often true, for example, with topaz. It must be noted, however, that if the refractive angle is too large, the refracted ray will not emerge (see Art. 303), the limit being when the angle of the prism is equal to twice the critical angle.

Special Optical Characters of Monoclinic Crystals,

398. Position of the Ether-axes and Optic Axial Plane.—In crystals belonging to the MONOCLINIC system one of the ether-axes always coincides with the orthodiagonal axis b.

and the other two lie in the plane of symmetry at right angles to this axis. Here obviously three cases are possible, according to which two of the axes, a, b, or c, lie in the plane of symmetry.

Corresponding to these three positions of the ether-axes, there may occur three kinds of dispersion of these axes, or dispersion of the bisectrices (Art. 377). This dispersion arises from the fact that, while the position of



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one axis is fixed, the position of the other two may be different for the different colors (wave-lengths), so that the bisectrices of the different colors may not coincide.

399. Stauroscope. —The position of the two axes in the plane of symmetry may be determined by use of the microscope described in Art. 380. A more accurate method, applicable in certain cases, involves the principle of the Ktanoscope, an instrument first devised by von Kobell (1855 •), and since then much improved. In its present form it is simply the orthoscope of Fig. 515 with a composition plate of calcite inserted at m in the tipper tube. This gives a peculiar interference-figure the form of which is altered if a doubly refracting section is placed on the stage below, unless its vibration-planes coincide with those of the crossed nicols. The adjustment which restores the normal figure can be made with great precision. To accomplish this, it is essential that the direction of the known edge of the crystal should be exactly parallel to the vibration-direction of one of the nicols. This condition, in the case of small crystals especially, is hard to

fulfill, and to accomplish it most satisfactorily

the holder shown in Fig. 548 is made use of.

A plate of glass, v, held in its present position

by a spring, has one edge polished, that which

adjoins », and the direction of this is made to

coincide exactly with the line joining the opposite

zero points of the graduation. The crystal section

is attached to this plate over the hole seen in v,

and with a plane of known crystallographic

position, either n, c or a face in that zone {or a

corresponding edge), coinciding with the direction

of the polished edge of the plate. Whether this

coincidence is exact can be tested by the reflecting goniometer.

After the adjustment of the section on the plate v, the latter is inserted in its place, the whole plate placed in position (Fig. 515), and the nicols so adjusted that the vibration plane of one coincides with the line 0° to 180° , The angle of revolution of the circular plate, θ , is obtained from the graduated scale on k.

It is not always easy to make the adjustment of the nicols alluded to, but the error arising when the vibration-plane of the nicol does not coincide with the line 0° to 180° is easily eliminated. This is accomplished by removing the plate V, and, without disturbing the crystal section, restoring it to its place in an inverted position. The measured angle, if before too great, will now be as much too small, and the arithmetical mean of the two measurements will be the true angle. (Cf. Groth, 1. c.)

400. Position of the Plane of the Optic Axia. —The investigation of a section of a monoclinic crystal parallel to the plane of symmetry determines the position of the two ether-axes lying in this plane, but it does not fix the relative position of the axes a and c, that is, the plane of the optic axia. To solve the latter point, sections normal to each of the three axes must be examined in converging polarized light, and one of them will show the characteristic interference-figure. The section parallel to the plane of symmetry is first to be examined, and if it does not show the axes even in oil, one or both of the other sections spoken of must be employed.

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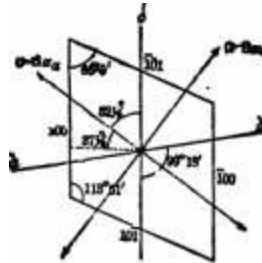
The position of the optic axial plane is described as ψ according as it is parallel or normal to the plane of symmetry, that is, to the face $h(010)$. In the former case, the position of the bisectrices may be defined according to the angle which the acute bisectrix (B_z) makes either with the normal to the face a (100) or that of c (001) or with the vertical axis i. The last method is

particularly convenient because the direction of the vertical crystallographic axis, i , is that marked by the prismatic zone (e^\wedge , in a section by cleavage lines), and still more since the

extremity of i is the middle point of the sphere of projection (Fig. 550) and the angle is either $-|-$ (in front, toward 100) or $-$ (behind, toward 100).

For example, Kypsum is optically positive, hence the $sxts$, e , h the acute bisectrix, Bxa . Farther (Figs. 6&, 650), the $pOBltloo$ of Be . U defined by the angle $Bxa Ai = + 6fH^\circ$. But $BlQce$ the $hkUI$ angle $/S, oi ik (001 A 100), = 80^\circ 43'$, it is ntra true that the normal

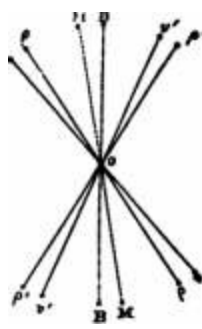
649. 660.



a and the planes e , a are as follows: $\ll = + 48' 13'$, $ac = 4- ST 80^*$, The poaitlone of the bisectrices are shown in the section parallel to $ft (010)$ of Fig. 646. and also Id the sphere of pvojecloii, Fig. 650,

551. 401, Siapereiou of the Biseotrioes.—1. Inclined

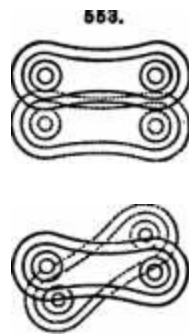
DiSPKRSiON.—In this first case the plane of the optic axes is parallel to the face $h (010)$; in other words, the two bisectrices (the axes a and C) lie in the plane o (symmetry, and the mean axis b coincides with the orthodiagonal axis. The optic axes may here suffer a dispersion in this plane of symmetry, and, as already stated, they then do not lie symmetrically with reference to the acute bisectrix. This is illustrated in Fig. 551, where JO is the bisectrix for the angle vOv' , and BB for the angle pOp' . This kind of dispersion was called by Dea Gloizeaux inclined (dispersion inclinee). The position of the two axial planes is further illustrated by Fig. 553 (from Schrauf), and corresponding to this the axial interference-figure, when the dispersion is considerable, shows a distribution of colors illustrated by Figs. 4a and 4b of



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tho plate (frontispiece), which should be compared with the Dormsl figures (Figs. 3a and Bb), where there is no dispersion of the bisectrices.



S. Horizontal. Ax. pl.X6; Bx,,±b.

' 8. Crossed. Ax. plX^; Bx^±b.

1. loclioed. Ax. pi. B b (010).

2. HoEizoNTAL DISPBBBiON. In the second case the acnte bisectrix and b lie in the plane of sjmmetrj, and the optic axial plane is heuoe normal to it. In other words, the plane of the optic axes for aJl the colors lies parallel to the orthodiagonal axis, but these planes may have different inclinations to the vertical axis. This was called horitovlal diaperaion by Des Cloizeaox.

The relative positions of the axial planes is illnBtrated by Fig. 553, and the resulting interference-figure is shown in Fig. 5 of the plate.

3. Crossed Dispersion. In the third case the obtuse bisectrix and b lie in the plane of symmetry, that is, the plane of the optic axes is normal lo this plane and the acuta bisectrix coincides with the orthodiagonal axis t. This was called crossed dispersion by Dea Cloizeaux (dispersion tonmante or crois^e). The relative positioiis of the axial planes is illstrnted by Fig. 554 and the correspoding interfereiice-figure is shown in Fig. 6 of the plate.

402. Axial Angle, Dispersion, etc.—The metlrod of measuring the axial angle has been

already explained, and if this is determined for the different colors it will determine the dispersion of the axes p^v .

The dispersion of the bisectrices has been shown to be in general indicated by the character of interference-figures; its amount, where considerable, may be determined by making the stauroscopic measurements for different colors.

The remaining points to be investigated, the refractive indices, the + or - character of the crystal, etc., need no further explanation beyond that which has been already given.

Special Optical Characters of Triclinic Crystals.

403. The crystals of the triclinic system are characterized by the absence of a plane of crystallographic symmetry, the position and inclination of the axes being thus arbitrary. It follows from this that there is no necessary connection between them and the rectangular ether-axes. More than one of the three kinds of dispersion mentioned in Art. 401 may occur in a single crystal, and the interference-figures will indicate the existence of both.

The practical investigation of triclinic crystals optically involves considerable difficulty; in general a series of successive trials are required to determine the position of the axes. When these are found, the axial sections can be prepared and the axial angle determined, and the other points settled as with other biaxial crystals. Cf. Fig. 341, p. 109, of chalcantite, where S repre-

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PHYSICAL CHARACTERISTICS.

resents approximately the position of Bx^v , or in other words is the pole or normal to the plane at right angles to the acute bisectrix. On the general relation between the extinction-directions, ether-axes and optic axes, see the authors referred to on p. 212.

404. Effect of Heat upon Optical Characters,—The general effects of heat upon crystals as regards expansion, etc., are spoken of later. It is convenient, however, to consider here, briefly, the changes produced by this means in the special optical characters. It is assumed that no alteration of the chemical composition takes place and no abnormal change in molecular structure. The essential facts are as follows:

(1) Isotropic crystals remain isotropic at all temperatures. Crystals, however, which like sodium chlorate (NaClO_3 , of Group 5, p. 51) show circular polarization, may have their rotatory power altered; in this substance it is increased by rise of temperature.

(2) Biaxial crystals similarly remain uniaxial with rise or fall of temperature; the only change noted is a variation in the relative values of m and e , that is, in the strength of the

double refraction. This increases, for example, with calcite and grows weaker with beryl and quartz. It is, further, interesting to note that the rotatory power of quartz increases with rise of temperature, but the relation for all parts of the spectrum remains sensibly the same.

(3) With Biaxial crystals, the effect of change of temperature varies with the system to which they belong.

The angle of biaxial crystals may be measured at any required temperature by the use of a metal air-bath. This is placed at 644° , and extends beyond the luminous spectrum on either side, so as to allow of its being heated with gas-burners; a thermometer inserted in the bath makes it possible to regulate the temperature as may be desired. This bath has two openings, closed with glass plates, corresponding to the two tubes carrying the gases, and the crystals, held as usual in the holders, is seen through these glass windows. Suitable accessories to the refractometer also allow of the measurement of the refractive indices at different temperatures.

In the case of orthorhombic crystals, the position of the three rectangular axes cannot alter, since they must always coincide with the crystallo-

graphic axes. The values of the refractive indices, however, may change, and hence with them also the optic axial angle; indeed a change of axial plane or of the optical character is thus possible.

For example, Des Clozeaux gives the following values for barite: $2S = 68^{\circ} 5' - 69^{\circ} 49'$ at 05.5° , $74'$ at 195° . Further, Aron has obtained the following values of the refractive indices of the same species for the D line:

$\frac{1}{S} \wedge SB$

With uniaxial crystals, one ether-axis must coincide at all temperatures with the axis of symmetry, but the position of the other two in the plane of symmetry may alter, and this with the possible change in the value of the refractive indices may cause a variation in the degree (or kind) of dispersion as well as in the axial angle.

With tridinic crystals, both the positions of the ether-axes and the values of the refractive indices may change. The observed optical characters may therefore vary widely.

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A striking example of the change of optical character with change of temperature is furnished by gypsum, its alteration being described by Des Clozeaux. At ordinary temperature, the dispersion is anomalous, the H and B rays are separated at 116° . At a higher temperature, the dispersion becomes normal, the H and B rays are separated at 130° . At still higher temperatures, the dispersion becomes anomalous again, the H and B rays are separated at 130° .

the IUES [or Ttd rays also seiiriiltt lu this p.aie (j. b) sod the disporsioa becomea horiional. Tbe uiullOD towiiTd the center of one red azia is mure ra))ld tliua tliut of the other, uamely betwuvD 30° aud 85° , $83^\circ 55'$ aud $Zi^\circ 88'$, respectively; lliua Bx, moves B $^\circ 88'$.

Auuther ialereatlDg case Is that (if glauberlte. Its optical cbaraclers under uoruial conditions are described as follows : Optically —, A*, pi. 1 6, Bx,, a A = - $31^\circ 8''$, Bx^, = - $30^\circ 46'$, Bi^w A A = - $80^\circ 10'$. The optical charactT (-) and the poaltion of the axes of elasticity remain sensibly consiinnl betwei-u 0° aud 100° . 'ihe ax. pi., however, at first X b with horizontal dlapeiBion nnd v <, fi becomes on rise of temperature fb with incliued dispersion and t> > p. Tbe axial auple accordingly diminisbes to 0° at a temperature depending upon the wave-lunglh aud then increaaea lu lbe new plane. In wblle light, therefore, the Iuierference-figures are abuurmal and cliange with rise in tempernture. Axial angles, L.iapeyrea:

red (Li) yellow (Na) green (Tl) blue

At S $^\circ 3E = 16' 0' 14^\circ 8 11^\circ 43' 8^\circ 81'$

$32^\circ '' = 18^\circ SO$ ir $8' 8^\circ 14' 0^\circ$ (at $18''$)

;'14' tlie feldsoni change in the position of the axet they do not return again to Ibelr o:

In addition to the typical cases referred to, it ia tobe noted that vhoii elevation of temperature is connected with change of chemical comjiosition wide changes in optical characters are possible. This is illnstrated bj the zeolites and related species, where the effect of loss of water haa been particularly investi-

Further, with some crystals, heat serves to bring about a change of molecular etructure nnd with that a total change of optical characters. For example, the

freeniah-jeHow (artificial) orthorhombic crystals of antimony iodide (Sbl,) on eating (to about li4'') change to red uniaxial ihexagonal crystals. Note also the remarks made hiter in regard to the effect of heat upon leucite and boracite (Art. 411).

iOS. BomaPseollaritleiln Axial InterfertnBS-flgurM.*—In Ibeciise of uniaxial cryatala, the cliarncteristic iili:rferauce-flgure varies but liLlle from one apeties to auuihor. such variation OS is ubaervecl being usually due In the thickness of the sucliou and the birefringence. In some ciises, however, peculiarities are noted. For example, the Inderferencu-flgure of npophyllite la aomewhat peculiar, sliico its liltefrlDgcnce la very weak, aud It may be optically positive for one part of the apeclruiu and negative for the other.

In the case of biaxial crystals, peculiarities are more common. The following are some typ'c-nt examples ■

BrKriklie Is optically-f nnd the acuti: bisectrix Is always normal to a (100). While, how-
cv.T. the fixial piano Is [a for red and yellow, willi 2ff, = 65°, 2Bi, = 80°. it ia commonly \b
for green anil blue, with 2S!^ = 34°. Hence n section la Iii the c<iuoscoi>e shows a figura
sompirhnt resembling llmt of a unlnxial cry-tnl but witli f.uir sets of byperboUc bands.

Titnnlte iil»o elves n pecnllnr [nteiferencp-figure with colored hyperbolas becniiise of the
high color-di«per-ion. o > n: I'uui r>es Oloizpaiix gives 2Br = 55° to 66°, 2E, = 84°; the
'linpendon of the Miiwtrirea 1». howi-vcv, very araall.

The mmt xtrikng canes of p'-oiillnr iirin! fli;iircB nrc affonted by twin cryatala (Art. 40T).

408. Relation of Optical Propertiea to Chemical Compoition.—The effect of

Taryiug chemical composition upon the optical characters haa been minutely

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studied in the case of many series of isomorphons salts, and with important results.* It is,
indeed, only a part of the general subject of the relation between oryatalline form and
molecular structure on the one hand and chemical composition on the other, one part of
which has been discussed in Art. 302. It was shown there that the refractive index can
often be approximately calculated from the chemical composition.

Amoug inluKTiils. Ilie most important examples of the relutlau between cotnpodliou iind
oplfcal uL&ntcturB are afforded by llie triclluic felilspurs or tUe albUe-atiortbite seiies.
Here, us uxpluJued !□ detail iu Ibe tlescrijiiiive part of this work, the relatloD is so close
lliut thecumpusltionof any tmermedlste member oF this isomorpboua group cau be
predicted from llie posltlou of ll« etlier-axes, or more aimply from (lie
TibTB^ondrectluns od the fuiidBmeulul clesTsge-dirsctlouB,]e (001) mid [b (010).

The effect of Tarjing urouDts of liOD protoxide (FeO) Is illiistrfiled Iu the case of lbs

orthorliombic species of the same group, eustatlte, MgSiOinlth llllle Iron, and
byper«tbeDe, (Hg,Fe}SiOi with iron to uearly SO p. c. With both of these species the mini
plsne ii

Etralle] to 6(010), but the former is optically + (Bv. = c)Ht[i] the liispt-rsioD/> < c: the
latter optically -' (Bx, = a) aud dispeision p > v. In other words, the optic axial angle
Increosea rapidly with the FeO percentage, Ijeinc about 60° for FeO = 10 p. o. Iu the cuse
of the chryeblitea, the epidotee, the species tdphyJite uud lithioplbtie, and others,
analogotu telatlonu have been mode out.

407. Optical PropertiH of Twin Cryatab.—The examination of sections of BBS. twin
crystals of any other than fhe isometric

sjBtem in polarized light serves to establish the compound character at once and also to show the :=Bx. relative orientation of the several parts. This is

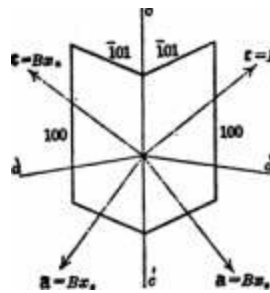
most distinct in the case of contact-twins, but is also well shown with penetration-twins, though here the parts are usually not separated by a sharp line.

Thus the exatninntioD of a section parallel to b (010) of a twin crystal of gypsum, ol the tym of Fig. 555, makes it easy not only to establish toe fact of the twinning but also to fix the relative positions of the ether-axes in the two parts. The measurement can in such cases be made between

the extinction-directions in the two halves, instead of between one of these and some definite crystallographic line, as the vertical axis.

The polygynthetic twiuning of ceilain species, as the tricllnlc feldspars, appears with.

great distinctness in poliirized light. For example, in the case of a section of alUte. parallel

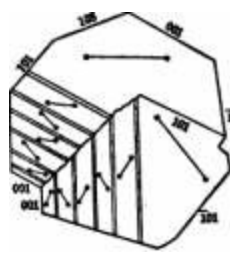


S&6.

to the basal cleavage, the alternate bands extinguish together sad assume the same tint wheu the quartz section U inserted. Hence the angle between these directions is easily menaured, and this Is obviously double the eitInc1ion'«ngle made with the edge b/e. A basal section of inicrocline Iu the same way shows its compound twinoing according to both tlie albite and pericline Inwa. ' iLe characteristic graliDg structure being t clearly revealed In polarized light. Fig. 556 of a section of r chondrodite (from Des Clolzeaux) abowa bow the compound [structure is shown by optical examination ; the position of the [sztat plane is indicated in the case of the successive polysyn-tbetic lamelle. The complex penetration-twins of right- and Isft-handed crystals of quartz (see the description of that species) also have their character strikingly revealed In polar Ized light.

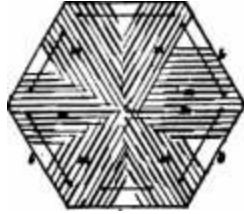
Still again, the true structure of complex multiple twins, exhibiting pseudo-eyminetty

' See a recent paper by Pockels, Jb. Mln., Beil.-Bd., 8,117. 1896.



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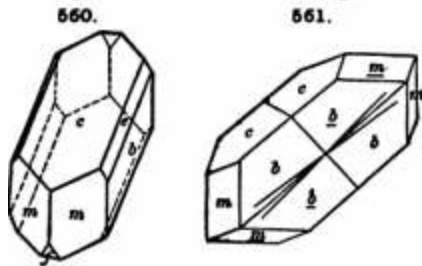
Id tliL'Jr exiernal form, can only be fully made out in this way. Tlila Is illiiastrated by Pie. A.17, a basal section of au appnraot bekaaf.'Otial pyramid of wllherlte (cf. Fig. 400. p. 1S8). Thu aualugoua Blx-aided pyramid of bromlite (Fig. 056) has a still more com['



Bromlite (E>ea Clolzeaux).

in polarized Itgbt. Other flluBtrations

rized light. Other illustrations are given in .



Art. 411. It will be understood that

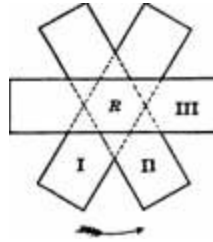
gMlbite (Laaauli). the Bzla) iaterference-flgurea of twla crystals, where the parts are superpoied, often show maoy peciillaritiea; the Airy spirals of quarlz (p. 208) will serve as an lilitgtrnilon.

408. A particularly interestlog caae, related to the subject digciisi-< i| in the preceding article, is that of the special propertiea of superpoeed cieavage-

663, sections of mica, as developed by Reusch.* It three or more of these, say of rectangular form, be superposed and so placed that the lines of the nxiial planes make equal angles of 60" (45°, etc.) with each other the effect is that

If polarized light which has passed through the center suffers circular polarization, with a rotation to right or left according to the way in which the sections are built up. The interference-figure resembles that of a section of quartz cut normal to the axis. This is illustrated in Fig. 563.

If the sections are numerous and very thin the imitation of the phenomena of quartz is closer. As shown by



theorie der Krystalstruktur, 1875.

; Bohncke, Abh., Erg.-Bd., 6, 16, 1876, and

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Bohncke and others, these facts throw much light upon the ultimate molecular structure of a crystallized medium showing circular polarization. Further, it is easy from this to understand how it is possible to have in sections of certain crystals {e.g., of clinoclone} portions which are biaxial and others that are uniaxial, the latter being due to an intimate twinning after this method of biaxial portions.

40V. Optical Properties of Crystals* Aggarwal.—The special optical phenomena of the different kinds of crystalline substances described on pp. 142, 148. and the extent to which their optical characters can be determined, depend upon the distinctness in the development of the individuals and their relative orientation. The case of ordinary granular, fibrous, or columnar aggregates needs no special discussion. Where, however, the doubly refracting minerals are extremely small, the microscope may hardly serve to do more than to show the aggregate polarization present.

A case of special interest is that of spherulites, which are spherical in form and radiated or concentric in structure; such aggregates occur with calcite, various chlorites, feldspars, etc. If they are formed of a doubly refracting crystalline mineral, or of an amorphous substance which has birefringent characters due to internal tension, they commonly exhibit a dark cross in the microscope between crossed nicols: further, this cross, as the section is revolved on the stage, though optically stationary, seems to rotate backward.*

A distinct and more special case is that of spherical aggregates of a mineral optically uniaxial (or biaxial with a small angle) Sections of these (not central) in parallel polarized light show more or less distinctly the interference-figure of a uniaxial crystal. The

objective must be focused on a point a little removed from the action itself, say on the surface of the sphere of which it is a part. In such cases the -|- or — character of the double refraction can be determined as usual.

410. Changes of Optical Character needed by Pressure.—As the difference between the optical phenomena exhibited by an isometric crystal on the one hand and a uniaxial or biaxial crystal on the other is referred to a difference in molecular structure modifying the properties of the ether, it would be inferred that if an amorphous substance were subjected to conditions tending to develop an analogous difference in its molecular structure it would also show doubly refracting properties.

This is found to be the case. Glass which has been suddenly cooled from a state of fusion, and which is therefore characterized by strong internal tension, usually shows marked double refraction. Further, glass plates subjected to great mechanical pressure in one direction show in polarized light more or less distinct interference-curves. Gemstone sections, also, under pressure exhibit like phenomena. Even the strain in a glass block developed under the influence of unlike charges of electricity of great difference of potential on its opposite sides is sufficient to make it doubly refracting.

In an analogous manner, as shown by Klein, Böcking, and others, the double refraction of a crystal may be changed by the application of mechanical force. Klein found that pressure exerted normal to the vertical axis of a section of a trigonal or hexagonal crystal which has been cut \perp to the c axis changes the uniaxial interference-figure into a biaxial, and with substances optically positive, the plane of the optic axes was parallel, and with negative substances normal, to the direction of pressure.

The quartz crystals in rocks, which have been subjected to great pressure, are often found to be in an abnormal state of tension, showing an undulatory extinction in polarized light. 411. Optical Anomalies. Since the early investigations of Brewster, Herschel, and others (1815 et seq.) it has been recognized that many crystals exhibit optical phenomena which are not in harmony with the apparent symmetry of their external form. Crystals of many isometric species, as analcite, alum, boracite, garnet, etc., often show more or less pronounced double refraction, and sometimes they are distinctly uniaxial or biaxial. A section examined in parallel polarized light may show more or less sharply

•Cf. Rosenbusch, *Monatsh. Phys.* 88:111-112, 1892.

Bertrand, *C. R.*, 94, 542, 1888; Mallard, *Bull. Soc. Min.*, 4, 67, 1881.

This subject has been discussed by various authors, among whom (in recent years) are the following: Klocke, *Jb. Min.* 2, 24, 1881; Böcking, *Zs. Kryst.*, 1, 58, 1888; Brauns, *Jb. Min.*, 1, 288, 1886; Klein, *Ber. Ak. Berlin*, 734, 1880; Pockels, *Wied. Ann.*, 37, 144 etc, 1889; 39, 440, 1890; *Jb. Hin., Bell.-Bd.*, 8, 317, 1698. See also literature on p. 581.

defined doubly refracting areas, or parallel bands or Umellee with varying extinction. Occasionally, as noted by Klein in the case of garnet, while most crystals are normally isotropic, others show optical characters which seem to be determined by the external bounding faces and edges; thus, a dodecahedron may appear to be made up of twelve rhombic pyramids (biaxial) whose apices are at the center; a hexoctahedron similarly may seem to be made up of forty-eight triangular pyramids, etc.

Similarly, crystals of many common tetragonal or hexagonal species, as vesuvianite, zircon, beryl, apatite, corundum, chabazite, etc., give interference-figures resembling those of biaxial crystals. Also, analogous contradictions between form and optical characters are noted with crystals of orthorhombic and monoclinic species, e.g., topaz, uatrolite, orthoclase, etc. All cases such as those mentioned are embraced under the common term of optical anomalies.

This subject has been minutely studied by many investigators in recent years (see literature), and important additions have been made to it both on the practical and the theoretical side. The result is that, though doubtful cases still remain, many of the typical ones have found a satisfactory explanation. No single theory, however, can be universally applied.

The chief question involved has been whether the anomalies are to be considered as secondary and non-essential, or whether they belong to the inherent molecular structure of the crystals in question. On the one hand, it has been urged that internal tension suffices (Art. 410) to call out double refraction in an isotropic substance or to give a uniaxial crystal the typical optical structure of a biaxial crystal. On the other hand, it is equally clear that twinning often produces pseudo-symmetry in external form, and at the same time conceals or changes the optical characters. From the simplest case, as that of aragonite (Fig. 397), we pass to more complex cases, as witherite (Figs. 400, 401, and 557), bromlite (Figs. 558, 559), phillipsite (Figs. 362, 433-424), which last is sometimes pseudo-isometric in form, though optical study shows the monoclinic character of the individuals.* Reasoning from the analogy of these last cases, Mallard was led (1876) to the theory that the optical anomalies could in most cases be explained by the assumption of a similar but still more intimate grouping of molecules which themselves without this would unite to form crystals of a lower grade of symmetry than that which their complex twinned crystals actually simulate.

In regard to the two points of view mentioned, it seems probable that internal tension (due to pressure, sudden cooling, or rapidity of growth, etc.) can be safely appealed to to explain the anomalous optical character of many species, as diamond, halite, beryl, quartz, etc. Again, it has been fully proved that the later growth of isomorphous layers of varying composition may produce optical anomalies, probably here also to be referred to tension. Alum is a striking example. The peculiarities of this species were early

investigated by Biot and made by him the basis of his theory of "lamellar polarization," but the present explanation is doubtless the true one. Fig. 564 (from Brauns) shows the appearance in polarized light of a section || o (ill) from a crystal in which the successive layers have different composition. Further, according to Brauna the optical peculiarities of many other species may be referred to this same cause. He includes here, particularly, those cases (as with some garnets) in which the optical characters seem to depend upon the external form, as

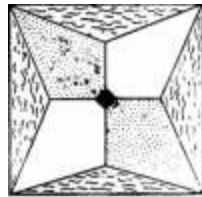
V called mimetic crj^tali

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PHYSICAL MINERALOGY.

noted above. Here belongs also apophyllite, a section of which (from Golden, Colo., by Klein) is shown in Fig. 565, The section has been cut || c {001} through the center of the crystal and is represented as it appears in parallel

polarized light.



Alum. I m.

Apophyllite || 001

Leucite, || 100.

Another quite distinct but most important class is that including species such as boracite and leucite which are dimorphous; that is, those species which at a certain elevation of temperature (300° for boracite and 500° to 600°

for leucite) become strictly isotropic. Under ordinary conditions, these species are anisotropic, but the fact stated makes it probable that originally their crystalline form and optical characters were in harmony. The relations for leucite deserve to be more minutely stated.

Leucite usually shows very feeble double refraction; $n_o \sim 1.508$, $n_e = 1.509$. The usual double refraction, early noted (Brewster, Biol), was originally explained. In 1873, Rati, on the basis of careful measurements, referred to the isometric crystal system, the tetragonal system, the trapezoidal face 113 being taken as 111, and 311, 121 to 431, 341, -respectively; also 101. Oil is 201, 031. Later Weibull (<]830), on the

same ground, made the orthorhombic; Mallard, however, referred them (1876, cited by
on optical grounds.

, to reappear again in the (110) ; in some cases a cubic form, the axial
Klein (Klein) to the later penetration of the

is usually or unequally developed; or there may be

position on cooling. Sections ordinarily show a bisectrix ($-|-$) is normal to what
corresponds to small. The structure corresponds in general crystals, in twinning; position
of the lamellae may be '■■' or '■■'

the fundamental individual lamellae

Still again, in a limited number of cases it can be shown that the inter-growth of lamellae
having slightly different crystallographic orientation is the cause of the optical
peculiarities. Prehnite is a conspicuous example of this class.

After all the various plausible explanations have been applied-there still remain, however,
many species about which no certain conclusion can be reached. This is true, for example,
of perovskite. To many of these species the theory of Mallard may probably be applicable.
Indeed it may be added that much difference of opinion still exists as to the cause of the
"optical anomalies" in a considerable number of cases.

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CHARACTERES DEPENDING UPON HEAT.

Optical Anomalies.*

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etc.

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1841; also full in *Hem. de l'Institut*, 18, 589.

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Launlx. Tridymite. Zs. Ery»t., 2. 268. 1878.

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IV. CHARACTERS DEPENDING UPON HEAT.

413. The more important of the special properties of & mineral species with respect to heat include the following: Fusibility; conductivity and expansion, especially in their relation to crystalline structure; change in optical characters with change of temperature; specific heat; also diathermancy, or the power of transmitting radiation, that is, ether-waves. The full discussion of these

" A complete bibliography is given in the memoir by Brauns (1891), see below.

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and other related subjects lies outside of the range of the present text-book. A few brief remarks are made upon them, and beyond these reference must be made to text-books on Physics and to special memoirs, some of which are mentioned in the literature (p. 233).

413, Fusibility.—The approximate relative fusibility of different minerals is an important character in distinguishing different species from one another by means of the blowpipe. For this purpose a scale is conveniently used for comparison, as explained in the articles later devoted to the blowpipe. Accuracy

. determinations of the fusibility are difficult, and though of little importance for the above object, they are interesting from a theoretical standpoint. They have been attempted by various authors, for example by Joly, who employed the "meldometer" for this end. This consisted of a strip of platinum in which the mineral in powder was inclosed; it was heated to the necessary point by an electrical current. He obtained the following values for the minerals used in von Kobell's scale (Art. 474): Stibnite, 525°; natrolite, 965°; almandite, 1296°; actinolite, 1296°; orthoclase, 1175°; bronzite, 1300°; also for quartz, 1430°.

414, Conductivity.—The conducting power of different crystallized media was early investigated by Senarmont. He covered the faces of the substance under investigation with wax and observed the form of the figure melted by a hot wire placed in contact with the surface at its middle point. Later investigations have been made by Routgen (who modified the method of Senarmont), by Jannettaz, and others. In general it is found that, as regards their thermal conductivity, crystals are to be divided into the three classes noted on p. 192. In other words, the conductivity for heat seems to follow the same general laws as the propagation of light. It is to be stated, however, that experiments by B. P. Thompson and O. J. Lodge have shown a different rate of conductivity in tourmaline in the opposite directions of the vertical axis.

415, Expansion.—Expansion, that is, increase in volume upon rise of temperature, is a

nearly universal property for all solids. The increment of volume for the unit volume in passing from 0° to 1° C. is called the coefficient of expansion. This quantity has been determined for a number of species. Further, the relative expansion in different directions is found to obey the same laws as the light-propagation. Crystals, as regards heat-expansion, are thus divided into the same three classes mentioned on p. 192 and referred to in the preceding article.

The amount of expansion varies widely, and, as shown by Jannettaz, is influenced particularly by the cleavage. Mitscherlich found that in calcite there was a diminution of $8' 37''$ in the angle of the rhombohedron on passing from 0° to 100° C, the form thus approaching that of a cube as the temperature increased. The rhombohedron of dolomite, for the same range of temperature, diminishes $4' 46''$; and in aragonite, for a rise in temperature from 21° to 100° , the angle of the prism diminishes $2' 46''$. In some rhombohedrons, as of calcite, the vertical axis is lengthened (and the lateral shortened), while in others, like quartz, the reverse is true. The variation is such in both cases that the birefringence is diminished with the increase of temperature, for calcite possesses negative double refraction, and quartz, positive.

It is to be noted that in general the expansion by heat, while it may serve to alter the angles of crystals, other than those of the isometric system, does not alter the zone-relations and the crystalline symmetry. In certain cases, however, the effect of heat may be to give rise to twinning (as in anhydrite) or to cause their disappearance (as in calcite). K^{ely} lieat servei

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to develop a new molecular structure; thus, as explained in Art. 411, bonioit and leucite, which are anisotropic at ordinary temperatures, become isotropic when heated, the former to 3UU~ the latter to SOU or tloo". The change in the optical properties of crystals produced by heat has already been noticed (Art. 404).

416. Specific Heat.—Determinations of the specific heat of many minerals have been made by Joly, by Oeberg, and others. Some of the results reached are, as follows :

417. Diathermancy.—Besides the slow molecular propagation of heat in a body, measured by its thermal conductivity, there is also to be considered the rapid propagation of what is called radiant heat through it by the wave-motion of the ether which surrounds its molecules. This is merely a part of the general subject of light-propagation already fully discussed, since heatwaves, in the restricted sense, differ from light-waves only in their relatively greater length. The degree of absorption exerted by the body is measured by its diathermancy, which corresponds to transparency in light. In this sense halite, sylvite, and fluorite are highly diathermanous, since they absorb but little of the heat-waves passing through them; on the other hand, selenite and, still more, alum are comparatively diathermanous, since while transparent to the slowest light-waves they absorb

the long heat-waves, transforming the energy into (most of) sensible heat. Measurements of the diathermancy were early made by Melloni, later by Tyndall, Langley, and others.

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PHYSICAL UINBRA.LOCIT.

V. CHARACTERS DEPENDING UPON ELECTRICITY AND MAGNETISM.

1. ELECTRICITY.

418. Electrical Conductivity. —The subject of the relative conducting power of different minerals is one of minor interest,* In general most minerals, except those having a metallic luster among the sulphides and oxides, are nonconductors. Only the DOnc-
conductofB can show pyro-electrical phenomena, and only the conductors can give a thermo-electric current.

419. Frictional Electricity. ^The development of an electrical charge on many bodies bi/
frictioti is a familiar subject. All minerals become electric by friction, although the degree to which this is manifested differs widely. There is no line of distinction among minerals, dividing them mtopositively electric and negatively (Aacina; for both electrical states may be presented by different varieties of tlie same species, and by the same variety in different states. The gems are in genera! positively electric only when polished; the diamond, however, exhibits positive electricity whether polished or not. It is a familiar fact that the electrification of amber upon friction was early observed (600 B, C), and indeed the Greek name (t/Xeierpov) later gave rise to the word electricity.

420. Pyro-electricity. ^The simultaneous development of plus and minus charges of electricity on different parts of the same crystal when its tempera-tnre is suitably changed is called pyro-electricity. Crystals exhibiting such phenomena are said to be pyro-electric. This phenomenon was first observed in the case of tourmaline, which is rhombohedral-hemimorphic in crystallization, and it is particularly marked with crystals belonging to groups of relatively low symmetry, especially those of the hemimorphic type. It is possible, of course, only witi non-conductors. This subject was early investigated by Eieess and Rose (1843), later by Ilaukel, also by C. Friedel, Kundt, and others (see literature).

In all cases it is true that directions of like crystallogrsptic symmetry show charges of like sign, while unlike directions may exhibit opposite charges. Substances not crystallized cannot show pyro-electricity. A few of the many possible examples will serve to bring out the most essential points.

Boracite (isometric-tetrahedral, p. 46) on heating exhibits + electricity on one set of tetrahedral faces and ~ electricity on the other. Cf. Fig. 567.

Tourmaline (rhoniboJiedral-hemimorphic, p. ^9) shows opposite charges at the opposite extremities of the vertical axis corresponding to its hemimorphic crystallization. In this

and in other similar cases, the extremity which becomes positive on heating has been called the analogous pole, and that which becomes negative has been called the antilogous pole.

Calamine and struvite (orthorhombic-hemimorphic, p. 95) exhibit phenomena analogous to these of tourmaline.

Quartz (rhombohedral-trapezohedral, p. 82) shows + electricity on heating at the three alternate prismatic edges and - electricity at the three remaining edges; the distribution for right-handed crystals is opposite to that of left-handed. Twins may exhibit a high degree of complexity. Cf. Figs. 568, 569.

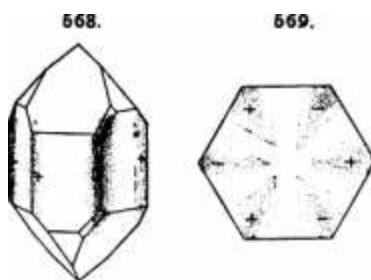
Axinite (triclinic, p. 107), when heated to 130° or 130° , has an analogous pole (Ries & Rose) at the solid angle rxM'' ; the antilogous pole at the angle $mr'M'$ near plane n .

A very convenient and simple method for investigating the phenomena is

- On the conductivity of minerals, see Beljers. *Jl*. *Hiii*. *Bcil*.-*Iid*. 11, 403, 1898.

CHARACTERS DEPENDENT UPON ELECTRICITY AND TEMPERATURE. 235

the following, which is due to Kundt: First heat the crystal or section carefully in an air-bath; pass it several times through the flame of an alcohol lamp and then place it on a little upright cylinder of brass to cool. While cooling, a mixture of red lead and sulphur finely pulverized and previously agitated is dusted over it through a fine cloth from a suitable bellows. The



positively electrified red lead collects on the parts having a negative charge, and the negatively electrified sulphur on those with a positive charge. This is illustrated by Figs. 568-569, and still better by the illustrations given by Kundt and others. (Of. Plate III of Orst, *Phys. Kryst.*, 1895.)

431, Piezo-electricity. —The name piezo-electricity has been given to the development of electrical charges on a crystallized body by pressure. This is shown by a cleavage-fragment of calcite, also by topaz. This phenomenon is most interesting where a relation can be established between the electrical excitement and the molecular structure, as is conspicuously true with quartz, tourmaline, and some other species.

This subject has been investigated by Hankel, Curie, and others, and discussed theoretically by Lord Kelvin (see literature). Hankel has also employed the term *actinohlectricity*, or, better, *photo-electricity*, for the phenomenon of calling out of an electrical condition by the influence of direct radiation ; fluorite is a conspicuous example.

422. *Actinohlectricity in Mineralogy.*— The power of different minerals to transmit the so-called X-rays, or Rontgen-rays, emitted from a suitable vacuum-tube during the discharge of an induction-coil has been investigated by Doelter.* He has found, for example, that sulphur, beryl, epidote, pyrite, etc., are nearly opaque; tourmaline less so; fluorite transmits the rays slightly, the feldspars and quartz better; corundum is nearly transparent and diamond and graphite are highly so. Diamond is easily distinguished in this way from its imitations, which are relatively highly opaque.

423. *Thermo-electricity.* —The contact of two unlike metals in general results in electrifying one of them positively and the other negatively. If, further, the point of contact be heated while the other parts, connected with a wire, are kept cool, a continuous current of electricity—shown, for example, by a suitable galvanometer—is set up at the expense of the heat-energy supplied. If, on the other hand, the point of junction is cooled, a current is set up in the reverse direction. This phenomenon is called *thermo-electricity*, and two metals so connected constitute a *thermo-electric couple*. Further it is found that different conductors can be arranged in order in a table—a so-

*Jb. Min., 2, 67, ISH; 1, SM, 1897. Also Quodvius, Nature, April 30, 1896.

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called *thermo-electric series*—according to the direction of the current set up on heating and according to the electromotive force of this current. Arsenic; the metals, bismuth (—) and antimony (—) stand at the opposite ends of the series; the current passes through the connecting wire from antimony to bismuth.

This subject is so far important for mineralogy, as it was shown by BiinBei that the natural metallic sulphides stand farther off in the series than bismuth and antimony, and consequently by them a higher electromotive force is

produced. The thermo-electrical relations of a large number of minerals were determined by Flight. >

It was early observed that some minerals have varieties which are both + and —. Bose attempted to establish a relation between the plus and minus pyritohedral forms of pyrite and cobaltite, and the positive or negative thermo-electrical character. Later investigations by Chruoff and Dana have shown, however, that the same peculiarity belongs also to glaucodot, tetradymite, Bkutterudite, danaite, and other minerals, and it is demonstrated by them that it cannot be dependent upon crystalline form, but rather upon

chemical composition.

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2. Magnetism.

424. magnetic Minerals. Natural Magnetite.—A few minerals in their natural state are capable of being attracted by a strong steel magnet; they are said to be metallic. This is conspicuously true of magnetite, the magnetic oxide of iron; also of pyrrhotite or magnetic pyrites, and of some varieties of native platinum (especially the variety called iron-platinum).

A number of other minerals, as hematite, franklinite, etc., are in some cases attracted by a steel magnet, but probably in most if not all cases because

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CHARACTERISTICS UPON ELECTRICITY AND MAGNETISM. 337

of admixed magnetite (but see Art. 428). Occasional varieties of the three minerals mentioned above, as the lodestone variety of magnetite, exhibit themselves the attracting power and polarity of a true magnet. They are then called natural magnets. In such cases the magnetic polarity has probably been derived from the inductive action of the earth, which is itself a huge magnet.

426. Paramagnetism. Diamagnetism.—In a very strong magnetic field, as that between the poles of a very powerful electromagnet, all minerals, as indeed all other substances, are influenced by the magnetic force. According to their behavior they are divided into two classes, the paramagnetic and diamagnetic; those of the former appear to be attracted, those of the latter to be repelled. For purposes of experiment the substance in question, in the form of a rod, is suspended on a horizontal axis between the poles of the magnet. If paramagnetic, it takes a position parallel to the magnetic axis; if diamagnetic, it sets transversely to it. Iron, cobalt, nickel, manganese, platinum are paramagnetic; silver, copper, bismuth are diamagnetic. Among minerals compounds of iron are paramagnetic, as siderite, also diopside; further beryl, diopside. Diamagnetic species include calcite, zircon, wulfenite, etc.

By the use of a sphere it is possible to determine the relative amount of magnetic induction in different directions of the same substance. Experiment has shown that in isometric crystals the magnetic induction is alike in all directions; in those optically uniaxial, that there is a direction of maximum and, normal to it, one of minimum magnetic induction; in biaxial crystals, that there are three unequal magnetic axes, the position of which may be determined. In other words, the magnetic relations of the three classes of crystals are analogous to their optical relations,

426. Corresponding to the facts just stated, that all compounds of iron are paramagnetic, it is found that a sufficiently powerful electromagnet attracts all minerals containing iron, though except in the cases given in Art. 424 a bar magnet has no sensible influence upon them; hence the efficiency of the electromagnetic method of separating ores,

Pliicker* determined the magnetic attraction of a number of substances compared with iron taken as 100,000. For example, for magnetite he obtained 40,237; for hematite, crystallized, 633, massive, 134; limonite, 71; pyrite, 15U.

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PHYSICAL CHEMISTRY.

VI. TASTE AND ODOR.

In their action upon the sense a few minerals possess taste, and others under some circumstances give off odor.

487. Taste belongs only to soluble minerals. The different kinds of taste adopted for reference are as follows:

1. Astringent: the taste of Tannin.

2. Sweetish astringent: taste of alum.
3. Saline: taste of common salt.
4. Alkaline: taste of soda.
5. Cooling: taste of saltpeter.
6. Bitter: taste of Epaom salts.
7. Sour: taste of sulphuric acid.

428. Odor.—Excepting a few gaseous and soluble species, minerals in the dry unchanged state do not give off odor. By friction, moistening with the breath, and the elimination of some volatile ingredient by heat or acids, odors are sometimes obtained which are thus designated:

1. Alliaceous: the odor of garlic. Friction of arsenical iron elicits this odor; it may also be obtained from arsenical compounds by means of heat.
2. Horse-radish odor: the odor of dehydrating horse-radish. This odor is strongly perceived when the ores of selenium are heated.
3. Sulphurous: friction elicits this odor from pyrite, and heat from many sulphides.
4. Bituminous: the odor of bitumen.
5. Fetid: the odor of sulphureted hydrogen or rotten eggs. It is elicited by friction from some varieties of quartz and limestone.
6. Argillaceous: the odor of moistened clay. It is obtained from serpentine and some allied minerals, after moistening them with the breath; others.

. pyrargillite, afford it when heated. 429. Feel ""

. Feel—The feel is a character which is occasionally of some importance; it is said to be smooth (sepiolite), rough [talc), hard, or meager, etc. Some minerals, in consequence of their hygroscopic character, adhere to the tongue when brought in contact with it.

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PAET m. CHEMICAL MINERALOGY.

(GENERAL PRINCIPLES OF CHEMISTRY AS APPLIED TO MINERALS.

430. Minerals, as regards their chemical constitution, are either the combined

elements in their native state, or definite compounds of these elements formed in accordance with chemical laws. It is the object of Chemical Mineralogy to determine the chemical composition of each species; to show the chemical relations of different species to each other where such exist; and also to explain the methods of distinguishing different minerals by chemical means. It thus embraces the most important part of Determinative Mineralogy.

In order to understand the chemical constitution of minerals, some knowledge of the fundamental principles of Chemical Philosophy is required; and these are here briefly recapitulated.

431. Chemical Elements.—Chemistry recognizes about seventy substances which cannot be decomposed, or divided into others, by any process of analysis at present known; these substances are called the chemical elements. A list of them is given in a later article (436); common examples are: Oxygen, nitrogen, hydrogen, chlorine, gold, silver, sodium, etc.

433. Atom. Molecule.—The study of the chemical properties of substances and of the laws governing their formation has led to the belief that there is for each element a definite, indivisible mass, which is the smallest particle which can play a part in chemical reactions; this indivisible unit is called the atom.

With some rare exceptions, the atom cannot exist alone, but unites by the action of what is called chemical force, or chemical affinity, with other atoms of the same or different kind to form the molecule. The molecule, in the chemical sense, may be defined as the smallest particle into which a given kind of substance can be subdivided without undergoing chemical decomposition. For example, two atoms of hydrogen unite to form a molecule of hydrogen gas. Again, one atom of hydrogen and one of chlorine form a molecule of hydrochloric acid gas; two atoms of hydrogen and one of sulphur form a molecule of the gas hydrogen sulphide.

483. Physical Molecules.—An important distinction must be made between the simple chemical molecules, regarded as made up of the smallest possible number of the atoms of each kind, united in the given proportion, and the actual physical molecules which together build up the structure of a particular

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mass of matter. These physical molecules may be much more complex, each being made up of a number of chemical molecules but necessarily containing the respective atoms in the same proportion. In the case of a gas, it is always possible to determine the constitution of the molecule, as is explained later, but in the case of liquids and solids this is in general impossible.

For example, it is certain that a molecule of water vapor consists of two atoms of

hydrogen and one atom of oxygen, but the molecules of water which unite to build up a snow crystal, though containing the atoms of the two elements hydrogen and oxygen in the same proportion, may be highly complex, as if made up of many gas molecules. Since it is in general impossible in the case of solids to fix the constitution of the actual molecule, it is usually better to regard it as a chemical molecule of the simplest possible form.

434. Atomic Weight.—The atomic weight of an element is the weight, or, better expressed, the mass of its atom compared with that of the element hydrogen taken as the unit. Thus the mass of an atom of oxygen is very nearly sixteen times that of the atom of hydrogen (exactly 15.96), and hence this number is called the atomic weight of oxygen. Of the methods by which the relation between the masses of the atoms is determined it is unnecessary here to speak; the results that have been obtained are given in the table on p. 241.

435. Symbol Formula.—The symbol of an element is the initial letter, or letters, often of its Latin name, by which it is represented when expressing in chemical notation the constitution of substances into the composition of which it enters. Thus O is the symbol of oxygen, H of hydrogen, Cl of chlorine, Fe (from ferrum) of iron, Ag (from argentum) of silver, etc. Further, this symbol is always understood to indicate that definite amount of the given element expressed by its atomic weight; in other words, it represents one atom. If twice this quantity is involved, that is, two atoms, this is indicated by a small subscript number written immediately after the symbol. Thus, Sb_2S_3 , means a compound consisting of two atoms of antimony and three of sulphur, or of 2 X 120 parts by weight of antimony and 3 X 32 of sulphur.

This expression, Sb_2S_3 , is called the formula of the given compound, since it expresses in briefest form its composition. Similarly the formula of the mineral albite is $NaAlSi_3O_8$.

Strictly speaking, such formulas are merely empirical formulas, since they express only the actual result of analysis, as giving the relative number of atoms of each element present, and make no attempt to represent the actual constitution. A formula developed with the latter object in view is called a rational, structural, or constitutional formula (see Art. 493).

436. Table of the Elements.—The following table gives a list of all the definitely established elements with their accepted symbols and also their atomic weights.*

Of the elements given in this list—about seventy in all—only a very small number, say twelve, play an important part in making up the crust of the earth and the water and air surrounding it. The common elements concerned in the composition of minerals are: Oxygen, sulphur, silicon, aluminium, iron, calcium, magnesium, sodium, potassium. Besides these, hydrogen is present in water, nitrogen in the air, and carbon in all animal and vegetable substances. Only a very few of the elements occur as such in nature, as

native gold, native silver, native sulphur, etc.

* These correspond in value to those commonly accepted, and are given accurate to one decimal place. In strict chemical literature the atomic weight of oxygen is 16, etc.

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GENERAL PRINCIPLES OF APPLIED TO METALS. '241

Of the elements, oxygen, hydrogen, nitrogen, chlorine, and fluorine are gases; bromine is a volatile liquid; mercury is also a liquid, but the others are solids under ordinary conditions.

Symbol Alum in turn. Alum is $Al_2(SO_4)_3$ Sh Argou A

(.Weight. , Byml:

37 ' Manganese Mn

130 ' Mercury (Hydrargyrum) Hg

96 ' Molybdenum Mo

ibol. At. Weight.

-Ba Be (or Gl) 91

Beryllium Be

J - Bismuth

' Bismuth Bi

Bromine Br 79-8

Calcium Ca 40-1

Cesium Cs 132-9

Calcium Ca 40-1 ■ « » »

Carbon C 12

Cerium Ce 141

CLlortae CI - 85*4

Clironium ^""Cr 52'5

CobfiU -Co 58-7

^ ColuDibluni.see JRoMum. ^ 69, t

k \A Copper (Ouprwwi) --Ca " -< o %

Diciymium Di 142

Erbium Er 160

Fluorine ^ f 191

ORtlum Ga 69-9

tieroiftDfuin Oe 78'8 Oliicinum, see Berj/IUvm.

Gold {Aumra) —An 194-7

Hdium He 4-4

Hvdrogen H 1

Imlium In IIS'4

Iodine I 130-6

^ Iridium Ir 192-5

J t Iron (FoTvm) --Fo 65-9

Untbaiium La 138

-Pb im^

_it 7

-wMg 24

- .KMolybde 74'9 - / i Nickel "" Niobium

I OiyRen ' Palladium

Phosphorus

Platinum

Polosalum (Saiiuit)

Rbndlum

Rubidium

Rutbenium

Scandium

Selenium

Silicon

Silver (Argentum)

Sodium (Xfalrium)

Strouium

Sulphur

Tantalum

Tellurium

Thallium

Tburium

Tin (Stannum)

lltauium

Tungsten (Woffiramiw

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Viinadium

Ytl erbium

Yttrium

Z{nc

Zirconium

1D4-I 85 9 108-e

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437. Metals and Non-metals. —The elements may be divided into two more or less distinct classes, the metals and the non-metals. Between the two lie a number of elements sometimes called the semi-metals. The metals, as gold, silver, iron, sodium, are those elements which, as described, possess to a more or less perfect degree the fundamental characters of the ideal metal, viz.: malleability, metallic luster (and opacity to light), conductivity for heat and electricity; moreover, chemically described, they commonly play the part of the positive or basic element in a simple compound, as later defined (Arts. 446-449). The non-metals, as sulphur, carbon, silicon, etc., also the gases, as oxygen, chlorine, etc., have none of the physical characters alluded to; they are, if solids, brittle, often transparent to light-radiation, are poor conductors for heat and electricity. Chemically expressed, they usually play the negative or acid part in a simple compound.

The so-called semi-metals, or metalloids, include certain elements, as tellurium, arsenic, antimony, bismuth, which have the physical characters of a metal to a less perfect degree {e.g., they are more or less brittle}; and, more important than this, they often play the part of the acidic element in the compound into which they enter. These points are illustrated later.

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It is to be understood that the distinctions between the classes of the elements named cannot be very sharply applied. Thus the typical metallic characters mentioned are possessed to a very unequal degree by the different substances classed as metals; for example, by silver and tin. Corresponding to this a number of the true metals, as tin and manganese, play the part of an acid in numerous salts. Further, the mineral magnetite, Fe_3O_4 , is often described as an iron ferrate; so that in this compound the same element would play the part of both acid and base.

438. Positive and Negative Element!.—It is common to make a distinction between the electro-positive and electro-negative element in a compound. The passage of a sufficiently strong electrical current through a chemical compound in many cases results in its decomposition (or electrolysis) into its elements or parts. In such cases it is found that for each compound the atoms of one element collect at the negative pole (the cathode) and those of the other at the positive pole (the anode). The former is called the electro-

positive element and the latter the electro-negative element. Thus in the electrolysis of water (H₂O) the hydrogen collects at the cathode and is hence called positive, and the oxygen at the anode and is called negative. Similarly, in hydrochloric acid (HCl) the hydrogen is thus shown to be positive, the chlorine negative. This distinction is also carried to complex compounds, as copper sulphate (CuSO₄), which by electrolysis is broken into Cu, which is found to be electro-positive, and SO₄, (the last separates into SO₂, forming H₂SO₄, and free oxygen).

For reasons which will be explained later, the positive element is said to play the basic part, the negative the acidic. The metals, as already stated, in most cases belong to the former class, the non-metals to the latter, while the semi-metals may play both parts.

It is common in writing the formula to put the positive or basic element first, thus H₂O, H₂S, HCl, H₂SO₄, FeS, As₂O₃, AlH₃, NiSb, FeAs₂. Here it will be noted that antimony (Sb) and arsenic (As) are positive in some of the compounds named but negative in the others.

439. Periodic Law.—In order to understand the relations of the chief classes of chemical compounds represented among minerals, as still more their further subdivision, down finally to the many isomorphous groups —groups of species having analogous composition and closely similar form, as explained in Art. 456—the fundamental relations and grouping of the elements must be understood, especially as developed of recent years and shown in the so-called Periodic Law."

Although the subject can be only briefly touched upon, it will be useful to give here the general distribution of the elements into Groups and Series, as presented in the Principles of Chemistry (Engl. Ed., 1891) of D. Mendeleeff, to whom is due more than any one else the development of the Periodic Law. A few remarks are added on the grouping of the elements as illustrated by mineral compounds; artificial compounds show these relations still more fully and clearly. For the thorough explanation of this subject, more particularly as regards the periodic or progressive relation between the atomic weights and various properties of the elements, the reader is referred to the work above mentioned or to one of the many other excellent modern text-books of chemistry.

The relations of some of the elements of the first group are exhibited by the isomorphism (see Art. 456, also the description of the various groups and species here referred to, which are given in Part IV of this work) of NaCl, KCl, AgCl; or again of LiMnPO₄, and NaMnPO₄, etc. In the second group,

and if the validity of the Periodic Law

THE PERIODIC LAW OF CHEMISTRY IS APPLIED TO MINERALS. 243

reference may be made to the isomorphism of the carbonates and sulphates (p. 250) of calcium, barium, and strontium; while among the sulphides, ZnS, CaS, and HgS are

doubly related. In the third group, we find boron and aluminium often replacing one another among silicates. In the fourth group, the relations of silicon and titanium are shown in the titanate-silicates, while the compounds TiO_2 , SnO_2 , PbO , (and MnO), $ZrSiO_4$ and $ThSiO_4$, have closely similar form. In the fifth group, many compounds of arsenic, antimony, and bismuth are isomorphous among metallic compounds, while the relations of phosphorus, vanadium, arsenic, also antimony, are shown among the phosphates, vanadates, arsenates, and antimonates; again the mutual relations of the niobates and tantalates are to be noted.

In the sixth group, the strongly acidic elements, sulphur, selenium, tellurium, are all closely related, as seen in many sulphides, selenides, tellurides; further, the relations of sulphur and chromium, and similarly of both of these to molybdenum and tungsten, are shown among many artificial sulphates, chromates, molybdates, and tungstates.

In the seventh group the relations of the halogens are too well understood to need special remark. In the eighth group, we have Fe, Co, Ni alloyed in meteoric iron, and their phosphates and sulphates are in several cases closely isomorphous; further, the relation of the iron series to that of the platinum series is exhibited in the isomorphism of FeS_2 , $BiAsS_3$, $FeAs_2$, etc., with $PtAs_2$, and probably RuS_2 .

440. Combining Weight.—Chemical investigation proves that the mass of a given element entering into a compound is always proportional either to its atomic weight or to some simple multiple of this; the atomic weight is hence also called the combining weight. Thus in rock salt, sodium chloride, the masses involved of sodium and chlorine present are found by analysis to be equal to 39.4 and 60.6 in 100 parts, and these numbers are in proportion to 33 : 35.4, the atomic weights of sodium and chlorine; hence it is concluded

that one atom of each is present in the compound. The formula is, therefore, $NaCl$. In calcium chloride, by the same method the masses present are found to be proportional to 39.9 : 70.9, that is, to 39.9 = 3 X 35.4; hence the formula is $CaCl_2$.

Still again, a series of compounds of nitrogen with oxygen is known in which the ratios of the masses of the two elements are as follows: (1) 28:16, (2) 14:16, (3) 28:48, (4) 14:32, (5) 28:80. It is seen at once that these must have the formulas (1) N_2O , (2) NO , (3) N_2O_3 , (4) NO_2 , (5) N_2O_5 . On the contrary, atmospheric air which constitutes about the mass of 76.8 to 282 parts of a chemical compound of nitrogen and oxygen since (aside from other considerations) these numbers are not in the ratio of $n \times 14 : m$: It is where n and m are simple numbers.

441. Molecular Weight.—The molecular weight is the weight of the molecule of the given substance, expressed in terms of the mass of the hydrogen atom as unit. The molecular weight of hydrogen is 2 because the molecule can be shown to consist of two atoms. The molecular weight of hydrochloric acid (HCl) is 36.4, of water vapor (H_2O) it is 18, of hydrogen sulphide (H_2S) it is 34.

Since, according to the law of Avogadro, like volumes of different gases under like conditions as to temperature and pressure contain the same number of molecules, it is obvious that the molecular weight of substances in the form of gas can be derived directly from the relative density or specific gravity. If the density is referred to hydrogen, whose molecular weight is 2, it will be always true that the molecular weight is twice the density in the state of a gas and vice versa. Thus the observed density of carbon dioxide (CO₂) is 22, hence its molecular weight must be 44. It is this principle that makes it possible in the case of a gas to fix the constitution of the molecule when the ratio in number of the atoms entering into it has been determined by analysis. In the case of Solids, where the constitution of the molecule in general cannot be fixed, it is best, as already stated, to write the molecular formula in its simplest form, as NaAlSi₃O₈, for albite. The sum of the weights of the atoms present is then taken as the molecular weight.

442. Valence. —The valence of an element is given by the number of its atoms which are required to unite with one unit atom, as of hydrogen or chlorine. Thus, using the examples of Art. 440, in NaCl, since one atom of sodium unites with one of chlorine, its valence is one; or in other words, it is said to be univalent, and is called a monad. Further, calcium (as in CaCl₂), also barium, etc., are bivalent; gold is (usually) trivalent; tin is tetravalent, etc. The valence may be expressed by the number of bonds by which one element in a compound is united to another, thus:

Na-Cl, Ba=Cl₂, Au=Cl₃, Sn=Cl₄, etc

The valences of the common elements, expressed by their symbols, are as follows:

Univalent: H, Cl, Br, I, F; Li, Na, K, Rb, Cs, Ag.

Bivalent: O, S, Se, Te; Be, Mg, Ca, Sr, Ba, Pb, Hg, Cu, Zn, Co, Ni.

Trivalent: B, Al; probably also Fe, Mn, Cr.

Tetravalent: C, Si, Ti, Zr, Sn.

Pentavalent: N, P, As, Sb, V, Bi, Nb, Ta.

The above list, though convenient for reference, is not to be taken as complete or final. A considerable number of the elements show a different valence in different compounds. Thus both SnO₂ and SnO are known; also FeS, Fe₂O₃, and FeO; Cu₂Cl₂, CuCl₂, and similarly Cu₂(SO₄) and CuSO₄,

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etc. In certain cases the composition of two compounds of the same elements may be made consistent with each other, by an assumption as to the possible grouping of the atoms. Thus in cupric chloride, CuCl₂, or cupric oxide, CuO, copper is bivalent as usual.

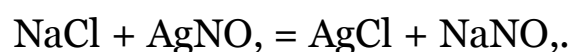
But the cuprous compounds, Cu₂O, and Cu₂Cl₂, also occur, and for them the formulas may be written



Again, the elements Al, Fe, Mn, Cr, which form the compounds AlCl₃, Al₂O₃, etc, are sometimes called tetravalent and the formula of the oxide written, for example, (Al = Al) = O₂.

443. Chemical Reaction.—When solutions of two chemical substances are brought together, in many cases they react upon each other with the result of forming new compounds out of the elements present; this phenomenon is called a chemical reaction. One of the original substances may be a gas, and in many cases similar results are obtained from a liquid and a solid, or even often from two solids.

For example, solutions of sodium chloride (NaCl) and silver nitrate (AgNO₃) react on each other and yield silver chloride (AgCl) and sodium nitrate (NaNO₃). This is expressed in chemical language as follows:



This is a chemical equation, the sign of equality meaning that equal weights are involved both before and after the reaction.

Again, hydrochloric acid (HCl) and calcium carbonate (CaCO₃) yield calcium chloride (CaCl₂) and carbonic acid (H₂CO₃); which last breaks up into water (H₂O) and carbon dioxide (CO₂), the last going off as a gas with effervescence. Hence



444. Radicals.—A compound of two or more elements according to their relative valence in which all their bonds are satisfied is said to be saturated. This is true of H₂O, or, as it may be written, H—O—H. If, however, one or more bonds is left unsatisfied, the resulting combination of elements is called a radical. Thus —O—H, called briefly hydroxyl, is a common radical, having a valence of one, or, in other words, univalent; NH₂ is again a univalent radical; so, too, (CaF), (MgF) or (AlO). Radicals often enter into a compound like a simple element; for example, in ammonium chloride, NH₄Cl, the univalent radical NH₄ plays the same part as the univalent element Na in NaCl. In the chemical composition of mineral species, the commonest radical is hydroxyl (—O—H) already defined. Other examples are (CaF) in apatite (see Art, 488), (MgF) in wagnerite, (AlO) in many basic silicates, etc.

445. Chemical Compound.—A chemical compound is a combination of two or more elements united by the force of chemical attraction. It is always true of it, as before stated (Art. 440), that the elements present are combined in the proportion of their atomic

weights or some simple multiples of these. A substance which does not satisfy this condition is not a compound, but only a mechanical mixture.

Examples of the simpler class of compounds are afforded by the oxides, or compounds of oxygen with another element. Thus, among minerals we have Cu_2O , cuprous oxide (cuprite); ZnO , zinc oxide (zincite); Al_2O_3 , alumina

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(corundum); SnO_2 , tin dioxide (stannic oxide); SiO_2 , silicon dioxide (quartz); As_2O_3 , arsenic trioxide (arsenolite).

Another simple class of compounds are the sulphides (with the selenides, tellurides, arsenides, antimonides, etc.), compounds in which sulphur (selenium, tellurium, arsenic, antimony, etc.) plays the same part as oxygen in the oxides. Here belong CuS , cuprous sulphide (chalcocite); ZnS , zinc sulphide (sphalerite); PbTe , lead telluride (altaite); FeS_2 , iron disulphide (pyrite); Sb_2S_3 , antimony trisulphide (stibnite).

446. Acids.—The more complex chemical compounds, the understanding of which is needed in a study of minerals, are classed as acids, bases, and salts; the distinctions between them are important.

An acid is a compound of hydrogen, or hydroxyl, with a non-metallic element (as chlorine, sulphur, nitrogen, phosphorus, etc.), or a radical containing these elements. In them the hydrogen atoms may be replaced by metallic atoms; the result being then the formation of a salt (see Art. 448). Acids in general turn blue litmus paper red and have a sharp, sour taste. The following are familiar examples:

HCl , hydrochloric acid, $\text{H} - \text{Cl}$.

HNO_3 , nitric acid, $(\text{HO}) = \text{NO}_2$.

H_2CO_3 , carbonic acid, $(\text{HO})_2 = \text{CO}$. H_2SO_4 , sulphuric acid, $(\text{HO})_2 = \text{SO}_3$. H_4SiO_4 , orthosilicic acid, $(\text{HO})_4 = \text{SiO}_2$. H_3SiO_3 , metasilicic acid, $(\text{HO})_3 = \text{SiO}_2$. H_2SiO_2 , pyrosilicic acid, $(\text{HO})_2 = \text{SiO}_2$.

The full explanation of the constitution of the different acids requires a more detailed discussion than is possible here. The second series of formulas given above must serve as suggestions in this direction.

It is to be noted that with a given acid element several acids are possible. Thus normal, or orthosilicic, acid is H_4SiO_4 , in which the bonds of the element silicon are all satisfied by the hydroxyl (HO). But the removal of one molecule of water, H_2O , from this gives the formula H_3SiO_3 , or metasilicic acid.

Acids which, like HNO_3 , contain one atom of hydrogen that may be replaced by a metallic atom {e.g., in KNO_3 } are called monobasic. If, as in H_2CO_3 , and H_2SO_4 , there are two atoms (e.g., in CaCO_3 , BaSO_4) the acids are dibasic. Similarly H_3PO_4 is tribasic, etc.

Most acids are liquids (or gases), and hence acids are represented very sparingly among minerals; $\text{B}(\text{OH})_3$, boric acid (sassolite), is an illustration.

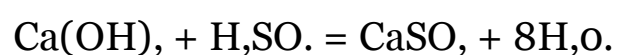
447. Bases.—The bases, or hydroxides as they are also called, are compounds which may be regarded as formed of a metallic element (or radical) and the univalent radical hydroxyl, $-\text{OH}$; or in other words, of an oxide with water. Thus potash, K_2O , and water, H_2O , form $2\text{K}(\text{OH})$, or potassium hydroxide; also $\text{CaO} + \text{H}_2\text{O}$ similarly give $\text{Ca}(\text{OH})_2$, or calcium hydroxide. In general, when soluble in water, bases give an alkaline reaction with turmeric paper or red litmus paper, and they also neutralize an acid, as explained in the next article. Further, the bases yield water on ignition, that is, at a temperature sufficiently high to break up the compound.

Among minerals the bases are represented by the hydroxides, or hydrated oxides, as $\text{Mg}(\text{OH})_2$, magnesium hydrate (brucite); $\text{Al}(\text{OH})_3$, aluminium hydrate (gibbsite); also, $(\text{AsO})_2(\text{OH})_2$, diaspore, etc.

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448. Salts.—A third class of compounds are salts; these may be regarded as formed chemically by the reaction of a base upon an acid, or, in other words, by the neutralization of the acid. Thus calcium hydroxide and sulphuric acid give calcium sulphate and water:



Here calcium sulphate is the salt, and in this case the acid, sulphuric acid, is said to be neutralized by the base, calcium hydroxide. It is instructive to compare the formulas of a base, an acid, and the corresponding salt, as follows:

Base, $\text{Ca}(\text{OH})_2$; Acid, H_2SO_4 ; Salt, CaSO_4 .

Here it is seen that a salt may be simply described as formed from an acid by the replacement of the hydrogen atom, or atoms, by a metallic element or radical.

449. Typical Salts.—The commonest types of salts represented among minerals are the following:

Chlorides: salts of hydrochloric acid, HCl ; as AgCl , silver chloride (cerargyrite).

Nitrates: salts of nitric acid, HNO_3 ; as KNO_3 , potassium nitrate (niter).

Carbonates: salts of carbonic acid, H_2CO_3 ; as $CaCO_3$ calcim carbonate (calcite and aragonite).

Sulphates: salts of sulphuric acid, H_2SO_4 ; as $GaSO_4$, calcim sulphate (anhydrite).

Phosphates: salts of phosphoric acid, H_3PO_4 ; as $Ca_3(PO_4)_2$, calcium phosphate.

Silicates: several classes of salts are here included. The most common are the salts of metasilicic acid, H_2SiO_3 ; as $MnSiO_3$, manganese metasilicate (rhodonite). Also salts of orthosilicic acid, H_4SiO_4 ; as Mn_2SiO_4 , manganese orthosilicate (tephroite).

Numerous other classes of salts are also included among mineral species; their composition, as well as that of complex salts of the above types, is explained in the descriptive part of this work.

450. Normal, Acid, and Basic Salts.—A neutral or normal salt is one in which the basic element completely neutralizes the acid, or, in other words, one of the type already given as examples, in which all the hydrogen atoms of the acid have been replaced by metallic atoms or radicals. Thus, K_2SO_4 is normal potassium sulphate, but $HKSO_4$, on the other hand, is acid potassium sulphate, since in the acid H_2SO_4 , only one of the bonds is taken by the basic element potassium. Salts of this kind are called acid salts. The formula in such cases may be written as if the compound consisted of a normal salt and an acid; thus, for the example given, $K_2SO_4 \cdot H_2SO_4$.

A basic salt is one in which the acid part of the compound is not sufficient to satisfy all the bonds of the base. Thus malachite is a basic salt—basic carbonate of copper—its composition being expressed by the formula $Cu_2(OH)_2CO_3$.

This may be written $CuCO_3 \cdot Cu(OH)_2$, or $(Cu)_2O \cdot CO_2 \cdot H_2O$. The majority of the name often given to the compound

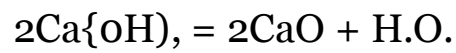
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minerals consist not of simple salts, as those noted above, but of more or less complex double salts in which several metallic elements are present. Thus common grossular garnet is an orthosilicate containing both calcium and aluminium as bases; its formula is $Ca_3Al_2(SiO_4)_3$.

451. Sulpho-salts.—The salts thus far spoken of are all oxygen salts. There are also others, of analogous constitution, in which sulphur takes the place of the oxygen; they are hence called sulpho-salts. Thus normal sulphureous acid has the formula $H_2S_2O_3$, and the corresponding silver salt is $Ag_2S_2O_3$, the mineral proustite. Similarly the silver salt of the analogous antimony acid is Ag_3SbS_4 , the mineral pyrargyrite. From the normal acids named, a series of other hypothetical acids may be derived, as $HAsS_3$, $H^2As_2S_5$, etc.; these

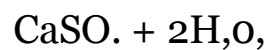
acids are not known to exist, but their salts are important minerals. Thus ziukenite, PbSb_2S_6 is a salt of the acid $\text{H}_2\text{Sb}_2\text{S}_6$, and jamesonite, $\text{Pb}_8\text{Sb}_7\text{S}_{21}$, of the acid $\text{H}_8\text{Sb}_7\text{S}_{21}$, etc.

462. Water of Crystallization,— As stated in Art. 447, the hydroxides, or bases and further basic salts in general, yield water when ignited. Thus calcium hydroxide $\text{Ca}(\text{OH})_2$ breaks up on heating into CaO and H_2O , as expressed in the chemical equation



So also the basic capric carbonate, malachite (formula given in Art. 460), yields water on ignition; and the same is true of the complex basic orthosilicates, like zoisite, whose formula is $(\text{HO})_2\text{Ca}_2\text{Al}_2(\text{SiO}_4)_2$. It is not to be understood, however, in these or similar cases, that water as such is present in the substance.

On the other hand, there are a large number of mineral compounds which yield water readily when heated, and in which the water molecules are regarded as present as so-called water of crystallization. Thus, the formula of gypsum is written



and the molecules of water ($2\text{H}_2\text{O}$) are considered as water of crystallization. So, too, in potash alum, $\text{KAl}(\text{SO}_4)_2 \cdot 12\text{H}_2\text{O}$, the water is believed to play the same part.

453. Formulae of Minerals.—The strictly empirical formula expresses the kinds and numbers of atoms of the elements present in the given compound, without attempting to show the way in which it is believed that the atoms are combined. Thus, in the case of zoisite the empirical formula is $\text{H}_2\text{Ca}_2\text{Al}_2\text{Si}_2\text{O}_{10}$. While not attempting to represent the structural formula (which will not be discussed here), it is convenient in certain cases to indicate the atoms which there is reason to believe play a peculiar relation to each other. Thus the same formula written $(\text{HO})_2\text{Ca}_2\text{Al}_2(\text{SiO}_4)_2$, shows that it is regarded as a basic orthosilicate, in other words, a basic salt of orthosilicic acid, H_2SiO_4 .

Again, the empirical formula of common apatite is $\text{Ca}_5(\text{F},\text{Cl})_3(\text{PO}_4)_3$; but if this is written $[\text{CaF}(\text{PO}_4)]_3\text{Ca}_2$, it shows that it is regarded as a phosphate of the acid H_3PO_4 , that is, $\text{H}_3(\text{PO}_4)_3$, in which the nine hydrogen atoms are replaced by four Ca atoms together with the univalent radical FOaF . In another kind of apatite the radical (CaCl) enters in the same way. Similarly to this the formula of pyromorphite is $(\text{PbCl})_2\text{Pb}_3(\text{VO}_4)_4$, of vanadinite $(\text{PbCl})_2\text{Pb}_3(\text{VO}_4)_4$.

Further, it is often convenient to employ the method of writing the formulas in vogue under the old dualistic system. For example.

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CaO.CO, for CaO, 3CaO.Al₂O₃.SiO₂, for Ca,Al,Si,O,

3Ag₂S.Sb₂S₃ for Ag,SbS₂, etc.

It is no longer believed, however, that the molecular groups CaO, Al₂O₃, etc., actually exist in the molecule of the substance. But in part because these groups are what analysis of the substance affords directly, and in part because so easily retained in the memory, this method of writing is still often used.

4f>4. Oxygen Ratio.—In the case of certain compounds, more especially the silicates, it is sometimes regarded as convenient to take note of the oxygen ratio, that is, the ratio in the number of oxygen atoms combined with the several elements, basic and acid. For example, the formula for grossular garnet given above is

Ca₃Al₂Si₃O₁₂, or 3CaO.Al₂O₃.3SiO₂

Here the oxygen atoms combined with the basic elements (calcium and aluminium) and the acid element (silicon) are

or again, for the basic elements combined,

3 + 3 : 6 or 1 : 1.

It must be noted that the oxygen ratio is in fact the ratio of the total valence of the elements of the different groups, the valence being measured by the combining power with hydrogen, while the oxygen ratio really notes the combining power with oxygen.

4S5. Calculation of a Formula from an Analysis.—The result of an analysis gives the proportions, in a hundred parts of the mineral, of either the elements themselves, or of their oxides or other compounds obtained in the chemical analysis. In order to obtain the atomic proportions of the elements:

Divide the percentages of the elements by the respective atomic weights; or, for those of the oxides: Divide the percentage amounts of each by their HOLECTULAR WEIGHTS; then find the simplest ratio in whole numbers for the numbers thus obtained.

Example.—An analysis of bournonite from Wolfsberg given by C. Bromela the results under (t) below. These percentages divided by the respective atomic weights, as indicated, give the numbers under (3). Finally the ratio of these numbers gives very nearly 1:8:1:1.

Hence

the formula derived is $\text{Cu}_8\text{S}_8\text{As}_2\text{S}_8$.

added under (4).

In the brief discussion of the periodic classification of the chemical elements of Art. 439, attention has been called to the prominent group among the elements which form analogous compounds. Thus calcium, barium, and strontium, and also lead, form the two series of analogous compounds, Aragonite Group. Barite Group.

CaCO_3 , aragonite. Also CaSO_4 , anhydrite.

BaCO_3 , witherite. BaSO_4 , barite.

SrCO_3 , strontianite. SrSO_4 , celestite.

PbCO_3 , cerussite. PbSO_4 , anglesite.

Further, the members of each series crystallize in closely similar forms. The carbonates are orthorhombic, with axial ratios not far from one another; thus the prismatic angle approximates to 60° and 120° , and corresponding to this they all exhibit pseudo-hexagonal forms due to twinning. The sulphates also form a similar orthorhombic series, and though anhydrite deviates somewhat widely, the others are close together in angle and in cleavage.

Again, calcium, magnesium, iron, zinc, and manganese form a series of carbonates with analogous composition, as shown in the list of the species of the Calciferous Group given on p. 353. This table brings out clearly the close relation in form between the species named. Incidentally, as an example of the deviation in form sometimes observed, it is to be noticed that dolomite (and perhaps others) are not normally rhombohedral like calcite, but belong to the prismatic type (p. 80).

This table also illustrates another essential point in regard to an isomorphous series, viz., the presence of intermediate members, or isomorphous mixtures of the simple compounds. These are viewed by most authors as due to the presence of both molecules crystallized together, usually in a certain definite ratio. Thus in normal dolomite, (CaCO_3) and (MgCO_3) are both present in the ratio of 1 : 1, and its formula is $\text{CaMg}(\text{CO}_3)_2$, or $\text{CaCO}_3 \cdot \text{MgCO}_3$. In mesitite (MgCO_3) and (FeCO_3) are present in the ratio of 3:1; its formula is $\text{Mg}_3\text{Fe}(\text{CO}_3)_4$, or $2\text{MgCO}_3 \cdot \text{FeCO}_3$. If it is not desired to express the ratio of

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the elements present, it is convenient to write the elements together in appearance, separated by a comma. Thus $(\text{Ca}, \text{Mg}, \text{Fe})\text{CO}_3$, would mean a carbonate in which calcium, magnesium, and iron are all present.

The Apatite Group forms another valuable illustration since in it are represented the analogous compounds, apatite and promorphite, both phos-

phosphates, but respectively phosphates of calcium and lead; also the analogous end compounds pyromorphite, mimetite, and vanadinite respectively lead phosphate, lead arsenate, and lead vanadate. Further, in all these compounds the radical (RCl) or (BF) enters in the same way (see Art. 493). Thus the formulas for the two kinds of apatite and that for pyromorphite are as follows:



Some of the numerous isomorphous groups are mentioned below. For a discussion of them, as well as of their properties, which might be mentioned here, reference must be made to the descriptive part of this work.

Isometric System.—The *Sphalerite* group, including *spinel*, $\text{Hg}_2\text{As}_2\text{O}_7$; also *magnetite*, *chromite*, *garnet*, etc. The *Orthoclase* group, as *galena*, PbS ; *argentite*, Ag_2S , etc. The *Garnet* group, as *grossularite*, $\text{Ca}_3\text{Al}_2(\text{SiO}_4)_3$; etc.

Tetragonal System.—The *Rutile* group, including *rutile*, TiO_2 ; *stibnite*, Sb_2O_3 . The *Beelite* group, including *beelite*, CaWO_4 ; *wolframite*, PbWO_4 ; *wulfenite*, PbMoO_4 .

Hexagonal System.—*Apatite* group, already mentioned, including *apatite*, *pyromorphite*, *mimetite*, and *vanadinite*. *Corundum* group, *corundum*, Al_2O_3 ; *emerald*, Fe_2O_3 .

Trigonal System.—*Calcite* group, already mentioned. *Phenacite* group, etc.

Orthorhombic System.—*Arsenic* group, and *Barite* group, both mentioned above. *Crysolite* group, $(\text{Mg}, \text{Fe})_2\text{SiO}_4$; *Topaz* group, etc.

Monoclinic System.—*Copper* group, including *malachite*, $\text{Cu}_2(\text{OH})_2\text{CO}_3$; *azurite*, $\text{Cu}_3(\text{OH})_2(\text{CO}_3)_2$; *pyroxene* and *amphibole* groups, and the *Mica* group.

Triclinic System.—*Feldspar* group.

457. **Isomorphism.**—It is important to note that the intermediate compounds in the case of an isomorphous series, such as those spoken of in the preceding article, often show a distinct gradation in crystalline form, and more particularly in physical characters (fl., sp., specific gravity, optical properties, etc.) This is illustrated by the species of the calcite group already referred to; also still more strikingly by the group of the triclinic feldspars as fully discussed under the description of that group. See further Art. 406.

The feldspars also illustrate two other important points in the subject, which must be briefly alluded to here. The triclinic feldspars have been shown by *Tschermak* to be isomorphous mixtures of the end compounds in varying proportions:



Here it is seen that these compounds have not an analogous composition in the narrow sense previously illustrated, and yet they are isomorphous and form an isomorphous series. Other examples of this are found among the pyroxenes, the scapolites, etc.

Further, the Feldspar group in the broader sense includes several other species, conspicuously the monoclinic orthoclase, KAlSi_3O_8 , which, though belonging to a different system, still approximates closely in form to the triclinic species.

458. Dimorphism. Isodimorphism.—A chemical compound, which crystallizes in two forms genetically distinct, is said to be dimorphous; if in three, trimorphous, or in general polymorphous. This phenomenon is called DIMORPHISM

or POLYMORPHISM.

An example is given by the compound calcium carbonate (CaCO_3), which is dimorphous: appearing as calcite and as aragonite. As calcite it crystallizes

in the rhombohedral system, and, unlike as its many crystalline forms are, they may be all

referred to the same fundamental rhombohedron, and, what is more, they have all the same cleavage and the same specific gravity (2.7), and, of course, the same optical characters. As aragonite, calcium carbonate appears in orthorhombic crystals, whose optical characters are entirely different from those of calcite; moreover, the specific gravity of aragonite (2.9) is higher than that of calcite (2.7).

Many other examples might be given: Silica (SiO_2) is dimorphous; appearing as quartz, rhombohedral, $G = 2.65$; as tridymite, hexagonal, $G = 2.3$, and perhaps in other forms. Titanium dioxide (TiO_2) is trimorphous, the species being called rutile, tetragonal ($G = 3.9$), orthorhombic, $G = 4.25$; octahedral, tetragonal ($G = 3.9$); and anatase, tetragonal, $G = 3.9$. Carbon appears in two forms, in diamond and graphite. Other familiar examples are pyrite and marcasite (FeS_2), sphalerite and wurtzite (ZnS), etc.

When two or more analogous compounds are at the same time isomorphous and dimorphous, they are said to be isodimorphic, and the phenomenon is called ISODIMORPHISM. An example of this is given in the Pyrite and Marcasite groups described later. Thus we have in the isometric Pyrite Group, pyrite, FeS_2 , smaltite, CoAs_2 ; in the orthorhombic Marcasite Group, marcasite, FeS_2 , safflorite, CoAs_2 , etc.

459. Chemical and Microchemical Analysis.—The analysis of minerals is a subject treated of in chemical works, and need not be touched upon here except so far as to note the convenient use of certain qualitative methods, as described in the later part of this chapter.

Of more importance are the microchemical methods applicable to sections under the

microscope and often yielding decisive results with little labor. This subject has been particularly developed by Boficky, Haushofer, Behrens, Strong, and others. Reference is made to the discussion by Rosenhnsch (Mikr. Phys., 1893, p. 259 et sea., also the list of authors on p. 313.)

460. Mineral Syntheaia.— '5^hB occurrence of certain mineral componnda (e.g., the chrysolites) among the products of metallurgical furnaces has long been noted. But it has only been in recent years that the formation of artificial minerals has been made the subject of minute Byatematic experimental study. In this direction the French chemists have been particularly anccessful, and now it may be stated that the majority of common minerals—

?uartz, the feldspars, ampbibole, mica, etc.—have been obtained in crystallized orm. Even the diamond has been formed in minute crystals by MoiEsan. These studies are obviously of great importance particularly as throwing light upon the method of formation of minerals in nature (e.g., the diamond). The chief results of the work thus far done are given in the volumes mentioned in the Introduction, p. 4.

461. Alteration of Minerals. Fsendomarplii.—The chemical alteration of mineral species under the action of natural agencies is a subject of great importance and interest, particularly when it results in the change of the original composition into some other equally definite rompound, A crystallized mineral which has thus suffered change bo that its form no longer belongs to its chemical composition has already been defined (Art. 262. p. 144) OS a pseiidomorpa It remains to deGcrihe more fully the different kinds of psendomorphs. Pseudomorphs are classed under several heads:

1. Pseudomorphs by auMUuimi.
3. Pseudomorphs by simple deposition, and either by (n) incrustation or (ft) infiltration.

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3. Psendomorphs by alteration; and tlieBe maj be altered (a) without a change of compositioD, by paramorphism; ib) by the)obb of an ingredient;

(c) by the aeBumptioiii of a foreign Bnbstance;

(d) by a partial exchange of constituents.

1. The first claea of paeudomorphs, by substitution, embraces those cases where there has been a gradual removal of the original material and a corresponding and simultaneous replacement of it by another, without, however, any chemical reaction between the two. A common example of this is a piece of fossilized wood, where the original fiber has been

replaced entirely by silica. The first step in the process was the filling of the pores and cavities by the silica in solution, and then as the woody fiber by gradual decomposition disappeared the silica further took its place. Other examples are quartz after fluorite, calcite, and many other species; cassiterite after orthoclase; native copper after aragonite, etc.

3. Pseudomorphs by incrustation form a less important class. Such are the crusts of quartz formed over fluorite. In most cases the removal of the original mineral has gone on simultaneously with the deposition of the second, so that the resulting pseudomorph is properly one of substitution. In pseudomorphs by infiltration a cavity made by the removal of a crystal has been filled by another mineral.

3. The third class of pseudomorphs, by alteration, includes a considerable proportion of the observed cases, of which the number is very large. Conclusive evidence of the change which has gone on is often furnished by a nucleus, of the original mineral in the center of the altered crystal. e. g., a kernel of cuprite in a pseudomorphous octahedron of malachite; also of chrysolite in a pseudomorphous crystal of serpentine, etc.

(a) An example of paramorphism—this is, of a change in molecular constitution without change of chemical substance—is furnished by the change of aragonite to calcite (both CaCO_3) at a certain temperature; also various pseudomorphs of rutile after brookite (both TiO_2) from Magnet Cove, Arkansas.

(b) An example of the pseudomorphs in which alteration is accompanied by a loss of ingredients is furnished by crystals of native copper in the form of cuprite.

(c) In the change of cuprite to malachite—e. g., the familiar crystals from Ghessy, France—an instance is afforded of the assumption of an ingredient—viz., carbon dioxide (and water). Pseudomorphs of gypsum after anhydrite occur where there has been an assumption of water alone.

(d) A partial exchange of constituents—in other words, a loss of one and gain of another—takes place in the change of feldspar to kaolin, in which the potash silicate disappears and water is taken up; pseudomorphs of limonite after pyrite or siderite, of chlorite after garnet, pyromorphite after galena, are other examples.

The chemical processes involved in such changes open a wide and important field for investigation. Their study has served to throw much light on the chemical constitution of mineral species and the conditions under which they have been formed. For the literature of the subject see the Introduction, p. 3 (Blum, Bischof, Roth, etc.). As typical studies of special cases the following articles are referred to:

Omtb. Corundum, etc. Am. Phil. Soc. Proc., 13. 861, 1878.

J. D. Dana. Serpentine pseudomorphia, Till; Foster mine. Am. J. Sc., B, 871, 1874.

Bnub aod Dana. Spodiimene, etc., BrancliTille, Ct. Am. 3. Sc. 20, aS7, 1880.

Sehraof. - Serpentine of BuhemU. Zs. Erysl., 6, 331. 1883.

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CHEMICAL EXAMINATION OF MINERALS.

462. The complete investigation of the chemical composition of a mineral includee, first, the identification of the elements present b; qualitative analysis, and, second, the determiDation of the relative amounts of each by quantitative analysis, from which last the formula can be calculated. Both processes carried out in full call for the equipment of a chemical laboratory. An approximate qnalitative analyBis, however, can, in many cases, be made quickly and simply with few conveniences. The methods employed involve either (a) the nae of acids or other reagents " in the wet way," or (b) the Qse of the blowpipe, or of both methods combined. Some practical instractionis will be given applying to both cases.

EXAMINATION IN THE WET WAY.

463. Reagents, etc.—The most commonly employed chemical reagents are ihe three mineral acids, hydrochloric, nitric, and sulphuric acids. To these may be added ammonia, also solutions of barium chloride, silver nitrate, ammonium molybdate, ammonium oxalate; finally, distilled water in a wash-bottle.

A few teat-tubes are needed for the trials and sometimes a porcelain dish with a handle called a casserole; further, a glass funnel and filter-paper. The Bunsen gas-bnrner (p. 256) is the best sonrce of heat, though an alcohol lamp may take its place. It is unnecessary to remark that the use of acids and the other reagents requires much care to avoid injury to person or clothing.

In testing the powdered mineral with the acids, the important points to be noted are: (1) the degree of solubility, and (3) the phenomena attending entire or partial solution; that is, whether (a) a solution is obtained quietly, without effervescence, and, if so, what its color is; or (b) a gas ia evolved, producing efFervescence; or (c) an insoluble constituent is separated out.

464. Solubility.—In testing the degree of solubility hydrochloric acid ie moat commonly used, though in the case of many metallic minerals, as the sulphides and compounds of lead and silver, nitric aoid is required. Less often sulphuric acid and aqua regia (iitrohydrochloric acid) are resorted to.

The trial is usually made in a test-tube, and in general the fragment of mineral to he

examined should be first carefully pulverized in an agate mortar. In most cases the heat of the Bunsen burner must be employed.

(a) Many minerals are completely soluble without effervescence; among these are some of the oxides, as hematite, limonite, goethite, etc.; some sulphates, many phosphates and arsenates, etc. Gold and platinum are soluble only in aqua regia or nitro-hydrochloric acid.

A yellow solution is usually obtained if much iron is present; a blue or greenish-blue solution (turning deep blue on the addition of ammonia in excess) from compounds of copper; pink or pale rose from cobalt, etc.

(i) Solubility with effervescence takes place when the mineral loses a gaseous ingredient, or when one is generated by the mutual reaction of acid and mineral. Most conspicuous here are the carbonates, all of which dissolve with effervescence, giving off the odorless gas carbon dioxide (CO_2), though some of them only when pulverized, or, again, on the addition of heat. In applying this test dilute hydrochloric acid is employed.

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Sulphureted hydrogen, or hydrogen sulphide (H_2S), is evolved by some sulphides when dissolved in hydrochloric acid; this is true of sphalerite, stibnite, etc. This gas is readily recognized by its offensive odor.

Chlorine is evolved by oxides of manganese and also chromic and vanadic acid salts when dissolved in hydrochloric acid.

Nitrogen dioxide (NO_2) is given off, in the form of red suffocating fumes, by many metallic minerals, and also some of the lower oxides (cuprite, etc.), when treated with nitric acid.

(c;) The separation of an insoluble ingredient takes place; With many silicates, the silica separating sometimes as a fine powder, and again as a jelly; in the latter case the mineral is said to gelatinize (sodalite, analcite). In order to test this point the finely pulverized silicate is digested with strong hydrochloric acid, and the solution afterward slowly evaporated nearly to dryness. With a considerable number of silicates the gelatinization takes place only after ignition; while some others, which ordinarily gelatinize, are rendered insoluble by ignition.

With many sulphides (as pyrite) a separation of sulphur takes place when they are treated with nitric acid.

Some compounds of titanium and tungsten are decomposed by hydrochloric acid with the separation of the oxides of the elements named (TiO_2 , WO_3). The same is true of salts of molybdic and vanadic acids, only that these oxides are soluble in an excess of the acid.

Compounds containing silver, lead, and mercury give with hydrochloric acid insoluble residues of the chlorides. These compounds are, however, soluble in nitric acid.

When compounds containing tin are treated with nitric acid, the tin dioxide (SnO_2) separates as a white powder. A corresponding reaction takes place under similar circumstances with minerals containing arsenic and antimony.

Insoluble Minerals. —A large number of minerals are not sensibly attacked by any of the acids. Among these may be named the following oxides: Corundum, spinel, chromite, diaspore, rutile, cerite, quartz; also cerargyrite; many silicates, titanates, tantalates, and niobates; some of the sulphates, as barite, celestite; many phosphates, as xenotime, lazulite, chondrodite, amblygonite; also the borate, boracite.

465. Examination of the Solution.—If the mineral is difficultly, or only partially, soluble, the question as to solubility or insolubility is not always settled at once. Partial solution is often shown by the color given to the liquid, or more generally by the precipitate yielded, for example, on the addition of ammonia to the liquid filtered off from the remaining powder. The further examination of the solution yielded, whether from partial or complete solution, after the separation by filtration of any insoluble residue, requires the systematic laboratory methods of qualitative analysis.

It may be noted, however, that in the case of sulphates the presence of sulphur is shown by the precipitation of a heavy white powder of barium sulphate (BaSO_4) when barium chloride is added. The presence of silver in solution is shown by the separation of a white curdy precipitate of silver chloride (AgCl) upon the addition of any chlorine compound; conversely, the same precipitate shows the presence of chlorine when silver nitrate is added to the solution.

Again, phosphorus may be detected if present, even in small quantity.

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in a nitric acid solution of a mineral by the fine yellow powder which separates, sometimes after standing, when ammonium molybdate has been added.

EXAMINATION BY MEANS OF THE BLOWPIPE.*

466. The use of the blowpipe, in skilled hands, gives a quick method of obtaining a partial knowledge of the qualitative composition of a mineral. The apparatus needed includes the following articles:

Blowpipe, lamp, platinum-pointed forceps, platinum wire, charcoal, glass tubes; also a small hammer with sharp edges, a steel anvil an inch or two long, a horseshoe magnet, a

small agate mortar, a pair of cutting pliers, a three-cornered file.

Further, test-paper, both turmeric and blue litmus paper; a little pure tin-foil; also in small wooden boxes the fluxes: borax (sodium tetraborate), soda (anhydrous sodium carbonate), salt of phosphorus or microcosmic salt (sodium-ammonium phosphate), acid potassium sulphate (K₂SO₄); also a solution of cobalt nitrate in a dropping tube or bottle; further, the three acids mentioned in Art. 463.

467. Blowpipe and Lamp.—A good form of blowpipe is shown in Fig. 570. An air-chamber, at a, is essential to stop the condensed

moisture of the breath, but the tip (b), of platinum or of

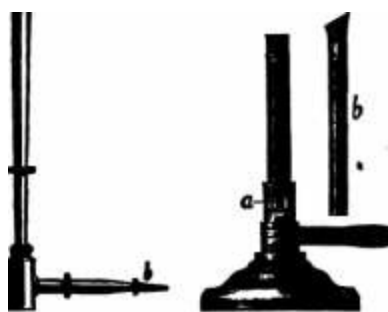
brass, though convenient is not essential, and many will prefer to do without the mouthpiece (c). The most convenient form of lamp is that furnished by an ordinary Bunsen gas-burner (Fig. 571), provided with a tube, J, which when inserted cuts off the air-supply at a; the gas then burns at the top with the usual yellow flame. This flame should be one to

one and a half inches high. The tip of the blowpipe is held near (or just within the flame, see beyond), and the air blown through it causes the flame to take the shape shown in Figs. 573, 574.

It is necessary to learn to blow continuously, that is, to keep up a blast of air from the compressed reservoir in the mouth-cavity while respiration is maintained through the nose. To accomplish this successfully and at the same time to produce a clear flame without unnecessary fatiguing effort calls for some practice.

* The subject of the blowpipe and its use is treated very briefly in this place. The student who wishes to be fully informed not only in regard to the use of the blowpipe and its instruments, but also as to all the valuable reactions practically useful for the identification of minerals, should consult a manual on the subject. The Manual of Descriptive Mineralogy, with an Introduction on Blowpipe Analysis, by George J. Brush; revised and enlarged by Samuel L. Penfield (New York, 1880), is particularly to be recommended. Another recent work is the Manual of Qualitative Blowpipe Analysis and Descriptive Mineralogy by F. M. Erdlich (New York, 1893).

\ Instead of this, a good stearia candle will answer, or an oil lamp with flat wick.



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When the tube, b, is removed, the gas burns with a colorless flame and is used for heating glass tubes, test-tubes, etc. An alcohol lamp will serve the same purpose. 672.

468. Porcelain wire.—The forceps (Fig. 573) are made of steel, nickel-plated, and should have a spring strong enough to support firmly the small fragment of mineral between the platinum points at d. The steel points at the other end are used to pick up small pieces of minerals, but must not be inserted in the flame. Care must be taken not to injure the platinum by allowing it to come in contact with the fused mineral, especially if this contains antimony, arsenic, lead, etc.

The platinum wire required should be of the size designated No. 37. A piece of platinum-foil is often useful; also a small platinum spoon.

469. Charcoal.—The charcoal employed should not snap and should yield but little ash; the kinds made from basswood, pine or willow are best. It is most conveniently employed in rectangular pieces, say four inches long, an inch wide, and three-quarters of an inch in thickness. The surface must always be perfectly clean before each trial.

Instead of charcoal a support of the metal aluminium, as suggested by Boss, is used by some workers with good results.

470. Glass Tubes.—The glass tubes should be of rather hard glass and say one-sixth to one-quarter of an inch in interior diameter. The smaller size is suitable for the closed tubes; these are simply made by beating a piece six inches long in the middle and then drawing the ends apart, the long ends being fused and pinched off. The larger size serves for open tubes, which may be five inches or so in length.

471. Blowpipe Flame.—The blowpipe flame, shown in Figs. 573, 574, consists of two cones: an inner of a blue color, and an outer cone which is nearly invisible. The heat is most intense just beyond the extremity of the blue flame, and the mineral is held at this point when its fusibility is to be tested.

The outer cone is called the oxidizing flame (O.F.); it is characterized by the excess of the oxygen of the air over the carbon of the gas to be combined with it, and has hence an oxidizing effect upon the assay. This flame is best produced when the jet of the blowpipe

is inserted a very little in the gas flame (see Fig. 573); it should be entirely non-luminous. The mineral is to be held at d.

The inner flame is called the reducing flame (R.F.); it is characterized by the escape of the carbon or hydrocarbons of the gas, which at the high temperature present tend to combine with the oxygen of the mineral brought into contact. 573. 574.

it (at d), or, in other words, to reduce it. The best reducing flame is produced when the blowpipe is held a little distance from the gas flame; it should retain the yellow color of the latter at its upper edge (see Fig. 574).

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472. Methods of Examination.—The blowpipe investigations of minerals include their examination, (1) in the platinum-pointed forceps, (2) in the closed and the open tubes, (3) on charcoal or other support, and (4) with the fluxes on the platinum wire,

1. EXAMINATION IN THE FORCEPS.

473. Use of the Forceps.—Platinum-pointed forceps are employed to hold the fragment of the mineral while a test is made as to its fusibility; also when the presence of a volatile ingredient which may give the flame a characteristic color is tested for, etc.

The following practical points must be regarded: (1) the mineral, especially those containing arsenic or antimony, which when fused may injure the platinum, should first be heated on charcoal; (2) the fragment to be held should be thin, and as small as can conveniently be held, with its edge protected well beyond the points; (3) when decomposition takes place, the heat must be applied slowly, or, if this does not prevent it, the mineral may be powdered and a paste made with water, thick enough to be held in the forceps or on the platinum wire; or the paste may, with the same end in view, be heated on charcoal; (4) the fragment whose fusibility is to be tested must be held in the hottest part of the flame, just beyond the extremity of the blue cone.

474. Fusibility.—All grades of fusibility exist among minerals, from those which fuse in large fragments in the flame of the candle (stibnite, see below) to those which fuse only on the thinnest edges in the hottest blowpipe flame (bronzite); and still again there are a considerable number which are entirely infusible {e.g., corundum}.

The exact determination of the temperature of fusion is not easily accomplished (cf. Art. 413, p. 332), and for purposes of determination of species it is unnecessary. The approximate relative degree of fusibility is readily fixed by referring the mineral to the following scale, suggested by von Eobell:

1. Stibnite. 4. Actinolite.

2. Natrolite. 5. Orthoclase.

3. Almandite Qamet. 6. Bronzite.

476. In connection with the trial of fusibility, the following phenomena may be observed
■ (a) coloration of the flame (see Art. 476); (b) swelling up (stilbite), or w/oKai'on of the mineral (Termiculite); or (c) glowing without fusion (calcite); and (d) intumescence, or a spitting out of the mass as it fuses (scapolite).

The color of the mineral after ignition is to be noted; and the nature of the fused mass is also to be observed, whether a clear or blebby glass is obtained, or a black slag; also whether the bead or residue is magnetic or not (due to iron, less often nickel, cobalt), etc.

The ignited fragment, if nearly or quite infusible, may be moistened with the cobalt solution and again ignited, in which case, if it turns blue, this indicates the presence of aluminium (as with cyanite, topaz, etc.); but note that zinc silicate (calamine) also assumes a blue color. If it becomes jinit, this indicates a compound of magnesium (as brucite).

Also, if not too fusible, it may, after treatment in the forceps, be placed upon a strip of moistened turmeric paper, in which case an alkaline reaction proves the presence of an alkali, sodium, potassium; or an alkaline earth, calcium, magnesium, barium, strontium.

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476. Flame Coloration.—The color often imparted to the outer blowpipe flame, while the mineral held in the forceps is being heated, makes possible the identification of a number of the elements.

The colors which may be produced, and the substances to whose presence they are due, are as follows:

Orange-red Lithium.

Green-red Lithium.

Purple-red Strontium.

Orange-red Calcium.

Yellow Sodium.

White or blue-white Barium.

Greenish-blue Boron.

Emtrald'grtan Oxide of copper.

Bluithgrttn Ptioepboric acid<pboBphates).

OrttnitA Hut AntimoDj.

WhUkb blue Anenic.

Ature-Uaa Chloride ot copper; also nleulum.

FioM Potatdum.

A yellowieh-green Qnme la also given by ibe oxide or sulphiile of molybdenam; a bluleh-greeD flame (In ttreaka) by eIdc; a pale greenlfb flame bj lelluiium; a blue flame b^ lead.

477. VotM.—Tbe preaece of soda, even In small quautltiea, produces a yellow flame, which (except In tbe spectroecope) more or le&B completely masks tbe coloration of the flame due to other siibBtancei, e.g., potassium. Tbe use of a wedge of blue glass then allows tbe cbarncterlslic violet color to be observed. Silicates are ofrec so diffliciilty decompuied that do distinct color Is obtained even wbea tbe substance U present; in such cases ('■o., potash feldspar) tbe powdered mineral may be fused on the platinum wire with au eqtuu volume of gypsum, wlieii tbe flame can be seeu (st least through blue glass). Assln, a siHcate like tourmaline fused witi a mixture of fluorlle and acid poiassium smphate yields Ihe characteristic green flame of boron. FhospLates and borates give the greeo flame in general best when they have been pulverized and motstened with sulphuric acid- Uolsteumg with hydrochloric acid makes the coloration In many cases (as with the carboatet of calcium, barium, strontium) mure dlstiuct.

3. Heating in thb Clossd ahd Open Tubes.

478. Tbe tnbes are nseful chiefly for examining minerals containing volatile ingredients, given off at the temperature of the gas flame.

Id the case of the closed tube, the heating goes on practically uninfluenced by the air present, since this is driven out of the tnbe in the early stages of the process. In the open twbe, on the other hand, a continual stream of hot sir, that is, of hot oxy^n, passes over the assay, tending to produce oxidation and hence often materially changing the result.

479. Cloud Tnbe.—A small fragment is inserted, or a small amount of the powdered mineral—in this case with care not to soil the sides of the tube— and neat is applied by means of the ordinary Bnnaen flame. The presence of

a volatile ingredient is ordinarily shown by the deposit, or aublimate, upon the , tnbe at some distance above the assay where the tube is relatiyely cool.

Independent of this, other phenomena may be noted, namely: decrepitation, as shown by fluorite, calcite, etc.; glowing, as exhibited by gadolinite; ^Aos»Aorfl«ce»ice, of which ilnorite is an example; chnng e o/'co7o7-(iimonite), and here the color of the mineral should be noted both when hot, and again after cooling; fusion; giving off oxygen, as mercuric oxide; yielding <irid or alkaline vapors, which should be tested by inserting a strip of moistened litmus or turmeric paper in the tube.

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Of the suMimaiet which form in the tabe, the following are those with which it ia most important to be familiar:

aab>tin;a. Snb'lnuM in fh* Closed Tabe.

Water |H,o| Colorleu liquid drops.

Sulpbur (8) Red to deep yellow, liquid; pale fellow, tolld.

Telluriuni dioxide (TeOi) Psle yellow to colorless, liquid; colorlen or white, solid.

Arsenic sulphide (Aa,B>) Dark red, liquid; reddish yellow, solid.

Auiimony oxysulphlde (BbtBiO) BIsck to reddish brown on coollDg, solid.

Arsenic (As) BIsck, brilliant metallic to gmy crjstalllDe, solid.

Mercury sulphide (HgS) Deep bluck, red when Tubbed very floe.

Mercury (Bjc) Omy metallic globules.

Id-addition to the above: Tellurium gives black fusible globules; seleniuni the same, but In part diirk red when very smull; the chloride ot lead and oxides of arsenic and aDtimouy give white solid sublimates.

480. Open Tube.—The small fragment is placed in the tabe about an inch from the lower end, the tube being stiehtly inclined (say 20°), but not enough to cause the mineral to slip out, and heat applied beneath. The cnrrent of air passing upward through the tube during the heating process has an oxidizing effect. The special phenomena to be observed are the formation of a sublimate and the odor of the eEscaping gases. The acid or alkaline character of the vapors is tested for in the same way as with the closed tube. Fluorides, when heated in the open tube with previously fused salt of phosphorus, yield hydrofluoric acid, which gives an acid reaction with test-paper, has a peculiar pungent odor, and corrodes the glass.

The more important sublimates are as follows:

Substance. above in the Open Tube.

Arsenic trioxide (As_2O_3) White, crystalline, volatile.

Antimony antimonate (Sb_2O_3) Straw-yellow, hot; white, cold. Infusible. non-volatile, amorphous. Obtained (from stibnite, also the sulph-antimonates (e.g., pyrite) as dense white fumes. Usually accompanied by the following:

Antimony trioxide (Sb_2O_3)... White, crystalline, slowly volatile. From native antimony and compounds not containing sulphur.*

Tellurium dioxide (TeO_2)... White to pale yellow globules.

Selenium dioxide (SeO_2) White, crystalline, volatile.

Molybdenum trioxide (MoO_3) Pale yellow, hot; white, cold.

Mercury (Hg), grey metallic globules, easily united by rubbing.

It is also to be noted that if the heating process is too rapid for full oxidation sublimation, like those of the closed tubes, may be formed, especially with sulphur (yellow

arsenic

(black), arsenic sulphide (orange), mercury sulphide (black), antimony trisulphide (black to reddish brown).

3. Heating with Charcoal.

481. The fragment (or powder) to be examined is placed near one end of the piece and this so held that the flame passes along its length; a slight ridge to prevent the mineral being blown off is sometimes useful. If the mineral decrepitates, it may be powdered, mixed with water, and then the material employed as a paste.

The reducing flame is employed if it is desired to reduce a metal (e.g., silver, copper) from its ores: this is the common case. If, however, the mineral

- The distinction here made is important; cf. Peacock's revised edition of Brush's Determinative Mineralogy, 1896.

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is to be roasted, that is, heated in contact with the air so as to oxidize and volatilize, for example, the sulphur, arsenic, antimony present, the oxidizing flame is needed and the mineral should be in powder and spread out. The points to be noted are as follows:

(a) The odor given off after short heating. In this W117 the presence of sulphur, arsenic (garlic or alliaceous odor), and selenium (odor of decayed horseradish) may be recognized.

(b) Fusion. —In the case of the salts of the alkalis the fused mass is absorbed into the charcoal; this is also true, after long heating, of the carbonates and sulphates of barium and strontium. (Art. 484.)

(d) The Sublimate. —By this means the presence of many of the metals may be determined. The color of the sublimate, both near the assay (K) and at a distance (D), as also when hot and when cold, is to be noted. The important sublimes are the following:

. Sublimates.

Antimony trisulfide (As_2S_3) White, very volatile, distills from the assay; also

gives fumes. Antimony sesquioxide (Sb_2O_3 and Sb_2O_5) Dense white, volatile; fumes near the assay.

Zinc oxide (ZnO) Can be yellow, hot; white, cold; moistened with cobalt nitrate and ignited (O.F.) becomes green.

Ulybdenum trioxide (HfO_3) Pale yellow, hot; yellow, cold; touched for a moment with the K.F. becomes blue. Also a copper-red sublimate (HfO_2) near the assay.

Lead oxide (PbO) Dark yellow, hot; pale yellow, cold. Also (from sulphides) dense white (resembling antimony), a mixture of oxide, sulphide, and sulphate of lead.

Bismuth trioxide (Bi_2O_3) Dark orange-yellow (N), paler on cooling; also bluish white (D). See further, p. 260.

Cadmium oxide (CdO) Nearly black to reddish brown (N) and orange yellow (D): often iridescent. To the above are also to be added the following:

Selenium dioxide SeO_2 , sublimate Steel-gray (N) to white (D); touched with R.F. gives an azirene odor. Same: also an offensive selenium odor.

Tellurium dioxide, TeO_2 , sublimate dense white (N) to gray (D); in B.F. volatilizes with green flame.

Tin dioxide. SnO_2 , sublimate from yellow hot to white cold; becomes bluish green when

moistened with cobalt solution and ignited.

Silver (with lead and antimony), sublimate reddish.

(e) The Infusible Residue. —This may (1) glow brightly in the O.F., indicating the presence of calcium, strontium, magnesium, zirconium, zinc, or tin. (2) It may give an alkaline reaction after ignition: alkaline earths. (3) It may be magnetic, showing the presence of iron (or nickel), (4) It may yield a globule or mass of a metal (Art. 482).

482. Bednotion on Charcoal.—In many cases the reducing flame alone suffices on charcoal to separate the metal from the volatile element present, with the result of giving a globule or metallic mass. Thus silver is obtained from argentite (Ag_2S) and cerargyrite (AgCl); copper from chalcocite (Cu_2S) and cuprite (Cu_2O), etc. The process of reduction is always facilitated by the use of soda as a flux, and this is in many cases (silpharsenites, etc.) essential. The finely pulverized mineral is intimately mixed with soda, and a drop of water added to form a paste. This is placed in a cavity in the charcoal, and subjected to a strong reducing flame. More soda is added as that present sinks into the coal, and, after the process has been continued some time, a metallic globule is often visible, or a number of them, which can be removed

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and separately examined. If not distinct, the remainder of the flux, the assay, and the surrounding coal are cut out with a knife, and the whole ground up in a mortar, with the addition of a little water. The charcoal is carefully washed away and the metallic globules, flattened out by the process, remain behind. Some metallic oxides are very readily reduced, as lead, while others, as copper and tin, require considerable skill and care.

The metals obtained (in globules or as a metallic mass) may be: iron, nickel, or cobalt, recognized by their being attracted by the magnet; copper, color red; bismuth, lead-gray, brittle; gold, yellow, not soluble in nitric acid; silver, white, soluble in nitric acid, the solution giving a silver chloride precipitate (p. 255); platinum, white, harder than silver, soluble in nitric acid with separation of white powder (PtO_2); lead, lead-gray (oxidized), soft and fusible. The coatings (see the list of sublimates above) often serve to identify the metal present.

The metals obtained may be also tested with borax on the platinum wire.

463. Detection of Sulphur in Sulphates. —By means of soda on charcoal the presence of sulphur in the sulphates may be shown, though they do not yield it upon simple heating. When soda is fused on charcoal with a compound of sulphur (sulphide or sulphate), sodium sulphide is formed, and if much sulphur is present the mass will have the hepatic (liver-brown) color. In any case the presence of the sulphur is shown by placing the fused

mass on a clean surface of silver, and adding a drop of water; a black or yellow stain of silver sulphide will be formed. Illuminating gas often contains sulphur, and hence, when it is used, the soda should be first tried alone on charcoal, and if a sulphur reaction is obtained (due to the gas), a candle or lamp must be employed in the place of the gas.

484. It is also useful in the case of many minerals to test their fusibility or infusibility with soda, generally on the platinum wire. Silica forms if not in excess a clear glass with soda, so also titanate acid. Salts of barium and strontium are fusible with soda, but the mass is absorbed by the coal. Many silicates, though alone difficultly fusible, dissolve in a little soda to a clear glass, but with more soda they form an infusible mass. Manganese, when present even in minute quantities, gives a bluish-green color to the soda bead.

4. Treatment on the Platinum Wire.

485. Use of the fluxes. —The three common fluxes are borax, salt of phosphorus, and carbonate of soda (p. 354). They are generally used with the platinum wire, less often on charcoal (see p. 360). If the wire is employed it must have a small loop at the end; this is heated to redness and dipped into the powdered flux, and the adhering particles fused to a bead; this operation is repeated until the loop is filled. Sometimes in the use of soda the wire may at first be moistened a little to cause it to adhere.

When the bead is ready it is, while hot, brought in contact with the powdered mineral, some of which will adhere to it, and then the heating process may be continued. Very little of the mineral is in general required, and the experiment should be commenced with a minute quantity and more added if necessary. The bead must be heated successively first in the oxidizing flame (O.K.) and then in the reducing flame (R.F.), and in each case the color noted when hot and when cold. The phenomena connected with fusion, if it takes place, must also be observed.

Minerals containing sulphur or arsenic, or both, must be first roasted (see p. 561) till these elements have been volatilized. If too much of the mineral has been added and the

bead becomes

opaque to iron the color, it may, while hot, be flattened out with the hammer, or drawn out into a wire, or part of it may be removed and the remainder diluted with more of the flux.

With salt of phosphorus the wire should be held above the flame so that the escaping gases may support the bead; this is continued till quiet fusion is attained.

It is to be noted that the colors vary much with the amount of material present; they are also modified by the presence of other metals.

486. Borax,—The following list emimerateB the difEerent colored beads obtained with borax, both in the oxidizing tO.F.) aud reducing flames (R.F.)> and also the metals to the preaence of whose oxidee the colore are dne. Compare further the reactions given in the list of elements (Art. 4B8).

Oolta In Boiax Bead. Bubstanoe.

1. OxiPiziNO Flauk. Colorlen, or opaque white... Silica, calcium, aluminlumi also silver, ^c, etc Iron, cold—(pule yellow, hot, if in smiill amount). Bed, red-brown to brown.... Chromium (CrOi), Lot—(yellowUb green, cold).

Manganese (Hn,Ot), amethysllDe-rM—{rlolol, hot). Iron (Fe,o,), hot—HJellow, cold)—It saturated. Nickel (NiO), red-brown to brown, cold—<Tlolet, hot). Uranium (UO,), hotHyellow, cold).

Green Copper (CuO), hot—(blue, cold, or bluish green if highly saturated). Chromium (CrOi), yellowish green, cold—(red, hot).

Tellow Iron (FeiOi). bot—fpale yellow to co1orl«M, cold)—but

red-brown aud yellow if saturated. Uranium (UOi), hot. If in small amount; paler on cooling. Chromium {CrOi), hot and in small amount—(yellowish greeu, cold).

Blue Cobnlt(CaO), Lot and cold.

Copper (CuO), cold If highly saturated—{green, hot),

Violet Nickel (NiOl. hot—(red-brown, cold).

Mangniieec (HuiOi), hot—(amethystine-red, cold),

S. RsDucufo Flamx (R F.).

Colorleai. Hanganese (UuO), or a faint rose color.

Red • Copper (CuiO. with Cu). opaque red.

Green Iron (FeO), bottle-green.

Chromium (CriOi), ememld-green.

Uranium (U,Oi), yellowish green If uturated.

Blue Cobalt (CoO), hot and cold.

Gray, turWd Nickel (NI).

487. Salt of Fhotphorni.—This flux dyes for the most part reaoctions similar to those obtained with borax. The only cases enumerated here are those which are distinct, and hence those where the flux is a good test.

With silicates this flux forma a glass in which the bases of the silicate are dissolved, but the silica itself is left insoluble. It appears as a skeleton readily seen floating about in the melted bead.

The colore of the beads, and the metals to whose oxides these are dne, are: Color. ,
Bubstanoe.

Red Chromium In O.F., hot—(fine green when cold).

Green,,,,,,., Chromium in O.F. and R.F., when cold—(red in O.F., hot).

Uolybdenum in R.F., dirty greeu, hot; fine green, cold—(yellow-cmen
in O.F.). Uranium In R,F., cold; yellow.greeo, hot.

Vanadium, chrome-green in R.F., cold—(brownish red, liot). In O.F. dark yellow, hot, paler
on cooling.

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Tallow HolTbdennm, jellowiih green lu O.F., liot, pkler on cooling—<Iti B.F.,
' dirty gr««ii, hot; floe Kreen, cold).

Uranium In O.F., hot; yellowish green, cold—[in K.F., yellowish green,
cooling—(in R,F.,

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498. The following list contains the most characteriBtic reactions, chiefly before tho
blowpipe and in some cases also in the wet way, of the different elements and their
oxides. It is desirable for every student to gain familiarity with them by trial with as m&oy
minerals as possible. Many of them have already been briefly mentioned in the preceding
pages. For a thoroughly full description of these and other characteristic tests (blowpipe
and otherwise) reference should be made to the rolhme by Brush and Fenfield referred to

It is to be remembered that while the reaction of a single substance may be perfectly distinct if alone, the presence of other substances may more or less entirely obscure these reactions; it is consequently obvious that in the actual examination of minerals precautions have to be taken, and special methods have to be devised, to overcome the difficulty arising from this cause. These will be gathered from the "pyrognostic characters" (Pyr.) given in connection with the description of each species in the Third Part of this work.

For many substances the most satisfactory and delicate tests are those which have been given by Bunsen in his important paper on Flame-reactions.* The methods, however, require for the most part so much detailed explanation, that it is only possible here to make this general reference to the subject.

Aluminium. —The presence of aluminium in most infusible minerals, containing a considerable amount, may be detected by the blue color which they assume when, after being heated, they are moistened with cobalt solution and again cooled {e.g., cynnrite, andalusite, etc.). Very hard minerals (as corundum) must be finely pulverized. The test is not conclusive with fusible minerals since a glass colored blue by cobalt oxide may be formed. It should be noted that the infusible calamine (zinc silicate) also assumes a blue color when treated with cobalt nitrate.

Antimony. —Antimonial minerals roasted on charcoal give dense white inodorous fumes; metallic antimony and its sulphur compounds give in the open tube a white sublimate of oxide of antimony (see p. 260). Antimony sulphide (stibnite), also many sulphoantimonates, give in a strong heat in the closed tube a sublimate of antimony oxysulphide, black when hot, brown-red when cold. See also p. 260.

In nitric acid, compounds containing antimony deposit white insoluble metantimonic acid.

Arsenic. —Arsenides, sulpharsenites, etc., give no fumes when roasted on charcoal, usually easily recognized by their peculiar garlic odor. In the open tube they give a white, volatile, crystalline sublimate of arsenic trioxide. In the closed tube arsenic sulphide gives a sublimate dark brown-red when hot, and red or reddish yellow when cold: arsenic and some arsenides yield a black mirror of metallic arsenic in the closed tube. In arsenates the arsenic can be detected by the garlic odor yielded when a mixture of the powdered mineral with charcoal dust and sodium carbonate is heated (R.F.) on charcoal.

Barium. — A yellowish-green coloration of the flame is given by all barium salts, except the silicates; an alkaline reaction is usually obtained after intense ignition.

* *Flammeoreactionen*, Lieb. Ann., 185, 557, 1866, or Phil. Mag., 33, 81, 1866.

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In a solution the presence of barium is proved by the heavy white precipitate (BaSO₄) formed upon the addition of dilute sulphuric acid.

Bismuth.—On charcoal alone, or mixed with soda. Bismuth gives a very characteristic orange-yellow brittle globule of the reduced metal which is also obtained (with soda). Also when treated with 8 or 4 times the volume of a mixture in equal parts of potassium iodide and sulphur, and fused on charcoal, a beautiful red sublimate of bismuth iodide is obtained; near the mineral the coating is yellow.

Boron.—Many compounds of boron (borates, also the silicates datolite, danburite, etc.) give the flame intense yellowish green, especially if moistened with sulphuric acid. For some silicates (as tourmaline) the best method is to mix the powdered mineral with one part powdered fluoride and two parts potassium bisulphate. The mixture is moistened and placed on platinum wire. At the moment of fusion the green color appears, but lasts but a moment,

A dilute hydrochloric acid solution containing boron gives a reddish-brown color to turmeric paper which has been moistened with it and then dried at 100°; this color changes to black when ammonia is poured on the paper.

Calcium.—Many calcium minerals (carbonates, sulphates, etc.) give an alkaline reaction on turmeric paper after being ignited. A yellowish-red color is given to the flame by some compounds {e.g., strontium after moistening with HCl}; the strontium flame is a much deeper red.

In solutions (not too acid) calcium is precipitated as oxalate by the addition of ammonium oxalate.

Cadmium.—On charcoal with soda, compounds of cadmium give a characteristic sublimate of the reddish-brown oxide.

Carbonates.—All carbonates effervesce with dilute hydrochloric acid, yielding the odorless gas CO₂ (e.g., calcite); many require to be pulverized, and some need the addition of heat (dolomite, siderite). Carbonates of lead should be tested with nitric acid.

Chlorides.—If a small portion of a mineral containing chlorine (e.g. chloride, also pyromorphite, etc.) is added to the bead of salt of phosphorus, saturated with copper oxide, the bead when heated is instantly surrounded with a very intense purplish flame of copper chloride.

In solution chlorine gives with silver nitrate a white curdy precipitate of silver chloride, which darkens in color on exposure to the light; it is insoluble in nitric acid, but entirely so in ammonia.

O⁺remium. —Chromium elves with borax abend which (O.F.I is yellow to red (hot) and yellowish Kreeo (cold) and R.F. a flue emerald-green. With salt of phosphorus In O.F. the bead Is dirty green (hot) aud clear green (coldj; In H.F, the same, Cf. Tauadlum beyond (also p. 3S8).

(MalL —A beautiful blue bend is obtained with boras in both flames from miDerats containing cobalt; tbe color may be obscured by considerable Iron or nickel unless These are first oxldted off [p. 182), Where sulphur or arsenic is present the mineral should first be thoroughly roasted on charcoal.

Copper.—On charcoal, at least with soda, metallic copper can be reduced from most of Its compounds. With borax It gives (O.P.) a green bead when hot. becoming blue when cold; also (R.F.), If saturated, an opaque red bead containing Cu₂O and often Cu is obtained.

Host nieialilic compounds are soluble in nitric add. Ammonia produces a green pre-dpltate in the solution, which la dissolved when an excess is added, the solution taking an intenHc blue color.

^uortn«.—Heated In the closed tube many fluorides give off fumes of hydrofluoric acid, which react acid with test-paper and etch the glass. Sometimes pota⁺um Uaulpbste must be added [see also p. 200).

Heated gently In a platinum crucible with sulphuric acid, many rompounda («.ff., fluorite) give off hydrofluoric add, which corrodes the exposed parts of a glass plale plnced over It which has been coated with wax and then scratched.

/nm.—With borax iron pives a bead (O.P.I which Is yellow to brownish red (according to quantily) while hot, but is colorless to yellow on cooling; R,F, becomes bottle-green (see p. 8G8). Minerals which contain even a small amount of Iron yield a magnetic mass when heated In tbe reducing flame.

Lead.—With soda on charcoal a malleable globule of metallic lead is obtained from lead compounds; tbe coating has a yellow color near tbe assay; the mlphide gives also a white coatini; (PbSOi) farther off (p. 201), On being touched with the redudng flnme the coating disappears, tingeing the flnme azure-blue.

In solutions dilate sulphnrlic acid gives a white predpltate of lead sulphate; when

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delfcBcj Is required an excess of tbe acid is added, tbe solulfon evaporated to dryoen, uid WHlur iiddtid; tLe lead lulpbule, If preseat. will lieu be lefl as a residue.

XifAiun).—Litbiutn gives an iutense canaine-red lu Uie outer tlaioe, Ibe color somewhat resemblIDg that of the stnintluiu dame but ia deeper; In veiy small quantities It b evident

in the spectroscope.

Magnamm. —Holsleoed, aflai heatlug, wltli coliat ultral« and axain Ignited, a pink color Is ubtained from some Infusible compounds of magnesium {'.g-, bruccllel.

Mangaiitae. —With borax manganese gives a bead Tiolet-red [O.F.). and colorless (R.F.). With soda (O.F.) it gives a bluish-green bead: this reaction is very delicate and may be relied upon, even Id presence of almoat any other meiat.

UeTmry.—la tbe closed tube a sublimate u(metallic mercury ia yielded when tbe mineral is heated with dry sodium carbonate. In the open tube the sulphide gives a mirror of metallic mercury; in the closed tube a black lUBlerleas sublimate of HgB, red when rubbed, la obtained.

UoMydenvm.—Oa charcoal molybdenum sulphide gives near the assay a copper-red Btalu (O.F.), and beyond a while coating of the oxide: the former become» azure-blue when for a moment touched with the R.P. The salt of phosphorus bead (O.F.) is yellowlsli green (hot) and nearly colorless (cold); also (RF.) a fine green.

Nickel. ^'W\l\i borax, nickel oxide gives a bead which (O.F.) is violet when hoi and red-brown on cooling; (R.F.) the glass becumes gray and turbid from the separation of metallic nickel.

il'ti'Mufn (Columbium).— Ad add sol u lion boiled with metallic tin gives a blue color. The relictions with the fluxes are not very satisfactory.

Nitrate*. —These detonate when heated on charcoal. Heated in a tube with sulphuric acid they give off red fumes of nitrogen dioxide (NOi).

iVuMjjA^fus.^Most phosphates impart a green color to the flame, especially after having been moistened wiith sulphuric acid, though this lest maybe rendered unsutisraclory by the presence of other coloring agents. If they are used in the closed tube with a fragment of memllic magnesium or sodium, and afterward moistened with water, phospUureled hydrugen is given oB, recognizable by lu disagreeable odor.

&. few drops of a nitric acid solution, containinr phosphoric add. produces in a solution of ammonium molybdale a pulverulent yellow preclptnte of ammonium phospbOr molybdate.

PoloMium. —Potaali imparts a violet cotoc to Uie flame when alone. It Is best delected In small quantities, or when soda or liihia Is present, by the aid of the specttoacope. See also p. Sw.

SeUniuTn. —On charcoal selenium fnses easily, giving oS brown funea with a peouliar disagreeable organic odor; the sublimate on charcoal is rolatlle, and when heated (ILF.)

gives a fine azure-blue flame.

Boron. — h. small fragment of a silicate in the salt of phosphorus bead leaves a skeleton of silica, the bases being dissolved.

If a silicate in fine powder is fused with sodium carbonate and the mass then dissolved in hydrochloric acid and evaporated to dryness, the silica is made insoluble, and when strong hydrochloric acid is added and then washed, the bases are dissolved and the silica left behind.

Many silicates, especially those which are hydrous, are decomposed by strong hydrochloric acid. The residue separates as a powder or, after evaporation, as a jelly (see p. 265).

Silver. — On charcoal in O.F. silver gives a brown coating. A globule of metallic silver may generally be obtained by heating on charcoal in O.F., especially if soda is added. Under some circumstances it is desirable to have recourse to cupellation.

From a solution containing any salt of silver, the insoluble chloride is thrown down when hydrochloric acid is added. This precipitate is insoluble in acid or water, but entirely so in ammonia. It changes color on exposure to the light.

Strontium. — Compounds of strontium are usually recognized by the brick-red color which they give to the blowpipe flame; many yield an alkaline reaction after ignition. (Cf. barium.)

Sodium. — Compounds containing sodium in large amount give a strong yellow flame.

Sulphur. Sulphide; S . — In the closed tube some sulphides give off sulphur; in the open tube they yield sulphur dioxide, which has a characteristic odor and reddens a strip of moistened litmus paper. In small quantities, or in sulphates, sulphur is best detected by fusion on charcoal with soda. The fused mass, when sodium sulphide has thus been formed, is placed on a clean silver coin and moistened; a distinct black stain on the silver is thus obtained (the precaution mentioned on p. 382 must be exercised).

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A solution in hydrochloric acid gives with barium chloride a white trace of barium sulphate.

Tellurium. — Tellurides heated in the open tube give a white or gray, fusible to colorless drope (p. 360). On charcoal they give a white coating and color the B.P. green.

Tin. — Minerals containing tin {e.g., cassiterite), when heated on charcoal with soda or

potasalum cjrualde, yield metallic tin in minute globules; these are malleable, but harder than silver. Dissolved in nitric acid, white insoluble stannic oxide separates out.

antimony.—Titanium gives in the R.P. with salt of phosphorus a bead which is violet when cold. Fused with sodium carbonate and dissolved with hydrochloric acid, and heated with a piece of metallic tin, the liquid takes a violet color, especially after partial evaporation.

Wulfenite.—Tungsten oxide gives a blue color to the salt of phosphorus bead (I.F.). Fused and treated as stannic oxide (see above) with the addition of zinc instead of tin, gives a blue color.

Uranium. —Uranium compounds give to the salt of phosphorus bead (O.F.) a greenish yellow bead when cool; also (H.F.) a deep green color (p. 255).

Vanadium.—With borax (O.P.) vanadates give a bead yellow (hot) changing to yellowish green and nearly colorless (cold); also (H.F.) dirty green (hot), fine green (cold). With salt of phosphorus (O.F.) a yellow to amber color (thus differing from chromium); also (R.F.) fine green (hot).

Zinc.—On charcoal compounds of zinc give a coating which is yellow while hot and white on cooling, and moistened by the cobalt solution and again heated becomes a fine green. Note, however, that the zinc silicate (calamine) becomes blue when heated after moistening with cobalt solution.

Zinc-ionium.—A dilute hydrochloric acid solution, containing zinc, imparts an orange-yellow color to turmeric paper, moistened by the solution.

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488. Determinative Mineralogy may be properly considered under the general head of Chemical Mineralogy, since the determination of minerals depends mostly upon chemical tests. But crystallographic and all the physical characters have also to be used.

There is but one exhaustive way in which the identity of an unknown mineral may in all cases be fixed beyond question, and that is by the use of a complete set of determinative tables. By means of such tables the mineral in hand is referred successively from a general group into a more special one, until at last all other species have been eliminated, and the identity of the one given is beyond doubt.

A careful preliminary examination of the unknown mineral should, however, always be made before final recourse is had to the tables. This examination will often suffice to show what the mineral in hand is, and in any case it should not be omitted, since it is only in this way that a practical familiarity with the appearance and characters of minerals can be gained.

The student will naturally take note first of those characters which are at once obvious to the senses, that is: crystalline form, if distinct; general structure, cleavage, fracture, luster, color (and streak), feel; also, if the specimen is not too small, the apparent weight will suggest something as to the specific gravity. The characters named are of very unequal importance. Structure, if crystals are not present, and fracture are generally unessential except in distinguishing varieties; color and luster are essential with metallic, but generally very unimportant with unmetallic, minerals. Streak is of importance only with colored minerals and those of metallic luster (p. 187). Crystalline form and cleavage are of the highest importance, but may require careful study.

The first trial should be the determination of the hardness (for which eud

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the pocket-knife U often sufficient in experienced hands). The Beoond trial should be the determination of the specific gravity. Treatment of the pow> dered mineral with acids may come next; by this means (see pp. 354, 255) a car- -bonate is readily identified, and also other resnlts obtained. Then should follow blowpipe trials, to ascertain t\io fusibility; the color given to the flame, if any; the character of the sublimate given oS in the tahea and on charcoal; the metal reduced on the latter; the reactions with thej?n»o«, and other points as explained in the preceding pages.

How mnch the observer learns in the above way, in regard to the nature of his mineral, depends npon hia knowledge of the characters of minerals in general, and npon his familiarity with the chemical behavior of the varions elementary substances with reagents and before the blowpipe (pp. 264 to 367). If the results of snch a preliminary examination are sufficiently definite to snggest that the mineral in hand is one of a small number of species, reference may be made to their full description in Part. IV. of this work for the final decision.

A number of tables, in which the minerals included are arranged according to their crystalline and physical characters, are added in the Appendix, They will in many cases aid the observer in reaching a coQcIusion in regard to a specimen in nand.

The first of these tables is intended to include all well-defined species, grouped according to the crystalline system to which they belong and arranged under each system in the order of their specific gravities; the hardness is also added in each case. The relative importance of toe individual species is shown by the type employed. Following tnis are minor tables enumerating species characterized by some one of the prominent crystalline forma; that is, those crystallizing in cubes, octahedrons, rhombohedrons, etc. Other tables give the names of species prominent because of their cleavage; structure of different cypee; hardness; luster; the various colors, etc The student i inded t

i to make frequent use of these tables, not simplv for aid in the identificntion of specimens, but rather because they will help him in the difficult task of learning the prominent obaractmra of tiie more important minerals.

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PART IT, DESCRIPTIVE MINERALOGY.

489. Soopo of DeMTIptin Kin«rftli^7. —It is the province of Descriptive Mineralogy to deecribe each miDeral species, as regards: (1) form and atructore; (3) physical characters;

(3) chemical composition and allied blowpipe characters; (4) occurrence in nature with reference to geographical distribution and association with other species; also in connection with the above to show how it is distinguished from other species. Further, to classify mineral species into more or less comprehensive groups according to those characters regarded as most essential. Other points which may or may not be included are the investigation of the methods of origin of minerals; the changes that they undergo in nature and the results of such alteration; also the methods by which the same compounds may be made in the laboratory; finally, the uses of minerals as ores, for ornament and in the arts.

490. Boheme of Classification.—The method of classification adopted in this work, and the one which can alone claim to be thoroughly scientific, is that which places similar chemical compounds together in a common class and which further arranges the mineral species into groups according to the more minute relations existing between them in chemical composition and in crystalline form and other physical properties.

Upon this basis there are recognized eight distinct classes, beginning with the Native Elements; these are enumerated on the following page. Under each of these, sections of different grades are made, also based on chemical relations. Finally, the mineral species themselves are arranged, as far as possible, in isomorphous groups, including those which have, at once, analogous chemical composition and similar crystallization (See Art. 466). It is unnecessary to take the space here to develop the entire scheme of classification in detail, since a survey of the successive sub-classes under any one of the divisions will make the principles followed entirely clear. A few remarks, only, are added for sake of illustration.

Under the Oxides, for example, the classification is as follows: First, the Oxides of silicon (quartz, tridymite, opal). Second, the Oxides of the semi-metals, tellurium, arsenic, antimony, bismuth, also molybdenum, tungsten. Third, the Oxides of the metals, as copper, zinc, iron, manganese, tin, etc. The third section is then subdivided into the anhydrous and hydrous species. Farther, the former fall into the four divisions: Protoxides, R_2O and RO ; Sesquioxides, R_2O_3 ; Intermediate oxides, RO, R_2O_3 ; Dioxides, RO_2 . Under each of these heads come finally the individual species, arranged so far as possible in isomorphous groups. Thus we have the Hematite group, the Rutile group, etc.

In regard to the various classes of salts it may be stated that, in general, they are separated into anhydrous, acid, basic and hydrate sections; the special subdivisions called for, however, vary in the different cases.

For an explanation of the abbreviations used in the description of species, see p. 4.

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DESCRIPTION OF THE MINERAL SPECIES.

SCHEME OF CLASSIFICATION.

I. HATIVX XlBMBim.

IL SVXFEIBla, SKLSmiU, TSILITRIIU, AUIHIDXt, AaTIMOnDXB.

III. Snlptao-saltB—BULFEABmna, SUIPHAVTIMOKITEB, SirLPHOBIIXDTE-tT».

IV. Haloids.—GELoan>i>, Baoumn, looniu; Pi.iroxiinB. V. OzmiA

VI. Oxygen Salts.

1. Casbosatib.

2. SnlCATXB, TiTABATIS.

3. HLOBATEB, TAjn'Al.ATtl.

1. PHOBPEATEB, AXSKHATXS, TAVADATU; ALTTIMOBATIS. Hl-

6. BOBATEB. UBABATBB.

6. S'ITLFEATEB, CHBOXATBI, TIILVBATEB.

7. TintCWIATBS, MDLTBSATXS.

Vn. Salts of Organic Acids: OxalateB, Mellates, etc. VIII. HYSBOOABBOH COKFOITBU.

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NATIVE ELEMENTS.

I. HATIVi: BT.FWmWTS.

The NATIVE ELEMENTS are divided into the two distinct sections of the Metals and the Non-metals, and these are connected by the transition class of the Semi-metals. The distinction between them as regards physical characters and chemical relations has already been given (Art. 487).

The only non-metals present among minerals are carbon, sulphur, and selenium; the last, in one of its allotropic forms, is closely related to the semi-metal tellurium.

The native semi-metals form a distinct group by themselves, since all crystallize in the rhombohedral system with a fundamental angle differing a few degrees only from 90°, as

shown in the following list:

Tellurium, $rr' = 93^\circ 3'$. Arsenic, $rr' = 94^\circ 54'$.

Antimony, $rr' = 92^\circ 53'$. Bismuth, $rr' = 93^\circ 20'$.

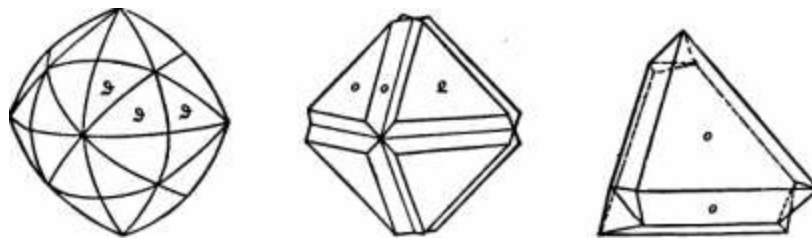
An artificial form of selenium is known with metallic luster and rhombohedral in crystallization, with $rr' = 93^\circ$. Zinc (also only artif.) is rhombohedral ($rr' = 93^\circ 46'$) and connects the semi-metals to the true metals.

Among the tetrahedral the isometric group is prominent, including gold, silver, copper, mercury, amalgam (AgHg), and lead.

Another related isometric group includes the metals platinum, iridium, palladium, and iron; further palladium is rhombohedral and also iridosmine (Iridosmine).

DIAMOND.

Isometric and probably tetrahedral, but the $\{100\}$ and $\{111\}$ forms not distinguished. Commonly in octahedrons, also hexoctahedrons and other forms; luster frequently rounded or striated and with triangular depressions (on $\{111\}$).



Twins common with tw. pi. o. Crystals often distorted. In spherical forms; massive.

Cleavage: $\{111\}$ highly perfect, Fracture conchoidal. Brittle. H. = 10. G. = 3.516-3.535 crystals. Luster adamantine to greasy. Color white or colorless; occasionally various pale shades of yellow, red, orange, green, blue, brown; sometimes black. Usually transparent; also translucent, opaque. Refractive and dispersive power high; index $n^D = 2.4195$. (See Art. 305.)

$V_{\{111\}} = 1$, Ordinary. In crystals usually with rounded faces and varying from those which are colorless and free from flaws (first order) through various shades of color, yellow the most common; often full of flaws hence of value only for cutting purposes.

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2. Black or Boort; rounded forms with rough exterior and radiated or confused crystallization without cleavage. Color black or grayish black. Opaque. Brazil.

Comp.—Pure carbon; the variety carbouado yields on combination a slight ash.

Pyr., etc.—Unaffected by heat except at very high temperatures, when (like an oxide) it burns to carbon dioxide (CO₂); out of contact with the air transformed into a kind of coke. Not acted upon by acids or alkalis.

Diaph.—Disintegrates («.f., from quartz crystal} but its extreme hardness and brilliant adamantine luster; the form, cleavage, and high specific gravity are also distinctive characters; it is optically isotropic; transparent to X-rays.

Occ.—The diamond occurs chiefly in alluvial deposits of gravel, sand, or clay, associated with quartz, gold, platinum, zircon, corundum, leucite, rutile, ilmenite, hematite, and also andalusite, chrysoberyl, topaz, corundum, tourmaline, garnet, etc.; the associated minerals being those common in granitic rocks or granitic veins. Also found in quartzose conglomerates, and further in connection with the laminated gneiss-quartz rock or quartzose hydrothermal schist, which in thin slabs is more or less flexible. The rock occurs at the mines of Brazil and the Urals; and also in Georgia and North Carolina, where a few diamonds have been found.

It has been reported as occurring in a pegmatite vein in gneiss at Bellary in India. It occurs further in connection with an eruptive peridotite in East Africa. It has been noted as Krayish particles forming one per cent of the meteorite which fell at Novo-Urel, Russia. Sept. 23, 1988; also in the form of black diamond (H. = 6) in the meteorite of Carcote, Chili; in the meteoric iron of Cañon Diablo, Arizona. It has been formed artificially by Moissan.

India was the chief source of diamonds from very early times down to the discovery of the Bazaruto mines; the yield is now small. Of the localities, the most important in India, in the Madras presidency, included the famous "Golconda mines." The diamond deposits of Ilmizil have been worked since the early part of the 17th century, but have yielded very largely, although at the present time the amount obtained is small. The most important region was that near Diamantina in the province of Minas Gerais; also from Bahia, etc.

The discovery of diamonds in South Africa dates from 1826. They were first found in the gravel of the Vaal river; they occur from Pekaapstroom down to the Junction with the Orange river, and along the latter as far as Hope Town. These mines are now comparatively unproductive, and have been nearly abandoned for the dry digging, discovered in 1871.

The latter are chiefly in Orange-Free State, south of the Vaal river, on the border of the Orange-Free State. There are here a number of limited areas approximately circular or oval in form, with an average diameter of some 200 to 300 yards, of which Klomdriv, De Beer's, Du Toit's Pan and Bultfontein are the most important. A circle 3½ miles in diameter encloses the four principal diamond mines. The general structure is similar: a wall of

nearly horizontal black carbonaceous shale with upturned edges enclosing the diamantiferous area. The upper portion of the deposit consists of a friable mass of little coherence of a pale yellow color, called the "yellow ground." Below the reach of atmospheric influences, the rock is more firm and of a bluish green or greenish color; it is called the "blue ground" or simply "the blue." This consists essentially of a ferruginous breccia: a base of hydrated magnesian silicate permeated by calcic and opaline silica and

highly magnesian

— —, monazite, pyrite, zircon, etc. The diamonds are rather abundantly distributed

through the mass, in some claims to the amount of 4 to 8 carats per cubic yard. The original rock seems to have been a peculiar type of peridotite. These areas are believed to be volcanic pipes, and the occurrence of the diamonds is obviously connected with the eruptive outflow, they having probably been brought up from underlying rocks. The South African mines in Griqualand up to June 1890 are estimated to have yielded 60 million carats (13 tons) of diamonds, valued at about 870 million dollars.

Diamonds are also obtained in Borneo, associated with platinum, etc.; In Australia, and the Urals.

In the U. S. a few crystals have been met with in North Carolina, Georgia, and Virginia; several have been found in Wisconsin, also in California and several points. Reported from Idaho and from Oregon with platinum.

Some of the famous diamonds of the world with their weights are as follows: the

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Orlov, which weighed when brought to England 186 carats, and was recut as a brilliant, 106 carats; the Regent or Pitt, 187 carats; the Florentine or Grand Duke of Tuscany, 183 carats; the Hancy, 68 carats. The "Star of the South," found in Brazil in 1852, weighed before and after cutting respectively 264 and 125 carats. Also famous because of the rarity of their color are the green diamond of Dresden, 40 carats, and the deep blue Hope diamond from India, weighing 44 carats. The history of the above named and of others is given in many works on gems.

South Africa has yielded some very large stones. Among these may be mentioned the following: The Victoria (or the Imperial) from one of the Kimberley mines weighed when found 4,096 carats; the Stewart weighed before and after cutting 288 and 120 carats respectively; the Tiffany diamond, of a brilliant golden yellow, weighs, cut as a double brilliant, 119 carats. The Excelsior from Jagersfontein weighed when found 971 carats and was 8 inches in its largest dimension; this is the largest ever known to have been discovered.

CWFTONITB.—Carbon in minute cubic crystals. H. = 3. S. Q. = 213. Color and streak, black; from the Youdegn, West Australia, meteoric Iron, found in 1884.

OHAPBITB. Plumbago. Black Lead.

Rhombohedral. In six-sided tabular crystals. Commonly in embedded foliated masses, also columnar or radiated; scaly or slaty; granular to compact; earthy.

Cleavage: basal, perfect. Thin laminae flexible, inelastic. Feel greasy. H. = 1-2, G. — 2.09-2.23, Luster metallic, sometimes dull, earthy. Color iron-black to dark steel-gray. Opaque. A conductor of electricity.

Comp.—Carbon, like the diamond; often impure from the presence of iron sesquioxide, clay, etc.

Pyr., etc.—At a high temperature some graphite burns more easily than diamond, other varieties less so. B.B. Infusible. Unaltered by acids.

Diagn.—Characterized by its extreme softness (soapy feel): iron-black color; metallic luster; low specific gravity; also by infusibility. Cf. molybdenite, p. 85.

Obs.—Graphite occurs in beds and embedded masses, in limestone, or scales. In granite, gneiss, mica schist, crystalline limestone. It is in some places a result of the alteration of heat of coal. Often observed in meteoric irons. A common furnace product.

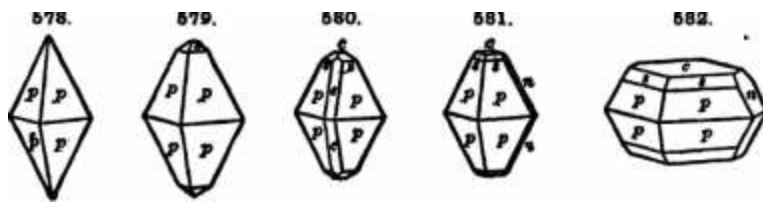
Occurs at Borrowdale in Cumberland; at Arendal in Norway, in quartzite; in the Ural, Finland; Passau in Bavaria. In Irkutsk, in the Tunguska Mts., in eastern Siberia, the Alibert graphite mine furnishes some of the best graphite of the world. Large quantities are brought from the East Indies, especially from Ceylon.

Forms beds in gneiss, at Sturbridge, Mass.; at Ticonderoga, N. T., with pyroxene and titanite; and at Hillsdale. Columbia Co., N. Y.; Byers. Chester Co., Pa.; Loudon Co., Va.; Wake Co., N. C. A graphitic earth is mined for paint in Arkansas. In California, in Alpine Co., Kern Co., etc. In Humboldt Co., Nevada; Beaver Co., Utah; Albany Co., Wyoming. A large deposit occurs at St. John, New Brunswick.

The name black lead, applied to this species, is inappropriate, as it contains no lead. The name graphite, of Werner, is derived from γράφω, to write, alluding to its use for "lead" pencils.

SCHCDOITB. Amorphous carbon observed in some schists.

Orthorhombic. Axes $S:l:h = 0.8131 : 1 : 0.9034$. $d = 2.678$. $S = 79$. $B = 68.1$.



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SBBCEIPTITB UINERALOOT.

Crystals commonlr acute pyramidal; aometimes tliick tabular | c, also ■phenoidal in habit (Fig. 583). See also Figs. 66, p. 30, and 303, p. 94. Also ""^ moaave, in reniform shapes, in-

,110 A 110 = 78* H'. ■ ■ ' '

001 A 101 = W 58'. 001 A Oil = 63° 17'. 001 A 118 = 40* 10'. 001 A HI = 71' 40'. Ill A 111 = 94'82'. , 111 A 111 = 78" 84'.



ciUBtiii^, stalstitic and Btalag-mitic; in powder.

Cleavage: c,m,p imperfect. Fracture conchoidal to uneven. Bather brittle to imperfectly soctile. H. = 1-5-2-5. G. = 2-05 - 2-09. Luater resinous. Color sulphur-yellow, straw- and ^_^ honey-yellow, yellowish brown,

greenish, reddish to yellowish gray. Streak white. Transparent to trans-Sicent, A non-conductor of electricity; by friction amatively electrified. Optically +. Double refraction strong. Ax, plane \b.. Bi.J_c. Dispersion p <v. 2H^, = 103° 16' Di. Refractive indices, see p. 208.

Comp., Tar.—Pure sulphur; often contaminated with clay, bitumen, and other impurities.

Sulphur may also be obtsinod Id the Uborotoiy Id other allotropio fomu; a tnonocllnic form U commoD.

Pyr., «to.—Helta at 108' o., and at 370* buroa with a blulah flame vIeldInK aulpfanr dioxide. iDBoluble in water, and not acted on by the acide, but soluble in carbon dtaulpbide.

Dlff.—Beadtty distinguished by tbe color, fusibility nnd combuBtibility. Oba,—The great

repositories of sulphur are either beds of Kypsuni and the amOclate rocks, or the regions of active and extinct volcanoes. In the valley of Nolo and Mazsaro, Id Sicily; at Conil, in Spain; fiei, Switzerland; Cracow, Poland, ft accura In tbe former sltuaiion; near Bologna, Italy, in flue crystals, embedded in bitumen. Slcily and the neighboring volcanic Isles; the Solfalara, near Naples; the volcanoes of the Pacific ocean, etc., are localities of the latter kind. It is also deposited from hot sprinn in Iceland: aod Is met with in certain metallic veins, thus with lead ores near HOsena at Monte Poni, Sardinia. The Sicilian mines at Qirgenti yield large quantities for commerce, including beautifully crystallized specimens.

Sulphur is found near tbe sulphur springs of New York, Virgiula, etc.. sparingly; In many coal deposits and elsewhere, where pyrites Is underiroing decomposition; lii minute crystals on cleavage surfaces of galena, PhenlKville, Pa. Some Important deposits occur Id the western UTS., as In Wyoming, In the Ulntsh Sits., 30 miles s.e. of Evan^toii; In Nevada, Humboldt county; Steamboat Springs. Waalioe Co.; Columbus, Esmeralda Co. In southern Utah Id large deposits, at Cove Creek, Millard county. In Oaliforniii, at the

Eysers of Napa vnley, Sonoma Co.; in Santa Barbara In good crystals; near Clear Lake, ike Co., a large deposit. In tbe Yellowstone Park, In deposits and about the fumarolee. S«leiiBilphar. Contains sulphur and selenium, orange-red or reddish brown; from the islands Vulcano and Lipari.

Rhombohedral. Generally granular massive; sometimes reticulated, reniform, stalactitic.

Cleavage: c highly perfect. Fracture uneven and fine granular. Brittle. H. = 3-5. G. = 5-63-5T3. Luster nearly metallic. Color and streak tin-white, tarnishing to dark gray.

Comp.—Arsenic, often with some antimony, and traces of iron, silver, gold, or bismuth.

Pyr.—B.B. on charcoal volalilizes without fusing, coats tbe coal with white arsenic trioxlde, and affords a garlic odor; the coating treated In R. F. volatilizes, tingeing the ^me blue.

Oba.—Occurs In veins in crystalline rocks and the older schists, often accompanied by ores of antimony, ruby silver, realgar, sphalerite, and other metallic minerals. Thus in tbe

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KATTVE ELEMNT8,

•llT«r mlDM of SasODV; alio AndreaBberr; Joachiauthsl. Bohemlft; tii Hungaiv; Norwar. etc. Abundant at Chatlai^tlo, Cbili- la the U. S. •parlaglr at HaverbilT and Jackaon N. H.; near LeadTtle, Colorado; Walaon Creek, Brltlib Columbia.

Allcmontlte. Arsenical Antimony, SbAsi. In rcQlforni uiasse*. O. = 6-208. LuBter metallic. Color tln-irfalte or reddish gray. From Allemont; PHbnun, Bohemia, etc.

TaUnrlnm. In prlamatic crystals (Fig. 14. p. 10): commonly columnar to flne-graoular mawive. Q. =63. Color lio-wblte. From TntoiyivMiIa and Colorado.

Rhombohedral. Generally tnaMive, lamellar and distinothly cleavable; also radiated; grsnrlar.

Cleavage: c liigblj perfect; also other cleayagee. Fracture nneTen; brittle. H. = 3-3-5. G. = 6'65-6-72. Luster metallic. Color and Btreak tin-white.

Gonp.—Antimony, containing sometimes silver, iron, or arsenic. Pyr.—B.B. on cbarcool fuses, gives a while coalur In bolb O. F. and R. F.; If the blowing be Intermitted, the globule continues to glow, giving off white fumea, until It Is flnnlly crusted over with prismatic ciystala of antimony iriozide. The white coating tinges the B. F. bluish green. Crystalllzs readily from fusion.

Oba.—Occurs near Sala in Sweden; Andreasberg lu the Bare; Allemont. Dauphin^; Pfibram, Bohemia; Mexico; Chili; Borneo. In the U. 8., at WarreD, TS. J., rare; in Kern Co., Cal. At Prince WillUm parish, York Co., N. Brunswick.

Rhombohedral. Usnally reticnlated, arboreBceot; foliated or granular. Cleavage: c perfect. Seotile. Brittle, but when heated somewhat malleable. H. = 2-3-5. G. = 9-7l>-9'83. Luster metallic. Streak and color silver-white, with a reddish hue; subject to tarnish. Opaque.

Comp., Var.—Bisiiinth, with traces of arsenic, sulphur, tellurium, etc.

Fyr., eto.~B.B. on charcoal fuses and entirely vulattllzea. giving a coating oranpe-yellow while hot, lemon yellow on cooling. Fuses at 266° C. DlBaoIres in nitric acid; Bub»^ueut dilution csiises n wliite precipitate. Ciystallizea readily from fusion.

Obi Occurs in vehia in gucisa and otljer cryBlalliije rocks and clay slate, accompanying various ores of silver, cobalt, lend and zinc. Thus at tbe mines of Saxony and Bohemia, etc.; HcTmnc. Coi-r^ze, France. Also at Mndum, Norway; at Falun, Sweden. In Cornwall ana Devonshire; near Copiapo. Cbili; Bolivia.

Occurs nC Monroe, Conn.; Brewer's mine, Cheaterfleld district, S. Car.; near Cummins City, and elsewliere In Colorado.

Zlno. Probably does not occur in the native Btate. In the laboratory It is obtained iu hexagonal prisms with tapciing pyramids; also in complex crystalline aggregates. It also appears to crystallize In the laomelrlc system, at least in various alloys.

Gold Group.

aoij>.

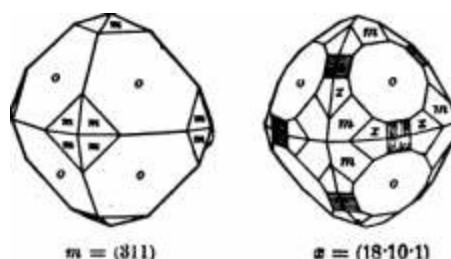
Isometric. Distinct crystals rare, o most common, also d (110), m (331), aud X (ISlO'l); crystals often elongated in direction of an octahedral axis, giving rise to rhombohedral forms (Figs. 452. 453, p. 135), and arborescent shapes; also in plates flattened | o, and branching at 60° parallel either to the edges or diagonals of an o face (see pp. 131, 132). Twins: tw. plane o. Skeleton crystals common; edges salient or rounded; in filiform, reticulated, dendritic shapes. Also massive and in thin laminae; often in flattened grains or scales.

Cleavage none. Fracture hackly. Very malleable and ductile. H. = 2-5-3. G. = 15-6-19-3, 19-33 when pure. Luster metallic. Color and streak gold-yellow, sometimes inclining to silver-white and rarely to orange-red. Opaque.

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DESCRIBTION MINERALOGY.

Comp^o Tkt. —Gold, but usually alloyed with silver in nature and sometimes containing also traces of copper or iron.



Tar.—1, Ordinary. Containing up to 16 p. c. of silver. Color varying accordingly from deep gold-yellow to pale yellow, and specific gravity from 19.3 to 16.5. The ratio of

Sold to silver of 8 : 1 corresponds to 15.1 p. c. silver. For Q. = 17.9, Ag = 16 p. c ; ■. = 16.8. Ag = 18.3; G. = 14.6, Ag = 38.4, Rose. The purest gold which has been described is that from Mount Huronian, in Quebec, which has yielded 89.7 to 88.8 of gold, the remainder being copper with a little iron; silver is present only as a trace.

2. Argentiferous: Uctum. Color pale yellow to yellowish white; G = 15.5-13.5. Ratio for the gold and silver of 1 : 1 corresponds to 86 p. c. of silver; 1 : 1, to 26 p. c; 2:1, to 31 p. c.; 3 : 1, to 18 p. c. The word in Greek means also amber; and its use for this alloy probably arose from the yellow color it has as compared with gold,

Mineralogists have also described containing palladium to 10 p. c. (—porphyry), from Porpez, Brazil; bismuth, including the black gold of Australia [—millerite, Ulrich]; also rhodium(?).

Pyritic Mo.—B.B. fuses easily (at 1100° C). Not acted on by acids. Insoluble in any single acid; soluble in hydrochloric acid (aqua regia), the separation not complete if more than 30 p. c. Ag is present.

Di£—Readily recognized {eg., from other metallic minerals, also from scales of yellow mic'i) by its malleability and high specific gravity, which last make it possible to separate it from the gangue by washing; distinguished from chalcopyrite and pyrite by insolubility in nitric acid, both of which are brittle.

Occurrence—Native gold is found, when in situ, with comparatively small exceptions in the quartz veins that intersect metamorphic rocks, and to some extent in the wall rock of these veins. The metamorphic rocks thus intersected are mostly chloritic talcose and argillaceous schist of dull green, dark gray, and other colors: also, much less commonly, mica and hornblende schist, gneiss, diorite, porphyry; and still more rarely granite. A laminated quartzite, called Itacolumite, is common in many gold regions as those of Brazil and North Carolina, and sometimes specular schists, or slaty rocks containing much foliated specular iron (hematite) or magnetite in grains. A quartzose conglomerate is sometimes richly auriferous as in Transvaal. Less frequently calcite is the vein material.

The gold occurs in the quartz. Irregularly distributed. In strings, scales, plates. In masses which are sometimes an agglomeration of crystals; and the scales are often invisible to the naked eye. The associated minerals are: pyrite, which far exceeds in quantity all others, and is generally auriferous; next, chalcopyrite, galena, sphalerite, arsenopyrite each frequently auriferous; often tetradymite and other tellurium ores, native bismuth, native arsenic, stibnite. In some cases magnetite, hematite, sometimes barite, apatite, fluorite, siderite, chrysocolla. The quartz at the surface, or in the upper part of a vein is usually cellular and rusted from the more or less complete disappearance of the pyrite and other sulphides by decomposition; but below, it is commonly solid.

The gold of the world was early gathered, not directly from the quartz veins (quartz reefs of Australia and Africa), but from the gravel or sands of rivers or valleys in auriferous regions, or the slopes of mountains or hills, whose rocks contain in some parts and generally not far distant, auriferous veins; In California the method of hydraulic mining (water digging*) has been carried on on a stupendous scale. Most of the gold of

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NATIVE ELEMENTS. 577

In the Urals, Brazil, Australia, and all other gold regions, the gold is obtained from alluvial washings. At the present time, however, the alluvial washings are much less depended upon, in many regions all the gold being obtained direct from the rock.

The alluvial gold is usually in fluted scales of different degrees of fineness, the size depending partly on the original condition in the quartz veins, and partly on the distance to which it has been transported and assorted by running water. The rolled masses when

of some size ore called tellurides; in rare cases these occur very large and of great value. The Australian gold region has yielded many large nuggets: one of these found in 1858 weighed 164 pounds, and another (1862) weighed 190 pounds. In the auriferous sands, crystals of zircon are very common; also garnet and cyanite grains; often also monazite, diamond, topaz, corundum, iridosmine, platinum. The zircons are sometimes mistaken for diamonds.

Besides the free gold of the quartz veins and gravels, much gold is also obtained from auriferous sulphides or the oxides produced by their alteration, especially pyrite, also arsenopyrite, chalcopyrite, sphalerite, marcasite, etc. The only minerals containing gold in combination are the rare tellurides (sylvanite, etc.).

Gold exists more or less abundantly over all the continents in most of the regions of crystalline rocks, especially those of the semi-crystalline schists; and also in some of the large islands of the world where such rocks exist. In Europe, it occurs with silver ores in Hungary; in Transylvania at Verespatak and Nagy[^]; in the sands of the Rhine, the Danube, and other rivers; on the southern slope of the Pennine Alps; in Piedmont; in many of the streams of Cornwall; in North Wales; in Scotland, near Leadhills; in the county of Wicklow, Ireland; in Sweden, at Edelfors; in Norway, at Kongsberg.

In Asia, gold occurs along the eastern flanks of the Urals for 500 miles, and is especially abundant at the Berezniki mines near Ekaterinburg; also at Fetropavlovski; Nizhni Tagilsk; Hias, near Zlatoust and Ut. Ilmen, etc. Ekaterinburg is the capital of the mining district. Siberian mines less extensive occur in the lesser Altai; at Nerchinsk, east of L. Baikal, including the Earn mines. Alluvial mines occur also in Little Tibet, Ceylon, and Malacca, China especially in the Amur district, Corea, Japan, Formosa, Sumatra, Java, Borneo, the Philippines, and other East India Islands; at numerous points in British India.

In Africa, gold occurs at Kordofan, between Darfur and Abyssinia; also, south of the Sahara in western Africa, from the Senegal to Cape Palmas. Also in Transvaal in southern Africa, at Lydenburg, with quartz veins and alluvial washings, and at Eeratereng; recently the Kaap gold fields in southern Transvaal have become very productive; the chief town of the region is Birmahon. The quartz reefs of Witwatersrand in the immediate vicinity of Johannesburg, farther west, are also very productive; here the gold occurs largely in a quartzose conglomerate.

In South America, gold is found in Brazil; in the U. S. of Colombia; Chile; Bolivia; sparingly in Peru. Also in Central America, especially in Honduras; also San Salvador, Guatemala, Costa Rica.

In Australia, the principal gold mines occur along the streams in the mountains of New South Wales and along the continuation of the same range in Victoria. Also obtained largely in Queensland. In Australia, particularly at Mt. Morgan, Rockhampton district. Gold occurs in Tasmania, New Zealand, and New Caledonia.

In North America, there are numberless mines along the mountain range of western America, and others along the eastern range of the Appalachians from Alabama and Georgia, to Labrador, besides some in portions of the intermediate Arcadian region about Lake Superior. They occur in many portions along the higher regions of the Rocky Mountains. In Mexico; in New Mexico, near Santa Fe, Cerillos, Avo, etc.; In Arizona, in the San Francisco, Yuma, and other districts; in Colorado abundant, the gold largely in auriferous pyrites, also in connection with tellurium minerals; the Cripple Creek mine in Colorado affords at present large quantities of gold; also in Montana: the Black Hills of Dakota; Idaho especially the Cassiar district, also Utah. Along ranges between the summit and the Sierra Nevada, in the Humboldt region and elsewhere. Also in the Sierra Nevada, mostly on its western slope (the mines of the eastern being principally silver mines). The auriferous belt may be said to begin in the Californian peninsula. Near the Tejon pass it enters California, and beyond for 180 miles it is sparingly auriferous, the slate rocks being of small breadth; but beyond this, northward, the slates increase in extent, and the mines in number and productiveness, and they continue thus for 200 miles or more. Gold occurs also in the Coast ranges in many localities, but mostly in too small quantities to be profitably worked. The regions to the north in Oregon and in Washington and Alaska, with British Columbia, are not many points auriferous, and productively so, though to a less extent than California. The Cariboo region on the Fraser river, and the Cassiar district on the Stickeen, have yielded considerable amounts. The Alaska quartz mines have

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DESCRIBE THE UNITED STATES.

has been worked to some advantage. In the alluvial gravels of the Yukon river and the tributaries; of the latter, the Klondike (now closed) reputed to be fabulously rich.

In eastern North America, the chief mines are mostly confined to the States of Virginia, North and South Carolina, and Georgia, or along a line from the Kappahannock to the Coosa in Alabama. In Canada, gold occurs to the south of the St. Lawrence, in the soil of the Chaudière and elsewhere; in Nova Scotia, at Deloro, near Hastings, Ontario (in arsenopyrite), also in the Port Arthur region, north of Lake Superior, and in the river-gravels of the Pacific slope, as before noted.

The world's production of gold was in 1897 about 9300,000,000, having considerably more than doubled since 1890. Of this amount the United States afforded about \$1,600,000,000, Africa about \$65,000,000, Australia nearly \$50,000,000, Russia about \$20,000,000, and other countries (China, Canada, India, So. America, etc.) the remainder. It is also interesting to note that in 1897 the production was nearly the same for the States of California and Colorado, the former (17,000,000, the latter a little less) of this amount. In 1890 Colorado produced only a little more than \$4,000,000.

shjVer.

Isometric. Crystals commonly distorted, in acicular forms, reticulated or arborescent shapes; coarse to fine filiform; also massive, in plates or flattened scales.

Cleavage none. Ductile and malleable. Fracture hackly, H. = 2.5-3. G. = 10.49, pure 10.5. Luster metallic. Color and streak silver-white, often gray to black by tarnish.

Comp^o Tar.—Silver, with some gold (up to 10 p. c), copper, and sometimes platinum, antimony, bismuth, mercury.

Pyr- ate.—B.B. on charcoal fuses easily to a silver-white globule, which in O.F. gives a faint dark-red coating of silver oxide; crystallizes on cooling; fusibility about 1050° C. Soluble in nitric acid, and deposited again by a solution of copper. Precipitated from its

solutions by hydrochloric acid in white curdy forms of silver chloride.

—Distinguished by its malleability, color (on the fresh surface), and specific gravity.

■

It occurs in arborescent and filiform shapes, in veins in rocks. Also occurs disseminated, but usually

invisible, in native copper, galena, chalcocite, etc.; rarely in volcanic ashes (Halite).

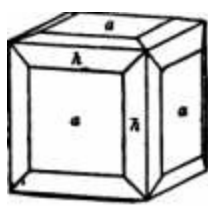
The mines of Kongsberg, in Norway, have afforded magnificent specimens; also the Baxton mines; occurs in Bohemia at Ibra and Joachimsthal at Andreasberg; Hiddgait: Alledont, Dauphin: In the Uml near Berezov; in the Altai, at Zmeov; and in some of the Cornish mines. In Durango, Sinaloa, and Sonora, in Mexico, are noted mines affording native silver; abundant in Peru.

In the United States disseminated through much of the copper of Michigan; at Silver Islet and at Port Arthur, Lake Superior. Occurs in Idaho, at the "Poor Man's Mine": in Nevada.

• aq rare: in California, especially in Silver Mountain district. Alpine

Co.: In the Maria vein, in Los Angeles Co. In Colorado, at many

localities, especially with argentiferous ores; in Montana, near Butte, Silver Bow Co., with manganese ores. In Arizona, at the Silver King mine, and with argentiferous ores elsewhere.



OOPPBR.

Isometric. The tetrahedron the most common form (Fig. 587); also in octahedral plates. Crystals often irregularly distorted and passing into twisted and wire-like forms; filiform and arborescent. Massive; as sand. $A = (\langle 10 \rangle)$ Twins: tw. pi. o, very common, often flattened or elongated

to spear-shaped forms. Cf. pp. 131,132.

Cleavage none. Fracture hackly. Highly ductile and malleable. $H. = 2-5-3$. $G. = 8.8-8.9$. Luster metallic. Color copper-red. Streak metallic shining. Opaque. An excellent conductor for heat and electricity.

Comp.—Pure copper; often containing some silver, bismuth, mercury, etc, Pyr., ato.—B.B. fuses readily; on cooling becomes covered with a coating of black

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oxide. Dissolves readily in nitric acid, evolving off red fumes, and producing a deep azure-blue solution with ammonia, fusing at $750^\circ C$,

Obs.—Copper occurs in beds and veins accompanying its various ores, especially cuprite, malachite, and azurite; also with the sulphides, chalcopyrite, chalcocite, etc.; often in the vicinity of dikes of igneous rocks; also in clay slate and sandstone.

Occurs at Timpanog, in the Ural, in fine crystals; at Nizhni Tagilak and elsewhere; Siberia. In Germany, at the Friedrichsberg mine, Saxony. Common in Cornwall, Brazil, Chili, Bolivia, and Peru. Found native copper. In Australia at Wallaroo; in New South Wales.

Occurs in the red sandstone region of the eastern United States, sparingly in Massachusetts, Connecticut, and more abundantly in New Jersey. Near New Haven, Conn., a mass was found in the drift weighing nearly 200 pounds; smaller isolated masses have also been found. The Lake Superior copper region, near Keweenaw Point, in northern Michigan, is the most important locality in the world. The copper is obtained practically all in the native state, sometimes in immense masses, and is obtained over an area 200 miles in length. It occurs in both amygdaloidal dolerite and andesite, near the junction of these two rocks; associated with quartz, prehnite, datolite, anatase, etc.; also distributed widely in grains through the sandstone. Occurs sparingly in California. In Arizona, common at the Copper Queen mine, Cochise Co.; also in Grant Co.,

N. Mexico, «t the Sauia Rita and olber mlues.

MBROURT. Quicksilver. Oediegeo Quecksilber Oerm,

In small fluid globules scattered through its gangne. G. = 13*596. Ltu-ter metallic, brilliant. Color tin-white. Opaque.

Comp.—Pure mercnry (Ug); with aometimes a little silver.

Pyr., atc—B-B. entirely vtilalile. vaporizing at 8So' C. Becomes solid at — 40* C, crystallizing in regular octahe<|TonB with cubic cleavage; G. = 14'4. Dissolves iu nitric acid.

Oba.—Mercury in the metallic stale is a rare mineral, and is usually associated with the sulphide ciunabar, from wblcli the supply of commerce is obtained. The rocks affording the metal and its ores are cblefly clay shales or scbisla of different geological ages. Also found in connection witli bot apriugs. See cinuabar.

ZiBAI).

Isometric. Crystals rare. TJaually in thin plates and small globular masses. Very malleable, and somewhat dactilo. H. = 1-5. G. = 11-37, Harstig mine. Lnster metallic Color lead-gray. Opaque.

Comp.—Nearly pure lead; sometimes contains a little silver, also antimony.

Pyr.—B.B. fusea easily, coating the charcoal with a yellow oxide whlcb, treated fn R.F., volatilizes, giving no azure-blue tinge to the flame. Fusibility 880° C. Disaolvei easily iu dilute nitric acid.

Ob«,—Of rare occumtnce. Found at Pajsberg, BarsiU. and Lftugban Iu Sweden; similarly at Nordmark; also In the gold washings of the Ural; reported elsewhere, but localities often doubtful. In the U. &., occurs at Breckinridge and Qunnison, Colorado; Wood Kiver district, Idaho,

AaOAIXIAM.

Isometric. Common faabit dodecahedral. Crystals often highly modified (Fig. 100, p. 39). Also massive in plates, coatings, and embedded grains.

Cleavage: d ia traces. Fracture conchoidal, uneven. Hather brittle to malleable. H. = 3-3-5. G. = 13-75-14-1. Luster metallic, brilliant. Color and streak silver-white. Opaque.

Comp.—(Ag,Hg), silver and mercury, varying from Ag.Hg, to Ag,,IIg. -

'Vmi.—Ordinary amalgam, AggHp, (silver S&i p. c.) or AgHg (silver 3G'oj; alto AgiRgi, etc.

Arquenal*. Ag., Hg (Silver 86-6); G. = 10-8; malleable and soft. Kongt-bergiu. Ag., Hg or Ag., Hg.

Pyr., ato.—B.B. on charcoal the mercury volatilizes and a globule of silver is left. In the closed tube the mercury sublimes and condenses on the cold part of the tube in minute globules. Dissolves in nitric acid. It is rubbed on copper it gives a silvery luster,

Oba.—From the Palatinate at Moschellandsberg; at Friedrichsaege, Nassau; from

by Google

Bala. Sweden; Eongaberg, Norway; AUemoDt, Dauphin^e; Almadeu, Bpoiu; Ctili; Vitale Creek, Br, Columbia (argentea).

Tin. Native tin has been reported from several localities. The only occurrence fairly above doubt is that from the watershed of the Clarence river, near Obiia, New South Wales. It has been found here in grayish-white rounded grains, with platinum, iridosmine, gold, coelestite, and corundum.

Platinum Group.

Isometric. Crystals rare; usually in grains and scales.

Cleavage none. Fracture hackly. Malleable and ductile. H. = 4-4.5. Q. = 14-19 native; 21-33 chem. pure. Luster metallic. Color and streak whitish steel-gray; shining. Sometimes magnetically,

Comp.—Platinum alloyed with iron, iridium, osmium, and other metals.

It is only slightly magnetic. O. = 16-6-18-0 molecular.

f. ! _ _ j '---'-' ■ ^ aa polarity,

B ,,,, _ . j p. , , ,,,, u-platinum,

platinum 6erm. \ although this description does not hold without exception.

Pyr., sto.—B.B. infusible. Not affected by borax or salt of phosphorus, except in the state of fine dust, when reactions for iron and copper may be obtained. Soluble only in heated nitro-hydrochloric acid.

Dist.—Distinguished by its color, malleability, high specific gravity, infusibility and

its grains, associated with iridium, T Pinto, in the district of Choco, in (platina del Pinto) from p'

—ial mulerial in the Ural at Ni

Tagiisk. and with chromlle in a Berpentlne probably derived from a perldotite; also ii Qoroblagodatsk district. Alao found on Borneo; In New Zealand, from a region cbarac-terized by a cbryaollte rocl with serpentine; in New South Wales, in the Broken Hill district, and in gold washings at various points.

Id California, In tbc KlamnlU region, at Cape Blanco, etc., not abundant; in the gold washings of Cherokee, Butte Co.; ut St. Francois. Beauce Co., Quebec; at several points in Briiteh Columbia.

Iridium. Platin-irldlum. Iridium with platinum and other allied metals. Occurs usually in aneular grains of a silver-white color. H. = 6-7. G. = 23'6-83-8. With tha platinum of the Urala and Brazil.

IRIDOBMINB. Osmlrldlnm.

Rhombohedral. Usually in irregular flattened grains.

Cleavage: c perfect. Slightly malleable to nearly brittle. H. = 6-7. G. = 19-3-21-12. Luster metallic. Color tin-white to light steel-gray. Opaque.

Comp., Tar.—Iridium and osmium in different proportions. Some rhodium, platinum, ruthenium, and other metals are uenally present.

Var.—1. Ifwj/arukiU. H. = 7; G. = 18-8-19-5. In flat scales; color tin-white. Over 40 p. c. of iridium. 2. BUvunkiU. In flat scaleii, often sii-Bided, color grayish while, steel-gray. G. = a>-8rS. Not over 80 p. c. of Indium. Less common than ihe light-colored

Dii—Distiaguiahed from platiuum by greater hardness and by Its lighter color.

Oba.—Occurs witli platinum in South America; in the Ural mountains; in 'auriferona drift in New South Wales. Rather abundant iu the auriferous beach-Bands of norUiem California,

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KATIVE ELEUIiNTS.

PaUadinm.—Pslilium, alloyed willi a little plallnum H.=4-5-5. G. = 118-11-8. ColorwL-'' '' bIso from tbe Urn is.

Isometric. Uaoftlly massive, rarely in crystals.

Cleavage: a perfect; also a lamellar etruotnre Q o and | d. Fracture hackly. Malleable. H, = 4-5. Q. = 7-3-7-8. Luster metallic. Color steel-gray to iron-black. Strougly magnetic.

Vkt. —1. TMTMtilal Iron. Found Id nuisaes, occ&aloDally of Kre&t'Bize, as well as Id small embedded panicles. Id busalt at Blaafjeld, Ovifak (or Ulfnk), Disko Isliiud, West Greeoliudi also elsewhere on the aame coast. Tills Iroo codIhiub 1 to 3 p. c. of Nl. Some oher occuireDces. usually classed ns meieoric, miy be la fact terrestrial; «.g., the Sitota Calhiirina iron of Brazil discovered lu t87S.

A nickellferoiis metallic iron (FkNIi) called awarut(« occurs lu Ihe drift ot tbe Gorge river, which empties into Awarua Bay on the west coast of the south Uluud of New Zealand; Associated with go!il, plalluum. ctissilerile, chromEle; probably derived from a partially serpentInteed peridotlte. Joatpkinilt is a uickel-Irou (FeiNii) from Oregon, occurring In stream gnivel. Native Iron also occurs sparingly lu some basalts; reported from gola or platinum washings at various points.

H. H«t»orio Iron. Nntlve iron a1>o occurs to most meteorites, forming in some cases (a) the entire mass (iron meUoritet-, also (b) os a spongy, cellular matrix in which are embedded grains of chrysolite or other sillicai miidtroliit*); (e)in grains or scales disB«mtuated more or leas freely [hrouglioul aalony miklrix {meUorie ...

tlone»). Rarely a meteoite consists of a single cryaial ^"-

Hue individuil (Briuuau) wlili nunieniis twlnulng lamellffi | o. Cubic cluiiVHT'e gometlmesobserved; also an oclahedral. less oflen dodecnliednil Inmellar alnic-iiire. Etlciliig wiil dilute nitric acid (or Iodine) commonly di'vrlopa a crystalline stnlnlure (called Wtdmanttdtten figuru) (F!g. o88); usually cousisiug of tines or bands crossing at various nnctts according lo the direction of the seeKon. at Go° if \o. W \ n. etc. They are formed by the edges of ervsinlline plates, usually I o. of tbe nickeliferoiis iron of different composition {kamaeiie, taniU, pUuiU), iia sliiiw[i by the fact that they are differently attacked by the acid. Irons with cubic structure Hiid with twinning lamells

\t.g., Brauuau) hare a aeries of fine lines correspond I

log lo those developed by etching (Neumann lint»\ /ii ^ . -u. -u u '

A d.n«K»™ lu,i., I. alio prod.fcil I., «.»i. cuei, "'»'<"* Ml- H." Mex., o

due to quadrilateral depressions. Borne irons show no distinct crystalline structure upon

The exterior of masses of meteoric iron is usually more or less deeply pitted with rounded IbumbUke depressions, and the surface at llie time of fall la covered wiib a film of iron oxide in fine ridges showing lines of flow due to the mellinc caused by iliu heat developed by the resistance of the air; this film disappears when the Iron Is exposed lo the weather.

Meteoric iron is always alloyed with nickel, whkh is usually present In amounts varving from 5 to 10 p. c.. sometimtES mucli more; small amounts of other metals, aa cobalt, manganese, tin. copper, chromium, are also often present. Occluded gases cnu usually be

detected. Graphite, in seams or nodules, also troilite (iron sulphide), schreiblerite (iron-nickel phosphide) are common in masses of meteoric iron; diamocd, daubreelite, etc., are rare. Cohenite, sometimes identified, is (Fe,Ni.Co)₈C In tin-white crystals

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DESCRIPTIVE MINERALOGY.

n. SULPHIDES, SKLENIDES, TELLURIDES, ARSENIDES, ANTIMONIDES.

The Sulphides, etc, fall into two groups according to the character of the positive element.

I. Sulphides, Selenides, Tellurides of the Semi-Metals. n. Sulphides, Selenides, Tellurides, Arsenides, Antimonides of the Metals.

I. Sulphides, etc., of the Semi-Metals.

This section includes one distinct group, the Stibnite Group, to which orpiment is related; the other species included stand alone.

HZ1AL.OAR.

Monoclinic Axes $a:b:c = 1.4403 : 1 : 0.9739$; $\beta = 66^\circ 5' . 88$ mm³, $110^\circ A \text{ to } 105^\circ 84' . rr'$, $012 \text{ a } 013 = 47^* 57$.

j^i Crystal short prismatic; striated vertically. Also granular,

// * l^l coarse or fine; compact; as an incrustation.). $^L^t$] Cleaves: b, rather perfect. Fracture small conchoidal.

; Y ;] 1 Sectile. fl. = 1⁵-2. Q. = 3⁵⁵⁶. Luster resinous. Color I I • I) aurora-red or orange-yellow. Streak varying from orange-red to aurora-red. Transparent—translucent. $^ ' / \}$ Comp.—Arsenic monosulphide, $A_2S = \text{Sulphur } 39 '9$, arsenic

$V' J' 70-1 = 100$.

$V . \sim \wedge \wedge \wedge 3^1$ ato.—In the closed tube, volatilizes, and gives a tins.

Natural parent red sublimes in the open tube (if heated very slowly) sulphurous

fumes, and a white crystalline sublimate of arsenic trioxide. B.B. on charcoal burns with a blue flame, emitting arsenical and sulphurous odors. Soluble in caustic alkalis.

Obs.—Orpiment associated with orpiment: occurs with ores of silver and lead, at Felsobány and Kapnik, Hungary; Joachimsthal: Schneeberg; Andreasberg; Blumenthal,

Switzerland, in dolomite: near Jul mine in the U. S., in Iron county, Utah; also in California, San Bernardino Co.; Trinity Co., in the Norris Geyser Basin, Yellowstone Park, as a deposit from the hot water. The native mineral is from the Arabic name *al-ghasr*, powder like the mint.

Monoclinic.* Axes $d : X : c = 1.2061 : 1 : 0.6743$, $\beta = 90^\circ$. approx.

Crystals small, rarely distinct. Usually in foliated or columnar but sometimes with reniform surface.

Cleavage: b highly perfect, cleavage face vertically striated; a in traces; gliding-plane c (001). Sectile. Cleavage laminae flexible, inelastic. H. = 5-6. Q. = 3-4-5. Luster pearly on b (cleavage); elsewhere resinous. Color lemon-yellow of several shades; streak the same, but paler. Subtransparent—subtranslucent.

Comp.—Arsenic trisulphide, As_2S_3 , = Sulphur 39.0, arsenic 61.0 = 100.

The fine crystals from Hercules, Utah, are distinctly

by Google

SULPHIDES, SELENIDES, TELLURIDES, ARSENIDES, ANTIMONIDES 283

Properties.—In the closed tube, fuses, volatilizes, and gives a dark yellow sublimate; other reactions as under realgar. Insoluble in aqua regia and caustic alkalis.

Diff.—Distinguished by its fine yellow color, pearl; luster. easy cleavage, and flexibility when in powder.

Obs.—Occurs in small crystals in veins at Tadjona, in Upper Bulgaria; in foliated and fibrous masses. In Moldavia in the Banat; in Kapnik and Felsöbány in the ironiferous veins; also in the Melfalura near Naples. Near Julmerek in the district of a large Turkish mine. Occurs with realgar in some compact clay beneath lava in Iron county, Utah; also finely crystallized at Hercules. Among the deposits of the Steamboat Springs, Nevada; also in the

Stibnite

Bismuthite

Native

Stibnite Group.

Bi₂S₃, Bi₂Se₃,

0-9926 : 1 : 10179

0-9679 : 1 : 09850

1 : 1 approx.

The species of the Stibnite Group crystallize in the orthorhombic system and have perfect brachydiagonal cleavage, yielding flexible laminae.

The specimen's appearance is in physical properties somewhat related to stibnite, but known to be monoclinic in crystallization. Orol notes that the oxide, As_2O_3 , is monoclinic like stibnite, while the corresponding compound, Sb_2O_3 (valentinite), is orthorhombic; further he remarks on the relation in form and physical characters between appearance and structure.

STIBNITE. Antimonite, Antimony Glance, Gray Antimony, Antimonoglauz Gmtn. Orthorhombic. Axes $a:b:c = 0.9936 : 1 : 1.0179$.

α , no $A_{110} = 89^\circ 84'$. Ill $A_{111} = 71^\circ 341'$. $118 A_{118} = 85^\circ 621'$. $118 A_{113} = 86^\circ 88'$.

$60, 010 A_{131} = 88^\circ 8'$.

fur. $010 A_{853} = 40^\circ 104'$.

$6r, 010 A_{848} = 46^\circ 88'$.

bp, $010 A_{111} = 51^\circ W$.

Crystals prismatic; striated or furrowed vertically; often curved or twisted (cf. p. 148), Common in con aoo. rbi koq

Massive aggregates or radiating groups of acicular crystals; massive, coarse or fine columnar, less often granular to impalpable.

Cleavage: b highly perfect'. Slightly sectile. Fracture small subconchoidal. H. = 2. G. = 4-5.2-4-6.2. Luster metallic, highly splendent on cleavage or fresh crystalline surfaces. Color and streak lead-gray, inclining to steel-gray: subject to blackish tanish, sometimes iridescent

Comp.—Antimony trisulphide, $9b, S = \text{Sulphur } 28'6, \text{ antimony } 71'4 = 100$. Sometimes auriferous, also argentiferous.

pyr[^] "te[^] Fuses very easily; (at 1), coloring the flame greenish blue. In the open tube

a>

California.

Hungary.

D-volatile. On charcoal fusae, spreads <

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fume fl. and cont. B. The color while with oxide of Antimony; this compound treated in R. F.

volatilizes and takes the name of blue. When pure perfectly soluble in hydrochloric acid: in dilute acid decomposed will separate a quantity of antimony peroxide.

Diff.—Distinguished {e.g., from stibic acid} by cleavage, color. Softness: also by its fusibility and other characters. It fuses like graphite. Resembles some lines certain of the rarer sulphurites of lead, but yields to lead on charcoal.

Obi.—Occurs in quartzite beds or veins in granite and gneiss, often accompanied with various other minerals produced by its alteration. Also associated in metalliferous deposits with sphalerite, galena, zinc, barite, quartz; sometimes accompanies native gold.

Occurs at Wolfenbuttel, in the Harz; Brunsdorf, near Freiberg; Freiberg; Caiparizecbe. near Arnsberg, Weglphalin; Fel»Sli6uya, Hungary; In Cornwall abundant also abundant in Borneo; in Victoria and New South Wales. Groups of large splendid crystals have come from the antimony mine in the Province of Iyo. island of Fihikobu. Japan.

In the United States occurs as a vein of some extent in Sevier county, Ark.; in California at San Emigdio. Kern county, and near Altu, Benito Co.; in the Humboldt mining region in Nevada; in Iron county, Utah. In New Brunswick in Prince William. York county, 30 m. from Fredericton; in Rawdon township, Hants Co.. N. S.

Materially. An amorphous brick-red deposit of antimony trisulphide, Sb₂S₃, occurring in small quantities and associated with sulphide upon siliceous sinter at Steamboat Springs, Washoe Co., Nevada.

BISUDTUINITB. Bismuth Glance. Wismuthglanz Qwm. Orthorhombic. Rarely in acicular crystals, $\alpha = 88^\circ 8'$. Usually massive, foliated or fibrous.

Cleavage: 6 perfect. Somewhat sectile. H. = 2. G. = 6.4-6.5. Luster metallic. Streak and color lead-gray, changing to tin-white, with a yellowish or iridescent tarnish. Opaque.

Composition.—Bismuth trisulphide, Bi₂S₃, = Sulphur 188, bismuth 81 = 100. Sometimes contains a little copper and iron.

Properties etc.—Fusibility = 1. In the open tube sulphurous fumes, and a white sublimate which fuses into drops, brown while hot and opaque yellow on cooling. On charcoal at first

gives sulphurous fumes; then fuses with spirting, and coats the coal with yellow bismuth oxide; with potassium iodide a bright red coloring of bismuth iodide is obtained. Dissolves readily in hot nitric acid, and a white precipitate falls on diluting with water,

Obs.— Found at Brandy Gill, Carrock Fells, in Cumberland; near Rednith, etc. In Prance at Meymac, Corrdze: at Jolianugeorgeuscadt, Schneeberg; at Wittichen, Baden; at Riildnrliyitan, Sweden; near Borata, Biilivla.

In the U. S., occurs with gold in Kowan Co., N. C., at the Bamhrdt vein; sparingly at Willmantic, Conn.; abundant in Beaver Co. Utah.

OnanaJuatiU. Frenzele: Selenwismnthtflanz Qerra. Bismuth selenide, Bi₂Se₃, sometimes with a small amount of sulphur replacing selenium. In acicular crystals; also massive, granular, foliated or fibrous. Cleavage: B distinct, H. = 2-5-8-5. G. = 6.25-6.63. Luster metallic. Color bluish gray. From the Santa Catarina mine, near Guanajuato, Mexico.

TRADYMIT. Tellurium bismuth Germ.

Rhombohedral. Crystals small, indistinct. Commonly in bladed forms, foliated to granular massive.

Cleavage: basal, perfect. Laminae flexible: not very sectile. H. = 1.5-2;

soils paper, G. = 7.2-7.6. Luster metallic, splendent. Color pale steel-gray.

Comp., Tar.— Consists of bismuth and tellurium, with sometimes sulphur and a trace of selenium; the analyses for the most part afford the general formula Bi₂(Te,S)₃.

Bi₂Te₃ =

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SULPHIDES, SULFIDES, TELLURIDES, ANTIMONIDES, ANTIUNITES. 285

SO 0 ~ 100. This is the more common variety and includes the Uiradymite of Uildinger. Id crystals from Schubkai.

Fusion.— In open tube a white sublimate of tellurium dioxide, which fuses to colorless drops. On charcoal fuses, gives white fumes, and entirely volatilizes; in the R.F. bluish green; coats the coal at first white (TeO₂), and finally orange-yellow (Bi₂O₃); some give sulphurous and arsenical odors.

Obs.— Occurs at Schubkau near Bismoltz; Rezbunya; Orwilza in the Banat; Telle-mark

in Norway; Bastuaes mine, near RiddHrhyliin, Sweden. In the U. S., in Virginia at the Whitelall gold mines, Spotsylvania Co.; in Davidson Co., N. C. and Idaho the gold washings of Burke and McDowell counties, etc.; similarly in Honlaur. At the Uonlgomery mine, Arizona. Named from its twofold, in allusion to complex twin crystals sometimes observed.

JoalUL—A bluish telluride ($\text{Te } 80 \text{ p. c.}$, also 8 and 8«). $Q. = 7'9$. San José, Bnull.

WshUte. A foliated bismuth telluride ($\text{Te } 80 \text{ p. c.}$) of doubtful formula. $O. = 8'4$. Deutsch-Pilsen, Hungary.

MOLTIUDDHmi. HoljbdadglDZ Oerm.

Crystals hexagonal in form, tabular, or short prisms slightly-tapering and horizontally striated. Commonly foliated, massive or in scales; also fine granular.

Cleavage: basal eminent. Laminae very flexible, but not elastic. Sectile. $H. = 1-1'5$. $G. = 4'7-4'8$. Luster metallic. Color pure lead-gray; a bluish gray trace on paper. Opaque. Feel greasy.

Coup.—Molybdenum disulphide, MoS_2 , = Sulphur 400, molybdenum 60*0 = 100.

Pyr., ato.—In the open tube sulphurous fumes and a pale yellow crystalline sublimate of molybdenum trioxide (MoO_3). B.B. In the forceps infusible, imparting a yellowish color to the flame; on charcoal the pulverised mineral gives in O.F. a strong odor of sulphur, and coats the coal with crystals of molybdic oxide, yellow while hot, white on cooling: the assay the coating is copper-red. and If the white coating be touched with an Intermittent R.F., it assumes a beautiful azure-blue color. Decomposed by nitric acid, leaving a white or grayish residue.

DUt—Much resembles graphite in softness and structure (see p. 273), but has a bluish trace on paper and readily yields sulphur on charcoal.

Oh«.—Generally occurs embedded in, or disseminated through, granite, gneiss, zircon-syenite, ferruginous limestone, and other crystalline rocks. At Numedal, Sweden; Arendal and Laurvik in Norway; Allenberg, Saxony; Zinnwald, Bohemia; near Miask, Urals; Chessy in France; in Italy, at Traversella; Corrock Fells, in Cumberland; at several of the Cornish mines.

In Uaint. at Blue Hill Bny; In Conn., at Haddam, in gneiss; In Vermont, at Newport; In S. Hampshire, at Westmoreland; in JV. Tark. two miles from Warwick; In Penn., In Chester, near Reading; near Concord, Cabarrus Co., N. C. In Canada, at St. Jérôme, Quebec: in large crystals in Renfrew county, Ontario; also In Aldfield township, Pontiac Co., Quebec.

Named from the oxide of lead; the name, first given to some substances containing lead, later included graphite and molybdenite, and even some compounds of antimony. The distinction between graphite and molybdenite was established by Scheele in 1778.

n. Sulphides, Selenides, Tellurides, Arsenides, Antimonides of the Metals.

The sulphides of this second section fall into four divisions depending upon the proportion of the negative element present. These divisions with the groups belonging to them are as follows:

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A. Basic Division. B. Monosulphides, Monotellurides, etc., MS , RS , etc.

1. Galena Group. Isometric, normal group.

2. Chalcocite Group. Orthorhombic

3. Sphalerite Group. Isometric-tetrahedral.

4. Cinnabar—Wurtzite—Millerite Group. Hexagonal and rhombohedral.

C. Interstitial Division. Embraces Melonite, Te_2S_3 ; also Bismite, $3Cu_2S \cdot Fe_2S_3$; Linnsite, $CoS \cdot 2FeS$; Chalcopyrite, Cu_5FeS_4 ; etc.

D. Bisulphides, Diarsenides, etc., BS_2 , BA_2 , etc.

1. Pyrite Group. Isometric-pyritohedral.

2. Arsenic Group. Orthorhombic.

A. Basic Division.

The basic division embraces several rare basic compounds of bismuth or copper chiefly with antimony and arsenic. Of these the crystallization of dyscrasite only is known.

DTBOHABTTB. Antimonite As_2S_3 .

Orthorhombic. Axes $a:b:c = 0.5775 : 1 : 0.6718$. Crystals rare, pseudo-hexagonal in angles ($\alpha = 60^\circ 1'$) and by twinning. Also massive. Fracture uneven. Sectile. H. = 3-5-4. G. = 9.44-9.85. Luster metallic. Color and streak silver-white, inclining to tin-white; sometimes tarnished yellow or blackish. Opatjae.

Comp.—A silver antimonide, including $Ag_3Sb = \text{Antimony } 27 \times 1, \text{ silver } 72 \times 9 = 100$, and

Ag.Sb = Antimony 15-7, silver 84-3 = 100, and perhaps other compounds.

Analyses vary widely, some conforming also to Ag₂S, Ag₂(Sb,As)₂, etc. By some authors classified with cinnabar.

Pyr., etc.—B.B. on charcoal fuses to a globule, coating the coal with white antimony trioxide and finally giving a globule of almost pure silver. Soluble in nitric acid, leaving antimony trioxide.

Obs.—Occurs near Wolfach, Baden; Wittichen; Andreasberg in the Harz; Allemont, France. Named from *dv<TK/jdtrti*, a bad alloy.

Homocite, a white massive copper antimonide. probably Cu₂Sb (Sb 24 p. C). G. = 8.8. Asia Minor, near Mytilene.

Homocite, Asiatic. Occurs from Silver Islet. Lake Superior, apparently contains silver arsenide (Homocite, Ag₂As) and perhaps also a silver antimonide (Homocite, Ag₂Sb?), the latter related to or identical with dyscrasite.

Homocite.—Copper arsenide, Cu₂As. Resinous and brittle; also massive, disseminated. G. = 7.2-7.7. Luster metallic. Color tin-white to steel-gray, readily scratched. From several Chilean mines; also Zwickau. Baxony. In N. America, with cinnabar at Michipicoten Island, L. Superior.

Homocite. Copper arsenide, Cu₂As (As 16-6 p. c); Q. = 7.12. Resembles domeykite. From Cinnabar; also L. Superior.

Homocite. Copper arsenide, Cu₂As (As 16-6 p. c). Q. = 7.12. Color pale reddish white. From Houghton Co., Michigan; Sonora, L. California.

Homocite. Perhaps Ag₂S. Copiapó, Chili.

Homocite*. A rare silver telluride (Ag₂Te?). Probably from Nagy-Hegy.

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SULPHIDES, SULPHIDES, TELLURIDES, ARSENIDES, ANTIMONIDES. 287

B. Non-sulphides, Non-sulphides, etc., B¹, Galena Group. Isometric or orthorhombic. A¹

Also,

Ag₂S

(Ag₂Cn)₈

Ag.Te

Ag.Se

(Pb,Cn,)8, (Cu,,Pb)S Jalpaite

Altaito PbTe Eeuit«

ClautliaUta PbSe .^TulUrtl* Hftomuuiitfl (Ag.,Fb)Se

The following, knoTD oni; in masaive form, probably also belong here: Benelianite Cn,Se
Zorgite (Pb,Cii.,Af,)Se f

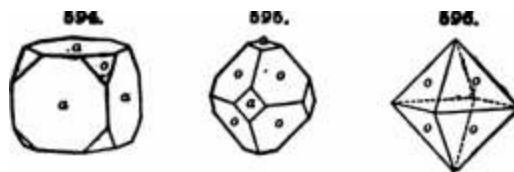
Lehrbaekite (Fb,Hg,)Se Crookedt* (Cn,Tl,Ag),Se

XooairiU Cn,Se.Ag,Se

The Galbna Qboup embrace a nnmer of monosnlphidea, etc., of the related metals, eilver, copper, lead, and mercnry. These crystallize in the normal group of the Isometric Bygtem, and seTeral show perfect cubic cleavage.

These cEaracters are most distinctly exhibited i QAIiBHA, or Oalehitb. Leul glance. Blelglan:

the type apeoiee, galena.



Isometric. Commonly in cubes, or en bo-octahedrons, lesa often ootahedral. Also in skeleton crystals, reticolated, tabular. Twins: tw. pi. o, both contact-and penetration-twins (Figa. 363, 366, p. 123), sometimes repeated; twin crystals often tabular || o. Also other tw. planes giving polysynthetic tw. lamelise. Massive cleavable, coarse or fine granular, to impalpable; occasionally fibrous or plumose.

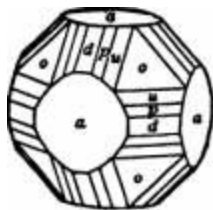
Cleavage : cubic, highly perfect; less often octahedral Fracture flat sub-conchoidai or even. H. = 2-5-2-75. G. = 7'4-7'6. Luster metallic. Color and streak pure lead-gray. Opaque.

Comp., Tar.—Lead sulphide, PbS = Sulphur 13-4, lead 8G'6 = 100. Often contains silver, and occasionally selenium, zinc, cadmium, antimony, bismuth, copper, as sulphides; besides, also, sometimes native silver and gold. Var.—1. Ordinary. (i) OrystnlllEed; (ft) umewhat flbrouB and pliimoBe; (c) cleavable, gniiiular coarge or fine; (d) ciypto-crysinlline. The variety wub ocUhedml clenavage la rare; the uBual cubic cleavage Is obtained readily after heating to 200° or

300*: the peculiar cleavage may be connected with the blamulh p(aBl), u(&S4)

uaiiallj preunl.

a. Argstiftrmt'. Ali galena li more or lesa argenUferous, and no eileinal characten BCTTe to dlstingulsi the kinds that are much so from those that are not. The silver la detected by cupellatioo, and may amount from a few thousandthB O* one per cent to one per cent or more; when mined for atlver tt ranks as a *iJwr on.



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8. CoDtalnlDK nnenlc, or antimoiiy, or a compouad of thiee metals, as Impurity. Here belong biniehv^ from CkiuLbal wIoi o-8*2 8I>, and tUinmanMU fnim Pflbram, wllh both

BTbeuiu uud uuiiiiioiiy.

Pyr,—lu the open tube etrea sulphuTOUS fumea. B.B. on charcoal fuses, emits sulphoi'ous fumes, coats the coal yellow near the assay (PbO) aud wblte vitb a bluish border at a distant'^ (PbyOt, cliieUy). and yields a globule of metallic lead. Decompoaed by stroDi; DJtrlc add with the separation of some sulphur and Che formatloD of lead sulphate.

TiJi. —Distinguished, except la very flue graular varieties, by Its cubic cleavage; the colui' aod the hiKh specitic gnvity are cbsracleristic; also the blowpipe reactions.

Oba.—One of the most widely distributed of the metallic sulphides. Occurs iu beds and veins, both in crystallue and un crystal line rocks. It is often assoctaied with pyrile, marcaste, sphalerite, chalcopyrite, arsenopyrlte, etc.. In a gangue of quartz, caldle, Wile or fliiorlte, etc.; bIho with cerussite, anglesile, and other salts of lead, whieh are frequent results of Its alteration. It Is also common with gold, and In veins of silver ores.

Al Freiberg In Saiony it occupies veins in gneissi in Spain, in Eranlte at Linares, also In Calalonia, Oreiiada, and elsewhere; at Clausthal and Neudorf In the Hai-z. and at Pflbram in Bohemia, It forms veins iu clay slate; similarly in Styria; at Sala In Sweden iu veins lu granular limesloue; through the graywacke of Leadhills and the klktus of Curuwall, In veins; SlUna cnvities In the Subcarbonlferous limestone In Derbyshire, Cumberland, and the northern districts of Eugland; also at Bleiberg, Cartntbia. In the English mines it Is associated with calcite, pearl spar, Quorite, barite, wiibeilte, calumlne, hucl sphalerite. Other localities are Joacbimstbal, Bohemia; Puullaouen and Huelgoet, Brittany; Sardinia;

Ifurchinsk, Bast Siberia; Australia; Chill; BoUvie, etc.

Extensive deposits of this ore in the United States exist in Missouri, Illinois, Iowa, and Wisconsin. The ore occurs not in veins but filling cavities or chambers in stratified limestone, of different periods of the Lower Silurian, especially the Trenton, also in part Silurian carboniferous. It is associated with sphalerite, smithsonite, calcite, pyrite. The Missouri mines are situated in the counties of Washington, Jefferson, Madison and others. Good crystals are obtained at Joplin, Jasper Co. Also occurs in New York, at Ross, St. Lawrence Co., in crystals with calcite and chalcopryite. In Maine, at Lubec, etc.; in Mass., at Soniampton, Newburyport, etc; in Penn. at Phenixville and elsewhere; in Virginia, at Austin mines in Wythe Co., and other places; in Tenn., at Hartsboro. near Nashville; in the Lake Superior copper district and on the N. shore of Lake Superior; in California, at many of the gold mines; in Nevada, abundant in the Eureka district; in Arizona, in the Castle Dome, Eureka, and other districts. In Colorado, at Leadville there are productive mines of argentiferous galena, also at Georgetown, the San Juan district and elsewhere. Mined for silver in the Coeur d'Alene region in Idaho; also at various points in Montana.

The name galena is from the Latin *galena* (from *galena*), a name given to lead ore or the dross from melted lead.

Cu₂PbOPLUUBiTR. A massive mineral, from Chill, varying in characters from galena to those of calcoelite and covellite; composition, $Cu_2Pb_3S_4$. Alunite is massive, deep indigo-blue quickly tarnishing; corresponds to $8Cu_2S \cdot PbS$. From Hina Grande. Cleavage. Whether these and similar minerals represent definite homogeneous compounds, or only ill-defined alteration-products, is uncertain, and if so it is not clear whether they should be classed with isometric galena or with orthorhombic calcoelite.

Altaita. Lead telluride, $AgTe$. Rarely in cubic crystals, ... cleavage. G. = 816. Color tin-white, with yellowish tinge tarnishing to bronza-jellot. From the Altai, with hessite; Coquimbo, Chill; California; Colorado

Olaosthalia. Lead selenide, Pb_3Se . Commonly in fine granular masses resembling

Bilena. Cleavage: cubic. G. = 7.8-8.0, Color lead-gray, somewhat bluish. From the Harz, at Clausthal, etc.; Coblenz, Mendoza. S. A. **TVUm** is a cobaltiferous variety. **Haunumita.** Silver-lead telluride (Ag_2PbSe). In cubic crystals; also massive, granular, in thin plates. Cleavage: cubic. G. = 8.0. Color and streak Iron-black. From Tilkerode in the Harz.

AROBNTITEI. Silver Olanze. Silberglanz Qerm.

Isometric. Crystals often octahedral, also a, o; often distorted, frequently grouped in reticulated or arborescent forms; also fibrous, massive; embedded; as coating.

Cleavage: a, d in traces. Fracture small subconchoidal. Perfectly sectile.

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BULPHIDBS, SBLBinDEB, TBLTTBIDBS, ABSEITIDES, AimUOlflDBS. 2l}9

H. = 3-3-5. G. = 7-30-7-36. Later metallic Color and streak black Uh lead-graji streak
Bhiiiiig. Opaqua

Comp.—Silver Sulphide, Ag.S = Sulphur 13-9, silver 87-1 = 100.

PyTi., 010.-^111 the open tube gives off sulphurous fume, B.B, and charcoal fuses with
Intumescence la O. F., umltliDg sulphurous fume, and yields a globule of silver.

DLE—Disintegrates from other sulphides by being readily cut with a knife; also by
yielding metallic silver on charcoal.

Obs.—Found at Freiberg, Joachimsthal, etc.; Schemnitz, Hungary; In Norway near
Eongsberg; in the Altai; in Cornwall: Peru; Chili; Mexico Huauajuato, etc.

Occurs in Nevada, at the Coalsack lode; at the Silver King mine, Arizona; at mines near
Port Arthur on the north shore of Lake Superior; with native silver and copper in
Michigan.

It is a cupreous arsenite from Jalpa, Mexico.

Hvadt. Silver telluride. Ag₂Te. Isometric. Usually massive, compact or fine-grained.
Cleavage indistinct. Brittle. E. = 3.5-S. Q. = 8-31-8 46. Color Between lead-gray
and steel-gray. From the Altai; at Ma.ET&g in Transylvania; Rezbanya, Hungary; Chili,
near Aiquero, Coquimbo. In the U. S. Calaveras Co., Cal.: Boulder Co., Colorado; Utah.
This species also often contains gold and thus graduates toward

Patsite. (Ag, Au)₂Te with Ag; Au = S: 1. Heavy; granular to compact, slightly sectile
to brittle. H. = 5-8. G. = 8-7-903, Color steel-gray to iron-black; tarnishing. From Nagy,
Transylvania; Colorado; California.

Aguarita. Silver argentide, Ag₂S and Ag₂(S, Se). In skeleton dodecahedral crystals. Sectile.
O. = 7.380. Color iron-black. From Guajuato, Mexico.

Benlianite. Copper selenide, Cu₂Se. In thin dendritic crusts and disseminated. G. = 0.71.
Color silver-white to tarnishing. From Skrikerum, Sweden; Leirbäck, in the Hue.

Laurensite. Selenide of lead and mercury, PbSe with Bi₂Se₃. Heavy, granular. O. = 7.8.
Color lead-gray to iron-black, From Leirbäck, in the Harz.

It is a silver selenide, massive, granular, Q. = 7.50, Color between silver-white and lead-

gray. From the Bkrkerum copper mine, Sweden; also Chill.

Zorgite, Selenide of lead and copper In varying amounts. Massive, granular. $O = 7-7'5$. Color dark or light lead'grsy. From the Harx; Cacheula, Argentina.

Orookedte. Selenide of copper and thallium, also silver (1-B p. c), $(Cu, Tl, Ag)_i Se$. Massive, compact. $O = 6'9$. Luster metallic. Color lead-gray. From the mine of Skrikerum, Sweden.

Umangita. $CuSe$ Cu, Se . Massive, flne-granular to compact. $H = 8$. $G = S'SaO$. Color dark cherry-Ted. From La Rloja, Argentina.

a. Chalco«lte Group.

The Bpecies of the Chalcocitb Group crystallize in the orthorhombio flystem with a prisiiatic angle approximating to 60° ; thej are hence pBendo-hexagonal in form especially when twinned. The group is parallel to the Oalena Gronp, since Cn , appears in isometric form in coproplnbite and Ag, S also in argentite. Some authors include dyscrasite here (see p. 286),

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DB8CEIPTIVE MINEBALOGY.

OHAiiOOOrrB. Copp«r Qlance Redruthite. Kupterglaoz Q«rm. Ortnorhombic. Axes (i:h:& = 0-5822 :1: 0*9701.

M, 031 A Oai = 135° aff. pif". in a hi = 68' 8f.

Crystals psendo-hexagonal in angle, also by twinning (tw. pi. m). Often massive, atracture grannUr to compact and impalpable.

Cleariige : m indistinct. Fracture conchoidal. Rather brittle. $H = 2-5-3$. $G = 5\ 5-5-8$. Luster metallic. Color and streak blackish lead-gray, often tarnished blue or green, dull. Opaque.

Comp.—Caprona sulphide, $Cu, S = Snlphur\ 20-2$, copper $79-8 = 100$. Sometimes iron in small amonnt is present, also silrer.

PyT,eto.—lu ttae open tubu glvea sulphurous fumes. B.B. on charcoiil melts lo a globule, which boila with spirtiug; ihe fine ponder roasttd at a low tempernture on charcoal, then healed iu R.F., yields a globule of melallic copper. Soluble In nitric actd.

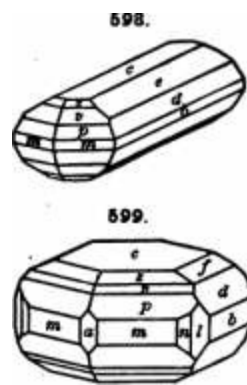
DiA—Ri-BembleB argeultc but much more brittle; bomfte , baa a different color au the fi'esh fmcture and becomes mag-uelli' B.B.

Obs,—Coruwall aSorda splendid ciyatals, esfielall; tbe

districts of Bnint Just, Camborne, and Bedruth (rtdruthiU).

Occurs at Juacblmstlial, Bohcmin; Tellemarken, Norwar;
compact arid mftssive varieties Iii Siberia; Saiouy; Ml. Catfnl
mines Id Tuacnny; Meiico; S. America.

In the U. S., Bristol. Conn., has afforded larjre and brilliant crysUls; also found at
Bimsbury and Cheshire; at Schuyler's mines, N. J.; in Nevada, in Washoe. Humboldt.
Churchill and Nye counties: in Moolana, massive at Butte City. Pouud in Canada, wllh
chalcopyrile and bornite at the Acton mines and elsewhere in the province of Quebec.
Stromarerita. (Ag,Cu),8, or Ag,a.Cu,8. Rarely In orthorbombc crystals, often twinned.
Commoulv mssaive. compiicl. H, = 2-5-8. G. = 615-6-8. Luster metallic. Color and streak
dark sleel-gmy. From the Zmeinoitorsk mine. Siberia; Silesia; also Chili; Zacatecas,
Mexico; Ihe Helntzelmaa mine in Arizona; Colorado.



BTERNBHRarm,

Orthorhombic Crystals tabular | c. Commonly in fan-like aggre^tions; twins, tw. pi. m-
Cleavage: c, highly perfect. Thin laminaa flexible, like tin-foil. H. = 1-1-5. G. = 4-215.
Lnster metallic. Color pinchbeck-brown. Streak black. Opaque.

Comp. AgFe,8, or Ag,S.Fe,S. = Sulphur 30-4, silver 34-2, iron 35-4 = 100.

Oba.—Occurs with pynirgyrlte and stepbanlte at Jonchlmsihal, Bohemia, and Johann-
georcenatadt, Saxony.

FRIB8B1TK. Near slernberglte. In thick tabular crystals: H. = 3-6; Q. = 4-32. Color dark
gray. Composition AgiFeiS>. Occurs with marcasite at Joachlmsthal.

Acaothit*. Silver sulphide, Ag,S, like argentite. In slender prismatic cryslals
(orthorhombic). Sectlle. G. = T3-7-3. Color Iron-black. Occura at Joachlmstbal; alsi> at

Freiberg, Saxony, and at Schneeberg.

It has been suggested that scanthlie may be only argentite in distorted Isometric crystals.

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8DLPH1DE8, SELENIDE8, TELLURIDES, ARSUNIUES, ANTIMONIDBS. 291

Sphalerite Group. RS. Isometric-tetrahedral.

Sphalerite ZnS Onofrite $Hg(S,Se)$

Hetaoimiibaiite HgS Alsbandite MnS

Guadalcazarite $(Hg,Zn)S$ Coloradoite $HgTe$ MaaeiYe

Tiemaimite $HgSe$

The Sphalerite Group embraces a number of sulphides, Belenidea, etc, of zinc, mercury, and manganese. These are isometric-tetrahedral in crystallization.

SPHAL. EIIUTB, Zinc Blende or Blende. Block-Jack, Hock-Lead. False Oaleiu Enfl. Zbkblenile Oerm.

Isometric-tetrahedral- Often in tetrahedrons. Twins common: tw. pL o; twinning often repeated, sometimes as polysynthetic lamellae. Commonly

600. Ml.



m = (311) massive cleavable, coarse to fine granular and compact; also foliated, sometimes fibrous and radiated or plumose; also botryoidal and other imitative shapes. Cryptocrystalline to amorphous, the latter sometimes as a powder.

Crystal: dodecahedral, highly perfect. Fracture conchoidal. Brittle. H. = 3-5-4. G. = 3-9-4-1; 4-063 white, N. J. Luster resinous to adamantine. Color commonly yellow, brown, black; also red, green to white, and when pure nearly colorless. Streak brownish to light yellow and white. Transparent to translucent. Refractive index high: $n_D = 2.362$ Xa.

Comp.—Zinc sulphide, $ZnS = \text{Sulphur } 33, \text{ zinc } 67 = 100$. Often containing iron and manganese, and sometimes cadmium, mercury and rarely lead and tin. Also sometimes

contains traces of indium, gallium and thallium; maybe argentiferous and auriferous.

V«r.—1, Ordinary. ContnlDlng little or no Iron; from colorless white to yellowish brown, sometimes green: O. ^J'CM'l. The red or reddish-brown tranipnrent crystnjllzed liitids are eometiraes called ruby bUiide or tTifty tine. The mnaslve cleavable forms are the most common, varying from coarse to fine granular; also cryptocrystailme. fkliaUabUrKU (Germ.) Is a closely compact variety, of a pale liver-brown color. In concentric layers with reiiiiform surface; galena and marcaslle nre often inters!ratified. Tbe fibrous forms (/iuitw< Zinkiiende Qemi.) are chiefly wurlzile. A aofl white amorphous forui of zinc sulphide occurs Id Cherokee Co., Kausas.

3. Ferriferovt; Marmatitr.. Contalninc 10 p. r. or more of lr"D: dark brown lo black; G = 8-9-4'05. The proportion of FcB to Zd8 varies from 1 : 5 to 1:2, and the last rstio is that of the chritlop/tiU of Brethnupt. a brilliant.black sphalerite from St. Chrstoplie mine, al Brelienbninii. havinc O. = 8-91-!t'(l23.

-t Cadm^ermu; PrUirimit'. Pititiramite. The amount of cadmium present in any sphalerite thus far analyzed Is less llian S per cent.

;;CoogIc

Pyr, etc—Difficultly [iiaible. Il the open tube sulphurous fumw, and generallj cbnnges colour. B.B. on cliarcoftl, In K.F.. tome vaiietles giTe At Sta » reddl^-brovrn coalIDE o[cadiiiiDi oxiJe, aud lal«r a cuatiDg of doc oxide, nhtch Is yellow nbllc hot and white after ciollug. With cobalt Bolution tbe zisc coating gives a green color when heMed In O.F. HoBt viirieilea, after roasting, slve with borei n reaction for Iron. With soda on charcoal iu B.F. a slroug grecD zlc name. Dissolves lu hydrochloric add with «TolutiOD of hj'drueu sulphide.

tUS. —Varies widely In color and appearance, but dlatlngulsbed by the resinous luster in all but deep bliicti varieties; UBually eiblbils dUlnct cleavage; much softer than garnet; neail'y Infusible B.B.i yields a zloc coating on charcoal.

Obs.—Occurs very commonly in both crystalline and sedlmentarj rocks, aod as a frequDi associate of galeoa; also associated with cbalcopyrlio, barlie, fluorite. siderlie: common in silver mines. Jt often forma lieda of cousidemblo magnitude fllling cavities In limestone.

Some of the chief localities for crystallized sphalerite are: Alston Moor In Cumberland, black variety; Derbyshire, St. Agnes and elsewhei'e in Cornwall; Oberlalinat«in in Nassau. £ms, red; Audreasberc, fellow aud brown; Neudorf iu the Harz, Freiberg, Brelteubrunn, aud other localities In Saxony, black and brown; Pi4bram, green or yellow, and Schlackenwald in Bohemia, black; Eapnik, Hungary, green or yellow; Nagyig, lYansylvania, brown; Bodna. black; the BInuenlhal In Switzerlaud, Isolated crystals of great beauty, yellow lo brown, In cavities of dolomite; Bula In Sweden; Nordmark, bliick, brown, also siiow-

while. A beautiful translucent variety yielding large cleavage masses is brought from Picos de Europa, Santander, Spain, where it occurs in a brown limestone. Fibrous varieties (see wurtzite) are obtained at Irbach; Gersroldseck in Baden; Rolbel; also in Cornwall. The original marmatite is from Harmato near Piave, Italy.

Abounds with the lead ore of Hillmourf, Wisconsin, Iowa, and Illinois. In New York, Sullivan Co., near Wurtzboro'; in St. Lawrence Co., at Mineral Point with galena; at the Aucrsm lead mine in Columbia Co.; in lime-toué at Lockport. In Massachusetts, at the Southampton lead mines. In N. Hamp., at the Calton lead mine; at Warren. In Maine, at the Lubec lead mines, etc. in Conn., at Roxbury. In S. Jersey, a white variety [cleioplakite] at Franklin Furnace. In Penn., at the Whealley and Perkiomack lead mines, in crystals; near Flederbville, Lehigh Co., a white waxy var. In Virginia, abundant at Austin's lead mines, Wythe Co. In Michigan, at Prince vein, Lake Superior, abundant. In Utah, near Rosclaire, with galena and calcite; near Deer's ditches, near Galena. In staurolite, with cryst. marmatite, and galena. In Wisconsin, at Mineral Point, in fine crystals. In Tenn. at Hajiboro', near Nashville. In Mittavri, in beautiful crystallizations with

Sienna, marcasite and calcite at Joplin and other points in the southwestern part of the state; 9 deposits here occur in limestone and are of great extent and value; also in adjoining parts of Kansas.

Named blende because, while often resembling galena, it yielded no lead, the word in German meaning blind or deceiving. Sphalerite is from vipulxioi, treacherous.

Hatacinuabari. Mercuric sulphide. HgS . In composition like cinnabar, but occurs in tetrahedral crystals; also massive. G. = 8.19. From the Bedlington mine. Lake county, California, with cinnabar, quartz and marcasite; also elsewhere sparingly.

Oodalouarite, near metacinnabarite. but contains zinc (up to 4 p. c), Guadalcazar, Mexico.

Tinnantite. Selonqueckite. Mercuric selenide. $HgSe$. Isometric-tetrahedral. Commonly massive-compact. G. = 8.19. Utah; SBO-847 Clausthal. Luster metallic. Color black to blackish lead-gray. Streak nearly black. Occurs in the Harz; California, in the vicinity of Clear Lake; Marysvale, Piute Co., Utah.

OnoMt*. $Hg(B.Se)$ with 8e = 4 to 6 p. c. Ssn Onofre, Mexico; Marysvale, Utah.

Coloradoite. Mercuric telluride, $HgTe$. Massive. G. = 8.8. Color iron-black. Colorado.

Alabandite. Manganese blend. Manganoglanz. Manganese sulphide, MnS . Isometric-tetrahedral; usually granular massive. Cleavage: cubic, perfect. Q. = 8.5-404 Luster submetallic. Color iron-black. Streak green. Occurs at Nagrftg, Transylvania; Karoluh, Hunpury; Mexico; Peru; crystallized and massive on Snake River, Summit county, Colorado; Tombstone, Arizona.

Oldhamite, Calcium sulphide, CaS. In pale brown spherules with cubic cleavage in the
Bust! meteorite.

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SULPHIDES, SELENIDES, TELLURIDES, ARSENIDES, ANTIMONIDES. 303

FEHTI.ANZ}mi. Else Dickel Uw Om A.

isometric. Mass, granular. Cleavage: octahedral. Fracture uneven. Brittle. H. = 3-5.6.
= 4-6. Luster metallic. Color light bronze-yellow. Streak light bronze-brown. Opaque. Not
magnetic.

CvMp.—A sulphide of iron and nickel, (Fe,Ni)S. In part, 2FeS.NiS = Sulphur 36%, iron
10%, nickel 22% = 100.

Occurrence.—Occurs in the Chatopite Mine, Lillhemmer, Horthay. The mineral from Sudbury,
Ontario, is mostly pure for nickel; it shows distinct octahedral cleavage (or
partly). The same locality also affords nickeliferous pyrite and pyrrhotite.

Properties.—Ferrous sulphide. FeS, occurring in nodular masses and in thin veins in mafic
meteorites. O. = 4.7483. Color to black-brown. By some authors regarded as identical
with pyrrhotite (p. 2M).

4. Cinnabar-Wurtzite-Heulandite Group. Rhombohedral or Hexagonal

Cinnabar HgS Rhombohedral-Trapezohedral 1-1453

Covellite CuS 1-1466

This fourth group among the monosulphides includes several subdivisions, as shown in
the scheme above, and the relations of the species are not in all cases perfectly clear. It is
to be noted that the sulphides of mercury and zinc, already represented in the sphalerite
group, appear here again.

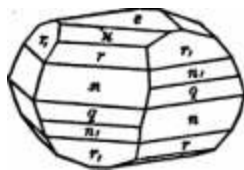
If, as suggested by Orin, the prominent pyramids of wurtzite, greenockite, etc., be made
pyramids of the second series (x = 1132, instead of 1011). then the values of h in the
second column are obtained, which correspond to Miller's. The form of several of these
species, however, is only imperfectly known. A rhombohedral form for greenockite has
been suggested.

CINNABAR. Zinnober, Schwefelquecksilber Q. Rhombohedral-trapezohedral. Axis c =
1.1453.

Crystals usually rhombohedral or thick tabular in habit, rarely showing trapezohedral

faces; also acicular prismatic. In crystalline in cru station a, granular, massive; so:..etimeB as an earthy coating.

Cleavage: nt perfect. Fracture sub- K-(:oi4), «(2(^1), y (0441) conchoidal, uneven. Somewhat sectile. H. = 2-2-5. Q. = 8*o-8'2. Luster adamantine, inclining to metallic when



o^

,ab,GoOgIc

dark-colored, and to dull in friable Tarieties. Color cochiiiiieal-red, often incliDiug to brownish red and lead-gray. Streak scarlet. Transparent to opaque. Optically -|- . Indices: $a^{\wedge} = 2-854$, $e^{\wedge} = 3'201$, Dx. See Art, S66.

Tm-~1. Ordinary: eilher {a) eryttatUted; {b) mauivt, grBDular embedded or compact; bright red to roddUli bronru 1q color; (e) wirtAy aud biighi red. 2. Hepatic. Of a liver-brown color, wltb soni«tini«s a browoUti alKak. occaaiuuually duty fu siructun, tliough coidiiiudIj grauular or coupact.

Comp.—Mercuric sulphide, HgS = Sulphnr 13-8, mercury Se-S = 100. Usually impnre from the admixture of clay, iron oside, bitumen.

Pyr.—In tbe closed lube aloue a black sublimate of mercuric sulphide, but with aodlum carbouate oue of muialUc mercury. CarcrullT beated bi tbe open tulie gives sulphuroua fuuiea and melallic mercury, whIcL coadenBea id mluule glubuleit ou Ibe cold walla of tlie tube. B.B. on cliarcnal wbolly volatile, but imly when quite free fri)in gaiiKue.

Dlft—CliaracterlEed by lis color and TermllfoD streak, bigli apeclQc gravUy (reduced, however, by tbe gangue usually present), softness; alau by the blowpipu cliatacters {e.g., la the closed tube). Resembles some varielleB of bematlte aud cuprite.

Oba.—Occura clifeHy in veins In slate rocks and sbales, and nirely in granite or porpLyry. Il has been observed in veins, wilh ores of iroD. Sometimes occurs in couiieclloii with bot apringa as the result of solfataric action. Pvrite and marcualle. siilpbtdes of co|iper. ;llb-nite, realgar, i^old, etc., are associated uilnerala; calcite, quariz or opil, niso burite, fluorite, are gangue mmemlei a biiimlnous mineral is common.

The most important European deposits are nt Almaden in Spain, and nl Idria in Car-nfola. where it is usualIv masiive; also nt Bakmut In southerti Russiil. Crystallizeil at Hoscbeltandsberg aud Wolfstein in the Palatinate and at the mines of Ht. Avala. tiear

fielgtade, Servia: at Rlpa In TuBcany; In the Urals; tbe Nerclilnsk region la Transbaikal; tu China: Japan; Meilco; Huaucaveltca, Peru; Chill.

Id the U. 8. forms extensive inines Iq California, Ibe moat Important at New Almadea and thevlcinlly, in Santa Clara Co.; also at Altoonn, TrinltyCo.; il Is now forming by eolfataric action at Sulpbur Bank, Cal., and Steamboat SprTogs, Nevada; also occurs In Boutbern Utah.

The name cinnabar is supposed to come from India, where It is applied 1o (he red resin, dragon's blood. The native cinnabar of Theopbnistus is true cinnnbar; he speaks of its affording quickailver. The Latin name of ciDDabar, minium, is now given to red leiui, a substance which waa early used for adulteratlag cinnabar, and so got at last the name.

OoTolUta. Eupferiodlg Oarm. Cupric sulphide. CuS . Rarely in hexagonal c^stals. Commonly maisive or spheroidal. $G. = 4'59$. Color indigo-blue or darker. From Baden-weller, Baden; Hansfeld, Thuriogia; Vesuvius, on lava; Chili, etc.

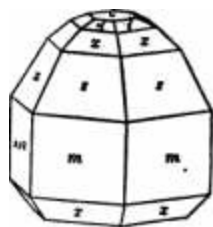
OREENOOKrm.

Hexagonal-hemimorphio. Rarely in Lemimorphio crystals; also as a coating.

OleaTage: a distinct, c imperfect. Fracture conchoidal. Brittle. $H.=3-3-5$. $\langle 0 \rangle$. $G. = 4-9-5-0$. Luster adamantine to resinous. Color

honey-, citron-, or orange-yellow. Streak-powder between orange-yellow and brick-red. Nearly transparent. Optically +,

Coup.—Cadmium sulphide, $CdS = \text{Sulphur } 22-3, ' \blacksquare 77-7 = 100$.



Pyr., etc.—In the closed tube assumes a carmine-rod color while hot, fading to the original yellow on cooling. In the open tube gives Hulpliirous fumes. B.B. on cbaTcoal. either alooe or with soda, gives in R.F. a reddish-brown coating. Soluble In hydrochloric acid, affording hydrogen sulphide.

Obi.—Occurs with pr«hnile at Bishoplton, Kenfrewahire, and elaewhere lu Scotland. At Pflbram In Bohemia, aa a coating on sphalerite; similarly at other points; so too In the U. 8. near Friedensville, I^, and In Ibe line region of south-westerti Missouri: in Marion Co.. Ark., It colore amlibBonlte bright yellow. Not un-a furnace product.

SULPHIDES, SELBHIDS8, TBLUBIDS8, ABERNISBS, ANTIUONIDES. 295

Wnzolte. Zinc sulphide, ZnS, \like spbiJerlU, but In hcmlmorplilc hexagonal cryaula; also fbroua and massive. O. = ii'98. Color brownUh black. From a siWer-mbe Dear OruTO lu Bolivia: PortugBt; Peru. la crystals nitb ipbalerite imd quartz at tUe "Orlglual Butte" mine. Butte City, MoDtuna.

TLe luu^lye fibrous forms of " ScbalenbleDde " occur at PHbram, Llskeard, etc. Other forniB, from Siolberg, Wieelocb, Attenberg, are iu part wuitiile, iu part spbalerits.

ERTTHROiUKCiTB is (Zu,Mu)S. From jjlberia.

MIZ.IJ1RITE. Capillary Pyrites. Eaarkies Germ.

Rhombohedral. TjBiially in very sleiider to capillary crystalB, often in delicate radiating groups; Bometimes interwoven like a wad of hair. AUo in columnar tnfted coatings, partly eemi-f^lobular and radiated.

Fracture nneven. Brittle; capillair crjetals elastic. H. = 3-3*5. G. = 5-3-5'65. Luster metallic Color brass-yeltow, inclining to bronze-yellow, with often a gray irideecent tarntsb. Streak greenish black. Comp.—Nickel sulphide, NiS = Sulphur 35-3, nickel 64-7 = 100.

Pyr., etc,—In the open tube sulpLiii-ous fumea. B.B. on charcoal fuses to a globule. When roiiisied, gives with borax and salt of phosphorus a violet bead in O.P., becoming gray in R,F. from reduced metallic nickel. On charcoal !u R.F. the roasted mineral gives a coherent metallic mass, attr«ctible by the magnet. Most varieties also show traces of copper, cobalt, and Iron witli the fluxes.

Oba.—Occurs commonly lu capillary crystals, In the cavities and among crystals of other mlnei-als. Fouod at Joachimslhui lu Bohemia: JobanngedStadl; rabram; Rlocbelsdorf; Audreasberg; Freiberg. Saxouy; Cornwall.

In (lie U. 8., nt Antwerp, N. T., iu cavities In hematite; in Lancaster Co.. Pa , at the Gap mine, lu thin velvety coatings of a radiated fbrous structure. "Wiih calcite. dolomite and fluorite, forming delicate tungled balr.like tufis, lu geodes in limestone, often pene- Iratiug the calcite crystals, at St. Louis, Mo.; similarly near Milwaukee. Wis. At Orford, Queb«j.

Betrichitb. NIS like mlllerlte witli also Co.Pe. From Westerwald. The relation of the two species is doubtful.

HARCECOKRiTK. Perhaps NI(Bi,Sb,S). In tabular tetragonal crystals. H. = o. Q, =: 6-4, Color light bronze-yellow. From Hamm a. d. Bleg.

NIOOOOTB. Copper Nickel. Eupfernickel. Rothnickelklea Qerm.

Hexagonal. Crystals rare. Usually mnaBiYe, astructure nearly impalpable; also reniform, columnar; reticulated, arborescent. Fracture uneven. Brittle. H. = 5-5'5. G. = 7-33-7'67. Luster metallic. Color pale copper-red. Streak pale brownish black. Opaque.

Comp.—Nickel arsenide, MiAs = Arsenic 56-1, nickel 439 = 100. Usually contains a little iron and cobalt, also sulphur; sometimes part of the arsenic is replaced by antimony, and then it graduates toward breitbanptite. The intermediate Varieties have been called ariie.

Pyr., etc.—In the closed tube a faint white crystalline sublimate of arsenic trioxide. In the open tube a sublimate of arsenic trioxide, with a trace of sulphurous fumes, the assay becoming yellowish green. On charcoal gives arsenical fumes and fuses to a globule, which, treated with borax glass, affords, by successive oxidation, reactions for iron, cobalt, and nickel: the antimonial varieties give also reactions for antimony. Soluble in aqua regia.

Ob.—Accompanies cobalt, silver, and copper ores in the Saxon mines of Annaberg. Schneeberg; etc.; also in Thuringia. Hesse, and Styria; at Allemont, Dauphine; in the Kongsberg mines in Norway; Sweden: at Balen in the Basses Pyrenees (Ariete); occasionally in Cornwall; Chili; abundant at Mina de la Rioja, Orlocha, Argentina. In the U. S., at Chatham, Conn., in Knelss; sparingly at Franklin Furnace, N. J.; Silver Cliff, Colorado; Tilt Cove, Newfoundland.

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Z96 DBSOBIPTIVE HIHERALOGT.

Br^uUuiqitlts. Antimonienickel ore. Nickel Antimonide, NiSb. natural crystals; usually massive. arborescent, disseminated. ~ red. From Andreasberg in the Hartz,

' PTBaHOTITD. HftgUftic Pfrites. Hagnelkfa Gem. Hexagonal. 6 = 0 8701.

a. OOOLAOil = «' 8*.

cu, 0001A4M1 = 78' ff.

cj/. OOOLA(200Sft8) = 81* »!'. Twins: tw. pi. s, with vertical axes nearly at right angles (Fig. 383, p. 125). Diatomic crystals rare, commonly tabular; also a few pyramidal with faces striated horizontally. Usually massive, with granular structure.

Parting: c, sometimes distinct. Fracture uneven to subconchoidal. Brittle. H. = 3'5-4'5. G. = 4'58-4'64. Luster metallic. Color between bronze-yellow and copper-red, and subject to speedy tarnish. Streak dark grayish black. Magnetic, but varying much in intensity; sometimes possesses polarity.

Comp.—A sulphide of iron, often containing also nickel; formula chiefly Fe_3S_4 ; analyses, however, vary from Fe_8S_8 , up to $Fe_{10}S_{10}$, while conforming to the general formula Fe_3S_4 . Percentage composition Fe_8S_8 = Sulphur 38.4, iron 61.6 = 100; Fe_9S_9 = Sulphur 39.6, iron 60.4 = 100; $Fe_{10}S_{10}$ = Sulphur 39.3, iron 60.7 = 100.

s fumes.

high oxide,

small

acid, with evolution of

—Distinguished by its peculiar reddish-bronze color; also by its properties.

Obs.—Occurs at Kongsberg, Hønefoss, etc., in Norway; Falun, Sweden; Adirondack; Breitenbrunn; Jáchymov, Bohemia; Nizhny Tagil, Siberia; Minas Geraes in Brazil, in large tabular crystals; the lavas of Vesuvius; Cornwall.

In N. America in Maine, at Standish with antimony: in Vermont, at Stafford, etc. In Florida, near Diana, Lewis Co.; Orange Co. In Pennsylvania, at the Gap mine. Laocaster Co. nickeliferous. In Tennessee, at Ducktown mines, abundant. In Canada. In large veins at St. Jerome, Elizabeth town, Ontario; at Sudbury, etc. This species is often mined for the nickel it contains.

Pyrrhotite is often present in disseminated particles or crystals in meteoric stones; the iron sulphide of meteoric irons is generally referred to troilite (p. 208).

Named from Fe_3S_4 ; i.e., Fe_3S_4 .

C. Intermediate Division.

Horobachite. An iron-nickel sulphide, perhaps $4Fe_3Si_2Ni_6S_{16}$. Crystalline, massive. Color pinchbeck-brown to steel-gray. O. = 443. From Ilorbach in the Black Forest.

Polydymite. A nickel sulphide, perhaps Ni_3S_4 . In octahedral crystals; frequently twinned. G. = 4.54–4.81. Color gray. From Grönau, Westphalia.

A nickel ore from Sudbury, Ontario, corresponds to Ni_3FeS_4 , conforming to the general formula of polydymite; another Sudbury ore agrees with pentlandite (p. 293), and still another is a nickeliferous pyrrhotite,

OrUnanite. Contains sulphur, bismuth, nickel, iron; perhaps a mixture. From Grönau.

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Halonlto. Tellunilckel Oerwt. A nickel telluride, Ni_3Te_2 , In Indistinct sraoular and (olUted particlea. Color r«cidl«h white, with matallc tiuter. From the StaaU&iu mioe, California; probably also in Boulder Co.. Colorado.

The following Bpeoiea are aoraetitnu regarded aa 8alpho-«alta, namely, Snlpho-femtes, etc

BORNim. Buntkuplererz Otrm. Purple Copper Ore. Variegated Copper Ore. Erubeacte.

Isometrio, Habit cubic, faces often rough or oarred. Twine: tw. pi. o, often penetration-twins. AtassiTe, atrnctnre grannlar or compact.

Cleavage: o in traces. Fracture Bmall conohoidal, uneven. Brittle. $H. = 3$. $G. = 4'9-5'4$. Luster metallic. Color between copper-red and pinchbeck-brown on freeh fracture, speedily iridescent from tarnuh. Streak pale grayish black. Opaque.

Comp., Tar.—A sulphide of copper and bon, but vjiryng in the proportions of these metals. The cryetallizea mineral agrees with $Gn_2FeS_3 =$ Sulphur 28-1, copper 55-5, iron 16'4 = 100; this may 1m written $30a_2S_3re_3S$, (Groth) or $Cn_2S.CuS.Fe_8$ (Rg.).

Analyses of massive Taretiea give from 50 to 70 p. c of copper and 15 to 6'S p. c. of IroD. The varIntloQ Ib due, in part at least, to mechanical admixture, chiefly of chalcecte.

Fjrr., etc.—In lbs closed tube gives a faint aubllmate of sulphur. In the open tube yields sulphurous fumes, but no sublimate. B.B. on charcoal fuses inR.F, to a brittle magnetic globule. The roasted mioeral gives with the fluzes the reaclioug of iron and copper, and with soda a metallic globule. Soluble in nitric acid with sepanttiou of sulphur.

DUt—DistiDeilshed (e.g. from chnlcoclle) by the peculiar reddish color on the fresh fracture aed by Its brtlliant tarnish; B.B, becomes stro'nffly magaetlc.

Ob*—Occura with other copper ores, and is a valuable ore of copper. Crjstallne varletieg are found in Cornwall, called by the miners " horse-flesh ore." Occurs massive St Ross Island. Killsrney, Ireland: Monte Catlnt, Tuscany; the Hansfeld district, Oer-maey: In Norway, Sweden, Siberia, Silesia, and Hungary. It is the principal copper ore at some Chilian mines; also common in Peru, Bolivia, and Heiico.

In tlio U, 8., found at the copper mine in Bristol, Cnna,; massive at Uaboopeny, near Wilkesbarre, Peon. A common ore in Canada, at the Actoo and other mloes.

Named after the minerali^ist Ignatius toQ Bom (1743-1791).

Xiinnalts. A sulphide of cobalt, $Co_2S_3 = CoS.Co_6i$, analogous to the spinel group. Also

contains nickel (var. Htgeniu). Commouiy in ociahedronsi also massive. H. = 55. o. = 4'S-5. Color pale steel-gray, tarnishing copper-red. Occurs at Basuaea, etc., Sweden; MQsen, near Sie^en, Pnissia; at Slegen (tie^tenitej. In ootahedrous. In tbe V. 8., at Hlne la Hotte, Hlssoui'i; Hiiieral Hill. Maryland.

OaobrSoIlta. An Iron-chromium sulphide, FeS.CriBi, occurring with tt^lite in some meteoric irons. Color black. G. = 601. <

Onbanita. An Iron-copper sulphide, perhaps CuFegBi = CuS.FeiSi. Isometric; massive. Color tietweeu bronze- and brass-yellow. 6. = 4'08-4'17. From Banacanao, Cuba; Tunaberg and Kafveltorp, Sweden.

CarroUlt*. A copper'Cobalt sulphide, CiiCo,Si = CuS.CoiBi. Isometric; rarely In octahedrons. Usually massive. G. = 485. Color light steel-gray, with a faint reddish hue. From Carroll Co., Maryland, near Finksburg.

OHAiiOOFTBITB. Copper Pyrites. Yellow Copper Ore. £upferkies o«rm. Tetragonal-sphenoidal. Axis d = 0*98525. pp'. 111 A ill = 108* 40'. pp,, m A III = 70* 7i'. «, 001 A 101 = 44* 84i'.

Crystals commonly tetrahedral in aspect, the sphenoidal faces p large

, small and brilliant. Sometimes both forms equally

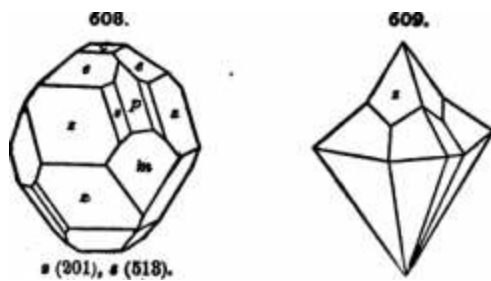
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DESCRIPTITE UINEBALOaT.

developed, and then octahedral in form. Twins: (1) tw. pi. p (111), reaembling spinel-twins (Fig. '679, p. 135); sometimes repeated ae a fiveling (Fig. 609). (2) Tw. pi. andcomp.-face e (Fig. 381, p. 125) often in repeated twins. (3) Iw. pi. m, tw. axis c, complementary penetration-twins. Often massive, compact.

Cleavage: 2 (201), sometimes distinct; c, indistinct. Fracture nneveD. Brittle. H, = 3-5-4. G. = 4-1-4-3. Luster metallic. Color brass-yellow; often tarnished or iridescent. Streak greenish black. Opaque.





Comp — A sulphide of copper and iron, CuFeS_2 , or $\text{Cu}^{\text{I}}\text{Fe}^{\text{II}}\text{S}_2$. = Sulphur 35.0, copper 34.5, iron 30.5 = 100. Analyses often show a ratio different from this formula, often due to mechanical admixture of pyrite.

Sometimes auriferous and argentiferous; also contains traces of selenium and tellurium.

Pyrite, Bto.—In a closed tube decrepitates, and gives a sulphurous sublimate, in the open tube sulphurous fumes. On charcoal fuses to a magnetic globule; with soda the roasted mineral gives a globule of copper containing iron. The roasted mineral reacts for copper and iron with the fluxes. Dissolves in nitric acid, excepting the sulphur, and forms a green solution; in excess changes the green color to a deep blue, and precipitates red ferric hydroxide.

DUE—Distill off from pyrite by its inferior hardness and deeper yellow color. Resembles gold when disseminated in minute grains in quartz, but differs in being brittle and in having a black streak; further it is soluble in nitric acid.

Obs.—A widely disseminated mineral in metallic veins and nests in gneiss and crystalline schists, also in serpentine rocks; often intimately associated with pyrite, also with alderite, tellurite, etc., sometimes with nickel and cobalt sulphides, pyrrhotite, etc. Observed coated with tetrahedrite crystals in parallel position, also as a coating over the latter.

Chalcopyrite is the principal ore of copper at the Cornwall mines: there associated with cassiterite, galena, bornite, chalcocite, tetrahedrite, sphalerite. At Falun, Sweden, it occurs in large masses embedded in gneiss. At Rammelsberg, near Qoslar in the Harz, forms a bed in argillaceous schist; occurs with nickel and cobalt ores in the Eupferschiefer of Mftcsfeld. The Kurprinz mine at Freiberg affords well-defined crystals; also Hor-hniisen, Dillenburg, Neidorf, Hesse; Schlackenwald in Bohemia. Common elsewhere as at Monte Callio in Tuscany; in New South Wales; Chill, etc.

It is found at the Lubec mines and elsewhere. In Vermont, at St. Albans, etc. In Massachusetts at the Southampton lead mines. In Conn., at Bristol. In New York, in crystals and massive at Allegheny, Ulster Co. In Pennsylvania, at Phenixville; at the French Creek mines, Chester Co., with pyrite, magnetite, etc. In Maryland, near Finksbury, Carroll Co., abundant. In Virginia, at the Pheolx copper mine*, Fauquier Co., and the Walton gold mine, Louisa Co. In N. Carolina, near Greensboro', abundant massive. In Tennessee, 30 miles from Cleveland, in Polk Co., in Missouri, with sphalerite at Joplin, Jasper Co.

In C<U.. In different mines along a belt between Mariposa Co. and Del Norte Co., on west side of, and parallel to the chief gold belt; occurring massive in Calaveras Co., and in crystals on Jomingo Creek, etc. Abundant in Montana, near Butte, with bornite,

pyrite, etc., also at other points, often argentiferous and auriferous. In Colorado, abundant in Olin, Boulder, Otero, and Gunnison counties, etc.; commonly associated with pyrite.

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SULPHIDES, SELENIDES, TELLURIDES, ARSENIDES, ANTIMONIDES. 399

tetrahedrite, sphalerite, and often highly argentiferous. Also mined in Arizona, Utah, but for most cases chiefly for silver and gold. Grant Co., New Mexico.

In Canada, in Penobscot and near Sherbrooke and at many points in the eastern part of the province of Quebec: in the Nipissing distr., Ontario, at various points; extensively mined at Sudbury; at the Bruce mines, on Lake Huron; at Fostat-au-Mines and elsewhere on Lake Superior.

Named from the affixes, bran, and prefix, by Hencel (1728).

D. Disulphides, Diarsenides, etc.

The disulphides, diarsenides, etc., embrace two distinct groups. The prominent metals included are the same in both, viz.: iron, cobalt and nickel. The arsenic present, therefore, in several cases of isodimorphism, as is shown in the lists of species below. These sulphides are all relatively hard, H. = 5-6; they become brittle with a steel, and this has given the familiar name pyrites applied to most of them. The color varies between pale brass-yellow and tin-white.

Pyrite Group. FeS , $FeAs_2$, $FeSb_2$. Isometric-pyritohedral.

Pyrite FeS , Oenontofite NiS , $NiAB$,

Hanerite MnS , Corynite NiS , $Ni(Aa, Sb)$,

Smaltite $CoAs_2$, also $(Co, Ni)As_2$, Ullmannite NiS_2 , $2NiSb_2$ (isometric-tetrahedral) Chloantfalte $NiAs_2$, also $(Ni, Co)AB$, Sperrylite $PtAs_2$ [tetrahedral]

Cobaltite CoS , $CoAs_2$, Lanrits BuS ,?

Marcaalte Group. BS_2 , RAs_2 , etc. Orthorhombic.

The Pyrite Group includes besides the compounds of Fe, Co, Ni, also others of the related metals Mn and Pt. The crystallization is isometric-pyritohedral.

The species of the Mabcasitb Group crystallize in the orthorhombio system with prismatic angles of about 70" and 110° and a prominent macro-dome of about 60° and 130°. Hence fivefold and sixfold repeated twins are common with several species, in the one case the prism and in the other the macrodome named being the twinning-plane.

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^1

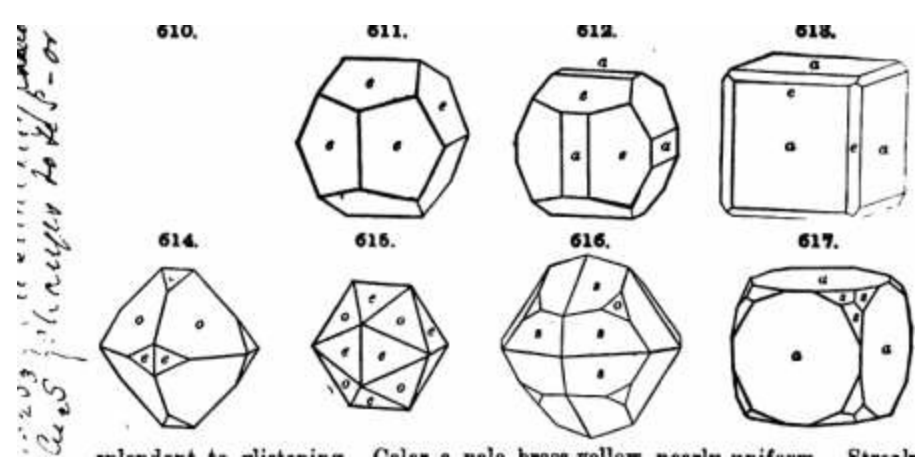
DESCBIPTITE HIKKRALOOT.

Pyrite Group.

PTKITB. Schwelelklcs, Eisenkies, Germ, Iron Pyrites.

iBoinetric-pyritoliedrttl. Cube and pTritobedron e (210) the common forms, tfae facee of both often with striations { edge a/e, due to oscillatory combination of these forms and tending to produce rounded faces; pjritohedral faces also striated J. to this edge; octahedroD also common. See Figs. 610-617, ako Figa, 117-123, pp. 44,45, Twins: tw. ai o, usuallypenetration-twins with parallel axes (Fi^ . 369, p. 124); rarely contact-twins. Freouently massive, fine granular; sometimes subfibrous radiated; reniform, globular, stalactitic

Cleavage: a, o indistinct. Fracture conchoidal to uneven. Brittle. H. = 6-6-5. O. = 4-95-5-10; 4-967 Traversella, 5-027 Slba. Luster metallic, «11. 612. 618.



splendent to glistening. Color a pale brass-yellow, nearly uniform. Streak greenish black or brownish black. Opaque.

Comp., Tar.—Iron disulphide, FeS, = Sulphur 53-4, iron 46-6 = 100, ^^ Nickel, cobalt, and ihallum, and aliiio copper !□ small qitautiities, sometitDes replace part

^vC of the iron, or else occur as mixtures; seleiiiuin is sometimes preaeal in traces, Qold U ,^ sometimes distributed iuvisibly tbrough it, auriferous pyrkte being an important source of -T gold. Arsenic Is rarely present, as in octahedral crystals from French Creek,

Pcdd. (O*3 p. c. As).

Pyr., etc.—In the closed tube a sublimate of sulphur (sni) a magnetic residue, B.B. on charcoal gives off sulphur, burning with a blue flame, leaving a magnetic residue which reacts like pyrrhotite (p. 290). Insoluble in hydrochloric, but decomposed by nitric acid,

Diff—Distinguished from chalcopyrite by its greater hardness and paler color; in form and specific gravity different from marcasite, which has also a violet color.

Obs.—Pyrite occurs abundantly in rocks of all ages, from the oldest crystalline to the most recent alluvial deposits. It usually occurs in small cubes, pyrohedrons, or in more highly modified forms; also (often with marcasite) in irregular spheroidal nodules and in veins, in clay slate, arenaceous sandstones, the coal formation, etc,

Fine crystals have been found in some of the Cornish mines; also in great variety with hematite on the island of Elba, and with magnetite at Traversella and Brossio in Piedmont. Other localities for crystals are Huelmo near Sleggen; Freiberg; Baxony; Schneeberg; Waldenstein in Carinthia; Příbram, Bohemia; Sebnitz, Hungary; Persberg, Falun, and Långban in Sweden; Kongsberg in Norway.

In Maine, at Peru, Waterville, etc. In N. Hampshire. at Unity, massive. In Massachusetts, at Rowe, Hawley, massive. In Vermont. at Shoreham. In limestone; Hartford. In Conn., at Roxbury. finely crystallized. In N. York, at Robb's, fine crystals; at Schoharie; at Chester, Warren Co.; in Otsego Co., at Warwick; massive in Franklin, Putnam, and Orange Cos.,

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SULPHIDES, SELBNIDE, TELLURIDE, ABSENIDE, ANTIMONIDE. 301

etc. In Pennsylvania, at Chester, Delaware Co.; at Kearsburg, Chester Co.; at French Creek mines, octahedra and other forms, sometimes tetragonal or orthorhombic in symmetry; Gorham, Lebanon Co.; in S. Car., near Greensboro', Guilford Co., in crystals. In Colorado, crystals near Central City, Clear Fork Co., and elsewhere. Auriferous pyrite is common at the mines of Colorado, and many of those of California, as well as in Virginia and the States south. In Canada, 2 miles N. W. of Brockville, Ontario, a cobaltiferous

Large quantities of massive pyrite are mined at the Rio Tinto and other mines in Spain, also in Portugal. Among important deposits in the U. S. are those at Rowe, Mass.; Herman, St. Lawrence Co., and Ellenville, Ulster Co., N. Y.; Tolersville, Louisa Co., Va.; Dalhart, Paulding Co., Ga.

The name pyrite is derived from *pyr*, fire, and alludes to the sparks from friction; hence the early name *pyrit* (p. 309).

Pyrite readily changes to an iron sulphate by oxidation, some sulphur being set free. Also to limonite on its surface, and upward throughout, by the action of a solution of bicarbonate of lime carrying the sulphuric acid as change proceeds, and from limonite to red iron oxide.

Hannite. Magnesian disulphide, MgS_2 . In octahedral or pyritohedral crystals; also massive. Color reddish brown or brownish black. From Elnka, Hungary; Kaddusa, Catania. Sicily.

SHALITZI-OHZ. OANTHIT. Speiskobalt Otrm.

Isometric-pyritohedral, common; massive; in reticulated and other imitative shapes.

Cleavage: none distinct; none in traces. Fracture granular and uneven. Brittle. H. = 5-5-6. G. = 6-4 to 6-6. Luster metallic. Color tin-white, inclining, when massive, to steel-gray, sometimes iridescent, or grayish from tarnish. Streak grayish black. Opaque.

Composition.—Shalitzite is essentially cobalt diarsenide, CoAs_2 , = Arsenic 71.8, cobalt 28.2 = 100. Chloanthite is nickel diarsenide, NiAs_2 , = Arsenic 71.9, nickel 28.1 = 100.

Cobalt and nickel are usually both present, and thus these two species graduate into each other, and no sharp line can be drawn between them. Iron is also present in varying amount; the variety of chloanthite containing iron has been called *chloanthite*. Further sulphur is usually present, but only in small quantities. Many analyses do not conform even approximately to the formula RAs_2 , the ratio H:As being from less than 1:3 to 1; the ratio is nearly 1:3, thus showing a tendency toward arsenic diarsenide, perhaps due to either molecular or mechanical mixture.

Much that has been called smaltite (speiskobalt) is shown by the high specific gravity to belong to the orthorhombic species safflorite.

Pyrite, etc.—In the closed tube gives a sublimate of metallic arsenic; in the open tube a white sublimate of arsenic trioxide, and sometimes traces of sulphur dioxide. On charcoal gives an arsenical odor, and fuses to a globule, which, treated with successive portions of borax-glass, affords reactions for iron, cobalt, and nickel.

Observations.—Usually occurs in veins accompanying ores of cobalt or nickel, and ores of silver and copper; also, in some instances, with niccolite and arsenopyrite. Found at the Saxon mines; Jonchiinstal, Bohemia; Wheal Sparrow, Cornwall; Riechelsdorf, Hesse; Tuoberg, Sweden; Allemont, Dauphiné. In the U. S. at Chatham, Conn., the chloanthite occurs in mica slate, with arsenopyrite and ulmannite; at Franklin Furnace, N. J.

COBALTTIT.

Isometric-pyritohedral. Commonly in cubes, or pyritohedrons, or combinations

resembling common forms of pyrite. Also granular massive to compact.

Cleavage: cubic, rather perfect. Fracture uneven. Brittle. $H. = 5.5$. $G. = 6-6.3$. Luster metallic. Color silver-white, inclined to red; also steel-gray, with a violet tinge, or grayish black when containing much iron. Streak grayish black.

CoMp.—Sulphureide of cobalt, $CoAs$ or $CoS, CoAs$, = Sulphur 19-3, arsenic 45-3, cobalt 35-5 = 100.

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Iron is present and in the variety *ferrobaule* in large amount.

Pyr., flt.—Unaltered to the closed tube. In the open tube gives sulphurous fumes, and a crystalline sublimate of arsenic trioxide. B.U. on charcoal gives off sulphur fumes and arsenic, and fuses to a magnetic globule; with borax a cobalt-blue color. Soluble in warm nitric acid, with the separation of sulphur.

Obs. Occurs at Tuusula and Håkås in Sweden; at the Nordmark mines; also at Sletterud in Norway; in Scladmlue, Kyrle; Bielefeld in Westphalia; Botallack in Cornwall, and St. Just, in Cornwall; Khetri mines. Rajasthan, India.

Greenockite. Sulphureide of nickel, $NiAs$ or Ni_3As_2 . Iron, and sometimes cobalt, replace more or less of the nickel. Isometric-rhombic: usually massive. $H. = 5-5.5$. $G. = 6-6.2$. Color steel-gray to blackish-gray. From Loos, Sweden; the Harz; Schladming, Styria, etc.

Orpiment is near greenockite, but contains also antimony. From Olsa, Carinthia.

Wu-yamite. $CoBi, NiBi, Co_8Bi, Ni_8Bi$. Cleavage cubic. Color tin-white to steel-gray. Broken Hill mines, New South Wales*.

Malachite. Sulphate of nickel, $NiSO_4$ or $Ni_3(SO_4)_2$; arsenic is usually present in small amount. Isometric-tetrahedral; both primitive and tetrahedral forms occur (see Figs. 146, 147, p. 61). Usually massive, granular. $H. = 5-5.5$. $G. = 6.2-6.7$. Color steel-gray to silver-white. Occurs in the mines of Freusberg, Nassau; Schleibach, Prussia; Lollig. Carinthia (tetrahedral); Montebello. Sicily, Sardinia (pyritoid).

Malachite. Sulphate of nickel, $NiSO_4$ or $Ni_3(SO_4)_2$. Massive, color light bluish gray. From the Friedrich mine near Schönbach a. d. Siegen, Prussia.

Platinochloride. Platinum chloride, Pt_2Cl_6 . In minute cubes, or cubo-octahedrons, $H. = 6-7$. $G. = 10-10.5$. Luster metallic. Color tin-white. Streak black. Found at the Vermillion mine, 33 miles west of Sudbury, Ontario, Canada; also in Mecklenburg Co., N. Carolina. This is the only known native compound of platinum,

LanrltA. Sulphide of ruthenium and osmium, probably essentially RuSi. In minute octahedrons; in gmlua. H. = 7-5. O. = 6-M. Luster metallic. Color dark Iron-black. From the platinum washings of Borneo. Also reported from Oregon.

Skuttemdit*. Cobalt arsenide, CoAs. Isometric-pyritohedral. Also typically granular. Cleavage: a distinct. H. = 6. Q. = 6-72-6'86. Color between tin-white and pale lead-gray. From Skuttenid. Norway.

Nickel. Bismuth arsenide. (Ni,Co,Fe)As. Hasslerite, granular. Color gray. From near Silver City. New Mexico,

Bismuth-telluride. Co(As,Bi)Te. A bismutefeldite containing bismuth. Color tin-white. o. = 603. Zschornau, near Schneeberg.

Marcafite Oront.

For the list of species and their relations, see p. 399.

MABOASTTB. White iron pyrites. Orthorhombic. $a:b:c = 0.7662 : 1 : 1.2343$. mm³, 110 Å $a:b = 74 : 57$, $c = 116.30$. Twins: tw. pi. m (Fig. 619), sometimes in stellate twinning {Fig. 406, p. 138, cf. Fig. 630}; also tw. pi. e (101), less common the crystals crossing at angles of nearly 60°. Crystals commonly tabular || c, also pyramidal; the brachydomes truncated || edge b/c. Often massive; in stalactites; also globular, reniform, and other imitative shapes.

Cleavage: m rather distinct; ? (Oil) in traces. Fracture uneven. Brittle. H. = 6-6.5. G. = 4.85-4.90. Luster metallic. Color pale bronze-yellow, deepening on exposure. Streak grayish or brownish black. Opaque.

Composition.— Iron disulphide, like pyrite, FeS, = Sulphur 53.4, iron 46.6 = 100. Arsenic is sometimes present in small amount.

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SULPHIDES, SELENIDES, TELLURIDES, ARSENIDES, ANTIMONIDES. 303

The varieties named depend mainly on state of crystallization. It is divided (Stibioferous): Radial: also the simple crystals. Cockcomb P. {Kammlite Gen.}: Aggregations of hexagonal in crystalline forms. Nipar P. (Sphalerite Uenii.): Twin crystals, with re-union angles a $\frac{1}{2}$ like the head of a spear in form. Capillary (Haarkite Qetm.); In capillary crystallizations.

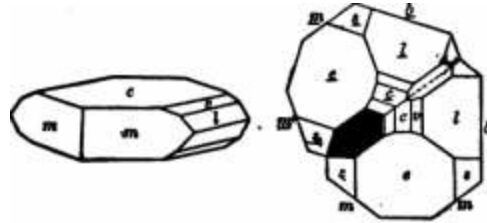
Pyrite, etc.—Like pyrite. Very liable to decomposition, more so than pyrite.

Specific gravity of pyrite, but has a lower specific gravity, and the color when fresh {e.g.

after treatment with acid) Is paler; when crystallized easily distinguished by the forms.

Joachimsthal, Bohemia, and in Saxony and the Harz. Occurs with galena and fluorite In Derbyshire in the chalk-marl between Folkestone and Dover; near Alston Moor, Cumberland;

818. SIB. «ao.



Common Form.

1 Galena, III., in stalactite with concentric layers

Wis., in fine crystals; on sphalerite at Joplin, Mo.

The word marcasite of Arabic or Moorish origin (and variously used by old writers, for

•sphalerite and galena; Mineral Point, Wis., In fl

: layers in, Mo.

bismuth, antimony), was the name of common crystallized pyrite among r . .

mineralogists in later centuries, until near the close of the last. It was first given to it by Scheuchter in 1645.

It is usually iron arsenide, FeAs₂, but passing into Fe^{As} (FeAsWpyrM);

■' ■ ■ - ..-.-.-.-.-.-----|, 50° . "

Occurs in the Joachimsthal district in Carinthia; with fluorite at Schladming, etc. In the U. S., marcasite occurs in Gunnison Co., Colorado, etc.

It is near marcasite, but contains sulphur; from Oeyers, Saxony.

ABSENORBIT, or Mispickel. Arsenic Oerm. Orthorhombic. Axes a:b:c = 0.6773 : 1 : 1.1882. mm³, 110 A liO = OS^{IV}.

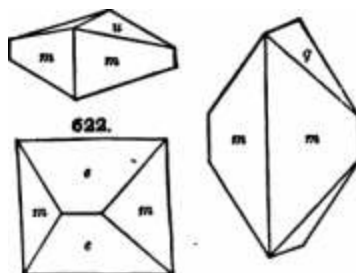
d, 101 A ioi = 100° 88'.

Mt. 014 A 01* = - 88° 6'.

$nn', 013 A \text{ o}i2 = 61^\circ 36'$,

^ . on A Oil = SQ" 50*. Twins: tw. pi. m, Bometimea repeated like marcasite (Fig. 407, p. 128) j e (101) crnciform twins, also trillings (Fiffs. 402, 403, p. 128). Crystals prismatic m, or flattened vertically by the oscillatory combination of brachydomea. Also columnar, straight, and divergent; grannlar, or compact.

Cleavage: m rather distinct; c in faint traces. Fracture nnerec. 1



H. = 5-5-6. G. = 5-9-6'2. Luster metallic. Color silver-white, inclining to steel-gray. Streak dark grayish black. Opaque.

Comp^ Tar.—Sulpbarsenide of iron, $FeAsS$ or $FeS, FeAB$, = ArBenic 46*o, sulphur 19'7, iroa 34'3 = 100. Part of the iron is sometimes replaced by cobalt, as in the variety danaiU (3 to 9 p. c Go).

Pyr., etc.— Id Ibe closed tube ftt flnt gives & red Hublfrm&te of aiaenfc trisulphlde. theo B black lustrous subllmiite of metallic aiMiitc. Ip the opeo tube givw eulpburous romeB sod a wbite autilimftte of arsenic trloxlde. B.B. on charcoal gives anenldu fumes aud a magoetic globule. The vu-letles containing cob>ll give, after the arsenic has been losMed ofl, a blue color with boTai-glass when fused in O.K. with successive portioos of flux until all the Iron is oildi>ed. OTves lite with steel, emitting an alliaceous odor. Decomposed by nitric acid with the separation of sulphur.

Dt£—Characterized bj Its hardness and tin-white color; closcl j resemblea some of the ■ulphlde and arsenides of cobalt and nickel, but identified, in most casea easllj, bj its blowpipe characters. LOIHngite does not gire a decided aulplur reaction.

Ob«,—Found principally ID crystalline rocks, lu usual mineral associates being ore* of silver, lead, and tin, also pyrite, calcopyrille, and sphalerite. Abundant at Freiberg, etc., InBaxonj: Relcbeniteln, Silesia, In serpentine: lu beds at Breitenbrunn, Andreatberg, JoacLltnethaU Tunabers, Sweden; Skntlerud. Norway; at several points In Cornwall.

In (he TJ. B., in N. SampthiM, In gneiss, at Frabconla (danatie). In Maine, at Blue Hill, etc. In Conn., at Chatham; at Mine HUl, Roiburj, wlib eideriie. In JIT. Fork. massive. In Lewis, Essex Co., with hornblende; near Edenville. and elsewhere Id Orange Co. In CVi^omia, Nevada Co.. Qrass valley. In crystals at St. Fniugoia, BeauceCo.. Quebec; large beds occur in quartz ore veins at Deloro, Hastings Co., Oniarlo, where It la mined for gold.

The name mlsplckel is an old German term of doubtful origin. DaTuiite is from J. Freeman Dana of Boston (1798-1887), who made known the Franconia locality.

Baffiorite. Like amantite, essentially cobalt diarsenide, CoAs_2 . Form near that of arsenopyrite. Usually massive. $H. = 4.5$. $O. = 69-78$. Color tin white, soon tarnishing. From Schneeberg, Saxony; Bleber, Hesse; Wittichen, Baden; Tunaberg, Sweden.

Rammelsborgite. Essentially nickel diarsenide, NiAs_2 , like chloanthite. Crystals resemble arsenopyrite; also massive. $G. = 6-8-78$. Color tin-white with tinge of red. Occurs at Schneeberg and at Rastatt.

Olancodot. Sulfarsenide of cobalt and iron, $(\text{Co,Fe})\text{As}_2$. In orthorhombic crystals (see, etc., p. 269). Also massive. $H. = 5$. $G. = 6-90-6-01$. Luster metallic. Color grayish tin white. Occurs in the province of Huasco, Chili; at Hakansjö, Sweden. Named from Yaxvito, Bolivia, because used (or mixed) as smelt.

Allochloite. Probably essentially Co(As,Bi)S with cobalt in part replaced by iron; or a glaucodot containing bismuth, commonly in columnar to hemispherical aggregates. $H. = 4.5$. $G. = 6-6$. Color steel-gray. From Oravitsa.

Tolfaite. Probably Ni(As,Bi)S , near arsenopyrite. In small crystals resembling arsenopyrite; also colorless radiated. $H. = 4-5$. $O. = 6-878$. Color silver-white to tin-white. From Wolfach, Baden.

The following species are tellurides of gold, silver, etc.

STLTAHITXI. Graphic Tellurium. Schrift-Tellur Germ.

Monoclinic. $h:c:t = 1.6339 : 1 : 1.1265$; $\beta = 89^\circ 35'$. Often in branching arborescent forms resembling written characters; also bladed and imperfectly columnar to fibrous.

Cleavage: none perfect. Fracture uneven. Brittle. $H. = 1.5-2$. $G. = 79-83$. Luster metallic, brilliant. Color and streak pure steel-gray to silver-white, inclining to yellow.

Comp.—Telluride of gold and silver $(\text{Au,Ag})\text{Te}$, with $\text{Au} : \text{Ag} = 1 : 1$; this requires: Tellurium 62.1, gold 245, silver $13 \times 4 = 100$.

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SULPHIDES, SELBMIDBS, TELLURIDES, ARSENIDES, ANTIMONIDES. 305

Pyr., etc. — If the open tube gives a white sublimate of tellurium dioxide which near the apparatus is gray: when treated with the blowpipe flame the sublimate fuses to clear transparent drops. B.B. on charcoal fuses to a dark gray globule, colorless; the coal with a while turning, which if treated in K.F. disappears, giving a bluish-green color to the flame; after

liiii)r blowing a yellow, malleable metallic globule is obtained. Musi Tarielies give u fuint cuaiiuf iif leiid oxiUe aud aniimonj oxide on cbarcoat.

Obi ^Witb gold, at Oifenb&ujra, TruuH^lvanla; also at Nagy^ . lu Californtn, Calaveras Co., at the HeUines and StuuilauB mines. lu Boulder Co., and elsewhere in Colorado. Named from Transylvauia, where fli-st found, and la allusioa to sjfjninium, one of the names at first proposed for the metal lelliirtum.

Krennarita. A telluride of gold and silTcr (Au,Ag)Tei like syWauite. In priamatlo orvsttlls (orthorhombtc), vertically sirialed. G. = 8853. Color silveT-white to braas-jellow. From Nagy&g, l'musylvauia; Cripple Creek Colorado.

C'ALAVERITB A gol^l-silrer telluride. Like syWaaite (Au,Ag)Tem with Au : Ag = 6 : 1 or 7 ; I. Massive. H. = 25. Q. = 9043. Color pale bronze-yellow. Occurs with l>etz<te nt the Sianislnus mine, Calaverus county, Callforuia. Also at the Red Cluud and other mines, Colorado,

Calaverite bat ihe same general formula ai syWanite but a much higher percentage of jfold. and miy belong with it; or, aa aeema prol»bl«, krennerlte may De the cryatallized form of calaverite.

Na^aglte. A sutpbo-lelluride of lead and gold; containing also about 7 p. c. iif antimony. Orthnrhombic. Cryatala tabular Tb; alau granular massive foliated. Cleavage: ft pevfecl; flexible H. = 1-1 5. G. = 6 85-7-3.. Luster metnllie, Bplendenl. Stirok and color blncklaU lead-gray. Opaque. From NagyAg, Transylvania; and at Offenb&uyu. Reported from Colorado.

Oxy sulphides.

Here are included Kermeaite, Sb,S,o, and Voltzite, Zn,S,o.

Hermaalte. Anlini on blende, Rotlispiassglanzerz o»rm. Pyroslibile. Antimony oxy-ulpbide. ab,8,o or 3Sb,Si.Sb,o.. Monoclinic. Usually in litfts of rapillary cryilals. Cleavage: a perfect. H. = 1-1 S. G, = 4 5-4 6. Luster adamantine. Color cherry-red.

Kesults from the alteration of stibnlle Oncurs at Malaczka, HiiDgary; Brflunedorf, Saxony: Allemoni, DaupbinS. At South Hnm, Wolfe Co., Quebec. Canmlii; with native anlimony and stibnito al the Prince William mine, York Co., Hew Brunswick.

Named from ktrne*, a name given (from the Persian gurmitg. crimson) in the older chemistry to red amorphous antimunv trisiilpliide, often mixed with antimony trioxide,

VoltiiU. Zineoxyeulphlde. Zn.S.O or 4Zn8,ZuO. In implanted spherical globules; n _< ..« n ^8-66-88o. Color dirty rose-red. yellowish. Occurs at RoBiSres, Puy —u,,i. «. _.,v— o y(Mtrilmtdt).

de Ddme; Joachlmsthal; Marleuberg. Saxony (I

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DX8SCRIPTITE UIXERA_{1.0aT}.

m. STTIf HO-SALTS.

L Snlphanenitei, SnlphantimoniteB, SnlphobUmuthltei. n. Snlpharsenatea, SnlphoBtauuates, etc.

X. Snlphanenitei, Snlphantlmonitei, etc

In these sulpho-salts, as further explained on p. 248, sulphur takes the place of the oxygen in the commoner and better understood oxy-acids (as carbonic acid, H₂CO₃, sulphuric acid, H₂SO₄, phosphoric acid, H₃PO₄, etc.).

The species included are salts of the sulpho-acids of trivalent arsenic, antimony and bismuth. The most important acids are the ortho-acids, H₃AsS₄, etc, and the meta-acids, H₂AsS₅, etc.; but HAsS₃, etc., and a series of others are included. The metals present as bases are chiefly copper, silver, lead; also zinc, mercury, iron, rarely others (as Ni, Co) in small amount. In view of the hypothetical character of many of the acids whose salts are here represented, there is a certain advantage, for the sake of comparison, in writing the composition after the dualistic method, RS.AS₃S₄, etc.

As a large part of the fifty species here included are rare and hence to be mentioned but briefly, the classification can be only partially developed. The divisions under the first and more important section of sulpharsenites, etc. with the prominent species under each, are as follows:

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BULPHO-SALTS.

D. Ortho- Division. BS - (As,Sb,Bi)₃S₄ = 3:1.

General formula: B₃AS₃S₄; B₃AS₃S₄^B₃S₄^ etc.

BouFDODlte Group. Bounonite 3(Ou.,Pb)S₃Sb₃S₄, Wittichenite 3Cti,S₃Bi₃S₄,

Aikinlte 3(Fb,Ou.)₃Bi₃S₄. XiUiuiito 3FbS₃Bi₃S₄^ etc

Pyrargyrite Group.

Pyn^yTite 3Ag₃S₃Sb₃S₄, Froystite 3Ag₃S₃As₃S₄,₉

E. Basic Division. BS : (As,Sb,Bi)₈, = 4 : 1, 5 : 1, 6 : 1, 9 : 1, 12 : 1 L Tetrahedrite Group.
4Cn, S.Sb, S, Tennantite 4Gu, 8.Aa, S.

Tordanite Group.

Tetrahedrite

A. Acidic Division.

Obitite. 3PbS.8Bi, S. Foliated massive. Color lead-gray. From Chivato, Peru.

Onproboscite. Probably 3Cu, S. 4Bi. Part argentiferous. Resembles bisulphide. G.
= 6-6-7. From Hill Talley, Park Co., Colorado.

Reichenbachite. 4PbS.5Bi, S. Fine-grained, massive. Color lead-gray. G. = 6-1-4-4. From
Reichenbach, Burgundy.

B. Meta- Division. E8.As, S, Sb, S, etc. Zinkite Group. Orthorhombic. ZINIKITE.
Bismutite (Oxide) Qerm. Zinkite.

Orthorhombic Bismutite: Bi_2S_3 . Crystals seldom distinct; sometimes in
nearly hexagonal forms through twinning. Cleavage faces are striated. Also
columnar, fibrous, massive.

Cleavage not distinct. Fracture slightly uneven. H. = 3-3.5. G. = 5-30. Luster metallic.
Color and streak steel-gray. Opaque.

by Google

Composition.—PbSb₂S₃ or PbS.Sb₂S₃, = Sulphur 22-3, antimony 41-8, lead 35.9 = 100. Arsenic
sometimes replaces part of the antimony.

Pyrolysis, —Decrepitate and fuse very easily: in the closed tube gives a faint sublimate
of sulphur and antimony trisulphide. In the open tube sulphurous fumes and a white
sublimation of antimony; the arsenical variety gives antimony trisulphide fumes. On
heating in a Henschel entirely volatilized, giving a continuous ring on the outer edge is white,
and in the center dark yellow; with sulphur R.F. yields globules of lead. In hot
hydrochloric acid with evolution of hydrogen sulphide and lead chloride.

Obs.—Occurs at Wolfsberg in the Harz; Eisingthal, Baden; Sevier County. Arkansas; Silver
Creek, Colorado.

Androsite. Ag₂S.2PbS.38Bi₂S₃. In prismatic, orthorhombic crystals. Color dark gray to black.
From Felspar. Waverite and Sundtite belong here.

Bastnaesite. Skierite. PbS.As₂S₃. In slender, tabular crystals. G. = 6-8. Color dark

lead-gray. Occurs In the doEouitte of Ibe BinoeDthai.

Emplactite. EupfiirwiBmutbgladZ Otrm. CuiS.Bl,S|. In thin striated prismi. O. = 6 3-6'5. Color grayUh white to tlu.white. Occui-a In quartz at Bchwarzenberg aod Anuaberg, tiaxouy.

ObalcoiUUta. Wo] fsbergite. Kupferant1 monglanz ofrm. Cu,8.Bb,B,. In amall aggre-gated prisma; also fine granulBT, mafisive. O, = 4'6-o. Color between lead-gray and on-gray. Frnm Wolfsuerg in tiie Rarz. Oufjaiite from Spaiu is the &ame species.

Oalenobiamutlte. PbS.Bl,8,; also with Ag.Cii. C'yElnlllne columuar to compact. Color lEini-gray io liii-wliile. Q. = B'B. From Noidmsrk, Sweden; Pougkkeepste Oulcb, Coloradi) laladkaite, nrgentlferouaj; Fainn, Bwedeu (Belenlftrous}.

Berthieite. Probably FeS Bb Bi. Fibrous massive, granular. 6. = 40. Color dark steel-gray. From Cbazelles and Bbirtouret, Auvergne; BrSunsdorf, Baiony, etc.

Matlldita. Ae.B.BI.S.. In slender, prismatic crystals. G. = 69 Color gray. From Morocfaoca, Peru; Lake City, Colorado. PLBNABaYBiTX, from Sebapbach, Baden, bas prubably the tame composition nnd may be identical.

Mlargyrita. Ag,B.ab,Si. In complex nionoclinic crjsuls. also masdve. H. = 3-3'5. G. = 01-6 SO. Luster metnllic-adamMiline Color iron-black to steel-gray, In thiti splints rs deep blood-red Streak cherry-red. From BrBunsdorf, Bs^iony; FersObfinya; PHbram, Bohemia; Clausthal, etc

Lonndite. A sulpbarsenide of tballlum, TlAsS^ MoDocloic. Color oochiDeal-red. .From Allchur, HacMdonla.

, C. Intennediate Division.

Plaglonlte. Perhaps 6Pb9.4Sh Si. Crystals tblk tabular, monocllnic; also maailTB, granular to compact. O. = G'4. Color blackish lead-gray- From Wotfaberg, etc

BcUnoerits. S(A7,,Pb)8.96l,Bi. Hastlve, granular. G. = «-7t- Color lead-gray. Treasury lode. Park Co., Colorado.

Klaprotbollta. SCuiS Bi,Bi. In furrowad prismatic cryital*. G. = 4'S. Color steel-gray. WItlicben, Baden.

Binnlte. Perhaps 8Cu,S.2Aa,8i Isometric-tetrahedral; also maaslve. Q. = 4'477. Color dark steel-gray to iron-black. From the dolomite of Ite Binnenthal.

Waireidte. Domlndte. 8PS.2Sb,8i- In wool-like aggregates of acicular crystals. Color

grayish black. Qunnlson Co., Colorado.

Jamesonite Group. $8\text{Pb}8.\text{As},\text{S},, 2\text{RS}8\text{b},\text{S},,$ etc. Orthorhombic. Prismatic angle about 80° . JAMESONITE.

Orthorhombic. Axes: $a:b:c = 0.8915 : 1 : 0.7840$. In cleavages; common in capillary veins; also fibrous massive, parallel or divergent; compact massive.

.dbyGoogle

Cleavage: basal, perfect; b, m cleaves so. Fracture uneven to conchoidal. Brittle. H. = 2-3. G. = 6.5-7.0, Luster metallic Color steel-gray to dark lead-gray. Streak grayish black. Opaque.

Comp.—Pb, Sb, S. or $2\text{Pb}8.8\text{b}, 8 = \text{Sulphur } 19.7, \text{antimony } 29.5, \text{lead } 50.8 = 100$. Most varieties show a little iron (1 to 3 p. c.), and some contain also silver, copper, and zinc.

Pyrite as found in zinc; $> 80\%$.

Occurs principally in Cornwall; also in Siberia; Hungary; at Vale of Leadhwaite in Yorkshire; at the old mine in the Bevkra Co., Arkansas. Named after Prof. Robert Jameson of Edinburgh (1791-1863).

The feather vein (Felspar vein) occurs at Wolfberg, etc. In the Harz; Freiberg. Bohemia; in the Tschibitz, near Bottschau.

Dufrenoyite. $3\text{Pb}8.\text{As}1\text{Bi}$. In irregularly modified orthorhombic crystals; always massive. Cleavage: e perfect. H. = 8. Q. = 6-7.5. Color blackish lead-gray. From the Binntal, Switzerland, in dolomite.

Ratbit. a Bismuth arsenite of lead resembling dufrenoyite. From the Blonenthal.

Jamesonite*. $8\text{Pb}8.\text{As},\text{S}$. Usually massive, fibrous or radiated. G. = 6.5-7.0. Color iron- or steel-gray. Coahuila, Province of Sinaloa, Mexico; Bjelke mine (Bjelle), Nordmark, Sweden) Colorado.

Kolmillerite*. $\text{Sn}8.(\text{Bi},\text{Sb}), 8$. Fibrous radiated or granular massive. G. = 6-7.5. Color lead-gray to steel-gray. From Hven, Sweden; Ouray, Colorado.

Brogniardite*. $\text{Pb}8.\text{As},\text{Sb},\text{Bi}$. In isometric octahedra and massive. G. = 6-7.5. Color grayish black. From Mexico.

Samsayite. Near Jamesonite, perhaps $2\text{Pb}8.\text{Sb},\text{Bi}$. Small tabular monoclinic crystals. Q. = 5.9. Color gray. From Felsőbánya, Hungary.

Bismuth arsenite. $\text{Pb}8.\text{As},\text{Sb},\text{Bi}$. Small tabular crystals and granular massive. Q. = 6-7.5. Color

lead-gray. ;n«m Schapbach. Baden.

MoDoclinio. Aiea d : 5 : <! = 0-5871 : 1 : 0-9277; fi = 87' 46'. Habit prie-matic. Q. = 6'2-6'4. Luster metallic. Color and streak light steel-gray inclining to silver-white, alao to blackish lead-gray.

Com[^]—(Pb,AeJ.Sb,8 or 5(Pb,AgJS.2Sb,S ."

Ob*.—From tht[^] HimmelaFUrst mine, at Freiberg, Saxony; Eapnlk. Hungary; b£nya; HiendeleQclna, Spain; also from Ibe Augusta Mt., Qunnison Co., Colorado.

Diaphorite. Like freleslebenlte in compodliou but orthorbombic In form, G. = Q-S. From PFibram, Bohemia.

BonxiANauKmi.

Orthorhombic. In crystalline plumose masses; graniilar, compact. H. = 2'5-3. G. = 575-60. Luster me[^]llic. Color blnian lead-gray; often covered with vellow spots from oxidation. Opaque.

CoMp.—Pb.Sb.S,, or 5PbS.2Sb,S, = Sulphur 18-9, antimony 25-7, lead 65-4 = 100.

Pyr.—Same aa for ifnkenlle, p. 807.

Obs.—Moliires, Depart, dii Gard, France; at Nerchinsk; Wolfaberg in Ibe H&rzt; Pfiham, Bohemia: near Bottino Tuscany: Echo District, Union conntv, Nevada.

Bn[^]riiJtUt and plumbotUb are from Nercblnsk; they correspond nearly to 10PbS.S8b(Si, but the material analyzed may not have been quite pure.

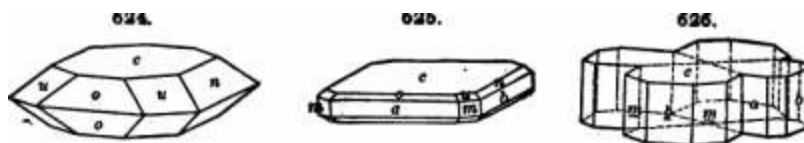
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DB8CR1PTITE HINEBALOQT.

D. Orlho- Division. 3RS.Ab,S,, 3Rg.Sb,S,, etc Boamonlte Group. Ortborhombic. PrtBmatic angle 86° to 87°.

BOITRNONITB. Radclerz Germ. Wheel Ore. Ortborhombic. Axes; £:i -.6 = 0-9380 : 1 : 0-8969.

sn, 001 A Oil = 41* 63' su. 001 A 112 = 88° 16'



Huz,

Nagyig.

Kapnik.

Twins; tw. pi. m, often repeated, forming cruciform and wheel-shaped crystals. Also massive; granular, compact.

Cleavage: b imperfect; a, c less distinct. Fracture subconchoidal to uneven.

Rather brittle. H. = 2-5-3. G. = 5.7-0-9. Luster metallic, brilliant. Color

and streak steel-gray, inclining to blackish lead-gray or iron-black. Opaque.

Comp.—{Pb,Cu,);Sb,S. or $3(\text{Pb,Cu,})\text{S}_8\text{b}_8 = \text{PbCuSbS}$, (if Pb : Cu, =

2 : 1) = Sulphur 19.8, antimony 34.7, lead 42.5, copper 13.0 = 100.

Pyr., etc.—In the closed tube decrepitates, and gives a dark-colored sublimate. In the open tube gives sulphur dioxide, and a white sublimate of antimony. B.B. on charcoal fuses easily, and at Urbs coats the coil while; continued heating gives a yellow coating of lead oxide; the residue, treated with soda in B.F., gives a globule of copper.

Decomposed by dilute acid, affording a blue solution, and leaving a residue of sulphur and a white powder containing antimony and lead.

Obs.—From Niedersdorf in the Harz; also Wolfsberg, Clausthal, and Andreasberg: Freiberg, Bohemia; Kapnik. Hungary; Bideford in the Wheal Boys, Cornwall; in Mexico: Huasco-Alto, Chihuahua.

In the U. S., at the Boggs Mine, Yavapai Co., Arizona; also Montgomery Co., Arkansas; reported from San Juan Co., Colorado. In Canada, in the townships of Marmora, Hastings Co., and Darlington. Lanark Co., Ontario.

Attn.— $8(\text{Pb,Cu,})\text{S}_8\text{b}_8$. Acicular crystals; also massive. G. = 8.1-6-8. Color blackish lead-gray. From Berezov near Ekaterinburg, Urals.

Wittichenite*. $3\text{Cu}_8\text{Bi}_3$. Rarely in crystals resembling boumonite; also massive. Q. = 4-6. Color steel-gray or tin-white. Wittichen, Baden, etc.

Butterfieldite. $3(\text{Cu,Ag,Pb})\text{S}_8\text{b}_8\text{Si}$. In orthorhombic crystals, in cruciform twins like boumonite. G. ~ 4.70. Color iron-black. Copiapó, Chile.

PbSbBiS and SPbS.BiS . Massive, crystalline. Color steel-gray. Evens, Sweden: Leadville, Colorado (argemiferous).

Qnltemunlta. Perhaps $3\text{PbS}\cdot\text{As}_2\text{Si}_2$. Massive, compact. $\rho = 5.1$. Color bluish gray. Zuill mine, Silverton, Colorado.

Tapalplta. A Sb_2Te_3 of bismuth and silver, perhaps $3\text{Ag}_2(\text{B},\text{Te})\cdot\text{Bi}_2(\text{S},\text{Te})$. Haaslve, granular. $G = 7.8$. Sierra de Tapalpa, Jalisco, Mexico.

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BULPHO-8ALTB.

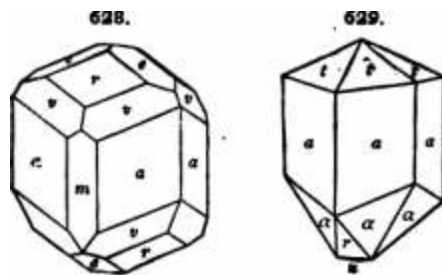
Pyrai^yrlte Gronp. Rhombohedrsl-lieinimorphic.

PTRASOTRTTB. RuLy Bllver Ore. Dark Red SlWer Ore. Duukles RotliKtLltlgerz, ADdmoudlberbleDde Qtrm.

Ehombohedral-hemimorphic. Axis: $(100) = 0.7893$; $(001) \perp (100) = 42^\circ 20'$.

«■. $(111) \perp (111) = 43^\circ 17'$, $(111) \perp (111) = 74^\circ 86'$

$(111) \perp (111) = 71^\circ 89'$, $(111) \perp (111) = 85^\circ 12-637$.



Crystals coniraoraly prismatic. Twins: tw. pi. a, very common, the axes \perp parallel; « (101) , also common. Also maeieve, compact.

Cleavage: r distinct; e imperfect, Fracturee onchoidal to uneven. Brittle. H. — 25. tr. — $5 \sim 7-5'86$; $5'85$ if pure. Luster metallic-sdamaDtine. Color black to grayish black, by transmitted light deep red. Streak porpligh red. Nearly opaque, but transparent in very thin splinters. Optically —. Refractive indices, $a > e = 2.881$ Fizeau.

Comp,—Ag,SbS, or $3\text{Ag}_2\text{S}\cdot\text{Sb}_2\text{S}_3$, = Sulphur ITS, antimony 22-3, silver 59-» = 100. Some varieties contain small amounts of a

Pyrr., etc.—In the clused tiile fuses and gives & reddish sublimOite at aDtlmoDj' oiyBul-

g'lidt;; Id the iJpeQ Cube sulplureua fumes uod a white Bublimate of oxide of autimony. .B, on cliirciml fuses with spfrlirig to a gliibiiiie. cotits the coal while, aod the assay U converted inio silver sulphide, whilcli. treated iu O.F.. or with soda in R.F.. gives n globule of silver. lu case ar^L'uic is present it may be detected by fusing the pulverized iniDeral with snda on chtrtoul In R.F. Decomposed by nilric acid with lue separallon of sulplur

and of intimacy trioxide.

011«. —Occurs at Andriessberg in the Harz; Freiberg, Saxony; Příbram, Bohemia; Schemnitz, etc. Huellegard; Kongsberg, Norway; Gaudischan, Spain; Cornwall. In Mexico it is worked at Quitova and elsewhere as an ore of silver. In Chile with proustite at Chantarrillo near Copalaco.

In Colorado, not in the district. Gunnison Co.: Lead sphalerite in the Leadville district. Ouray Co., etc. In Nevada, at the Daney Mine; about Austin, in the Colorado river; at Poorman lode. Idaho, in the Snake River valley with cerussite. In New Mexico, Utah, and Arizona with silver ores at various points.

Named from Lead, and also from the color.

Light Red Silver Ore. Lichtes Rothsilbertz, -Kbombohedral-hemimorphic. Axis $a = 0.8039$; $c/a = 42^\circ 63'$. $\alpha = 42^\circ$ iff

$\beta = 74-89^\circ$ w, $\gamma = 35-18^\circ$ *

$\delta = 70^\circ 13'$. Crystals often acute rhombohedral or scalenohedral. $h(10\bar{1}4)$ and r . Also massive, compact.

tw. pi.

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DESCRIPTION OF THE MINERAL.

Color: red distinct. Fracture conchoidal to uneven. Brittle. $H. = 2-2.5$, $G. = 5.5-5.64$; 5.57 if pure. Luster adamantine. Color scarlet-vermilion; streak same, also inclined to aurora-red. Transparent to translucent. Optically negative. Double refraction strong. $a > r = 2.979$.

Coap.— Ag_2S , or $3Ag_2S$, = Sulphur 19.4, arsenic 15.2, silver 65.4 = 100.

Prp., etc.—In the closed tube burns easily, and gives a faint blue flame of arsenic trisulphide; in the open tube sulphurous fumes and a white sublimate of arsenic trioxide. B.B. on charcoal fuses and emits an odor of sulphur and arsenic; will solder. In R.F. gives a globule of tellur. Decomposed by nitric acid, with separation of sulphur.

Obs.—Occurs at Freiberg, Jáchymov, etc. in Saxony; Joachimsthal, Bohemia; Chantarrillo in the Colorado; and in Mexico; Chile, at Chantarrillo in the silver-bearing crystalline rocks.

In Colorado, Ruby distr., Ouray Co.; Bheridan mine, San Juan Co.; Vanhook mine, Ouray Co. In Arizona, with silver ores at various points. In Nevada, in the Daney

mice, and Id Comstock lode, rare; Idabo, at the Pnorniiu lode.

Named after the French chemist, J. L. Proust (1765-1826).

SangoUilta. Near prouatite in composition. In glittering scales, hexagonal or rhombohedral. From Cbafiarctlo. Cbili.

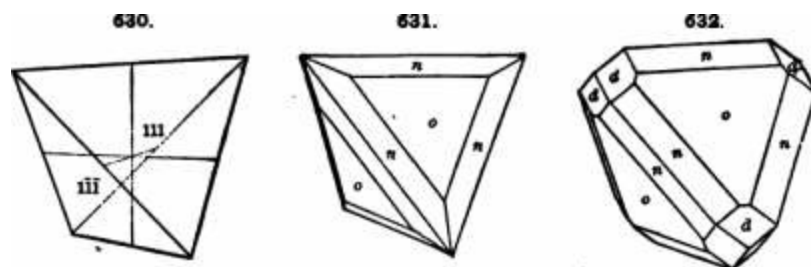
Falkenhaynita. Perhaps $SCuS \cdot BbBi$. Massive. resembles galena. From Joachimthal, Bohemia.

Fyroitilpnlte. Like pyrrholyte, $8AgBi \cdot BbBi$. In tufts of slender (moderately) crystals. $Q. = 4 \frac{1}{2}$. Color lilac-red. From Andreasberg in the Harz, Freiberg; Pfibram.

Sttngaiita. Contains arsenic, selenium, and silver. In small tabular crystals. $Q. = 5-6$. Color blackish-brown to black; hyacinth-red by transmitted light. Streak orange-yellow. From Chafiarctlo, CbilL

E. Basic Division. Tetrahedrite Group. Isometric-tetrahedral. TSTRABEDRITB. Gmy Copper Ore. Fnhlerz Otrm. Isometric-tetrahedral. Habit tetrahedral. Twins: tw. pi. o; parallel axes (Fig. 354, p. 121, Fig. 370, p. 124). Also massive; coarse or fine; compact.

also with granular.



Cleavage none. Fracture subconchoidal to uneven. Rather brittle. $H. = 4$. $G. = 4 \frac{1}{4}-5$. Luster metallic; often splendid. Color between flint gray and iron-black. Streak like color, sometimes inclining to brown and cherry-red. Opaque; sometimes translucent (cherry-red) in very thin splinters.

Comp. Formula.— Essentially $Cu_3Sb_2S_7$, or $4Cu_3Sb_2S_7 = S_{28}Cu_{12}Sb_8$, = Sulphur 23.1, antimony 24.8, copper 52.1 = 100.

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Antimony and arsenic are usually both present and in a tetrahedrite structure. Into the allied species tennantite. There are also varieties containing bismuth, chiefly at the Harnack end of the series. Further the copper may be replaced by iron, zinc, silver, mercury, lead, and rarely cobalt and nickel.

Vu.—Ordinary. Contains little or no silver. Color steel-gray to dark gray and iron-black.

G. = 4 75-19.

Argentiferous; FrtilMrgiU, Weiugiltigerz Qerm. Coutalus 3 lo 30 p. c. o[silTcr. Culo^r ^ usually sluel'gmy. lighter Ihaa Lbe oidloary TKrletles; aomellDies irou-blacki'streak oficu reddish. G. = 4-8B-S-0.

Mereuriul; Hchaattitt. Conlaln 6 to IT p. c at mercury. Color davk gray to Iron-black. Lu9ter often dull. O. = 5-10 ciletly.

Maliruneakite. from Peru and a similar variety from Arizjna, cootaln 18-16 p. c. of kad.

Pyr., etc.—Differ in ilie diS[^]reut varieties. Ju ilie closed lube all Ilie antimoiial kinds fuse uud give n dark red BUbilnialu of aDilmony oiysuljibide: if much sTsenlc is present, a sublimate o[arBenic irisulpliidu tlrst forms. lu the oprn lube fuses, gives sulphurous futne* and a while subliuMle of aotimouy oiide- if arsenic is piesenl, a crysialline volalile subli-male cod deuses with ihe aaimony; if the ore cuniuiiij inerrury it condenses in miniile nietullic globules. B.B. on cliarcoal fuses, givesa coating of the oxides of antimony and SQiiietimea arsenic, ziuc, ami lead; Hrseuic is delected by the odor wbeu the coating is treated ID R.F. The roasted mineral gives with the fluxes reactlous for Iron and copper; with soda yields a globule of metallic copper. Decomposed by nilric acid, with separation of sulphur aod antimony trioxide.

Diff.—Distinguished by its form, when crysllized, by its deep black color on fracture ADd brilliaut metallic tusler. It is harder than bournouite and much tofler than u:sguelbe: the blowpipe cbaraciers are usually distinctive.

Obs.—Ofen associaled viib chalcopyrlLe, pyrlie. splalerite. pslena. and VBri(.u> other •liver, lead, and copper ores; also sldente. Occurs at many Coruish mitits; tijiiti ai the Berod!<fi>ot mine. Liskenrd in tetrahedral crjsiils often coaled wTib iiidtcci-nt cIih1< i>'Viie; theXevaut mine near St. Just. From Andteisberg and Cluusthsl in lbe Barz; t'rcWrg. Saxony; Dlllenburg and Horbnusen iu Nassau; at Altlfeci; larioiis mines in ihe Black Foreii; Pfibram. Bobemin; Eogel near Brixlegg in Tyrol; Kiipuik, Heriengnind. Hungary. In Mexico, at Durango. Quaaajualo; Chili; Bolivia, eic. The argntiie^<w* verieiy occurs esprcinllv al Freiberg; I^ibnun; Hiiallanca in Peru, and elsewhere. The mrreuriul variety at Sclijit'Olniiz. Hungary; Schwstz. Tyrol; valleys of Angina and Casello. Tuscany.

In Ihe U. 1} , letrahedrlte occurs at lbe Kellogg mines, Arksntas. In Colorado, iu CIrar Creek and Summit Cos.; the Uluy mine, Luke Co.; witli pyrargyrie in Ruby district, Gunnison Co.. etc. Much of Ihe Colorado "gray copper" is leuuiinlile (see below). In Nevada, iibundant lu Humboldt Co.; near Ausliu in Lander Co.; I»ibel1a mine, Reese river. Iu Arizona at the Heinzelman mine; at various points Id British Columbia.

TE HNAN TITS.

Isometric-tetrahedral. Orystalg often dode'Ahedral. Also ciassiTe, compact. H. = 3-4. G. =

4-37-4-49. Color blackish lead-gray to iron-black.

Oomp.—Essentially Cu.As.S, or $4\text{Cu}, \text{S.A8}, \text{S}$, = Sulphur 25-5, arsenic 17.0, copper 57-5 = 100.

Var.—Often contains Hutmony and thus graduates into the trailled rille. The original Tennantite from Cornwall contains only copper and iron. In crystals, habit dodecahedral.

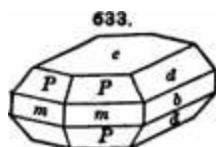
.Sandbergerite. Kupferbunde Oerm., contains 7 p. c. of zinc. Fredricite from Sweden has, besides copper, also iron, lead, silver, and tin.

Found in the Cornish mines, particularly at Wheal Jewel in Gwennap, and Wheal Unity at Gwinnear: first discovered at the Wilhelmine mine in the Pessarl: Bitterud, Norway. Near Central City and elsewhere in Colorado, AlCspeltou, Pr. Quebec, Canada. Named after the chemist. Smithson Tennant (1761-1815). See further above.

Jordanita. 4Pb8.Asi8 . Monoclinic: often pseudohexagonal by twinning. G. = 6.5. Color lead-gray. From the Binnenthal; Nagyfig. Transylvania.

Meagherite. 4Pb8,Sb,B . Orthorhombic. In slender prismatic crystals; also in fibrous. G. = 6.5. Color blackish lead-gray. From Botino, Tuscany; Marble Lake, Ontario.

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STDPANmi. Melaoglaz Owm. Brittle Bltver Ore. Orthorhombic. Axes A: $1.6 = 0.6292 : 1 : 0.6851$.

mm", $110 \text{ A liO} = 64^\circ 31' \text{ ed}$. $001 \text{ A } 031 = 68^\circ 62' \text{ eff}$. $001 \text{ A } 101 = 47^\circ 36' \text{ eA}$. $001 \text{ A } 113 = 83^\circ$ iff ek. $001 \text{ A Oil} = 84^\circ 25' \text{ eP}$. $001 \text{ a } 111 = 53^\circ 9'$ Crystals usually short prismatic or tabular. Cleavage: i, cf imperfect. Fracture subconchoidal to uneven. Brittle. H. = 2-2.5. G. = 6.2-6.3. Luster metallic. Color and streak iron-black. Opaque. Comp.—Ag.SbS, or SAg.S.Sb.S , = Sulphur 16.3, antimony 15.2, silver 68.5 = 100.

Pyx.—In the closed tube it decrepitates, fuses, and after long heating gives a faint blue color of antimony sulphide. In the open tube it fuses, giving off antimony and sulphur fumes. B.B. on charcoal fuses with projection of small particles, coats the coal with oxide of antimony, which after long blowing is colored red from oxidized silver, and a globule of metallic silver is obtained. Soluble in dilute heated nitric acid, sulphur and antimony tetroxide being deposited.

Obs. — In veins, with other silver ores, at Freiberg, Schneeberg, etc., in Saxony; Pibmtn, Bobemla: Bdicmiiiiz. Himgdrj: Andreiisberjir], the Hiira; Koiigsberg, Norway; Wheal Newton, Cornwall; Mexico; Peru; CbaBarclllo. Chili.

In Kevndu, In the Comstocct lode, Reese riier, etc. In Idabo, at the silver mines at Tanhee Fork. Queen's River district.

Named after the Aicbdiike BlepliHQ, Miuiig Director of Austria. a«ocroiilt«. $6\text{PbS} \cdot \text{BiS}$. Rarely in orthorhombic crystals; usually massive, granular. $a = d - 4$. Color lead-gray. From \wedge la, Sweden, etc.

Boagerlts. $e\text{PbS} \cdot \text{Bi}, \text{Bi}$. Hasslvc, indistinctly crystallized. $Q = 7 - \text{ST}$, Color light to dark gray. From Psrk Co., Colorado.

Eubrikanlte. $\text{Pb}^{\wedge\wedge}\text{S} \cdot \text{Sb}, \text{Si}$. Massive. $O = 6 - 41$. Color lead-gray. From Kilbrlcken, Co. dure, Ireland.

Polybasite Group. $\text{gRS}, \text{Ae}, \text{S}, \text{9ItS}, \text{Sb}, \text{S}$. Monoclinic, pseudo-rhombic.

POLTBASTTB. BprOdglnserz , Eugeiiglanz Oerm. \bullet^{\wedge} Monoclinic. Axes $a : b : c = 1 : 1.5796 : 1$, $\angle a : b = 90^{\circ} 0'$. Prismatic angle $60^{\circ} 2'$. In short six-sided tubular prisms, with beveled edges; c faces with triangular striation; in part repeated twins, tw. pi, m.

Cleavage : e imperfect. Fracture uneven. $H = 2 - 3$. $G = 60 - 62$. Luster metallic. Color iron-black, in thin splinters cherry-red. Streak black. Keai'ly opaque.

Coibp.-Ag, SbS, or $9\text{Ag}, \text{S}, \text{Sb}, \text{S}$, = Sulphur 15-0, antimony 9-4, silver 75-6 = 100. Part of the silver is replaced by copper; also the arsenic by antimony.

Pyr., etc.—In the open tube fuses, gives Sulphur and arsenical fumes, the latter forming a white sublimate, sometimes mixed with crystalline arsenic trioxide. B, B, fuses with spitting in a globule, gives off sulphur (sometimes arsenic), and coats the coal with antimony trisulphide: with long-continued blowing some varieties give a faint yellowish-white coating of zinc oxide, and a metallic globule, which with salt of phosphorus reacts for copper, and cupelled with lead gives pure silver. Decomposed by nitric acid.

Oba.—Occurs in the mines of Ouaosjusto, Mexico; at Tres Puotos. desert of Atacama; at Freiberg and Pibram. In Nevada, at the Reese lodes and at the Comstock Lode, ia, Colorado, at the Terrible Lode, Clear Creek Co. In Arizona, at the Silver King mine.

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Named from itoXdt, many, and fiiitrti, base. In allusion to the base character of the compound.

Fearcelte. $9\text{Ag}_2\text{S}, \text{As}_2\text{S}_3$. HoDOclulc, pseudo-rLombohedral. From Aspeo, Colorado; HaryRTille, Iewis and CIsrbe Co., HodIrdo.

XI. Sulpharsenate, Solphantimonate; Sulpho-antimonate, etc.

Here are included a few minerals chiefly Sulpho-salts of antimony and arsenic; also several simple oxides and rare alpha-germanates.

ENARGITE.

Orthorhombic. Axes: $d:c:b = 0.8711 : 1 : 0.8248$.

Crystals usually small; prismatic faces vertically striated. Twins: tw. pi. X (520) in star-shaped twillings. Also massive, granular, or columnar.

Cleavage: m perfect; a, b distinct; c indistinct. Fracture uneven. Brittle. H. = 3. G. = 4.43-4.45. Luster metallic. Color grayish black to iron-black. Streak grayish black. Opaque.

Comp.— $\text{Cu}_3\text{As}_2\text{S}_7$, or $3\text{Cu}, \text{S}_7, \text{As}_2$, = Sulphur 32.6, arsenic 19.1, copper 48.3 = 100. Antimony is often present, cf. feniattinite.

Pyr.— In the closed tube decrepitate. It gives a sublimate of sulphur; it also gives a sublimate of antimony. In the open tube, it fuses, the powdered mineral gives off sulphurous fumes, the latter changing to iron pyrites and some antimony oxide. B.B. on charcoal fuses, and gives a faint coating of iron pyrites, antimony, and zinc; the roasted mineral yields the usual globule of metallic copper. Soluble in aqua regia.

Occ.— From Morococha, Cordilleras of Peru; in Chile and Argentina: Mexico; Halzenknapp, Brixlegg, Tyrol; Maocayan, Luzon.

In the U. S., at Brever's gold mine, Cheslerfluß Co., 8, Colorado; at mines near Central City, Ojipin Co.; in Park Co. at the Missouri mine, etc. In Utah: also in the Tintic district: Butte, Houma.

Character.— Cerimpite identical with enargite. From the Clara mine. Schapbach, Baden.

Enargite. Composition of enargite, but superoxide in crystallization. Massive. O. = 1. Color black. From the island of Luzon.

Enargite. $8\text{Cu}, 8\text{S}, 8\text{As}$, in orthorhombic enargite. G. = 4.57. Color gray with tinge of copper-red. From the Sierra de Famatina, Argentina.

Zasthoconite.— $8\text{Ag}_2\text{S}, \text{As}_2\text{S}_3$. In rhombic rhomboidal crystals; also massive, reniform. G. = 5. Color orange-yellow. From Freiberg.

ZlplboD]aiigeril«.—8PbS.Bb,Bi. Id striated prismatic needles and granuhr. G. = 6'31. Color dark bluUh gray to black. From Altenberg.

Bplgsnils.—Perhaps 4Cu,8.IIFe8.A!t,S.. In short prisms resembling areeDopyrlte. Color sieel-gny. Prom Wittichen, Baden.

STAHHTm. TlnPyriles. Bel I-metal Ore. Zinnkles <7mn. Isomotric-tetrahedral. Massive, granular, and disseminated. . Cleavage: cubic, indistinct. Fracture uneven. Brittle. H. = 4. G. = 4*3-4-522; 4-506 Zinnwald. Luster metallic. Streak blackish. Color steel-gray to iron-black, the former vrhen pure; sometimes a bluish tarnish; of ten yellowish from the preaence of chalcopyrite. Opaque.

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SLG DESCRIPTIVE UIKEKALOQT.

Comp.—A Giilpbo-stannate of copper, iron and sometimes zinc, Cu,FeSiiSj or Cu,S.FeS.SriS, = Sulphur 299, tin 27-5, copper 295, iron 13-1 = 100.

Py., etc.—In tbe closed tube decrepiUtes, Bud giFes a faitut sublimaie; In lUe open tube Hulpburous fumeH. B.B. on cbarcoil TiiseB lo a globule, nbicli iu O h. gives oil aulpliur, Aud cuals Uii; coil wllh tiu dioxide; (Ije roasleu mineral treated willi borax gives reactlouB for Iron ftud cop|>er. Decomposed by uUric acid, iiSoiding a blue Boluiiou, wiili separaliou of aulpliur and liu dioxide.

Ob«.—Pormerly fouud at Wlieal Rock, Coruwall: nud at Cam Breii: more reteutly iu grraulle at St. HlcWl's Mourn: also al Steuna Gnyuu, etc.; at tlie Crunebane mine, Co. Wlcklow, lu Ireland: Ziuuwald, iu the Eizgeliirge. From llie Black HiHb, 8u. Dakota.

Argyrodlt*. A silver eulplio-gvrmatiate, Ag>OeSi or 4Ag,B.Ge8,. Isouielric, crysiuls usually ludUtluct; also niusslve. compact. H. = 2 5. G. = e'o8&-6-lll. LuBter metallic. Color steel-gray on a fresh frHtlure, wilb a tinge oC red turning to violet. From llie HimmelsrOrst mlue, Freiberg. Siixony.

Oanfialdlte. AgiSnSi or 4Ag,S.SiiS,, tbe lio iu part replaced by germanium. Isomelric,

Griapsteiratiedral; iu ocluhedroDBwItbddlO). o. = 6'28. Luster metallic. Colorblnck. Paz, Bolivia.

PranckaiU. Perbaps PbtSb,8uiS,, Groth. HaadTo. O. = SBS. Color blackish gray to black. Boltrla.

OyUndriU. Eyllodrlte. Perhaps Pb.SbtSn.S,, Groth. H. = S-6-& Q. = 5-42. Luster metallic. Color blackish lead-gtay. Poop6, BoUrla.

HALOIDS.—CHLORIDES, BROMIDES, IODIDES ; FLUORIDES.

IV. HALOIDS—CHLORIDES, BROMIDES, IODIDES; FLUORIDES.

I. Anhydrous Chlorides, Bromides, Iodides; Fluorides. n. Oxychlorides; Oxyfluorides. ni. Hydrous Chlorides; Hydrous Fluorides.

The Fourth Class includes the haloids, that is, the compounds with the halogen elements, chlorine, bromine, iodine, and also the less closely related

I. Anhydrous Chlorides, Bromides, Iodides; Fluorides.

Calomel. Horn Quicksilver. Chloride of Mercury, Hg_2Cl_2 . Tetragonal. Axis $a = 1.7299$; $001 \text{ A } 101 = 59^\circ 52'$. Crystals sometimes tabular; also pyramidal; often highly complex.

Cleavage: a rather distinct; also rhombohedral. Fracture conchoidal, Sectile. H. = 1-2. G. = 4.89. Luster adamantine. Color white, yellowish gray, or ash-gray, also grayish, and yellowish white, brown. Streak pale yellowish white. Translucent—subtranslucent. Optically +.

Comp.—Mercurous chloride, Hg_2Cl_2 , = Chlorine 15.1, mercury 84.9 = 100.

Pyr., etc.—In the closed tube volatilizes without fusion. Residue in the cold part of the tube as a white sublimate; with sodium gives a sublimate of metallic mercury. B.B. on charcoal volatilizes, leaving the colorless residue. Insoluble in water, but dissolved by nitric acid; blackens when treated with alkalis.

Obs.—Usually associated with cinnabar. Found at Mercur, in the Palmyra; at Idria in Carniola; Almaden in Spain; at Mt. Arenal near Beltrán in Bervin.

Calomel is an old name of uncertain origin and meaning, probably from *cal* (white) and *mel* (sweet), and hence the name being sweet, and the compound the most sweet of early chemistry; or from *cal* (white) and *mel* (sweet). Hack.

Mercurous chloride. Cu_2Cl_2 . Greenish, massive. Cleavage cubic. H. = 3-4, brittle. Colorless to white or grayish. From Nanloko. Chill; New Bolivia. W. Lea.

Mercuric iodide, Cu_2I_2 , Isometric-tetrahedral. Color oil-brown. Broken Hill lodes. New South Wales.

Halite Group. HCl, BBr, RI. Isometric. Halite NaCl Embolite $\text{Ag}(\text{Cl}, \text{Br})$

BylTite KCl Bromyrite AgBr

Bal Ammoniw (NH₄)Cl Iodobromite Ag{Cl.Br,I}

Cerargyrite AgCl Mierute AgI

The Halite Group includes the halogen compounds of the closely related metals, sodium, potassium, and silver, also ammonium (NH₄). They crystallize in the isometric system, the cubic form being the most common. Sylvite and sal-ammonia are trigonal, and the same may be true of the others.

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BAUTB. Common or Rock Salt. Sierrita, Bergaalt Oerm. Isometric. Usually in cubes; crystals sometimes distorted, or with cavernous faces. Also massive, granular to compact; less

often columnar.

Cleavage: cubic, perfect. Fracture conchoidal. Hard brittle. H. = 2-5. G. = 2.166; pure crystals 2.165. Luster vitreous. Colorless or white, also yellowish, reddish, bluish, purplish. Translucent to translucent. Soluble; taste saline. Refractive index 1.5442 Na. Highly diathermanous.

Comp.—Sodium chloride, NaCl — Chlorine 35.46, sodium 22.99 = 100. Commonly mixed with calcium sulphate, calcium chloride, magnesium chloride, and sometimes magnesium sulphate, which render it liable to deliquesce.

Prp., txc. — 111 lbs closed tube fuses, will decrepitate; will fuse on lie

Eklinum wire coil (lie flint clep yellow. Aditt tu u Mil of phosphorus bead which as beo saluted with oxide of copper, it colors the flint a deep azure-blue. Dissolve gradually in three parts of water

" lissoluble (taste), so Uessa. perfect cubic cleavage.

It is soluble in water, and soluble; also in the water of the ocean and in seas.

The principal salt mine in Europe is at Stassfurt, near Magdeburg; Wieliczka, in Poland; in Tyrol; and in the range; through Helmen in the Ba-Taru, Hailein in Salzburg, Hallstadt, Leoben, and Ebensee, in Upper Austria, and Austria in Styria; in Runenry. at Harmoros and elsewhere: Transylvania: Wollschick. Galic. in Upper Silesia: Vic and Dieuze in France; Valley of Cadun and elsewhere in Spain; Bex in Switzerland; and Nertschinsk in Siberia. At the Austrian tolls, where it con-

ina much clay, the salt is dissolved in large clauibers. and the clay thus preclialcd. After A time ihe water, saturated with Ihe silt, is rouved by nqutducls lo evaporatiiig; houses, and the chambers, after beioiE cleiired out, are agln filled.

Suit also occurs, foriilng liilla and covering exieuled plidus. near Lake Uruoila, the Caspian Sea, eic. In Alceria: in Abywinia. In India In enormous deposlm iu the Salt Rani^ of Ihe Punjnb. In China and AsixElc Russia; in South America, in Peru, and at Zlpaqiiera and Nemiicnn. the former ii Inree nilnc lore exploieil in the Cordilleras of U. S. of Ciiliimbini; clenr aalt is obtained Trom the Cerro ile Bnl, San Domingo.

Tn the United Stales, anlt has been frund iu laree amount In cenlral and western New Torlt. Salt wells had long been woikeH in this legion, but rock salt is now known lo exist over a large area from Ithnca iil. the bend of Cnyufra l^ake, Tompkiiiiis Co.. and Cannntlaigiia Ijike, Oiiiiirln Co.. Ihronuh Livinqsiiin Co., nish Genesee, Wjouiug. and Erie Cos. The aalt Is found in lieds with nn iivernire th<ckne)>» of 7S feet, but sometimes much iblcker, and at vniTing depths from 1000 to 2000 feet and nioie; the depth inreases soulhwanl wlih the dip of tbe airala. Ilic rocks bidong lo tbe Sallna period of ihe Upper Silurian Salt bus aiso been found near Cleveland. Ohio, asiociated with Eyp«um; in Washnglon Co., West Virginia, in the Ilolsion and Kanawha vallevs: in Kansas; at Pettle Anse. LoulFiana; along Ihe Rio Virgin in Lliie«ln Co.. Nevada, iii extensive beds of great purity: in Utah. near Nephi, Juab Co., and Salinn, Sevier Co ; in Arlmiia, on the Hio Verde, with Ibeuard-ile etc.: In Cnlifomla, San Dtego Co.

Brine 'springs are very numerous In the Mi'ldle and Weaiern Sinles. Vast lakes of mlt water exist bi many parta of the world. Tbe Great Sail Lake in Utah is 3,000square mi1«s ill area; L. Oale found in thia water 20-]96 per cent, of so<liim chloride. The Dead and Caspian seas are salt, and tbe waters of tbe former ronlain 20 to 26 parts of solid matter in 100 parts. Biidium chloride is the prominent salt present in thie ocean.

Huantajayita. $20\text{NaCl} + \text{AgCl}$. In cubic crystals and na an Incniatation. H. = 2, not aectlle. Color white. From Huantajaya, Tarapnca, Chill.

STX>V1TJB.

Tsometric-pUgihedral. Also in granular crystalline masses; compact. Clesvage: cubic, perfect. Fracture uneven. Brittle. H. = 8. G. = 1'97-

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1-99. Laster vitreouB. Colorleaa, white, bluish or yeilowiah red from inclnsions. Soluble ; taste resembliog that of common salt, but bitter.

Comp.—Potassium chloride, KCl = Chlorine 47-6, potsasinm 52-4 = 100. Sometimes contains sodium chloride.

Pyr., etc.—B.B. Id (he plaioum Iwp fuaea, Kod riveB a vtolet color to the outer lUme. DisaolTCH completely In water. Heated with sulpltuic add rivei off bvdroclilortc acid eas.

Ob«.—Occurs aiVeauTius. about tbe fumarolee of the Tolcauo. Alsoat Staaafurt' at Leopoldsball (UopoUite); at Kalusz lu Qajicla.

Bal Aumoniao. Ammonium chloride, NH_4Cl . Obteired u a wille facruBtalioa about Tolcaaoea, as at Etaa, VesuTiua, etc.

OBRAROTSITB. Silberhorners, Honullber Germ. Horn Bllver.

Isometric. Habit cubic. Twins: tw. pt. o. Usually massive and resem-blitig wax; sometimsa columnar; often in cruBts.

Cleavage none. Fracture Bomewhat conchoidal. Highly Bectile. H. = 1-1'5. G, = 5'552, Lnster reeinouH to adamantine. Color pearl-gray, grayish green, whitish to colorless, rarely violet-blue ; on esposure to the light turns violet-brown. Transparent to translucent. ludei, w, = 20611

Comp.—Silver chloride = Chlorine 24-7, silver 75-3 = 100. Some varieties contain mercury.

Pjnr., etc.—lu the closed tube fusea wltbout decomposition. B.B. on cliaicoiYl gives a globule of metallic silver. Added lo a tiead of salt of pbospLorus, previotiHly saiurati d nitli oxide of copper nnd beated ie o.Y., imparla an inteise azure blue to the flnme. Insoluble in nitric acid, but soluble in ammonia.

Ob».—Occurs in veins of clay slate, accomriinying oilier ores of silver, nnd usually only in Ibe higher piitis of iLi-sc veins. Ii hm also been observed with oclitrous VRileliea of bripwQ iron ore: ulso witlj several cupper ores, ailecite, bsrite, etc.; upon slibicoiile.

The Inrgesl masses are brought from Peru, Cbili. and Mexico, wbre it occurs vitb native silver. Also once oblslued from Johaungeoi^enstadt and Freiberg; Andi-easljerg (eaniiy var.. ButUrmilchtrt); occurs in tbe Altai; at Kongsbergin Norway; in Alsnce.

In tbe V. B., in Colorado, near Leadville, Lake Co.; near Breckenridge Summit Co., nnd e1s"wlere. In Nevada near Austin, Lander Co.; at mines of Comslock lode. la Idnbo. nt ilie Poorman mine, in crysials; also at various otlier mines. In Utah, in Beaver, Summit and Salt Lake counliee.

Named from K^tiai, horn, and anyvooi. tilner.

Bmbollta. Silver chloro-bromide. Ag(Br.Cl), the ratio of cblorine to bromine varying widely. TTainnly massive. Besenibles cerargyrite, but color gtayiBb green lo yellowM green and yellow. Abundant In Chili, leas so elsewbere,

Bromyrite. Silver bromide, AgBr. G. = 5B-.«. Color bright yellow lo amber-yellow; slightly greenish. From Mexico: Chili; Huelgoet in Brittany.

Iodobromite. SACl.SAeBr.Agl. Isometric; o with a. Q. = S-718. Color sulphur-yellow. greenish. From near Dernbach, Nassau.

Boloidalite. Silver iodide, AgI. crystallizing in the Isometric system ; probably tetrahedral-like mineral (p. 817). In bright yellow crystals from the Broken Hill Silver Mines. New South Wales.

Onnioidarffyrilite. A copper-silver iodide, occurring as a sulphur-yellow mineral. Huancabamba, Peru.

Iodite*. Silver Iodide, AgI. Hexagonal-hemimorphic (Fig. 211, p. 71); usually in thin plates; pale yellow or green. Ck = 5-6-5'7. From Mexico, Chill, etc. Lake Valley. Sierra Co., New Mexico.

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DBCKIPTITB MINERALOGY.

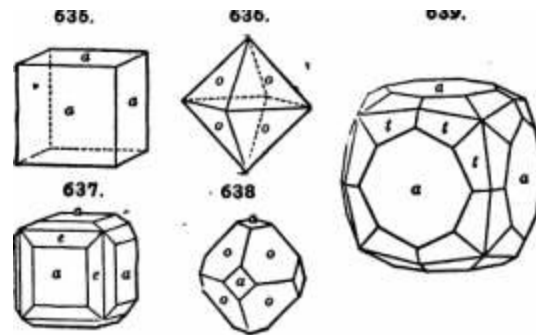
Fluorite Group. UF_2 .

The species here included are Fluorite, CaF_2 , and the rare $CaCl_2$. Both are isometric, habit cubic.

Hydrophilite,

Fluorite

Fluorite Sp. Fluorite Q. 636.



Isometric. Habit cubic; less frequently octahedral or dodecahedral; forms (310), e (310) (fluproids) common; also the vicinal form C(32'1'0?), producing a striation on a (Fig. 640); hexoctahedron (431) also common with the cube (Fig. 639). Cubic crystals sometimes grouped in parallel position, thus forming a pseudo-octahedron. Twins: tw. pi. o, commonly penetration-twins (Fig. 640). Also massive; granular, coarse or fine; rarely

columnar; compact.

Cleavage: 0 perfect. Fracture flat-conchoidal; of compact kinds splintery. Brittle. H. = 4. G. = 301-3-25; 3-18 cryat. Luster vitreous. Color white, yellow, green, rose- and crimeon-red, violet-blue, aky-blue, and brown; wine-yellow, greenish blue, violet-blue, most common; red, rare. Streak white. fraraparent—subtranslucent. Sometimes shows a bluish fluorescence. Phosphoresces when heated (p. 191). Refractive index: n, = r4339 Na.

Comp.—Calcium fluoride, CaF, = Fluorine 48'9, calcium 51'1 = 100. Chlorine is sometimes present in minute quantities.

Id colon: <fi> flbroua to SDtor (e) coHrse lofine yields a greeD pbo*-

gniDUlul-; {d} earthliy, dull, aud pboresceit llgllt.

Fyr., eto —In llie dosed tube decrepitates imd iilioaplioresces, on cbnruiil fuses, coloring tlie flume red. lo an enamel wlilcli reac . . -

Fused iu ail open tube with fiiwd BLlt oi phoapbonis jrive* the reaction for fluoHDe. Treated witii anlpliurip Rciil (fives fumes of lijdronuoric acid which elcb rlass.

Diff.—DistiDgiiUhe'l l>y ita crystidlice form, octaliedrai clntvage, relative softness (as compared with certain preoinus slonea, nlso with tlie feldspars); etching power when treated with siilpburic acid. Does not effervesce with neid like cnlcite.

Obs.—Sometimes In beds, but generally in veins, in gneiss, mica slate, clay slate, and also in limesiones. both crystalliDe aod uncrjslnlld. and sandstoDes. Often occurB as (he gangue uf metallic ores, especially nf lead. In the Norili nf England, it is the giogue of the lead veins, which intersect the coat formalion In Nnrthumberland, Cumlwrland, Durham, and Yorkshire. In Derbyshire It U abundant, aind nlsn In Cornwall, where ihe vdna intersect metamorpliic rocks. The Cumberland and Derbyshire localities cspeclBly iiave afforded magniflcent specimens. Common in the mining disirict of Saxony; floe near

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EougBberg In Noma;. Id the dolomitea of SU Oolliard occurs in pluk ocUUedront. Baruly in toIcuuIc re^oua. hh in tbc VesuTiaD lava.

Sunie iuculities id llie U, S. are, Tniinbui], Conn. (cAlorophaiu); Muscolouge Lake. Jellir-iiu Co.. N. Y., and Matoiub, St. Lawrence Co.. boib lu very large sen-ErMD cubei' FriUkliiiFuruuce. N. J.; Amelia Courl Houw), V(u; QallaUn Co., III.; St. Louis, Mo.

BydzopliUito. Clilorocalcite. Calcium cbloiidi^ CaCl. lu whiw cubic crfsiats or as uu iucrusiaiimi at Vesuvius.

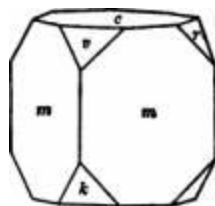
The (olovrlug are froui Vesuvius: OUorotnunaalta, HeCli: Softocbita, MnCli - OhloraUainiilU, AlCl. + 3H,o; Molydt«, FeCl..

Sallaita. Huguesium fluoride, MgFi. In prisinallc tetrngonal cryslals. H. = 6. ■Q. = 3-9?-S']3. Colorless. From, the moraine of ilie OebroulnE glacier Id Mavoj.

LaWToioit*. Ferrous chloride, FeCli. Occurs iu meteoric iron.

CotatuiiU. Lead chloride, PbCl.. In acicular crysbils (orthorhombic) and la semi. <;rystM.>iue uiibbi.-;- Soft. Q'=:B-24. Color wliile, yellowish. From Vesuvius; also Tunpiii'a. Cbili.

Tysoolte. Fluoride of the cerium melsK (Ce.Lii.DljFt. Iu tbclk bexagonal prisms, and massive. Cleavage: e pertect, H =4-3-6. G. = BIS. Color pale vrBZ-yenow, cbaDgiDg to yelluwiiJi and reddish bruwu. From Uiegranile of Pike's Peak, Et Paso Co., Colorado.



ORTOUTE. Eisatein Germ.

Moiioclinic. Axes d : t : i = 0-9663 : 1 :1-3882 ; ft mm'-, 110 A liO = 88° 2'. ck, 001 A lOl = 55* IT.

em. 001 A 110 = 89° W. tr. 001 a Oil = M° 14".

eo, 001 A 101 = 55° 2". op. 001 a HI = 88* 18".

Crystals often cubic in aspect and gronped in parallel position; often with twin lamellEe. MaasiTe, cleavable.

Cleavage: c most perfect, alEO m, it (lOlj somewhat less so. Fracture uneven. Brittle. H. = 25. G 2-95-3-0. Luster vitreous to greasy; somewhat pearly on V. ColorleBS to snow-white, sometimes reddish or brownish to hrick-red or even black. Transparent to tranBlucent.

Comp.—A fluoride of sodium and aluminium, Na.AIF, or SNaF.AIF, = Fluorine 54-4, aluminium 13-8, sodium 3S-8 = 100. A little iron geBt|nioxide is sometimes present as impurity.

Pyr., etc.—Fnstble in small fnigments in the flame of a candle. B.B. In the open lube faented so that the flume enters Ihe lube gives off hydrofluoric xcid. etching tiie glds. In the Torceps fuses very easily, coloring the Bnme yellow. On charcoal fuses easily to u

clear bead, which on cooling becomes opaque; after long blowing, the assay spreads out. The fluoride of sodium is absorbed by the coal, a suffocating odor of fluorine is given off, and a crust of alumina remains, which, when heated with cobalt solution in O_2 , gives a blue color. Soluble in sulphuric acid, with evolution of hydrofluoric acid.

DiS.—Disintegrated by its extreme fusibility, and its yielding hydrofluoric acid in the open tube. $\text{Al}_2\text{Si}_2\text{O}_7$ by its cleavages (resembling cubic cleavage) and angles.

Obs.—Occurs in a bay in Arksukfjord, Id West Greenland, at Ivluui (or Evigtok). About 12 m. from the Danish settlement of Arksuk, where it constitutes a large bed in a granitic vein in a gray gneiss. Cryolite and its alteration products, pectolite, thomsonite, proeopite, etc., also occur in limited quantity at the southern base of Pike's Peak, El Paso county, Colorado, north and west of Mount Peter's Dome.

Named from the Greek $\alpha\lambda\upsilon\sigma$, salt, hence meaning ice-salt, in allusion to the translucent cleavage in crystals.

Orthoclase. $5\text{NaF} \cdot \text{Al}_2\text{Si}_2\text{O}_7$. In small pyramidal crystals (tetragonal); also massive granular.—H. = 2–4. $\rho = 2.84$ –2.90. Color snow-white. From near Mts. in the Himalayas; also with the Greenland cryolite.

Albite. A fluoride of potassium and silicon. In grayish tabular crystals; isometric. From the fumaroles of the crater of Vulcano.

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n. Oxychlorides, Oxyfluorides.

ATAOAMTTB.

Orthorhombic $a:b:c = 0.6613 : 1 : 0.7515$.

643.

$m = 110^\circ$, $A = 67^\circ$. $\alpha = 63^\circ$, $\beta = 45^\circ$,

$\gamma = 78^\circ$, $\delta = 51^\circ$. $\rho = 1.10$, $H = 18$.

Commonly in slender prismatic crystals, vertically striated. In fused crystalline aggregates; also massive, fibrous or granular to compact; as sand.

Cleavage: 6 highly perfect. Fracture conchoidal. Brittle. H. = 3–5. G. = 3.75–3.77. Luster adamantine to vitreous. Color bright green of various shades, dark emerald-green to blackish green, W_4 Streak apple-green, transparent to translucent. X^2 Comp.— $\text{Cu}_2\text{Cl}(\text{OH})_2$, or $\text{CuCl} \cdot 3\text{Cu}(\text{OH})_2$, Chlorine 16.6, copper 19.9, cupric oxide 55.8, water 127 =

Pyr., «tc.—la the closed tube Kives off much water, and forma a eray iubllmate, B.B. cbai-coal fuses, coloring the O.F. azure-blue, wfilia green e^, and giving two coaliLss, B brownlsli and tbe other grayish white; couliiiued blowing yiuldB a. globule of melafiic iper; the contiug*. touched with the R.F., volatilize, coloring the flame azure-blue. In as easily soluble, -ba —Orlginalli . . and Bolivia; ? the U. S., with cuprite, etc., at the United Vei-de mine, Jerome, AHs

Peroylita. A lead-copper oxychloride, perhaps $PbCuO, HtCl$, Iu sky-blue fubes. From Sonora, Meilco; Atttcama; Bullvla, ■'''' "■-'■ » t. i. _ i. r. __ ., ■.

California, U a percylyte containing a ing to be tetragonal.

HaUooldt«. Lead oxychloride, $PbOCl$,. In tabular telragonal crystals. G. = 7-21. Luster adamantine to pearly. Color yellowish or eligbly greeDlab. From Cromford, near Haticli, Derbyshire.

Mendlplte. $Pb O-Cl$, or $PbCl_{1..2}PbO$. In flbrons or columnar masaes: often radiated. H = 3-5-8. G. = 7-7-1. Color white. From the Heudip Hills, Somersetsbire, England; near Brilon, Weatphalia.

Laorioulte. $PbClOH$ or $Pba, Pb(OH)$,. In minute prismatic colorless cryctals (orthorhombic), in ancient lead slags at Liurioti, Qreece. Fiedlerlte, associated with laudonite. Is probably also a lead oxychloride: In colorless monocloic crystals.

Panfioldlle. $Pb.OCl$, or $Pbo.2P\&Cl$,. In white hexagonal crysUls. Laurion, Greece. Davleilta. A lead oiychloride of uncertain composition. In minute colorless prismatic crystals (orthorhombic) from the MIna Beatriz. Sierra Qorda, Atncama.

Bohwartaembarglta. Probably $Pb(I,Cl)i.2PbO$. In druses of small ciyslals; also in crusts. Q. -= 6'3. Color honey-yellow. Desert of Atacama.

PlnooBrita. $(Ce.La.Di),OP$,. Heiagonal. H. = 4. Q. = B-7-6-9. Color reddish yellow. From Oaterby In Dnlarne. Sweden.

Noosrits. Perhaps $3(Ca,Mg)F,(Ca,Mg)o(7)$. In white hexagonal adcular crystals fwHO bombs In the tufa of Nocera, Italy.

Danlwedt*. An earthy yellowlah oqrcUwIde of bismuth. From Bolivia,

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UL Hydrou OUoridei, Hydxoiu Fluorides, etc: OARNAUJTB.

Orthorhombic. Crystals rare. Commonly massive, granular. No distinct cleavage. Fracture conchoidal. Brittle. H. = 1. G. = 1.60. Luster shining, greasy. Color milk-white, often reddish. Transparent to translucent. Strongly phosphorescent. Taste bitter. Deliquescent.

Comp.— $\text{KMgCl}_2 \cdot 6\text{H}_2\text{O}$ or $\text{KCl} \cdot \text{MgO} \cdot 6\text{H}_2\text{O}$ = Chlorine 38.3, potassium 14.1, magnesium 8.7, water 39.0 = 100.

Occurs.—Occurs at Blasfurt, in beds, alternating with the beds of common salt and kieserite.

Associated with camallite, is said to be $\text{SECl} \cdot \text{FeCl}_2 \cdot 5\text{H}_2\text{O}$. Böhmforsite. $\text{HgCl}_2 + 6\text{H}_2\text{O}$. Crystals granular; colorless to white. From Leopoldsball, Prussia.

Kramanite. $\text{KCl} \cdot \text{NH}_4\text{Cl} \cdot \text{FeCl}_2 + \text{H}_2\text{O}$. In red octahedrons. From Veraylus. Erythroriderite. $\text{KCl} \cdot \text{FeCl}_2 \cdot \text{H}_2\text{O}$. In red tabular crystals. Vesuvius. Taohbydilite. $\text{CaCl}_2 \cdot 2\text{MgCl}_2 + 12\text{H}_2\text{O}$. In wax- to bone-yellow masses. From Bressfurt.

Praepositite. $\text{CaF}_2 \cdot 3\text{Al}(\text{PO}_4)$. In monoclinic crystals, or granular massive. H. = 4-6. O. = 2-88. Colorless white, grayish. From Altenberg, Saxony; St. Peter's Dome near Pike's Peak, Colorado.

Practolite and Thomsonite, occurring with cryolite in Greenland and Colorado, have the same composition, $\text{NaF} \cdot \text{CaF}_2 \cdot \text{AlF}_3 \cdot 4\text{H}_2\text{O}$. Both occur in monoclinic prismatic crystals: prismatic angles for practolite, 98° 86', crystals twins, orthorhombic in aspect. For Thomsonite, 88° iff, crystals often resembling cubes (tetragonal), also prismatic distinguished by its basal cleavage; also massive.

OMTentite. $\text{C}_8\text{F}_{16} \cdot \text{Al}(\text{F},\text{OH})_8 \cdot \text{H}_2\text{O}$. Earthy, clay-like. Occurs with cryolite.

Haltonite. $(\text{Na}^+ \text{Mg}^{2+})\text{F}_8 \cdot 8\text{Al}(\text{F},\text{OH})_8 \cdot 2\text{H}_2\text{O}$. In colorless to white, isometric, octahedrons. H. = 4-6. G. = 2.56-2.62. With the Greenland cryolite.

Tallngite, a hydrated copper chloride from the Botallack mine, Cornwall; in blue globular crusts.

Foatite. A hydrous oxychloride of copper occurring in deep blue prismatic crystals (monoclinic) at the Copper Queen mine, Bisbee, Arizona.

Ttroomite. $(\text{T}, \text{Er}, \text{Ce})\text{F}_6 \cdot 6\text{CaF}_2 \cdot \text{H}_2\text{O}$. Has cleavage to granular and earthy. H. = 4-6. O. = 8-4. Color violet-blue, gray, reddish brown. From Falun, Sweden, etc.

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DESCRIPTIVE HISTORY,

V. OXIDES. L Oxidei of Silicon, n. Oxidea of the Semi-Metala: TeUarlunii, Ammiio, Antimcmy, Bismuth; also Molybdenum, Tungsten. m. Oxidei of the Metali.

The Fifth Class, that of the Oxides, is subdivided into three sections, according to the positive element present. The oxide§ of the non-metal silicon are placed by themselves, bat it will be noted that the compounds of the related element titanium are included with those of the metals proper. This last is made necBSSary by the fact that in one of its forms TiO, is isomorphouB with MnO, and PbO,.

A series of oxjgeu compounds which are properly to be viewed as salta, e.g., the B|>ecies of the Spinel Group and a few others, are for conTenience also included in this class.

I. Oxides of Silioon.

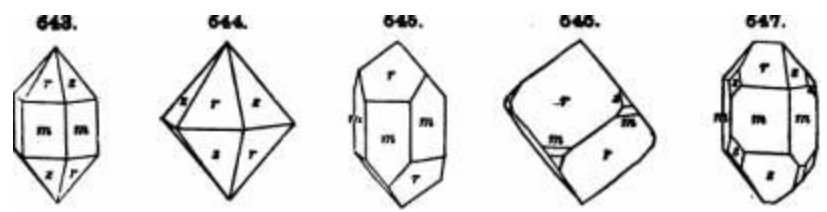
QUARTZ.

Rhombohedral-trapezohedral. Axis: i =109997.

rf. lOil A ilOl = 85' «'. rt, lOll A Olil = «° 16'. mr, lOiO A lOll = 88' Iff.

WW. lOiO A Olil = «* W. m#, lOlO A llSl = 8r58'.

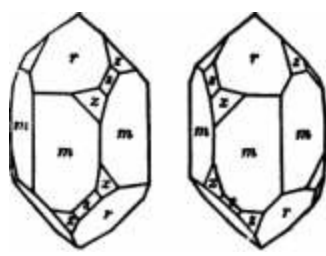
mi. loio A 5i6i = ia° r.



Orystals commonly pnsmatic, with the m faces horizontally striated; ^^^ ^^^ terminated either by both rhombonedrons, or

** by one only; the nredominating rbombo-

hedron is in almost all cases r (lOll). Often in double six-sided pyramids or quartzoids through the equal development of r and z; when ;■ is relatively large tlie form then has a cubic aspect {rr ~ t'^" 46'). Crystals fre-q^uently distorted, when the correct orientation may be obscure except ns shown by the striatrons on m. Crystals often elongated to acicular forms, and' tapering through the oscillatory combination of snrcessive rhombohedrons with the prism. Occasionally twisted or bent. Frequently in radiated maaaeswith a surface of pyramids, or in dmsee. Simple crystalH nre elllier ridit- or IcFi-lmnded. On a right /laiuUd crystal (Fie. S48). «, It preient. lie* lo lbe right of the m fnce, which b below the predomlDaling pliii rnombo-



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bedroD r, sad wllh this belong Ibe plus right trapezohedrouE, aa x, and minus left trapezohedrOQS (cf. Fig. B70, p. 831, alwitdBft). On a itflhandad cryual (Fig. 649). * lies lo lIhe Ipfl of iliem below r.Hudiritb il Uie plus teFt and mlDUH right trapezohedrous, also a (right). The right- uud lefl-lianded forniB occur together ouly Id twin*. In the absence of trape-zobednl faces the Btrialions on * || edge r/m), if disiluci, serre to distlnKiilsb the faces r aoil (, and heuce sliuw the right- and left-baadcd character of the crjstals. The tlght-ftpd U-ft-haudcd character is also revealed bj etching (Art. 865) and by pyro-electridty (Art. ISOj.

Twins: (1) tw. bxIb i, axes pumllel, (3) Tw. pi. a, Boroetimee called the Brazil law, usually as irregular penetratiou-twins (Fig, C50). (3) Tw. pi, S (1122), contact-twins, the axes ci-osBing at an angle of $84^{\circ} 33'$. See further p. 137 and Figs. 393-394. Massive forms common and in great variety, passing from the coarse or fine granular and crystalline kinds to thoBe which are flint-like or cryptoci'yatalline. Sometimes mammillary, stalactitic, and iu concretionary forms; as sand.

Cleavage not distinctly observed; sometimes fracture surfaces (| r, z iind m\ developed by sudden cooling after being heated (see Art. 258). Fracture conchoidd to sub-conchoidal in crystallized forms, uneven to splintery in some massive kinds. Brittle to tough. If. = 7. G. = 2'653-2-(i60 in crystals; cryptocrystalline forms somewhat lower (to 2-60) if pure, hut impure massive forms {e.g.,

jasper) higher. Luster vitreous, sometimes Basalaectioulnpolarlz«illight.ahow-greasy; splendent to nearly dull. Colorless . "" """"""-""■'~"TM ""1"1.,,,,"h when pnre ; often vurions shades of yellow, red, brown, green, blue, black. Streak

white, of pure varieties; if impure, often the same as the color, but much paler. Transparent to opa»]ue.

Optically +. Double refraction weak. Polarization circular; right-handed or left-handed, the optical character corresponding to right- and left-handed character of crystals, as defined above; in twins (law 2), both right and left forms soraelimes u'lited, sections then often showing Airy's spirals in the polariscope (cf. Art. 366, p. 202, and Fig. 650). Rotatory power proportional to thickness of plate. Refractive indices for the D line, $\mu = 1'54\text{-H8}$, $\mu_e = 1\text{-}^{\wedge}5328$; also rotatory power for section of I""""- thickness, $\alpha = 31\text{-}71$ (D line). Pvroelectric ; also electric by preBsnre or piezo-electi-ic. See Arts. 420, 421. On etchiug-

figure B, see Arts. 265, 266.

Comp.—Silica, or silicon dioxide, SiO_2 , = Oxygen 53.3, silicon 45.7. In its varieties often mixed with a little opal silica. Impure varieties contain iron oxide, calcium carbonate, clay, water, and various mineral inclusions.

Var.—1. *Phenocrystalline*: Crystallized, vitreous. Luster 3. *Cryptocrystalline*: Flint-like. *Massive*.

The first division includes the ordinary vitreous quartz, whether having crystalline structure or not. The varieties under the second are in general formed upon somewhat more by attrition, and by chemical agents, as hydrofluoric acid. These are those of the first. Some are made up of layers, as agate, successive layers are unequally eroded.

A. *Phenocrystalline* or *Vitreous* Varieties. Ordinary *Crystalline*; *Colorless* or nearly so, whether in distinct crystals or not. Here belong the *Bristol diamonds*, *Lake George diamonds*, *Brazilian pebbles*, etc. Some variations from the common type are: (a) cavernous crystals; (A) cap.

; *Coagulated*

quartz made up of separable layers or caps; (e) druse quartz, & crust of water or mineral crystals; (d) *interstitial* quartz, often separable into mediate parts having pyramidal crystalline structure; (e) *Quartz*, rarely delicately such as a kind from *Quintana Roo*, *South Africa*, derived from *crocidolite* (see *calcite* below, also *crocidolite* p. 4M).

Striated; *Star-shaped*.—*Crystalline* within the crystal *colorless* or colored *radiation* (long thin *interstitial* plates. Occasionally exhibits distinct *asterism*).

Ametrine; *Amethyst*.—Clear purple, or bluish violet. Color perhaps due to *maugueite*.

Variegated.—Honey-red or pink, but becoming paler on exposure. *Common massive*. Luster *subvitreous* a little iridescent. Color perhaps due to *titaniite*.

Yellow; *Faltt* *opac* or *Citrine*.—Yellow acid pellucid; resembling yellow *topaz*.

Smoky; *Smoky* *Stone*.—Smoky-yellow to dark smoky brown, luster often *translucent*; *vitreous* to brownish black. Color is probably due to some organic compound (*Forsterite*). *Common* from the locality at *Cairngorm*, S.W. of *Banff*, in *Scotland*,

is given to nearly black

Milk—*White* *stone* *is* *given* to nearly *opaque*. Luster often greasy,

Situla, or *Sapphire*-*guar*.—Of *indigo* or *Berlin-blue* color; a *rare* variety.

tiag riilie. —Iuc1r>sii)gacicular irysials of rutile. Olbcr included minerals in acicular .Tuniii arc; bluck toiirmnllue; eOlbile: alibnile; asbestiis; actiiiollte; horDblende; epidote. -Cat! fi'y«(KalseQHUgeG'r'nn,,ffil de Chat .FV.) eibibils opalescencL-, but without prismalic •coluis, espocially when cut ea eaboc/uin, uii ellect sometimes due to fibcn of nsbcslus. Also present lu the siliceous paeudomorpha, aflercrocidolite, called Uger-rynaee crocidolite). ' Tile liiylylily-prized Uriculal cal's-eye is u variety of clirysoberyl.

ATtiituriie. —Spangled witli sculoa of mica, bumalite, or other mineral.

Jiiipart from Vt« prtntentt of di*tiiirt mCii^raU dUlribuleiJ densely ibrongh the mssG. The more common kiuita ai'e those in which the Imporitiea are: {a)/«TUfJno>u(BisenkicBcl Oerm.), either red or yellow, from aithydrous or hydroua iron sesquioside; (6) ehion'tk, /roiii some kind of chlorite; {c) atUnolitie; i^d) mieaecont; {e) artnaceen*, or saidt.

Containing liquid* in eutitiet. The liquid ubusIiv water (pure, or a mineral w>lulion), or Fome petroleum-like compound. Quarts, especially unoky quartz, also often contains luclusioDs of bolL liquid and gaseous carbon dioxide.

B, Crtptocrtbtaluhb YAKiKTiza.

(7Aaic«tony.—Having the luster nearly of vrai, and eflher transparent or tnostncent.

O, —26-284. Color while, grayish, blue, pale brown to dark brown, black. Also of other shades, and then having other names. Often mammillnry. botryoldiil, stalacilic, and occurring lining or filling cavities in rocks. Il often coDlalna some disseminated opal- ■ silica. The name tnliydot is given to nodule" of chalcedony containing water, someliroea

in Inrge amount. Embraced under the general nam« ^ clialcertony i-lhecrysbiKna fotm of silica miilch forms concretirmary ma.ssea with nidfal-flbrous md concentric slrwlare. and which IV? shown by Roaenbusc'i, is optically impntfw; unlike trne quartz. Ithasn^s l-liS7;

*o = 3 58-3-64. Often iu spheruliles. showing the sphemlltie interference-Bgnre. l.vfo- lite of Mallard has a like siriicture. but Is opifefttly + and has the apectfic gtsrlty and refractive hidei of opal. See also quartzine, p. 928. , , v

Cnmaian. 8ard.—k clear rod chalcedony, pale to deep in ahade; atoobrownfsb red

to brown. , ,.

OArystoprow.—An apple green chalcedony, the color due to nickel oxide-

i>r(is«.—Translucent and dull leek-green.

R(7«mo.—Rather bright green to leek-green, and also someB nearly emerald-green, and subtranslucent or feebly translucent. Heliotrope, or Btoodtone, is the aanae stone essentially, with small spoU of red jasper, looking like drops of blood.

Agate.—A. variegated chslcedonv. The colors are either (ol banded; or (ft) irregularly clouded; or (e) due to visible impurities as in mo« agate, which has brown moss like or dendritic forms, as of manganese oxide, distributed through the mass. The bands are delicate parallel lines, of white, pale and dark brown, bluish and other shades; they are sometimes straight, more often waving or zigzag, and occasionally concentric circular. The bands are the edges of layers of depmilion. the agate having been formed by a deposition of silica from solutions intermittently supplied, in irregular cavities in rocks, and deriving their concentric waving courses from the irregularities of the walls of the cavity. The layers differ in porosity, and therefore agates may be varied in color by artificial means, and this is done now to a large extent with the agates cut for ornaments. There is also a natural wood: wood perforated with rounded agate.

Onyx —Like agate in consisting of layers of different colors, white and black, white and red, etc., but the layers in even planes, and the banding straight, and hence the name for cameos.

;CoogIC

g;

Siiiiont/x.—Like opal in structure, but includes layers of carnelian (sard) along with others of white or bluish, and brown. and anmetloieg Slack colors.

Affilejaiper.—A agate consisting of jasper with veins of chalcedony.

BiUceoiu ti'nleI.—Irregularly cellular quartz, formed by (luposUIoo from nstera coulaiD-iug silica or bitubular glasses in solution. See also under opal, p. 339,

FliAt. —Souiewbul allied to chalcedony, but more opaque, and of dull color, usually gray, and brown, and brownish black. The exterior is often whitish. from mixture with lime or barytes, in which it is embedded. Luster barely glassy, but vitreous. It exhibits a deeply conchoidal fracture, and a sharp conchoidal edge. The fluid of the -liulk formalioo consists largely of lime remains of diatoms, sponges, and other marine ~ >ducti<m8. The coloring matter of the common kinds is mostly carbonaceous matter.

nt implements play au iiiiporliiut piirl amouj; Ibe reiles of early man.

lioruntont. —Resembles flint, but \a m(>re brittle, the fracture more apliolery. Chert ts ft icrrii iifleii applied to borniituue. uud to iiuy impure tibily rock, including the Jaspers.

li'itaiiie; Lydim l>toit«. or '1 oueh»toiiie.—A. veWei-black siliceous stciue or dbiiy Jasper, used oil a<:couoi of lit bardiiess anil black color for tryiug llie purity of the pi'ecious melala. The color left on ibe stoue after rubbing the metal aci-oas li indicates to the expcrluced e>c the amount of alloy. It is uot spliilery like bornstone.

Jatper. —Impuie opuque colonsl qunriz; couiuionly red, also yellow, dark greeu and gtuyiaii blue. Striped or riband jatper lias tbe colors in broiid stripes. I'oreelain jatper is uotbing but balled clay, uud diilers from true jasper iu biiug B.B. fusible on the edges.

C. Besides Ibe nbove thiere lire also:

OrtiH'ii'ir Quiirlt, QuarltToek. or (^wruite.—A rock consietinp of qianz gi-alsn ycry Brndy compacted; the grnius often bardly distiuct. (tvarliote Sandttuna, Qvartx-con-glomerate. —A rock made of pebbles of tjuarlK wllb xitud. Tlie pebb'l'-s M^metiines are Jnaper and cha1ce<lony, aud muke u beault'ul stone wIjcd polfsbed, JtarolvtHile, at Fler^ Saniittone. —A friable snnd-rock, coDsislIng niHlnly of quanz-sand, liui conlaiuiug a little
—.,_ . .j , — ; . ^ _ , . _ y j ^ , . . . , . . . ,

. ind possesBiDga decree of dexiblilty wbeu in tbiin lamtDie. JJuliriUtu, ot hi —A cellular, Sfiily rock, buying tbe untuie iu purl of course cbulcedony.

PtmidomoTphouM Quarts. —Qusrls appears aliio under Ibe forms of many of Ibe mineral apectes. wbieli It has taken Ibroagh either the alteration or replacement of crystals of those speries. The most common quartz pseudomorpbs are those of calcite, harlte, fluorite, and siderile. Bilirijtrd wood is quartx pseudomorph after vood (p. 25S).

Pyr.. etc.—B.B. unnlled; with bomx disaolTes slowly to a clear glass; with soda dissolTCB with effervescence: unacted upon by salt of phosphorus. InEolubte In liydro-' cbloric acid, and only sligljlly acted upon by soludons of Hied caustic alknlies. tbe crypto-cryBtallne varieties to tbe greater extent. Siiliibleonly In hydrofluoric acid. When fused aud cooled it becomes opal-silica having G

Diff.—Churacteriied ' . • - ■

also <ti geuerul by bardm

Eiisily recognized in rock sections by its low refracion (" low relief," p. 170) and low birefringence (^ — oi := OOOB): Ibe interference colors In ^ood sections uot rising above yellow of Ibe lirst order; also by Its limpidity and the positive uniaxial cross yielded by axial sections (p, 203. note), wlncb remain dark when revolved lietwceu crossed uicols.

ConiDionly in formless grains (grunile). also wilb crystal outline (porphyry, eta).

Oba.— Quart* is an essenlinl coiuponent of certain igneous rocks, as granite, granite-poipbyry, quartz-porphyry and rbyollte in the granite group; in such rocks it la commonly In formless grains or masses filling the interstices between the feldspar, as tlic last product of orystnllixiitton. Fuither it Is an essential consllluent in quarlzdiiorlle. tjiiartz-diorlte porphyry and dnciles in the djorlte group; in the iwirphyrles frequently in distinct crystals. Iioccur'silsoasan accessory in other feldspnlhic igneous rocks, such as syenite aud trachyte, Amonir tbe meiamorpliic rocks it is un essential componeul of certain varieties of gneiss, of qunrtzlte. etc. It forms ttie mass of common sandstone. It occura lu tbe vein.stone in varioua rocks, and for a large pari of mineral veins; as a foreign mineral in some liflieslones, etc., making gecklesof crystals, or of chalcedony, agate, carnelian. etc.; ns embedded nodules or masses In vailous limestones, constituting Ibe liint of tlie Cbiilk formnlion. tbe bornstone of other Ibnestnnes—tthese nodnle!> aomctimes hpcondug continuous layers; as masses of jasiier occasionally in limestone. It is tbe principal material of Ibe pebbles of gravel-beds, and of the sands of the seaabore. and sandleds eveijwhere. In graphic graniti! {jvgrmaiita) (he quartz individuals are ari-anged In parallel position in feldspar, Ibe angulnr piirlicles rewmbing written cbumclera. The qunriz gniins in a fragmentai saudsione are often found lo have undergone a secondary growth by tbe deposition of crystallized silica with like orientation to tbe original nucleus.

Switzerland Dauphne, Piedmont, the Carrara quarries, and numerotis olber foreign

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S28 DESCRIPTIVE MINKRALOQT.

localities, afford floe epecimeDB oF rockcryHtnl; also Jspsn, whence tbe beftutlful crystal spheres, !□ rare cases up lo 6 ladies Id dUuieter. Smoky quarit crystals of great beauty, and ofteu bigbly complex iu form, occur at many points Id tbe central A-lpa, also at Calrn^rm. Scotland. Tbe most lieauliful atuethysla are brouglil from India, CtyloD, and Perslii, also from Bruzll; inftrlorspeclii: ens occur iu Transylvania. The flneBl air7t«f wiu and agnta axe found Iu Arabia, India, Brszil, Surinam, also fonneriy at Obersteiu and Saxony. BiMtliiDii affrjrds smaller but handsome speclineDB (fjcotch pebbles). Tbe banks of the Nile uOord tbe Egyptian jiisper; the striped jssper is met with in Siberia, Bazony. uiid DevoDsbire.

In New York, quartz crystals are abundant in Herkimer Co.. at Middleville, Little Fulls, etc.. liMise Iu cavltlts Iu tbe Calciferous siind-rock, or embedded lo loose earth. Flue quarizoida, at the beds of hematite Iu Fowkr, Herman, and Edwards, St. Lawrence Co., alsoatAitnei'p. JeSerwD Co. On tbe baaksof Laldlaw Lake, Rossie, large Irapliioted crystals; at Bllenville lead mine, Ulster Co , In fine groups. At Paris, He., handsome crysUils of brown or smoky quartz. Beautiful colorless crystaU occur at Hot Springs, ArkacsHS. Alexander Co., N. C, has aSorded great numbeia of blgiily complex crystals,

with rare modfications. line crystals of smoky quartz come from the granite of the Pike's Peak region, Colorado. Inclusions of quartz crystals, also enclosing calcite, sphalerite, etc, are common in the Keokuk limestone of the west.

Common quartz occurs at Albany and Paris, Me.; Acworth, N.; H. Bouthbuiy, Conn. Amethyst in trap, at Konveuwaw Point, Lake Superior; Specimen Hill, Yellowstone Park. In Pennsylvania. In East Bradford, Chester, and Providence (one fine crystal over 1 lbs. in weight), in Chester Co. - at the Prince vein. Lake Superior; Large crystals, near Greensboro. N. Y.; crystallized green quartz, in talc, at Providence, Delaware Co., Penn. Chalcedony and agate abundant and beautiful on N. W. shore of Lake Superior. Red jasper is found on Sugar L'mf Mt., Maine; In pebbles on the banks of the Hudson at Troy; yellow, with chalcedony, at Chester. Moss agate and jaspered wood of great beauty and variety of color is obtained from the petrified Forest called Chalcedony Park, near Carrizo. Apache Co. Arizona; also from the Yellowstone Park: near Florissant and elsewhere in Colorado; Ainetbyst Mt. Utah; Napa Co., California. Moss agates from Humboldt Co., Nevada, and many other points.

The word quartz is of German provincial origin. Agate is from the name of the river Aclintes, to Sicily, whence specimens were brought, as stated by Theophrastus.

Quartz is a name which has been given to a form of silica which is present in chalcedony and is inferred to be identical in crystalline structure. Lutaetle belongs here.

TRM TMTTB.

Hexagonal or pseudo-hexagonal. $a = 1.6530$. Crystals usually minute, thin tabular | c; often in twinning; also united in fan-shaped groups.

Cleavage: prismatic, not distinct; parting | c, sometimes observed. Fracture conchoidal. Brittle. H. = 7. G. = 2.65-2.66. Luster vitreous, on c pearly. Colorless to white. Transparent. Optically - . Often exhibits anomalous refraction phenomena.

Comp.—Pure silica, SiO₂, like quartz.

Pyr., etc.—Like quartz, but soluble in boiling sodium carbonate.

Obs.—Occurs chiefly in acidic volcanic rocks, trachyte, andesite, granite, less often in dolerite: usually in cavities. Often associated with epidote, also hornblende, and gneiss. Some resemble opal. First observed in crevices and druses in an andesite from the Cerro Sau Cristóbal, near Pachuca, Mexico; later proved to be rather generally distributed. Thus in the vicinity of the Siebengebirge; of Bugneau Hills in N. Italy: Putnam (Mont-Dore) in Central France, etc. In the ejected masses from Vesuvius consisting chiefly of sandstone. With quartz, feldspar, fayalite in lithophyses of Obsidian cliff Yellowstone Park. In the andesite of Mt. Rainier. Washington.

Niined from rriiSvfioi, thre^otd. in allusion to the common occ-urrence In trillings.

AsMANiTB, A form of silica found in the meteoric iron of Breitenbach, In very minute gnis. probably Identical with tridymite; by some referred to the onhorbomble svstein.

Cristobalite. Silica in white octahedrons (pseudo-isometric?) $Q = 2-27$. With tridymite in andesite of the Cerro 8. Cristobal, Pacifica. Mexico.

Melakophilite. In minute cubes and spherical aggregates. Occurring with calcite and celestite implanted upon an incrustation of opaline silica over the sulphur crystals of Girgipii. Sicily. Consists of SiO_2 , with 5 to 7 p. c. of SO_2 , The mineral turns black superficially when heated B.B.

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Amorphous. Masses; sometimes small reniform, stalactitic, or large tuberoses. Also earthy.

H. — 55-65. G. = 1.9-3.3; when pure 2.1-2.5. Later vitreous, frequently enbureo; often inclining to resinous, and sometimes to pearly. Color white, yellow, red, brown, green, gray, blue, generally pale; dark colors arise from foreign admixtures; sometimes a rich play of colors, or different colors by refracted and reflected light. Streak white. .Transparent to nearly opaque. $n = 1.44-1.45$.

Often bluish double refraction similar to that observed in colloidal substances due to leucis. The cause of the play of color in the precious opal was investigated by Brewster, who ascribed it to the presence of microscopic cavities. Brewster, however, in his paper on the subject (Ber. Ak. Wien, 41, 1871), and based on that this explanation is incorrect; he refers the color to thin curved lamellae of opal whose refractive power may differ by 0.1 from that of the mass. These are conceived to have been originally formed in parallel position, but have been changed, bent, and finally cracked and broken in the consolidation of the groundmass.

Comp.—Silica, like quartz, with a varying amount of water, $\text{SiO}_2 \cdot n\text{H}_2\text{O}$. The water is sometimes regarded as non-essential.

The opal condition is one of lower density of hardness and specific gravity, and, as generally believed, of incapability of crystallization. The water present varies from 3 to 18 p. c. or more, but mostly from 3 to 9 p. c. contains small quantities of ferric oxide, alumina, iron, magnesia, and alkalies are usually present as impurities.

Varieties.—*Opal*. — Exhibits a play of delicate colors. .

Fire opal. — Hyacinth-red to honey-yellow colors, with fire-like reflections, - somewhat

Oiiiol.—'B\u\s\ while, tmnslucent, with reddish reflectiona is ii bright light.-

Common Opal.—In part Iranslucent; (aj milk-opal, milk-white to greenish, yellowish, bluish; (b) Btiinopai, wax-r honey- to ocher-yeliow, with a resiuous luster; (e) dull oiive-Kreen and mountain-green; (<^ brick-ri-d. Includes i3«7uc]>a{; (f) Hydrophane, a variety which becomes more tmnalucnt or Irnnsparent In water.

Cadtolong. —Opaque, bluish while, porcelaln-white, pale yellowish or reddish.

Opal-agaU.^K%ate.\yi.e In strucluie, but consisting of opal of different shades of color.

Meniliie. —In concretionary forms; opnque, dull grayish,

Jatp-opal Opal-jasper —Opiil conlainng some yellow iron oxide and other impurities, and haviii<: the color of yellow jaspir. with Che luster of common opal.

Wood-opal Holz-opal Berm.—Wood petrified by opal.

Uyalil". Huller's Glass.—Clear as glass and colorless, constituting globular concretioDS, and crusta with a globular or botryoidal surface; also passing into Iranslncent, and whitish. Less readily dissolved In caustic alkalies than other varieiee<:.

Morilg, SUietout Sftifer.—Includes translucent to opaque, gtayish. whitish or brownish Incrustations, porous to firm in texture: somelimps fibrous-like or filamentous, and. wbeni ■o. pearly in luster {then called Pearl-niiter); deposited from the siliceous waters of hot 'springs.

i7ep*«r>(«.—Conslltu'esconcretlornry deposits about the geysersoftlie Yellowstone Park, Icehind, and New Zealand, preseuting white or grayish, porous, stolactitic. filamentous, canllltower like forms, often of great beauty; also com pact-massive, and scaly massive.

Float itont. —In light porous concretionary masses, while or grayish, sometimes cavernous, rouch in fmclure.

TripoliU. —Formed from the siliceous shells of diatoms (hence called dialomiit) and other microscopic species, and occurring in fxtcnMve depositn. Includes Infiiforinl Barth, or Earthy Tripelitt. a very tine jcraincd earth Inokins often like an earthy chalk, or a clay, but hais lo the fpe1. and scratching Klass when rubbed on it.

Pyr., etc.—Yields water. B.B. Infnailile. but becomes opaque. Some vellow varieties, contalninc iron oxide, liirn red. Soluble in hydrofluoric acid somewhat more readily than quartE; also si'luble in caustic alkalies, hut more readily in some varieties than in others,

ObB.—Occurs filling cavities and fissures or Reams In igneous rocks, as trachyte, porphyry, also in aome metallic veins. Also embedded, like flint, in limestone, and

siliceous concretions, in argillaceous beds; formed from the siliceous waters of

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Some hot springs; often resulting from the mere accumulation, or accumulation and part

of the shells of foraminifera, of epiphytic sponges, etc., of opal-silica. The most important is the hydrothermal source of the opal of Limburg and Hartz (its source is in the mountain: locks), and of part of the igneous rocks. It occurs in most basalts; and in

Opal occurs in porphyry at Cierwenitz, near Kaslau in Hungary; in Grochu at Dios in Huautla de Queretaro in Mexico; a beautiful blue opal in Bulli Creek, Queensland. Fire-opal occurs at Zimapan in Mexico; the Fire Opal; near San Antonio, Honduras. Common opal is abundant at Telkebenya in Hungary; near Penatepec, etc., in Moravia; in Bohemia; Silesia; and in Iceland. Hyaline opal occurs in amygdaloid at Tschermitz, Böhmen; in diorite at Wallmühl. Boljemi; at San Luis Potosí, Mexico.

In the hyaline opal occurs sparingly in connection with the trap rock of New Jersey and Connecticut. A water-worn specimen of fire-opal has been found on the John Davia river, in Crook Co., Oregon.

Common opal is found at Corvallis, Lehigh Co., Penn.; at Aiguas Calientes. Idaho Springs, Colo.; a white variety at Mokelumne Hill, Calaveras Co., Cal., and on the Mt. Diablo range. Opal occurs in great abundance and variety in the Yellowstone region (cf. above); also in the area of the Nevada. Other localities are given by Kunz, Gems and Precious Stones of N. A., p. 130.

II. Oxides of the Semi-Metals; also Molybdenum, Tungsten.

Arsenolite*. Arsenic trioxide. As_2O_3 . Isometric octahedron; in crusts and effluvia. Colorless or white. $\rho = 5.8$. Occurs with arsenical ores.

Andradite, Also $Ca_3Fe_2Si_2O_{12}$, but monoclinic in form,

Andradite. Antimony trioxide. Sb_2O_3 . Isometric octahedron: in crusts and granular masses. $\rho = 5.8$. Colorless, gummy. Occurs with ores of antimony.

Valentinite. Also Sb_2O_3 , but in prismatic or rhombic crystals.

Blimitz. Bismuth trioxide. Bi_2O_3 . Pulverulent, earthy; color straw-yellow.

TaUnrit«. Telliiriitm dioxide, TeO,, In while to yellow alender jiriaitiaUc eryalals.

Holybdit*. Molybdenum trioxlile, Mod. In capillary tnf ted forms and earthy. Color straw-yellow.

TungaUta. Tungsten trioiide, WO.. Pulverulent, earthy; color yellow or yellowbh green.

Oarvantita. Sb.Ot.Sb,o,. lo yellow to while aclcular crystals; also massive, pulverulent.

BtJbloonltv. H,Sb,Oi. Hasdve, compact. Color pale yellow to yellowish white.

ITT. Oxides of the Metals. A. Anhydrous Oxides.

I. Protoxides, R,o and RO.

II. Sesquloxides, B,o,.

III. Intermediate, BR,o, or RO.B,Oi, etc. iV. Dioxides, RO,.

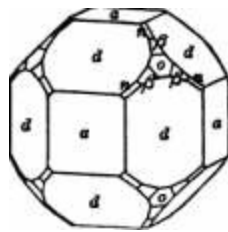
The Anhydrons Oiidee include, ae shown above, three distinct divisions, the ProtoiideB, the Seaquioxides and the Dioiides. The remaining Inter-mediate division embraces a number of oxygen compounds which are properly

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to he regarded chemically as salts of certain acids (alnminates, ferrates, etc.); here is included the wellcliarncterized Spinel Group.

Among tlie Protoxides the ously distinct aroup is the Periclase Gkoup, which iiiiiciiies the rare species Periclase, MgO, Manganosite, MnO, and Jiinsenito, NiO. All of these ure isometric in crystallization.

Tiie Scs(iDioxides include the well-cliaracteriKed Hematite Group, E,o,. T'lie Diosides include the prominent Kutile Group, RO,. Both of these groups are further defined later.



I. Protoxides, R,o and RO.

OUPRTTE. Ited Copper Ore. Rotlikiipferer?. Oerm.

Isometric-plagiJiedral. Commo)iily in octahedrons; also in cubes and ^lodecahedrouB,

often highly modified. Plagiolithic ^ci

faces sometimes distinct (see pp. 50, 51). Also massive, granular; some times earthy.

Cleavage: none interrupted. Fracture conchoidal, uneven. Brittle. H. = 4-5. G. = 5.85-6.15. Luster adamantine or submetallic to earthy. Color red, of various shades, particularly cochineal-red, sometimes almost black; occasionally crimson-red by transmitted light. Streak several shades of brownish red, shining. Sub transparent to subtranslucent. Refractive index, $n_r = 2.8-2.9$ Fizeau.

Vm.— 1. Occurs in the ... Arizona,

in the ... bed rolls, ... the crystals often with a crust of malachite; (f) ...

2. ... In capillary or ... crystallizations, which are sometimes cubes ... in the direction of the cubic axis.

8. ... or ... brown until earthy, often mixed with red oxide of iron; ... nearly black.

Comp.—Cuprous oxide, Cu_2O = Oxygen 11.3, copper 88.7 = 100.

Fyr.. etc.— ... in the closed tube B.B. in the forceps fuses and colors like flame emerald green. On ... fuses, and is reduced to metallic copper. With the fluxes gives reactions for copper. Soluble in concentrated hydrochloric acid, and a strong solution when cooled and diluted with cold water yields a heavy white precipitate of cuprous chloride of copper.

Uf.—Distinguished from hemimorphite by inferior hardness, but is harder than cinnabar and proustite and differs from them in the color of the streak; reactions for copper, U, B., are conclusive,

Oba.—Occurs at Kamsdorr in Thuringia; in Cornwall, in fine crystals, at ... and ... mine; in Devonshire near Tavistock: in isolated crystals, more or less altered ... lit Chesay, near Lyons, France; in the Ural; South Australia; also abundant in Bolivia.

It is observed at ... N. J.; at Cornwall. Lebanon Co., Pa.; in the Lake Superior region. With malachite, ... the Copper Queen mine. ... sometimes in fine crystals: beautiful ... Morenci; at Clifton, Graham Co., in crystals, and massive.

Ice*. B.O. Hexagonal. Familiarly known in six-rayed snow crystals; also coating ponds in winter, further as glaciers and icebergs.

Pe rid use Group. Fmlolaae. Magnesia, HgO. In cut>es or ociaheiiroons. and In gralna
Cleavage cubic. ^ r^ " -s Id while limestone al Mt. Bomnia; at the Kittein m&nga-

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DESCBIPTITE UINERALOQT.

d-green, becoming black od exposure. From L&agbua &>id Nordmark, Sweden.

Nickel protoxide, NfO. In green oclabL-drons. From Jobsnogeurgeustall L

ZINOITB, Red Oxide of Ziuc Bed Zink Ore. RoUiziukerz Germ. Hexagonal-hemiroorphic.
Axis 6=[^] 1'6319. Natural crjatala r^re (Fie. 50, p. IS); usually foliated massive, or in
coarse particles and graine; also witli granular structure.

CleavB^: c perfect; priBmatic, sometimes distinct. Fracture subconchoidal. Brittle, fi. =4-
46. G. = 5-43-5'7. Luster Bubftdaraantine. Streak orange-yellow. Color deep red, also
orange-yellow. Translucent to snbtrauslucent. Optically +.

Corap.—Zinc oxide, ZnO = Oxygen 197, zinc 80-3 = 100. Manganese proto:^ide is
sometimes present.

Pjrr., Btc—B.B. infusible: wltb tbe fluxes, on tlie plaliuuni wire, gives reactions for
manganese, and ou obarcoal in It.F. gives h coaling of ziuc oxide, yellow wblle bot, and
while on tooling. Tlie coating, moistened with cobalt solution and treated InO.F.,
asBumes a green color. Solulile in acids.

Dlff.—Cbanicteni^ by its color, particularlj that of tbe streak; by cleavage; bj

Obs.—Occurs with franklinlle and wiUeinite, at Sterling Hill near Ogdensbiirg, and at
Mine Hill, Franklin Furnace. SuEssex Co., N. J., sometimes in lamellar masses Id pink
calcile. A not uncomniiin furuacc product.

HsBSicot. Lead mnuoiide. PliO. Ha-^sive, eculy or earthy. Color'yellow, reddish.

Tenorlte. Ciipric oxide. CuO. In minute black scales with metallic luster; from Vesuvius.
Also black earthly mttssivc {melaeoniie); occurring wltb ores of copper as at Ducklown,
Tenn.. and Keweenaw Point, Lake Sni>erlor.

Paramelaconita is essentially ciipric oxide, CuO, occurring in black pyramidal crystals
referred lo the tetragonal system. From tbe Copper Queen mine, Bisbee, Arizona,

Hematite Group. E,Oi. Rhombohedral.

it' i

Corundum Al₂O₃. 93° 56' 13630

Hematite Fe₂O₃. 94° 0' 13656

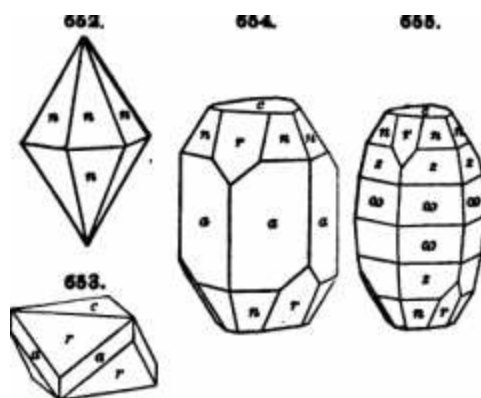
Ilmenite (Fe,Mg)O.TiO₂, Tri-rhombohedral 94° 29' 1-3846

Pyropluuiite MnO.TiO₂, " 94° 51' . 13C92

The Hematite Group embraces the sesquioxides of aluminium and iron. These compounds crystallize in the rhombohedral system with a fundamental rhombohedron differing but little in angle from a cube. Both the minerals belonging here. Hematite and Corundum, are hard.

To these species the titanate of iron (and magnesium) and manganese, Ilmenite and Pyrophanite, are closely related in form though belonging to the tri-rhombohedral group (phenacite type); in other words, the relation between hematite and ilmenite may be regarded as analogous to that between calcite and dolomite. It is to be noted, further, that hematite of ten contains titanium, and an artificial isomorphous compound, Ti₂O₃, has been described. Hence the ground for writing the formula of ilmenite (Fe,Ti)₂O₃, as is done by some authors. It is shown by Fenfield, however, that the formula {Fe,Mg}TiO₂ is more correct.

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Rhombohedral. Axis i = 1-3630. $\text{er. } 0001\text{AlO} = S7^*84'$. «fig.

$\text{en. } 0001\text{aS}2 \ll = 61^*11'$.

$\text{rt'. } 101\text{AlO} = W6 \ll'$. $\text{nn', } 3i4\text{aAS}423 = \text{Sr } S8'$. $\text{w-. } 4488\text{A}4848 = \text{ST}'' \text{ Sff. } b'$,
 $2341\text{a}2421 = :S8^{\circ}65'$.

Twins: tw.pl.r; sometimes penetration-twins; often poly-synthetic, and thus producing a laminated structure, Crystals usually rough and rounded. Also massive, with nearly rectangular parting or pseudo-cleavage; granular, coarse or fine.

Parting: c, sometimes perfect, but interrupted; also due to twinning, often prominent; a

less distinct. Fracture uneven to conchoidal. Brittle, when compact very tough. H. = 9. G. — 395-410. Lustre adamantine to vitreous; on cleaves sometimes pearly. Occasionally showing asterism. Color blue, red, yellow, brown, gray, and nearly white; streak uncolored. Pleochroic in deeply colored varieties. Transparent to translucent. Normally uniaxial, negative; for sapphire $\omega = 1.7626$ to 1.7682 and $\epsilon = 1.7594$ to 1.7598 . Di. Often abnormally biaxial.

Varieties.—There are three subdivisions of the species principally recognized in the series, but (referring only to the principal and simple of crystallized or amorphous).

Variety 1. Sapphirine. Recrystallized.—Includes the purer kinds of the blue colors, transparent to translucent, useful as gems in the same way as the blue sapphire; but their colors: Sapphirine blue: true blue. or Oriental Ruby, red; O. Topaz, yellow; O. Emerald, green; O. Amethyst, purple. A variety having a yellowish color when viewed in the direction of the vertical axes of the crystal, is the Amethystine Sapphirine or Smoky Sapphirine.

Variety 2. Colorless.—Includes the kinds of dark or dull colors and not transparent; but their colors: brown, black. The original amorphous variety is a brownish greenish or bluish yellow, but greenish or bluish yellow, light, in

the emerald. Schmalzgrün Oerme — Includes the granular corundum, of black or grayish black color, which contains magnetite or limonite intermixed. Sometimes associated with iron spinel or hercynite. Feels and looks much like a block fine-grained iron ore, which it was long considered to be. There are gradations from the evenly fine-grained emery to kinds of which the crystalline are distinct crystals.

Composition.—Alumina, Al_2O_3 , = Oxygen 47.1, aluminium 52.9 = 100. The crystallized varieties are essentially pure; analyses of emery show more or less impurity, chiefly magnetite.

Properties.—Sapphirine is slowly dissolved in borax and salt of phosphorus to a clear mass, which is colorless when free from iron; not acted upon by acids. The finely pulverized mineral, after long heating with cobalt solution, gives a beautiful blue color. Not acted upon by acids, but converted into a soluble compound by fusion with potassium bisulphate.

Diagnosis.—Characterized by its hardness (scratching quartz and topaz), by its adamantine luster, high specific gravity and infusibility. The massive variety with rhombohedral parting resembles cleavable feldspar but is much harder and denser.

Occurrences.—Usually occurs in crystalline rocks, as granular limestone or dolomite, gneiss, mica slate, chlorite slate. The associated minerals often include some species of the chlorite group, as prochlorite, comendophilite, margarite, also tourmaline, iolite.

Found in India.

DESCRIBTIYS HIKERALOGT.

granite, diaspore, and a Berieg of alumioous mlueralH, in part produced from Its
altemliod. OccBsiously found Id ejected musses encloaed lu juunger volcanic rocks, u at
KOiii|,'a-wiiter, Niedermendig, etc. Rarely observed as n contact-mineral. The tint-
sapphires are usually oblataed from i)ie beds of rivers, cillier in modified beiaoul
prisms (ft- iu rolled mas-ses, accompaied by ({iiiiia of mu^ neiite, aud aevei-al kloda of
gems, iia fpiiiel, etc. The emery of Asia Miur, Dr. Smith slates, occui-s <□ graaular
limestone.

The boat rubies come from the miiies !□ Upper Burma, oorlb of Maudaliiy. iu au arva
coverluK 23 Co 30 square miles, of which Hogok ia the center. The rubies occur tn titu iu
crystalline llmestoe, also lu the soil of tbe hillsides uud iu gem-bearitig gravel. Blue
sapphires are brought from Ceylou, oftea at rolle<l pebbles, also as well-presfrviKl
crystals. Coruudum occurs In the Caniatic o:i the Malabar const, ou Ibe CLauibuu hllU iu
ESmiii, aod elsewhere in the East Indies; also near Cantuu, Chlua. At til. Quthard, It
occurs uf & red or blue tinge In dolomite, and near Huzzo in Pledmoul, in white compact
feldspar. Adamantine spar is met with in lur^^ B loarae. hexagonal pyramids In Gellivara,
Bwt-den. Other localities are iu Boliemlii. ui'ar PetKchau: iu the Ilmen mountains, not far
fnun Hiask; In tbe gold-waabiugi uorihia-t of Zliitoust. Corundum, sapphires, and le^a
often rubies occur in roiled puliblus in thi^ diiimojd gnivels on the Cudgeeong river, at
Mud^e« Uid other points in New South WoIot, Einerj- is found iu larye bowlders at
Naios, Nicaria, and S iDios of the Grecl in islands; al-io in Asia Minor, 13 m. E. of
Ephesus, near Gumucli-dagh and near Smyrna, nstocinled wit li iiiargiLrile, chloriloid,
pyrlte.

Ia N. America, in l£ut-tthuielii, at Ch.'^ <tr. with magnetite, diaspore, ripldolite, mnr-
gartts, etc.. mined for use at emery. In C-maecUiml. near Litchfield. In Nea Tork, at
Warwick, bltilsh and pink, with spiiiel: Auiity, In gran, limestone; emery with magucttle
and green spinel (hercyuite) iu We.-tiche-ter Co., uear Cruger'a Station, and elsewhere. In
New Jirtey, at Newton, blue crystals in grnn. lltuestoie; iit Vernon. Iu Ptntuyltauia. In
Delaware Co., In Aston, neiir Village Qruue. In large crystals; at Mineral Hill, lu loose
CTjst.; in Chester C >, at Uninnvllle, abuudant in crystals: in laree crystals Ioo~e In the
soil at Bhimersvllle, Lehigh Co. In Virginia, in Che mica schiala of Bull Mt., Patrick Co.

Common nt inauy points along a belt extending froin Virginia across wentem North nnd
South Carolina and Georgia to Dudleyville, Alabama; especially in Madison, Buncombe.
Haywooil, Jackaou. Macou, Clay, ivud Gaston counties In North Carolina. Tbe localiiles at
which most work has been done are Che Cuia:igce mine, Corundum hill, near Franklin,
Macon Co., N. C aud 36 Rkile^ 8. B. of this, at Laurel Ci-eeK, Ga. Thi' coruudum occui-s In
betis in chrysolite (and seriieiiilue) aud borublendlc gneiss, aisociated with a species of
tlic chlorite group, also spiuBl. etc.. and here as elsewhere with many minerals resniing

from its alteration. Boiue fine rubles liave been found. Flue pink crystals of corunduiit occur aC Hiawaasee, Towns Co., Georgia. In Colorado, small blue crystU occur in tiitcn ■cbist near S.ilida. ChaSee Co. Gem sapphires are found near Helena, Montana In gold-washings and in bars in the MIsKuiiri river, especially the.Eldorado bar; at Togo Gulch ou tbe Judith river and at other points In Montana. Tn Calijornia, in Los Angeles Co., in the drift of San Franslsqueco Pass. In Canada, at Burgess, Ontario, red and blue crystals.

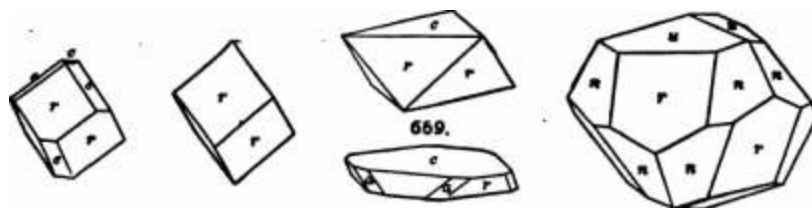
HEMATITB. EEsenglaiiE, Oerra. Bhombohedral. Axis i = 13656. er. 0001 A lOil = 57° dT. n', lOil A iioi = W o*.

a. oiia A iois = 64"Gi'.

uu'. 1014 A ilo4 = 87* y. »n', 3548 A 3438 = 61 ■■ Wf. en, 0001 A 2^48 = 81° 18".

658.

6S7.



Twins: tw. pi, { } c, penetntdon-tvius; (8) r, leas common, uanfill; u

D,,i,:cc.;;CoogIC

poljeynthetic twinning lamellfe, producing a fine Btriutioa on c, and giving rise to a diatiuct parting or pseudo-cleivage y r. Crystals often thick to thin tiibnlar H c, and grouped in parallel position or in rosettes; c faces striated II edge c/d and other forms due to oscillatory combination; also in cube-like rfaombohedrons; rhombohedral faces u (1014) horizontally striated and often rounded over in convex forma. Also columnar to granular, botryoidal, and stalactitic sliapes; also lamellar, Jaminte joined parallel to c, and Tariosly beut, thick or thin; also granular, friable or compact.

Parting: c, due to lamellar structure; also r, caused by twinning. Fracture subconcboidal to uneven. Brittle in compact forms; elastic in thin lamiuffi; soft and unctuous in some loosely adherent scaly varieties. H. = 5-5-6'5. G. = 4.9-53; of crystals mostly 520-5'25; of some compact varieties, as low as 4*3. Luster metallic and occasionally splendid; sometimeB dull. Color dark steel-gray or iron-black; in very thin particles blood-red by irans-mitted light; when earthy,red. Streak cherry-red or reddish brown. Opaque, except when in very thin laminae.

Vtir. 1. Speeul'ir. Luster inetanic, nnd crytala often aplendeot, whence the UBine apeailar iron (GlauMistceiz Germ). Whtn llie structure ie foliated or mfciceous, the ore is called

mioaetout hematite (BUenglommer Grrm.): some of the mlcnceousvarletieB are Boft iiJid unctuous (Eiieerahm Germ.j. Some vaneties are mnguelic, but probably from adiniid mngiieille (Arts. 424, 426).

2, Compact evlumnar; ur fibroiii. The masses oflen loar rsdiadng; luster submetallic to ineiuilic; color btowiiish red to Iron-black. Somellmea called red hematite, to contrast ft wllb limouiteaiKllurglte. OfieiiIn reniform masses with smooth fracture, c»] \etl kidney ore.

S, Ihd Ocheroui. Red auii earthy. Beddle and red ehaUc are red ocher, mixed with more or less clay.

4. Clny Iron-itone; Argillaceinia hematite. Hard, biownlah black to reddish brown, often in part deep red; of eubmetalljo to uimetatluc luateri and affording, like all Ibe preceding, a red slreiik. It conaiBts of oxide of jiOD wllb clay or sand, and sometimes olber

Comp.—Iron sesquioiide, Fe_2O_3 , = Oxygen 30, iron 70 = 100. Sometimes contains titanium and magnesium, and is thus closely related to ilmenite, p. 336.

Pyr., «tc.—B.B. infusible: on charcoal In B.F. becomes magnetic; with bornx givcft tlie iron reactions. With soda on cLarconl in R.F. is reduced to a gray magueilc melalljo powder. Soluble in concentrated hydrochloric acid.

' Diff.—Distinguished from magnetite by Its red ttreak, also from limonlte by the same means, as weH as by its not containing water; from lurgite by lis greater hstDnees aud by Dot decrepitiiliig- B.B. It ie hard in all but some micaceous viirleties (hence easily disllnguisbed frnni the black sulphides); also infutible, and B.B. becomes strongly majineic.

Obs.—This ore occura in rocks of all sgea. The specular variety is mostly confined to crystnlline or inetamorphic rocks, but is also a result of igneous action about some Tolcanoes, us at Vesuvius. Many of the geoiogical fonnations contain the argillaceous variety or ciay iron-stone, which Is mostly a marsh-formation, or a deposit over the bottom of shullow, s'A^nant water; but llils kind of clay iron-stone (tbnt giving a red powder) Is less common than tlie corresponding vaiety of limcmite. The beds that occur in metftmorphic rockt are aomellmes of very (Treat thickness, nnd, like thoxe of miignetle in the same situation, have resulted from the slteralion of sirnlfied beds of ore, orlglnnly of marsli origin, which were formeil at the same time with the enclosing rucks, and underwent meiamorphlFm, or a chanj;^ to the crystnlline condition, at the same time.

Beanlifui cryslallzallons of this species are brought from the Islnd of Elba, which baa nftordod it from a very remote period; The siirfncea of the crystals often present an Irlsed tamish and brilliant Insler. St. Gothard nfforris heautiful specimens, composed of crys-

tnllzed tnbles grouped in the form of rosettes (Eiteji/roeen): near Limoges, France, In large crystals; fine crystals are the result of volcanic action at Etna and Vesuvius. Arendal 1"

pendled specimen*.

Red hematite occurs in the form of fibrous conchoidal crystals, in the Bohemian, Saxonian, and Harz mountains,

In N. America, widely distributed, and sometimes in beds of various thicknesses in rocks of the Archæozoic age, as in the upper part of the Michigan district, also in the Michigan and west of Lake Superior in Gogebic county; further through northern Wisconsin. In Florence, Ashtabula and Dodge Cos., and in the Adirondacks near Vermilion, St. Louis Co.; in Hudson, at the Potosi Knob and the Iron Mt.

In New York, in Oneida, Herkimer, Madison, Westchester Cos., a lenticular argillaceous variety, consisting of one or two beds in the Clinton group of the Upper Silurian; also some in Pennsylvania, and as far south as Alabama; and in Canada, and Wisconsin to the west; in Alabama there are several beds; prominent ones are near Birmingham. Besides these regions of enormous beds, there are numerous others of workable value, either crystallized or argillaceous. Some of these localities, interesting for their specimens, are in northern New York, at Gouverneur, Antwerp, Henson, Edwards, Fowler, Canton, etc.; Woodstock and Aroostook, Me.; at Hawley, Mass. a micaceous variety; in North and South Carolina a micaceous variety in schistose rocks, constituting the so-called iron ore.

Named hematite from a Greek word, blood.

Hematite. Iron sesquioxide under an isometric form, occurring in octahedrons or dodecahedrons like magnetite, and believed to be pseudomorphous after magnetite; perhaps in part also after pyrite. Fracture conchoidal. H. = 5-7. G. = 4.9 (Brazil) to 5.2 (Monroe). Luster metallic. Color iron-black, sometimes with a bronzed tarnish. Streak reddish brown or purplish brown. Not magnetic, or only feebly so. The crystals are sometimes embedded in the massive hematite. They are distinguished from magnetite by the red streak, and very feeble, if any, action on the magnetic needle. Found in the Marquette iron region south of Lake Superior, where crystals are common in the ore; Monroe, N.Y.; Digby Co., Me.; at the Cerro de Mercado, Durango, Mexico, in large octahedrons; in the schists of Minas Geraes, Brazil; near Rittersgraben, Germany.

ILHBNITB or Hematite. Titanic Iron Ore. Titaniferous. Tri-rhomboidal; Axis $a = 1.3846$.

663. er. 0001 A 1011 = 57* 081'.

Crystals usually thick tabular; also acute rhombohedral. Often in thin plates or laminae. Massive, compact; in embedded grains, also loose as sand.

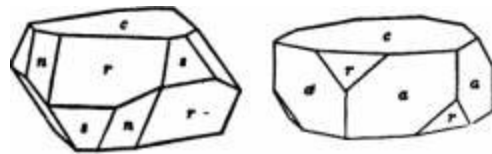
Fracture conchoidal, H. = 5-6. G. = 4.5-5, Luster submetallic. Color iron-black. Streak submetallic, powder black to brownish red. Opaque. Influences slightly the magnetin needle,

Comp., Yr.—If normal, FeTiO_3 , or $\text{FeO} \cdot \text{TiO}_2$, = Oxygen 31.6, titanium 31.6, iron 36.8 = 100. Sometimes written $(\text{Fe}, \text{Ti})_2\text{O}_3$, but probably to be regarded as an iron titanate. Sometimes also contains niobium (pyrochlore), replacing the ferrous iron; hence the general formula $(\text{Fe}, \text{Mg})\text{O} \cdot \text{TiO}_2$, (Penfield).

Pyro. etc.—Infusible in O.F., although slightly softened on the edges in R.F. With borax and salt of phosphorus reacts for iron in O.F. and with the latter flux assumes a more or less intense brownish-red color in R.F.: If treated with HCl on charcoal changes to a violet-red color when the amount of titanium is not too small. The pulverized mineral, heated with hydrochloric acid, is slowly dissolved to a yellow solution, which, altered from the decomposed mineral and boiled with the addition of tin-foil, assumes a beautiful blue or violet color. Decomposed by fusion with bisulphate of sodium or potassium.

Diff.—Resembles hematite, but has a submetallic, nearly black, streak; not magnetic like magnetite.

Obs.—Occurs, as an accessory component, in many igneous rocks in general, especially in gabbros.



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the place of magnetite, especially in gabbro and diorite. In these occurrences, it is often found in veins or large aggregates near the borders of the igneous rock where it is supposed to have formed by local differentiation or fractional crystallization of the molten mass. Some principal European localities are St. Christophe, Dauphiné (var. enstatite); Ulnafjell in the Fensås Mis. (Umenille); in the form of sand at Heden, Cornwall (titanite); Gaslein in Tyrol (titanite); Iserwiese (titanite). One of the most remarkable is at Eger, Norway, where it occurs in veins or beds in diorite, which sometimes afford crystals weighing over 16 pounds. Others are Egersund, Arendal, Ijnarum in Norway; St. Gervais, etc.

Five crystals, sometimes an Inch in diameter, occur (at Warwick, Amherst, and Union, Orange Co., N. Y.; Litchfield, Conn. {washingtonite). Vast deposits or beds of titanite ore occur at Bay St. Paul in Quebec, Canada, in syenite: also in the Beisnor of St. Francis, Beauce Co. Grains are found in the gold sand of California.

The Uranium of massive rocks is extensively altered to a dull white opaque substance, called leucite by G. Ueber. This for the most part is to be identified with titanite.

Pyrophanite. Manganese titanate, MnTiO_3 . In thin tabular rhombohedral crystals and scales. near titanite in form of TiO_2 . U. = G. O. = 4587. Luster vitreous to sub-siliceous. Color deep blood-red. Streak ochre-yellow. From the Harstgö mine, Pajsberg, Sweden.

111. Intermediate Oxides.

The minerals here included are retained among the oxides, although chemically considered they are properly oxygen-salts, silicates, phosphates, manganates, etc., and hence in a strict classification to be placed in section 5 of the Oxygen-salts. The one well-characterized group is the Spinel Group.

Spinel Group. $\text{RO} \cdot \text{R}'_2\text{O}_3$. Isometric.

Spinel $\text{MgO} \cdot \text{Al}_2\text{O}_3$,

Ceylonite $(\text{Mg}, \text{Fe})\text{O} \cdot \text{Al}_2\text{O}_3$,

Chlorospinel $\text{MgO} \cdot (\text{Al}, \text{Fe})_2\text{O}_3$,

Picotite $(\text{Mg}, \text{Fe})\text{O} \cdot (\text{Al}, \text{Cr})_2\text{O}_3$,

Hercynite $\text{FeO} \cdot \text{Al}_2\text{O}_3$,

Orthite (Atomelite) $\text{ZnO} \cdot \text{Al}_2\text{O}_3$,

Dyalnite $(\text{Zn}, \text{Fe}, \text{Mn})\text{O} \cdot (\text{Al}, \text{Fe})_2\text{O}_3$,

Kreittelite $\{\text{Zn}, \text{Fe}, \text{Mg}\}\text{O} \cdot (\text{Al}, \text{Fe})_2\text{O}_3$,

Magnetite $\text{FeO} \cdot \text{Fe}_2\text{O}_3$,

$\{\text{Fe}, \text{Mg}\}\text{O} \cdot \text{Fe}_2\text{O}_3$.

Kagnedoferrite $\text{MgO} \cdot \text{Fe}_2\text{O}_3$,

Franklinite $(\text{Fe}, \text{Zn}, \text{Mn})\text{O} \cdot (\text{Fe}, \text{Mn})_2\text{O}_3$,

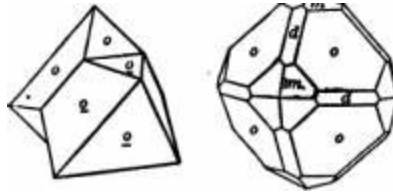
Jaoobite $(\text{Mn}, \text{Mg})\text{O} \cdot (\text{Fe}, \text{Mn})_2\text{O}_3$,

Chromite $\text{FeO} \cdot \text{Cr}_2\text{O}_3$,

$(\text{Fe}, \text{Mg})\text{O} \cdot (\text{Cr}, \text{Fe})_2\text{O}_3$,

The species of Spinel Group are characterized by isometric crystallisation, and, further, the octahedron is throughout the common form. All of the species are hard; those with nonmetallic luster up to 5.5-8, the others from 5.5 to 6.9.

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isometric. Usually in octahedron, rarely cubic. Twins: tw. <pl, and oomp.-face common (Fig. 663), hence ***••• often called spinel-twins; also re-

peated and polysynthetic, producing tw. lamellae.

Cleavage: none imperfect. Fracture conchoidal. Brittle. H. = 8. G. = 3.5-4.1. Luster vitreous; splendent to nearly dull. Color red of various shades, passing into blue, green, yellow, brown and black; occasionally almost white. Streak white. Transparent to nearly opaque. Refractive index: $n_x = 1.7155$ Na, Dz.

Comp., Tar.—Magnesium aluminate, MgAl_2O_4 , or $\text{MgO} \cdot \text{Al}_2\text{O}_3$, = Alumina 71.8, magnesia 28.2 = 100. The magnesium may be in part replaced by ferrous iron or manganese, and the aluminium by ferric iron and chromium.

Var.—Kunzt Spinel or Magnesia Spinel.—dark red or reddish; transparent to translucent; translucent to subtransparent. G. = 3.63-3.71. Composition normal, with little or no iron, and sometimes chromium oxide to which the "red color has been ascribed. The varieties are: (a) net-Sul, deep red; (b) Balat-B, rose-red; (c) SvHeel, yellow or orange-red; (d) J, inandin, violet.

Chloantite or Pseudobrookite, Iron-Magnesia Spinel. —Color dark green, brown to black, mostly opaque or nearly so; G. = 3.78-3.86. Contains iron replacing the magnesium and perhaps also the aluminium, hence the formula $\text{Fe}(\text{Mg}, \text{Fe})\text{O} \cdot (\text{Al}, \text{Fe})_2\text{O}_3$.

Chloantite or Magnesia-Iron Spinel. —Color grass-green, owing to the presence of copper; G. = 3.681-3.694, Contains iron replacing the aluminium, $\text{MgO} \cdot (\text{Al}, \text{Fe})_2\text{O}_3$.

Pseudobrookite or Chrome-Spinel.—contains chromium and also has the magnesia largely

replaced by Iron. $(Mg,Fe)O \cdot (Al,Cr)_2O_3$, hence lying between spinel proper and chromite. $Q = 4-08$. Color dark yellowish brown or greenish brown. Translucent to nearly opaque. Pyr., etc.—B.B. alone infusible. Slowly soluble in borax, more readily in salt of phosphorus, with which it gives a reddish blue while hot, becoming faint chrome-green on cooling. Black varieties give reactions for iron with the fluxes, Soluble with difficulty in concentrated sulphuric acid. Decomposed by fusion with potassium bisulphate.

DIS.—Distinguished by its octahedral form, hardness, and infusibility; zircon has a higher specific gravity; the true ruby (p. 835) is harder and is distinguished optically; garnet is softer and fusible.

Obs.—Spinel occurs embedded in granular limestone, and with calcite in serpentine, gneiss, and allied rocks. Ruby spinel is a common associate of the true ruby. Common spinel is often associated with chondrodite. It also occupies the cavities of masses ejected from some volcanoes. Spinel (common spinel, also picotite and chromite) occurs as an accessory constituent in many basic igneous rocks. Especially those of the peridotite group; it is the result of the crystallization of a magma very low in silica, high in magnesia and containing alumina; since, as in many of the peridotites alkalis are absent, feldspars cannot form, and the Al_2O_3 and Cr_2O_3 (also FeO perhaps) are compelled to form spinel (or corundum). The serpentines which yield spinel are altered peridotites

In Ceylon, in Siam, and other eastern countries, occur of beautiful color, as rolled pebbles: in upper Burma with the ruby (cf. p. 834). Pleonaste is found at Candy. In Ceylon; at Aker, in Sweden, a pale blue and pearl-gray variety in limestone; small black splendid crystals occur in the ancient ejected masses of Monte Somma; also at Pargaa, Finland, with chondrodite. etc.; in compact geodes at Honzoni, in the Fassa valley.

From Amity, K. Y., to Andover, N. J., a distance of about 80 miles, is a region of granular limestone and serpentine. In which localities of spinel abound; color, green, black, brown, and less commonly red, along with chromite and other minerals. Localities are numerous about Warwick, and also at Monroe and Cornwall: Gouverneur, 8 m. N. and 1 m. W. of Somerville, Bt. Lawrence Co.; green, blue, and occasionally red varieties occur at Bolton, Boxborough, etc., Haas, Franklin, N. J., yields crystals of various shades of black, blue, green, and red; Newton, Vermont: Sparta, Hamburgh and Vernon, N. J., are other localities. With the corundum of North Carolina, as at the

;Coogic

Cullagee mine, near Franklin, Hancock Co.; Bimlily at Dudleyville, Alabama. Spinel ruby
lit Gold Bluffs, Humboldt Co., Cal.

(ivod black spinel first found at Burgess, Ontario; a bluish spinel variety); a rough cubic form occurs at Wakefield, Ottawa Co.; blue spinel at Daillebout, Joliette Co., Quebec.

Heroynite. Iron Spinel, FeAl_2O_4 . laoineiric; massiTe, fine granular. H. = 7.5-8. G. = 3.91-3.97. Color black. From Roneberg, at the eastern foot of the Bohmerwald. A related iron-alumina spluel. nrltb about 9 p. c. HgO , occurs with magnetite and corundum in Cortlandt township, Westchester Co., N. Y.

aAHKITEI. Zinc-Biphenyl.

Isometric. Habit octahedral, often with faces striated | edge d/o; also less commonly in dodecahedrons and modified cubes. Twinning: tw.-pl. o.

Cleavage: o indistinct. Fracture conchoidal to uneven. Brittle. H. = 7.5-8. G. = 4.0-4.6. Luster vitreous, or somewhat greasy. Color dark green, grayish green, deep leek-green, greenish black, bluish black, yellowish brown, or grayish brown; streak grayish. Subtransparent to nearly opaque.

Comp., Tar.—Zinc aluminate, ZnAl_2O_4 , = Alumina 55.7, zinc oxide 44.3 = 100. The zinc is sometimes replaced by manganese or ferrous iron, the aluminium by ferric iron.

Var.—AlyroMOLiTB. or Zinc GoAnifc.— ZnAl_2O_4 , with sometimes a little iron. G. = 4.0-4.6. Colors as above given,

Dtbloitk. or Zinc-Manganite-Ironite.— $(\text{Zn}, \text{Fe}, \text{Mn})\text{O}(\text{Al}, \text{Fe})_2$. Color yellow-olive to grayish brown, O. = 4.0-4.6,

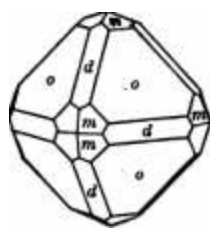
Krbittontk, or Zinn. Iran Gafti (».) $(\text{Zn}, \text{Fe}, \text{Mg})\text{O}(\text{Al}, \text{Fe})_2$. In crystals and granular massive. H = 7-8. Q. = 4.48-1.50. Color variable—black to greenish black; powder grayish brown. Opaque.

Pyr., etc.—Gives a coating of zinc oxide when treated with a mixture of borax and sodium carbonate on charcoal; otherwise like spinel.

Obs.—Occurs in talcoae schists at Falun, Sweden (nuttomoltr); at Tiriola, Calabria; at Bundenmais, Bavaria (kreiUoitiU); Hinas Geraes. Bmzll. In the Franklin Furnace, N. J., with franklinite and willemite; also at Sterling Hill, N. J. (dy»ii«to); with pyrite at Itowe. Mass.; near a field!par quarry in Delaware Co. Pa.; especially at the Deake mica mine, Mitchell Co., N. C; at the Canton Mine, Georgia; with galena, chalcopyrite, pyrite at the CoKipaxi mine, Cripple Creek, Colo.

Named after the Swedish chemist Gahn. The name Aulomotile. Either, is from avTOftoXoi. a deteler, alluding to the (act of the zinc occurring in an unexpected place.

Magnetic Iron Ore. Magnetite, Fe_3O_4 . Isometric. Most commonly in octahedrons, also in dodecahedrons with faces striated \ edge d/o from oscillatory combination; in dendrites between 668. 667. sea.



plates of mica ; crystals sometimes highly modified ; cubic forms rare. Twins: tw.-pl. o, sometimes as polysynthetic twinning lamellae, producing

Si" DESORIPrrVE MIMBRALOGY.

striation on an octahedral face and often a pseudo-cleavage (Fig. 456. p. 136). UsBtive with laminated structure; granular, coarse or fine; immiipiible.

Oleayage not distinct; parting octahedral, often highly developed. Fracture subcouchoidal to uoeven. Brittle. H. = 65-60. G. = 5*168-5-HJl> crystak. Lnater metallic and eplendent to submetallic and rather dull. Color iron-black. Streak black. Opaque, bat in thin dendrites in uicca nearly transparent and pale brown to black. Strongly magnetic; sometimes possess-iug polarity (lodestoe).

Comp, Tar.—FeFe.O. or FeO.Fe,o, = Iron sesquioxide 69-o, iron protoxide 31-o = 100 ; or, Oxygen 27'6, iron 72 4 = 100, The ferrous iron sometimes replaced by magnesium, and rarely uickel; also sometimes contains titanium (up to 6 p. o. TiO,).

Vmt. — Ordinarj/. —(a)

coiiaae or flue, (e) As lot . , , „ „

Is attracted by a magnet but bus uo power of uUraeiing paniclea of iron iiself. Tie proprtry or polarity wbfcb dtstluguislies tbe lodtttont (less pro|>«rly writteti loadsione) U exceptioiml.

Magnettian. Talk-Eisenerz.—Q. ■= 4'4l-4'42; lutler submetallic; wciik roiignt-tic; In crystals from Sparta. N. J., and elsewhere.

Manganetian. —Containing B'8 to 6 3 p. c. muuganesc (Manganmagnttilt). From Vesier Bilfberj, Sweden.

Pyr., etc.—B B. very difficultly fusible. In OF. loses Us influence on the innngct. Willi the Uuxes reacts like bemslile. Soluble in hydrocblHc acid.

Siff—Distinguished from otlier members of the spinel graiit>. as also from garnet, bj ils being attrac led by thomnguet, as well nsbyita high apeciflc gravity; rraDkliuiteandcbromite are only feebly maKnetic (if at all), and luive a brown or blackish-brown streak; nl^^o, n-bcn massive, by its black streak from hematite and liionite; much harder than letrHbethiu-.

Obi.—Magnetite is mostly confined to crystalline rocks, and is most abundant in metamorphic rocks, though widely distributed also in veins in eruptive rocks. In the Archean rocks the beds are of immense extent, and occur under the same conditions as those of hematite. It is an ingredient in most of the massive variety of iron called emery. The earthy magnetite is found in bogs like bog-iron ore. Occurs in meteorites, and forms the crust of meteoric irons.

Present in deuterite-like forms in the mica of many localities following the direction of the lines of the percussion-figure, and perhaps of secondary origin. A common alteration-product of minerals containing iron protoxide, e.g., present in veins in the BGR resulting from altered chrysolite.

The beds of ore at Arendal, Norway, and nearly all the celebrated iron mines of Sweden, consist of massive magnetite, as at Dannemora and the Taberg in Småland. In Sweden, and Corsica, small octahedral crystals, embedded in chlorite. Splendid dodecahedral crystals occur at Nordmark in Wernland. The most powerful native magnets are found in Siberia, and in the Harz; they are also obtained on the island of Elba. Other localities for the crystallized mineral are Traversella in Piedmont; Achmatovsk in the Ural; Balmuccia, near Predazzo, in Tyrol, also Rothenkopf and Wildkreuzjoch; the Blumenthal, Switzerland.

In N. America, it constitutes vast beds in the Adirondack region, Warren, Essex, and Clinton Cos., in Northern N. York, especially in Bt. Lawrence Co. the iron ore is mainly hematite; One crystal and masses showing broad partings; surfaces and yielding large pseudocrystals are obtained at Port Henry, Essex Co.; similarly in New Jersey; in the Caledonia, in Hull. Grenville. Madoc, etc.; at Oimwall in Pennsylvania, and Magnet Cove, Arkansas. It occurs also in N. York, in Saratoga, Herkimer, Oranfre. and Putnam Cos.; at the Tilly Foster iron mine, Brewster, Putnam Co., in crystals and masses accompanied by chondrodite, etc. In N. York, near Hamburg, near Franklin Furnace and elsewhere. In Penn., at Quakertown, Chester Co., and at the French Creek mines; sometimes forming hexagonal figures in mica at Pennsylvania. Good lodes are obtained at Magnet Cove, Arkansas. In California, in Sierra Co. abundant, massive, and in crystals; in Plumas Co.; and elsewhere. In Washington, in large deposits

Named from the loc. Magnesia, bordering on Macedonia. But Pliny favors Nicander's derivation from Hagnai, who first discovered it, as the fable runs, by finding, on taking his herds to pasture, that the nails of his shoes and the iron ferrule of his staff adhered to the ground.

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FRANXUNCTB.

Isometric. Habit octahedral; edges often rounded, and crystals passing into rounded

grains. Massive, granular, coarse or fine to compact.

Pseudocleavage, or parting, octahedral, as in magnetite. Fracture conchoidal to uneven. Brittle. H. = 5.5-6.5. G. = 5.07-5.22. Luster metallic, sometimes dull. Color iron-black. Streak reddish brown or black. Opaque. Slightly magnetic.

Comp.—(Fe,Zn,Mn)O.(Fe,Mn),₂, but varying rather widely in the relative quantities of the different metals present, while conforming to the general formula of the spinel group.

Pyr., «to.—B.B. fusible. With borax in O.F. gives a reddish translucent bead (manganese), and in R.F. this becomes bottle-green (Iron). With soda gives a bluish-green manganate, and on charcoal a faint coating of zinc oxide, which is much more marked when a mixture with borax and soda is used. Soluble in hydrochloric acid, sometimes with evolution of a small amount of chlorine.

DIA—Resembles magnetite, but is only slightly attracted by the magnet, and has a dark brown streak; it also reacts for zinc on charcoal B.B.

Obs.—Occurs in cubic crystals near Eibach in Nassau: in amorphous masses at Astenberg, near Aistipelle. Abundant at Mine Hill, Franklin Furnace, N. J., with willemite and luicite in granular limestone; also at Sterling Hill, two miles distant, associated with willemitite.

Hagnaiolite. Manganoferrite. $MgFeO_2$. In octahedrons. H. = 6-6.5. G. = 4.968-4.654. Luster, color, and streak as in magnetite. Strongly magnetic. Formed about the fumaroles of Vesuvius, and especially those of the eruption of 1669.

Jacobite. $(Mg,Fe,Mn)_2O_3$. Isometric; in distorted octahedrons. H. = 5.5. G. = 4.75. Color deep black. Magnetic. From Jakobsberg, in Kordmark, Wermland, Sweden; also at Långban.

OHROMITTE.

Isometric. In octahedrons. Commonly massive; fine granular to compact

Fracture uneven. Brittle. H. = 5.5. G. = 4.32-4.57. Luster submetallic to metallic. Color between iron-black and brownish black, but sometimes yellowish red in very thin sections. Streak brown. Translucent to opaque. Sometimes feebly magnetic.

Comp.— $FeCr_2O_4$, or FeO, Cr_2O_3 , = Chromium sesquioxide 68.0, iron protoxide 32.0 = 100.

The iron may be replaced by magnesium; also the chromium by aluminium and ferric iron. The varieties containing but little chromium (up to 10 p. e.) are hardly more than varieties of spinel and are classed under picotite, p. 338.

Pyr., etc.—B.B. in O.F. infusible: In R.F. slightly rounded on the edges, and becomes magnetic. With borax and salt of phosphorus gives beads which, while hot, show only a reaction for iron, but on cooling become chrome-green: the green color is heightened by fusion on charcoal with metallic tin. Not acted upon by acids, but decomposed by fusion with potassium or sodium bisulphate.

IWff.—Distinguished from magnetite by feebly magnetic properties, streak and by yielding the reaction for chromic acid with the blowpipe.

Oba.—Occurs in serpentine, forming veins, or in embedded masses. It assists in giving the variegated color to verde-antique marble. Not uncommon in meteoric irons, sometimes in nodules as in the Cobalt Iron. less often in crystals (Lodran).

Occurs in the Oulsen mountains near Kranbat in Syria: in crystals in the Islands of Tristan da Cunha; in tin- province of Trondhjem in Norway: in the Department du Var in France; in Silesia and Bohemia; abundant in Asia Minor: in the Eastern and Western Urals; in New Caledonia, affording ore for commerce.

At Baltimore, Md., in the Bare Hills. In veins or masses in serpentine; also in Mont-

Smery Co., Pa. In Pennsylvania, Chester Co., near Unionville, abundant: at Wood's Mine, near Texas, Lancaster Co., very abundant. Usual and in crystals at Hoboken, N.

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DESCRIPTIVE MINERALOGY.

ORTSOBBRTI. Cinnabarine. Orthorhombic. Axes $d:h:i = 0.4701:1:0.5800$. $e \ll \gg$. 670.

Twins: tw.pl./sfol), both contact-and penetration-twins; often repeated and forming pseudo-hexagonal crystals with or without re-entrant angles (Fig. 357, p. 123). Crystals generally tabular || a. Face a striated vertically, in thin a feather-like striation (Fig. 670).

Cleavage: 1 (01) quite distinct; imperfect; none more so. Fracture uneven to conchoidal. Brittle. H. = 8.5, G. = 3.84. Luster vitreous. Color asparagus-green, grass-green, emerald-green, greenish white, and yellowish green; greenish brown; yellow; sometimes raspberry- or columbine-red by transmitted light. Streak uncolored. Transparent to translucent. Sometimes a bluish opalescence or chatoyancy, and asteriated. Pleochroic, vibrations lb (=h) orange-yellow, c (= d) emerald-green, a (= d) columbine-red. Optically-I-. Ax. pi. | (J. Bi ± «. $\gamma = 1.7484$. $2E = 84^\circ 43'$.

Vat. 1. On Knary.—Color pale green, being colored by iron; Blauwand transparent and used as a gem.

3. Alexandrite.—Color emerald-green, but columbine-red by transmitted light; valued as

a gem. G. \wedge 3-644, meaD of resiiltg. Supposed to be colored by chromiuin. Crystals often very large, nnd !u twins, like Fig. 857. either six-sided or sixrayed.

8. C'ai'i-eye.—Color greeuiah and esbibiiing a floe clialoynt effect; from Ceylon. Comp.—Beryllium aluminate, $BeAl_2O_3$, or $BeO \cdot Al_2O_3$, = Alumina 802, glucina 19-8 \wedge 100.

Pyr., etc.—B.B. alone unaltered; with soda, tbe surface 1b merely rendered dull- With borax or gaU of pboapltarus fuses with great difficulty. With cobalt solution, the powdered minenil gives a bluish color. Not attacked by acids.

DIB. —Distinguished by Its extreme hardness, grea llity; also characterized by Its tabular crystallization. _..

Ob«,—In HIInaa Oeraes, Brazil, and also io Ceylon. In rolled pebbles; at Hanchendorf In Mora via, i[i the Ural, 8G verste from Ekaterinburg, In mica slate with beryl and pbenacte, tlie variety alexandrite; fn the Orenburg district, B. Ural, yellow; in the Hourne Uta., Ireland.

In the U. 8.. at Haddam, Ct.. in cranite traverilng gneiss, with tonrmallue. garnet, beryl; al Greenfield, near Saratoga, N, T.. with tourmaline, garart, and apatl«; Norway, He,, in granite with garnet; also al Stoneham, with flbrolfle, etc,

Ghrytdberyl is from j'pi^froi, golden, ff^fivXXoi, ierj/l. Cj/mophane, from Kvua, wave, and ipaivo), appear, nlliidt-s to n peculiar opalescence the crystals sometimes exhibit. , Al^anArile Is after Ihe Czar of Russia, Alexauder I.

HaDunamiite. HniO, or $HnO \cdot M, Oi$. In tetrngonal octahedrons and twins (Fig, 876, p, 128); also graoular massive, parlieles strongly coherent. H. = IS-5'5. O. = 4*858. Luster anbmelalllc. Color brownish black. Streak chestnut-brown. Occurs near Ilmenau in Tburiugia; Ilefeld in the Ilarz; Fillpsiad, L&ngban, Nordmark, In Sweden.

;,CoogIC

. Uennfge Germ. Pb_3O_4 , or $2PbO \cdot PbO$,. PuWeruleDt. u cryElalltne scAlea. Q. = 4'o. Color ylvid red, mixed with yellow; streak orange-yellow. Occurs ul BleEalf lu tbe Ejfel; Badenweiler In Baden, etc.

Orednarit*. CuiHa,U, or $8CuO \cdot 2M[i, Oi$. FolliLted cryatalline. H. = 4'5. O. = 4-9-S'l. Luster metallic Color iron-bluck to Bteel-gi-ny. Slrenk block, biuwilsl. From Fried riclia rode.

PsendobrooUta. ProbnblT Fei(T10,)i. TJBually Id minute ortborbombic crystola, tsbuUr I It kod often prismaLfc | S. G. = 4-4-*-l«t. Color dark browu to black. Streak ocher-yullow. Fouod with Lyperstbene (szaboie) Id cavitlee of the andeslte of AraDver Bvrg, TnusylTaDlu, aod eiaewliere: ou recent lava (1872) from Vesuvius; at Havredal, Biimle. Korway, embedded in kjerulflne (wagueritej altered to upatlie.

Tetragonal. Axis $a = 0.9850$. Commonly in octahedrons, nearly isometric in angle ($\alpha = 70^\circ 7'$). Also massive.

Cleavage: perfect. Fracture uneven to subconchoidal. Brittle. $H = 6 \sim 6.5$. $G = 4.75 - 4.82$. Lustre metallic. Color dark brown to black to steel-gray. Streak same.

Comp. — $3Mn, O, MnSiO_3$ — Silica 10.0, manganese protoxide 11.7, manganese sesquioxide 78.3 = 100,

P.T. — B.B. Infusible. With borax and salt of phosphorus forms a transparent bead in O.F. becoming colorless in R.P. With soda gives a blue-green bead. Dissolves in hydrochloric acid evolving chlorine, and leaving a residue of gelatinous or flocculent silica (Rg). Microscopically gelatinous with acids.

Obs. — Occurs in veins traversing porphyry, at Oehrenstock, near Ilmenau; near Uefeld in the Harz; Si. Marcel in Piedmont; at Elba: at Botvedal. Upper Telemark. In Norway; in the manganese mines of Jakobsberg, Sweden, also at Aagban, and at the Sjömla, Örebro. Ore. Microcline from Si. Marcel, Piedmont, is impure braunite.

Habit. Essentially $FeO.MnO$, in massive isometric crystals. $H = 6 - 6.5$. $O = 4 - 5$. Occurs with topaz in cavities in rhyolite; from Utah.

IV. Dioxides, RO_2 . Rutile group. Tetragonal.

The Rutile Group includes the dioxides of the elements tin, manganese, titanium, and lead. These compounds crystallize in the tetragonal system with closely similar angles and axial ratio; furthermore in habit and method of twinning there is much similarity between the two best known species included here.

With the Rutile Group is also included Zircon. $ZrO_2.SiO_2$; $h = 0.404$. In this work, however, Zircon is classed among the silicates, with the allied species Thorite, $ThO_2.SiO_2$. $h = 0.408$.

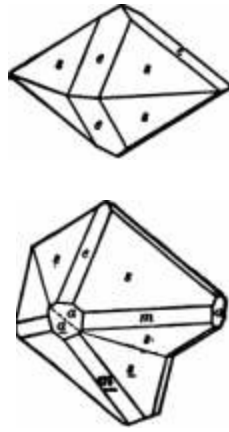
A tetragonal form, approximating closely to that of the species of the Rutile Group, belongs also to a number of other species, as Sellaite, HgF_2 ; Topolite, $Fe(Ta, Nb)_2O_7$; Xenotime, $TbPO_4$, etc.

It may be added that Zircon, as the species Baddeleyite, crystallizes in the monoclinic system.

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DESCRIPTIVE MINERALOGY,

OASSITBRITE. Tin-Ore. ZlonaWii Germ. TetragoDal. Axis i = 0'6723.



«*, 101 A Oil = 46° 38'. nu, 110 A 111 = 48' Sr.

«". 101 A 101 = er W. ff", 331 A 281 = 20' Mi'.

m", 111 A ill = 58* IV. B'l', 821 A 831 = 61' 42".

m". Ill A in = 87° r.

Twins common: tw. pi. e, both contact- and penetration-twins (Fig. fl73, also Fig. 373, p. 134); often repeated. Crjstale low pjrimalal; also prismatic and acutely terminated. Often in reriform shapes, structure fibrous divergent; also maesive, granular or impalpable; in rolled grains.

Cleava^: a imperfect; s {ill) more so; »i hardly distinct. Fracture Bubconchoidal to uneven Brittle. U. = 6-7. G. = 6-8-71. Luster adamantine, and crystals usually splendid. Color brown or black; sotnetimea red, gray, white, or yellow. Streak white, grayish, brownish. Nearly transparent to opaque. Optically -j-. Indices: oOj = 19966, €j = 2*0934.

Tar.— Ordinary. Tin.alODe. In ciTBtals and massive.

Wood-lln (Holzztniierz Oerm.). In bolryoldal and renlform shapeB, coDcentrJc in slructure, and ladinled flbrouu iutemallj, althlioughli very com pad, will) the color browuilli, of mixed shades, looking somenljal like dry wood in ila c<>lorB. ToaiTt-tj/e tin U llie same, on a smaller scale. Stream tin Is the ore In liie slute of sand, as It occurs along llie beds of streams or tn gmvel.

Comp.—Tin dioxide, SiiO, = Oxygen 21-4, tin 78-6 = 100. A little Ta,o, is sometimes present, also Fe,o,.

Pyr., etc.—B.B. alone nnalteicd. On cbarconl wllb sods reduced to metallic tin, and fdvea a wliitc costing. With the fluxes aomelimes gives reactfoos for fron and manganese. Only slightly acted upon by acids.

Dlff.—DisllniTulstied by Itn hiph specific fravily, bnrndnesa, infuaiblilllv. and by its yielding metallic tin B B ; resembles some varieties of garnet, spbaltrrite, aod black tourmaline. Specific gravity (8'5) higher than thai of i-utile (4): wolframite is easily fusible.

Oba.—Occurs In veins tnvverstng granite, gneiss, mica schist, ehloi'lle or clay schist, and porphyry; also in finely rellculaled veins forming the ore deposits called slockworks, or simply impregnating the enclosing mck. The cotnmonly associated minerals tire quartz, wolframite, sclieellte; also mica, tnpa/. tourmaline, npalile, fiuorlte; further pyrllte, urseuo-pyrllte, sphalerite; molybdenite, nllve bismnth, etc.

Formerly very abundant, now less sii. in Cornwall, in fine crystals, and nls0 as wood-tin and stream-tin ; in Devonshire, near Tavistock nri elsewhere; In pseudomoiphs after feldspar at Wheal Coates. near 8t. Apnea. Cornwall; in fine crystals, often twins, at Bcblackenwald, Graupen, Joachimsthal, Zinnwald, etc., in Bobemiu, and at E liven fried ers-dorf. Allenberg, etc., in Baiony; at Limoges lu splendid crystals; Sweden, at FInbo; Finland, at PitkOranta.

In the E. Indies, on the Malay peninsula nf Malacca and the neJi-Lboring islands. Banca. and Bilitng near Borneo, Tn New South Wales abnndanl over an area of 8500 sq. n ilea. also In Vlclona. Qiiei-Tislnd and Tasmania. In Bolivia; Hc:iico, in Durango; alio Guanajuato, Zacalecos, Jalisco.

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OXIDES.

345

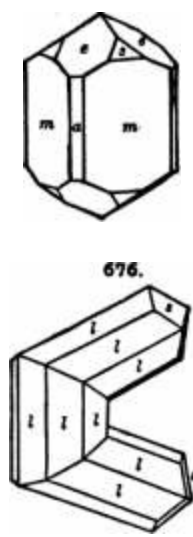
In the United States, lo Maine. spariDgljr at Paris, Hebron, etc. Id Matt., &t Cliesterfeld nud Gi)SbeD, rare. In N. Hamp,, at JucksoD. In Virginia, oa Irisli Creek, Rockbri(lt'<! Co.. with wol fremfte, etc. In Altibama. io Coosa Co. In 3. Dakota near Haroey Peak iind near Custer Citj \a tbe Black Hills, where It has been mined. Id Wyoming, in Crook Co, Id Montana, Denr Dillon. In California, iu San Bernardiuo Co., at Temescui.

FoliauIt«, Uanganese dioxide, HnOi. In compotiile parallel groupings of minute ciyslais: also forming tbe outer sUei) of crystals liaving Ihe form of mnnganite. H, = (t-6'5. G. = 4'993. Luster metallic. Color light steel-gray or iron-gray. Streak black. From Platten, Bohemia. It Is diatingulshed from pyrolu^te by its bardnesa and Ita aubydrous cLaracier. Like pyrolustle It is often a paeudomorp b after manganlie.

RUTILB.

Tetragonal. Axis 6 = 0*64415.

474. 675.



$U^{\circ}K 810 A 310 = 86^{\circ} 54'$. ft^* . $Ill A ill = 66^{\circ} 52';$.

$\bullet e''$. $101 A Oil = 45^{\circ} 2''$. \ll'' . $Ill A 111 = 84^{\circ} 40'$.

\ll'' , $101 A iOl = 65^{\circ} U_i'$. $If. 818 A 188 = 20^{\circ} 6'$.

Twine: tw. pi. (I) e, often geniculated (Fig. 676); also contact-twins of very varied habit, sometimes sizlings and eigbtlings (Fig. 361, p. 122; Fig. 375, p. 124). (2) r(301) rare, contact-twins (Fig. 377, p. 125), Crystals commonly prismatic, Tertically striated or furrowed; often slender acicular. Occasionally compact, massive.

CleaTage: a and m distinct; e in traces. Fracture subconchoidal to uneven. Brittle. $H. \wedge 6-6'5$. $G. = 4-18^{-25}$; also to 6-2. Luster me tall ic-ttdamantine. Color reddish brown, passing into red; sometimes yellowish, bluish, violet, black, rarely grass-green; by transmitted light deep red. Streak pale brown. Transparent to opaque. Optically +. Refractive indices high: CO, = 2-6158, e, = 2-9029 for Na. Birefringence very high. Sometimes abnormally biaxial.

Oomp., Tar.- \wedge Titanium dioxide, TiO_2 , = Oxygen 40-0, titanium 600= 100. A little iron is usually present, sometimes up to 10 p. c.

Var.—Ordinary. Brownish red and other shade*, not black. $Q. = 4-18(-35)$. Trans.

Birent tjuarlz (tagenite) is somellmes penetrated thickly with acicular or capillary crysliila. rirk smoky quartz peneiraied with the acicular rutile is the FlScbes d'amour Fr. (or Venus hair-stone). Acicular cryatala often implanted Id parallel position on tabular crystals of hematite: atao somewhat similarly on magnetite.

Ftrrif«rov». ia) Niffrint la black in coloi'. whence the name: ronlnins 2 to 3 p. c. Fr,Oi (S) limenor-ttlile is n bbick varltty frim the Ilmeii Mis., cimlalnng up Io 10 p. or more of Pe_2O_3 . $G = 5.07 - BIS$

Pyx., ato.—B B. infuaible. With salt of phnsponis pivea a colorless bead, which In R.F.

assumes a violet color on cooling. Most varieties contain iron, and give a brownish-yellow or red bend in R.F.. The violet only disappears after treatment of the lead with metallic tin on charcoal. Insoluble in acids; made soluble by fusion with an alkali or

; CoogIC

DESCRIPTION OF UINEBALOOT.

alkalide carbonate. The solution with an excess of acid, with the addition of tin-foil, gives a beautiful violet color when concentrated.

DiS.—Characterized by its peculiar sub-adamantine luster and brownish-red color. Differs from tourmaline, vesuvianite, augite in being entirely unaltered when heated alone. B.B. Specific gravity about 4, of caesitite 6.6.

Obs.—Rutile occurs in granite, gneiss, mica schists, and syenitic rocks, and sometimes in

serpentine and dolomite; common, as a secondary product, in the form of microlites in many schists. It is generally found in embedded crystals, often in masses of quartz or feldspar, and frequently in acicular crystals penetrating quartz; also in biotite (wh. see), and has been observed in diamond. It has also been met with in hematite and ilmenite, rarely in chromite. It is common in grains or fragments in many auriferous sands.

Prominent localities are: Arendal and Kragerø in Norway; Horraliberg, Sweden, with lazulite and cyanite; Saualpe. Carinthia: in the Urals; in Tyrol; at St. Othard and Biiuen-thai, Switzerland; at Trioux, near Limoges in France; at Oblaplan in Transylvania, nigra

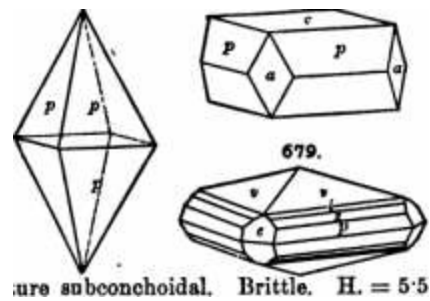
northern Vermont, acicular. Some specimens of great beauty in transparent quartz. In Massachusetts, at Barre, in gneiss; at Belbume, in mica slate. In Virginia, in Orange Co., Edenville; Warwick; E. of Amity. In Pennsylvania, at Budsbury, Chester Co., and the adjoining district in Lancaster Co.; at Parkburg, Concord, West Bradford, and Newlin, Chester Co.; at the Poor House Quarry, Chesler Co. In New Jersey, at Newton, with spinel. In South Carolina, at Camden's Mountain; at Stony Point, Alexander Co., in splendid crystals. In Georgia, in Habersham Co.: in Lincoln Co., at Graves' Mountain, with lazulite in large and splendid crystals. In Arieantat, at Harnet Gove, commonly in twins, with brookite and perovskite, also as pumice after eruption.

Plattnerite. Lead dioxide. PbO₂. Rarely in prismatic crystals, usually massive. H. = 3-5. G. = 5.5. Luster metallic. Color iron-black. Streak chestnut-brown. Found in Leadhills and Wanlockhead, Scotland. Also at the "As You Like" mine, Ullai, Caucasus Mountains, Idaho.

Baddeleyite. Zircon dioxide, ZrO₂. In tabular monoclinic crystals. H. = 6. O. = 5-6.

Colorless to yellow, brown and black. From Geyloa; also Jacupiranga, Brazil (tyratUiit) where it la associated with wirkdiU, (Ca,I^}o.3(Zr,Tl,Th)O.,

OOTAEBDBini. Anatase. TetragonaL Asia i = 17771. «79. 618.



ture subconchoidal. Brittle.

Commonly octahedral in habit, either acnte {p, 111), or obtuse {v, 117); also tabular, e predomi-Tiating; rarely prismatic crystals; freqaeutly highly modified.

m", 101 A on = 76° 5'.

et/\ 101 A lOl = lai* 18'.

j>^. Ill A 111 = 82° 9'.

Pt*'. niAlil = 136° 86'.

•f", 118aI13= 84° r.

mf ii7Aii7= 2ri».

Cleavage: c and p perfect. Prac-" G, = 3 82-3 95; sometimes 411-

416 after heating. Luster adamantine or metallic-adamantine Color variotis shades of brown, passing into itidigo-blue, and black; greenish yellow by transmitted light. Streak uncolored. Transparent to nearly opaque. Optically —. Birefringence rather high. Indices: oOj = 2'554, e, = 2*493. Sometimes abnormally biaxial.

Comp.—Titaninm dioxide, TiO, = Oxygen 40'o, titanium GO-O = 100.

Pyr., etc.—Same as for rutile.

Om.— Host abundaut at Bourg d'Olsnns, In DaupblnS, witli feldspar, asintle, and

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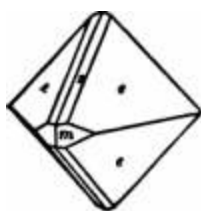
OXIDES.

Ilmenite; near Hof in the Fichteleebirge: Norway; in the Tynall; in chlorite in Devonshire, near Tynallstock; with brookite at Tremadoc in North Wales; in Cornwall, [near Liskeard and the Tullaghd Cliffs; in Brazil in quartz, also in dolomite in Switzerland in the Bienna; in the variety vanadate, long supposed to be titanite; also in Cavimdi, Tavetsch; Rauris, Salzburg, in the Kasteru Alps; also at Pfaffers Juch.

In the U. States, at the Dexter lime rock, Smithfield, R. I. In dolomite in the Washington at Brudenriew, Burke Co., N. C. In titanite in tabular crystals.

Orthorhombic. Axes $a:h:c = 0.8416 : 1 : 0.9444$. 680.481 .

$\alpha:J$



$\alpha, 110 \text{ a } 110 = 80^\circ$ IV. $112 \text{ a } 112 = 08^\circ 48'$. $112 \text{ A } 118 = 44' 46'$.

$M^* . 123 \text{ A } 123 = 44^\circ 23'$. $e \ll m$. $182 \text{ A } 132 = TS'' ST. AM, 110 \text{ A } 123 = 4S^\circ 42'$.

Fracture subconchoidal to un-lustrous metallic-adamantine

Only in crystals, of varied habit. Cleavage: m indistinct; c still more so. even. Brittle. H. = 5-5-C. G. = 3-87-4

to siliceous. Color hair-brown, yellowish, reddish, reddish brown, translucent; also brown to iron-black, opaque. Streak uncolored to grayish or yellowish. Optical characters, see p. 225.

Comp.—Titanium dioxide, TiO_2 , = Oxygen 40.0, titanium 60.0 = 100.

Pyr.—Same as for rutilite.

Obs.—Occurs at Binn; in the Oisans in Dauphin: at St. Gothard with albite and quartz; in the Maderaoer Thal, Switzerland: in the Ural, district of Zinlouel, near Allask, and in the gold-washings in the Binnurka river and Elhewbere; at Fronlen, near Tremadoc, Wales.

In the U. S., in thick black crystals (arkansite) at Magnet Cove, Ozark Uta., Arkansas, with anatase, black garnet, schmelzite, rutile, etc.; in small crystals from the gold-washings of North Carolina; at the Imid mine near Ellenville, Ulster Co., N. Y., on quartz, with chalcopyrite and galena; at Paris, Maine,

Named after the English mineralogist. H. J. Brooke (1871-1867).

FTSOLITSm.

Orthorhombic, but perhaps only pseudomorphous. Commonly columnar, often divergent; also granular massive, and frequently in reniform coats.

Soft, often soiling the fingers. H. = 2-2.5. G. = 4.7-4.86. Luster metallic. Color iron-black, dark steel-gray, sometimes bluish. Streak black or bluish black, sometimes submetallic. Opaque.

Comp.—Manganese dioxide. MnO_2 , like polianite (p. 345). Commonly contains a little water (9 p. c), it having had usually a pseudomorphous origin (after manganite).

It is uncertain whether pyrolusite is an independent species, with a crystalline form of its own, or only a secondary mineral derived chiefly from the decomposition of manganite also from polianite (Breith.). Pseudomorphous crystals having distinctly the form of manganite.

by Google

Fry., etc.—Like polianite, but more variable in some respects in the closed tube.

U.S.—Hardness less than that of pyrolusite. Derived from iron ores in its reaction for manganese. Easily distinguished from pyrolusite by its inferior hardness, and usually by its color; its streak is black; that of pyrolusite is more or less brown.

Obs.—This ore is extensively worked at Elgersberg near Ilmenau, and other places in Thuringia; at Vorderehrewald near Mühlhausen; at Platten in Bohemia, and elsewhere; near Jobanurg near Gera; at Hirschberg in Westphalia; Hatzkii, Traubylvanu; in Australia; in

Occurs in the United States abundantly in Vermont, at Braintree, etc.; at Plainfield and West Stockbridge, Mass.; Augusta Cu., Virginia; Pope, Pulaski, Montgomery Cos., Arkansas. In New Brunswick, 7 m. N. of Buxton. In South Carolina, at Tenney Cape; at Walling, etc.

The name is from its color, and *Xanthos*, to oath, because used to discharge the brown and green (FeO) tints of glass; and for the same reason it is whimsically entitled by the French *U tawne à verrou*.

B. Hydrated Oxides.

Among the hydrated oxides the Diaspore Group is well characterized. Here belong the hydrates of aluminum, iron and manganese. The general formula is properly written $BO(OH)_2$. The three species here included are orthorhombic in crystallization with related

angles and axial ratios; this relation is deviated from by manganite in the prismatic zone.

Another less prominent group is the Brucite Group, including the rhombohedral species Brucite, $Mg(OH)_2$, and Pyrochroite, $Mn(OH)_2$.

Oibbsite, $Al(OH)_3$, and Sassolite, $B(OH)_3$, are also related, and further Hydrotalcite and Pyroaurite.

Diaspore Group. $RO(OH)_2$ or $B_2O_3 \cdot H_2O$. Orthorhombic.

Diaspore $Al_2O_3 \cdot H_2O$ 09372 : 1 : 06039 or 0-6443

Gfithite $Fe_2O_3 \cdot H_2O$ 0-9185 : 1 : 0-6068 or 0-6606

Manganite $Mn_2O_3 \cdot H_2O$ 0-6411 : 1 : 0-5448 or 0-6463

DIASPORE.

Orthorhombic. Axes: $a:b:c = 0.9373 : 1 : 0.6039$. Crystals prismatic, $mm^2 = 86^\circ 17'$; usually thin, flattened; sometimes acicular. Also foliated massive and in thin scales; sometimes stalactitic.

Fracture: 6 eminent; cleavage (210) cleavage perfect, Fracture conchoidal, very brittle. H. = 6-5-7. G. = 3-3.5, Luster brilliant; pearly on cleavage-face, elsewhere vitreous. Color whitish, grayish white, greenish gray, hair-brown, yellowish, to colorless. Pleochroic. Transparent to subtransparent. Optically +. Birefringence high. Ax, pi. $\parallel 5$, Bx J. a. Dispersion $p < v$, feeble. $2H^{\wedge} = 103^\circ 53'$. $\mu = 1.722$.

Comp.— $Al_2O_3 \cdot H_2O$ or $Al_2O_3 \cdot H_2O =$ Alumina 85.0, water 15.0 = 100. Pyr., Btc. In the closed tube abundantly decrepitates strongly, separating into white pearly scales, and at a higher temperature yields water. Infusible; with cobalt solution gives a deep blue color. Not attacked by acids, but after ignition soluble in sulphuric

Diff.—Distinguished by its hardness and pearly luster; also by its decomposition and yield of water; by the reaction for alumina with cobalt solution. Resembles some varieties of hornblende, but is harder.

Obs.—Commonly found with corundum or emery. Occurs near Kossobrod, in the Ural; at Bchemuiz, Hungary; with corundum in dolomite at Campolungo, Tschudi, in

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Switzerland; Greiner in the Zugerthal. In the U. S., with corundum and niargarite at Newell. Also in Cu., Va.; in the emerald mines of Cbeuer, Mass.; in cavities in massive corundum in the Culsaic Mine. Dear Fruklid, Hacon Co., N. Carolina; wjui uliite

forming ruck masses at Ht. Kobinsod, Boeita UUs, Colorado.

Named by HaQy from Siatmfipeiv, lo««it(«r, alluding to the usual decrepitation before the blowpipe.

Orthorhombio. Axes $i:l:i = 0.9185 : 1 : 0.8$

In priams vertical; striated, and often flattened into scales or tables y i. Also fibrous; foliated or in scales; massive, reniform and stalactitic, with concentric and radiated structure.

Cleavage: b very perfect. Fracture uneven. Brittle. H. = 5-5-5. G. = $\sqrt{0.1}$ Luster imperfect adamantine. Color yellowish, reddish, and blackish brown. Often blood-red by transmitted light. Streak brownish yellow to ocher-yellow.

Tar.—In tbln scale-like or tabular crystals, usually attached by one edge. Also in adcular or capillary (not flexible) cryslala, or slender pnsmg, often railately grouped; (he NeedU-Ir<m*tom. It passes luto a variety with a velvety surface: the Prt^nnmite {fktmnut-bleiidd) of Pribram la of this kind. Also columnar, fibrous, etc., aa above.

Comp—FeO(OH) or Fe₂O₃.H₂O = Oiygen 37.0, iron 62.9, water 10.1 = 100, or Iron aeaqnoxiide 89.9, water 10.1 = 100.

Pyr., etc.^ In the closed tube gives off water and Is converted Inio red Iron sesqu-oxide. With the fluxes like hematite; most varieties give a manKsneae reactioD. and some, treated in the forceps in O.F., after moisteniug in nulphuric acid. Import a bluish-green color to tlie flame tphosphoric acid). Soluble In liydrocblnric acid.

Diff.—Distinguished from hematite by Its yellow atreak; from Ilmonlte by crystalline nature; It also contains less water than Ilmonlte.

Obs.—Found with the other oxides of iron, especially hematite or Ilmontla. Occurs at Eiserfelil near Siegen, in Nassau: at Clifton, near Brslol, England; la Cornwall. In ilie U. S., at the Jackson Iron mine, Negaunee, L. Superior; In Conn., at Salisbury; [u Pepn., near Eulon; inothe Pike's Penk region, Colorado. Named G^iie (Goethe) after the poet.philosopher Goethe (1749-1882).

MANOAMTTB,

Orthorhombic, Axes $/i:h:i = 0.8441 : 1 : 0.5448$.

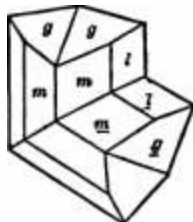
684. 086. hh" \ 410 A 4To = S8' SV.

mm", 110 A 110 = 80° aO*.

ee', 305 A 205 = 28' 57. «', Oil A Oh = B7" IC. Plf, 111 A ill = 59' BJ'.

Crystals commonly prismatic, the faces deeply striated vertically; often grouped in bundles. Twins; tw. pi. e (Oil). Also columnar; stalactitic.

Cleavage: b very perfect; m perit. = 4. G, = 4-2-4-4. Luster sub-metallic. Color dark steel-gray to iron-black. Streak reddish brown, some-



feet. Fracture uneven. Brittle.

times nearly black. Opaque; in minute particles becomes brown by transmitted light.

Comp.— $MnO(OH)$ or $Mn_2O_3 \cdot H_2O$ = Oxygen 27.3, manganese 62.4, water 10.3 = 100, or Manganese Sesquioxide 89.7, water 10.3 = 100.

Pyr., - - . . . _

Occurs in places; also in Cumberland, etc.

See also; Dyer's Head, Douglas,

Occurs in New Brunswick, at Brierley Mountain, Albert, etc.

LIMONITE. Brown Hematite. Brauneisenerz.

Not crystallized. Usually in stalactitic and botryoidal or mammillary forms, having a fibrous or subfibrous structure; also concretionary, massive;

and occasionally earthy.

H. = 5-5.5. G. = 3-6-40. Luster silky, often submetallic; sometimes dull and earthy. Color of surface of fracture various shades of brown, commonly dark, and sometimes bright; sometimes with a nearly black tarnish-like exterior; when earthy, brownish yellow, ochre-yellow. Streak yellowish brown. Opaque.

Uses.—(1) Compact. Sometimes siliceous to silky in texture; often stalactitic, botryoidal, etc. (Incl. Brauner Glaskopf Erz). (2) Earthy, brownish yellow to ochre-yellow, often impure from the presence of clay, sand, etc. (3) Bog ore. The ore from many places, generally loose or porous in texture, often petrifying leaves, wood, nuts, etc. (4) Brown clay-iron ore, in compact masses, often in concretionary nodules.

Comp $2\text{Fe}_2\text{O}_3 \cdot 3\text{H}_2\text{O}$ - Oxygen 25.7, iron 59.8, water 14.5 ^ 100, or Iron

Beauiquioxide 85.5, water 14.5 = 100. In the bog ores and ochers, sand, clay, phosphates, oxides of manganese, and humic or other acids of organic origin are very common impurities.

Pyr., etc.—LUTE color. Some varieties leave a alliceous skeleton In the salt of phosphoric acid, and a siliceous residue when dissolved in acids.

Diff.—Distinguished from hematite by its yellowish streak, inferior luster, and its reaction for water. Does not decrepitate B.B., like limonite. Not crystallized like goethite and yields more water. , , ...

ObB —In all cases a result of the alteration of other ores, or minerals containing iron, through exposure to moisture, air, and carbonic or organic acids; derived largely from the change of pyrite, magnetite, siderite, ferriferous dolomite, etc.; also various species (as mica, pyroxene, hornblende, etc.) which contain iron in the ferrous state (FeO). It consequently occupies, as a bog ore, various places, into which it has been home by streamlets from the hills around it. It is often associated with manganese ores. Limonite is a common ore in Bavaria, the Netherlands, Scotland, Sweden, etc.

Abundant in the United States. Extensive beds exist at Salisbury and Kent, Conn., also in the neighboring towns of New York, and in a similar situation in Berkshire Co., Mass., and in Vermont; in Pennsylvania widely distributed; also in Tennessee, Alabama, Ohio, etc.

Named Limonite from Xiphidion mendoza.

TRIOXIDE. Hydrohematite. $\text{Fe}_2\text{O}_3 \cdot \text{H}_2\text{O}$, or $2\text{Fe}_2\text{O}_3 \cdot \text{H}_2\text{O}$. Resembles limonite but has a red streak. G. = 4.14-4.8. Decrepitates B.B. From the Turginsk mine in the Ural, etc.; also from Salisbury, Conn. Intermediate between hematite and limonite.

Xanthopside. Gel bei sen stein. $\text{Fe}_2\text{O}_3 \cdot 2\text{H}_2\text{O}$. In fine needles or fibers, stellate and concentric; also as an ocher. Color golden yellowish, brown to brownish red. Associated with manganese ores at Umenau, etc.

BADSM. Beauxite.

In rounded concretionary disseminated grains. Also massive, oolitic; and earthy, clay-like. G. = 3-5.5. Color whitish, grayish, to ocher-yellow, brown, and red.

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Ymer.—1. In cordierite, talc, or olivine; bauxite. 1. Clay-like, like the purpur kind grayish clay-like. can be a little iron oxide; also red iron oxide present.

Comp.—EgaeBtiallj $\text{Al}_2\text{O}_3 \cdot 2\text{H}_2\text{O} = \text{Alumina } 73.9$, water $36.1 = 100$; some analyses, however, give $\text{Al}_2\text{O}_3 \cdot \text{H}_2\text{O}$ like diaapore.

Iron (esquioxide ie ueually present, sometlines in large amount, in part leplaciug alumina, iu pan only an Impurity. Silica, pbospboric acid, caibonic acid, lime, magnesia are common Impurities.

Oba.—From Bauil (or Beaux), near Aries, and elsewhere In France, disseminated Iu gratia Id compact limestone, and also oOUtlc. WoehMiiitt occurs In Carciola, between Felstrilz and Luke Wocbelu. The purest bauxite 1b uaed for the manufacture of alumlu-ium (aluminum), and is called aluminium ore. In the U. S., bauxite occurs in Saline and Pulaski Cos., Arkansas; also Id Cherokee and Calhoun Cos., Alabatua, and in Floyd, Barton aad Walker Cos., Georgia.

Brucite Oroup. $\text{B}(\text{OH})_3$. RhombohedraL BRUOTTB.

Rhombohedral. Axis $6 = 15208$; $\text{cr} = 60^\circ 20'$, $\text{rr}' = 97^\circ 37'$.

Crystals aaaally brood tabalar. Also commonly foliated massive; fibrous, fibers separable and elastic.

H. = as. G, — $2'38-2'4$, Cleavage: c eminent. Folia separable and flexible, nearly as in gypsum. Sectile. Luster | c pearly, elsevtaere waxy to vitreous. Color white, luclining to gray, blue, or green. Transparent to translucent. Optically+. Indices: a), = $1'059$, e, = 15795 .

Comp., yar.—Magnesium hydrate, $\text{Mg}(\text{OH})_2$, or $\text{MgO} \cdot \text{H}_2\text{O} = \text{Magnesia } 69.0$, water $31.0 = 100$. Iron and manganese protoxide are sometimes present.

Var. ^OnKnory, occurring in plales, white to pale greenish in color; slrong pearly luster on the cleavage surface. NemaUte is a Hbrous variety containliig 4 to 5 p. c. iron protoxide, with O. = $S'44$ Nattall. Manganbrvcite contains mauganese; occurs granular; color honey-yellow to brownish red.

Pyr., etc.—In Ihrf closed tube gives off water, hecoming opaque and friable, sometimes turning gray to brown: the mangaeuaian variety becotnes dark brown. B.B. infusible, glows wub a bright light, and the Igniteil mineral reacts alkaline to test-paper. With cobnlt solution gives the pale pink color of magnesia. The pure mineral is soluble in acids without enervesceuce.

Ditt—Distinguished by its infusibllity, Boflneas, cleavage, and foliated structure. Ia harder than talc und differ? In its solubility In acids; the magnesia test and optical cbaraclers separate it from gypsum, which U also somewhat softer,

Oba. —Accompanies other raagneslan minerals In serpentine, also found In limestone. At Swinauess in Unat, Shetliind Isles; at the iron mine of Cogne, Ansta, Italy: nearFilipsiadit

in Sweden. At Holioken. N. J , in serpentine; at the Tlrv Foster iron mine. Brewster, N, T., well crystallized; Richmond Co.. N. Y.; at Wood's mine, Texas, Pa., in large pieces or masses, and often crystallizations several inches across: at Low's mine with hydromagnesite. The fibrous variety, occurs at Hoboken. and at Xetles In the Vosges, Mn-juganbrucite occurs with hausmannite and other manganese minerals in the gneiss of Jakolwberg. Nordmark, Sweden.

Named after the early American mineralogist, A. Bruce (1777-1818). Pyrochroite. Manganese hydrate. $Mn(OH)_2$. Usually foliated, like brucite. Luster pearly. Color white, but growing dark on exposure. Occurs in magnetite at Pajsberg, Sweden; also at Nordmark; and at Franklin Furnace N J

Crystallography. Monoclinic.

Monoclinic. Axes $a : b : c = 1.7089 : 1 : 1.9184$; $\beta = 85^\circ 29'$, Crystals tabular on c , hexagonal in aspect. Occasionally in spheroidal concretions. Also stalactitic, or small mammillary, sometimes with smooth surface, and often a faint fibrous structure within.

Cleavage: none. Tough. H. = 2-5-3-5. G. = 2-3-2-4. Color white,

grayish, greenish, or reddish white. Luster of c pearly; of other faces vitreous; of surface of stalactites faint, translucent; sometimes transparent in crystals. A strong argillaceous odor when breathed on.

Comp.—Aluminium hydrate, $Al(OH)_3$, or $Al_2O_3 \cdot 3H_2O =$ Alumina 65.4, water 34.6 = 100.

Pyro. etc.—In the closed tube becomes white and yields water. Burns to form Al_2O_3 , and does not impart a green color to the flame. With cobalt solution gives a deep blue color. Soluble in concentrated sulphuric acid.

Obf.—The crystallized gibbsite (hydrargillite) occurs in the Bjihlmsk mountains near Zlatoust in the Ural; also in crystals filling cavities in strolite in the Langesundfjord. Norway: Ouro Preto, Minas Gerais, Brazil, in the U. S., in stalactitic form at Richmond, Hasbrouck, in a bed of limonite; at the Clove Mine, Union Vale, Dutchess Co., N. Y., on limonite; in Orange Co., N. Y.

Named after Col. George Qibba.

Saisoite*. Boric acid, $B(OH)_3$. Crystals tabular on c (triclinic). Usually small white, pearly scales. O. = 1.48. From the waters of the Tuscan lagoons of Hnue Rotondo and Castelnuovo. Exist also in other natural waters, as at Clear Lake, in Lake Co., California. Occurs also abundantly in the crater of Vulcano, Lipari Isles.

Hydroxide. Perhaps $Al(OH)_3 \cdot 8H_2O$. Lamellar-massive, or foliated, somewhat fibrous. H. = 3-4. D. = 2.3-2.4. Color white. Luster pearly. Occurs in the mines of

Shlshimsk, dlBtrlet oi Zlstouat, Ural; at Snarum, Norway, iu serpentine {hydrataleite).

Pyroaurite. Perhaps $\text{Fe}(\text{OH})_3\text{Mg}(\text{OH})_8\text{H}_2\text{O}$. Occurs at the Langbau iron-mine, Wermland, Sweden, in gold-like submetallic scales ($\text{Fe} > \text{Au}$). In thin seams of a silvery white color in serpentine In the island Haaf-Gruuay, Scotland (igelttromitt).

Ohalophanite. Hydrofranklinite. $(\text{Mn}, \text{Zn})_{0.3}\text{MnO}_2 \cdot 2\text{H}_2\text{O}$. In druses of minute tabular rhombohedral crystals; sometimes octahedral In aspect. Also in foliated aggregates; tabular and plumose. G. = 8-10. Luster metallic, brilliant. Color bluish black to iron-black. Streak chocolate-brown, dull. Occurs at Sterling Hill, near Ogdensburg, Sussex Co., N. J.

pyrrhotite.

Massive and botryoidal; reniform; tabular. H. = 6. Q. = 3-4-7. Luster submetallic, dull. Streak brownish black, shining. Color iron-black, passing into dark steel-gray. Opaque.

Comp.—A hydrous manganese manganate in which part of the manganese is often replaced by barium or potassium, perhaps conforming to H_2MnO_4 . The material is generally very impure, and the composition hence doubtful,

Pyrr., ato.—In the closed tube roasts varieties yield water, and all lose oxygen on ignition; with the fluxes reacts for manganese. Soluble in hydrochloric acid, with evolution of chlorine.

Obs.—A common but impure ore of manganese; frequently in alternating layers with pyrolusite. From Devonshire and Cornwall; Ilfeld in the Harz; also at Ilmenny, Siegen, etc. Forms mammillary masses at Brandon, etc., Vt. In Independence Co., and elsewhere in Arkansas. With pyrolusite at Douglas, Hants Co.. Nova Bocola. Named from $\text{Fe}_2\text{Si}_2\text{O}_7$ or $\text{Fe}_2\text{Si}_2\text{O}_7$, and $\text{Fe}_2\text{Si}_2\text{O}_7$, black.

The following mineral substance here included are mixtures of various oxides, chiefly of manganese (MnO , also H_2O), cobalt, copper, with also iron, and from 10 to 30 p. c, water. These are results of the decomposition of other ores—partly of oxides and sulphides, partly of manganese carbonates, and can hardly be regarded as representing distinct mineral species.

Wad. In amorphous and reniform masses, either earthy or compact; also encrusting or as stains. Usually very soft, sometimes the fluore; less often hard to H. = 5. O. s = 8-10-26: often loosely aggregated, and feeling very light to the hand. Color dull black, bluish or brownish black. ,

Boo Manquitsk consists mainly of oxide of manganese and water, with some oxide of iron, and often silica, alumina, baryta

AsBOUTh, or Earthy Cobalt, contains oxide of cobalt, which sometimes amounts to

"IS,

I. or Oapremt* Manganett, Is a wad Ciontalnlng 4 to 18 p. c. of oxide of copper, and often oxide of cobalt also.

VI. Oxyjreu-salts.

The Sixth Class includeB the salts of the various ozjgen acids. These fall into the foUowiDg seven sections: 1. Carbonates; 2, Silicates and Titanates; 3. Niobstee and Tantalatea; 4. Phosphates, Arsenates, etc.; also the Nitratesj 5. Borates and Urauates; 6. Sulphates, Ghromates and Tellnrates; 7. Tung-states and Molybdates.

1. OASBOHATES. A. Anhydrous Carbonates. The ADhydrons Carhonates include two distinct isomorphoua groups, the CALCITE Qkoup and the Aragonite Oboup. The metallic elements present in the former are calcium, magnesium, iron, manganese, zinc and cobalt; in the latter, they are calcium, barium, strontium and lead. The species included are as follows:

Monganocalcite pt. (MD,Ca)CO, Smithsonian ZnCO, 72° 20' 0-8063

Monheiraite {ZE,Fe)CO,

Spherocobaltite CoCO,

This list gives DOt only the promineat Bpecles of thla group, but also some of the isomornhous fntermedlate compounds.

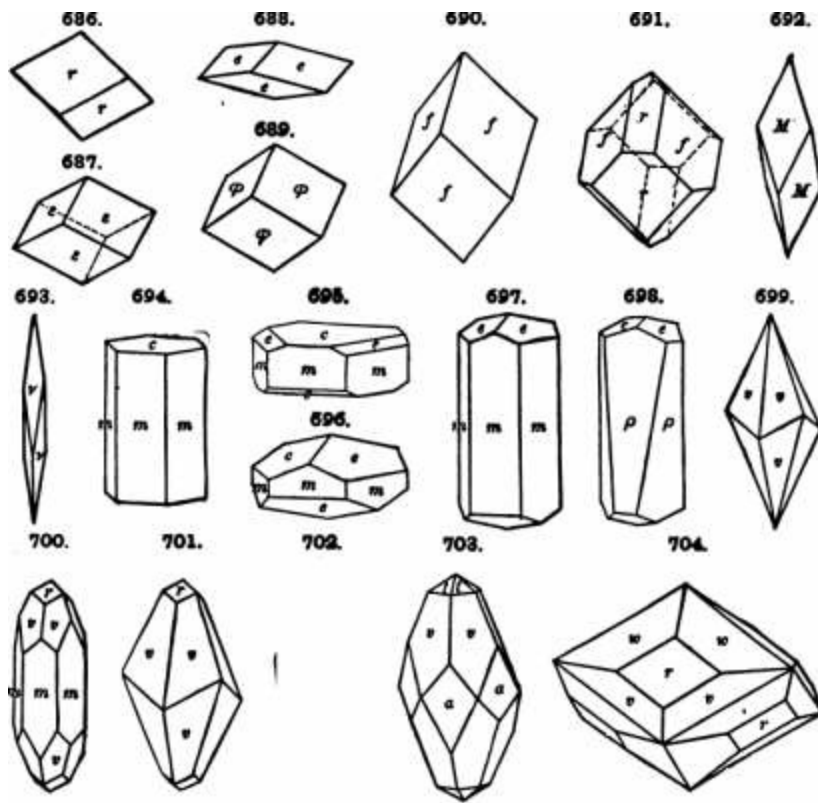
The Calcite Gboitp ie characterized by rhombohedral crystallization. All the species show, when distinctlr crystallized, perfect rhombohedral cleavage, the angle varying from 75° (and 105°) in calcite to 73° (and 107°) in siderite. This is exhibited in the table above.

DEBCRIPTtVE HINEKALOGY.

The species oi the Aragon:te Group crvstallize in the orthorhombtc avHU-m, but the relation to those of the Calcite Group is made more close by the fact that the prismatic angle varies a few degrees only from 60° (and 120^) and the twinned forms with the fundamental prism as twinning-plane are pseudo-hexagonal in character.

1. Calcite Oroap. BCO,. Bhombohedral. OALOITB. Ealkspftth Oerm. Calc Bpar; Calcsreotu Bpor. Rhombohedrah Axis i — 0-8543.

<8«. 6BB. CM).



• See the apherical projfcilon, Fig. 252, p. 70.

Habit of crystals Tery varied, as shown in the figures, from obtuse to acato
Thombohedral; from thin tabular to long prismatic; and scalctiobedral of many types,
often highly modified.

Twins (see Figs. 383-390, p. 126): (1) Tw. pi. c, common, the crystals having the same
vertical axis, (2) Tw. pi. e (oLl2), very common, the vertical axes inclined $137^{\circ} 29'$ and $52^{\circ} 30'$; often producing 706.

twinning lumellie as in Iceland Spar, vbicb are, in many cases of secondary origin as in
granular limestones (Fig. 705); this twinning may be produced artificially (see p. 148). (3)
Tw. pi. r, not common; the vertical axes inclined $90^{\circ} 46'$ and $89^{\circ} 14'$. (4) Tw. pi. / (o3Sl),
rare; the axes intersect at angles of $bZ'' 46'$ and $126'' 14'$.

Also fibrons, both coarse and fine; sometimes lamellar; often granular; from coarse to
impalpable, and compact to earthy. Also stalactitic, tuberosc, nodular, and other imitative
forms. •

Cleavage: r highly perfect. Farting \ e (oll2) doe to twinning. Fracture concboidaf,
obtained with difficulty. H. = 3, but varying with the direction on the cleavage face; earthy
kind^s softer. G. = 2'714, lu pure crystals, but varying somewhat widely in impure forms,
as in those containing iron, manganese, etc. Lnster vitreous to aubvitrooue to earthy.
Color white or colorless; also various pale shades of gray, red, green, blue, violet, yellow;
also brown and black when impure. Streak white or grayish. Transparent to opaque.

Optically—. Birefringence very high. Befractive indices for the D line: a = 1-65849, c =

148625.

Comp. —Calcium carbonate, CaCO_3 , = Carbon dioxide 44%, lime 56% = 100. Small quantities of magnesium, iron, manganese, zinc, and lead may be present replacing the calcium.

Varieties.—The varieties are very numerous, and diverse in appearance. They depend mainly on the following points: crystallization and structural conditions, presence of impurities, etc., the extreme being perfect crystals and earthy massive forms; also on composition as affected by isomorphous replacement.

A. Varieties based chiefly upon crystallization: Accidental impurities.

1. Ordinary. In crystals and cleavable masses, the crystals varying very widely in habit as already noted. Dog-head Spar is a scute-shaped form; Nail-head Spar, an acorn.

Usually having the form of a bug by the name. The transparent variety from Iceland, used for polarizing prisms, etc., is called Iceland Spar or Doumyrtz Spar (Iceland Spar). As regards color, crystallized calcite varies from the kinds which are perfectly clear and colorless through yellow, pink, purple, blue, to brown and black. The color is usually pale except as caused by impurities. These impurities may be pyrite, native copper, malachite, sand, etc.: they are sometimes arranged in symmetrical form, as depending upon the growth of the crystals and hence produce many varieties.

Fonlaine variety from Pontainebleau and Nemours, France, contains a large amount of sand, some 60 to 68 per cent. Similar forms occur at other localities,

3. Fibrous and lamellar kind. Satin Spar is fine fibrous, with a silky luster; resembles fibrous gypsum, also called satin spar, but is much harder than gypsum and effervesces with acids.

Argentine is a pearly lamellar calcite. the lamellae more or less undulating; color white,

pearly, yellowish, in its harder and more sparry variety (Schavmerde). is a pearly white calcite. near argentine; like the softer kinds (Schavmerde) it is like chalk, but much harder, pearly in luster, silvery white or yellowish in color, soft and greasy to the touch, and more or less scaly in structure.

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S. Oranular varieties to erythraean kind: Limalone, Marble, Chalk.

Oranular lime or Saeharoidal limestone, so named because like loaf sugar in fracture, varies from coarse to very fine grained, and Leace to compact limestone; color varies, as white, yellow, reddish, green; usually they are cloudy and give a

handsome effect when the material is polished. When such limestones are fit for sculpture, or for architectural or ornamental use, they are called marble. Many have special names. Safford-marble contains largely of fossil shells; Lumachelle or Jire-matt is a dark brown shell-marble, with brilliant green or chatoyant interreflections. Roman-marble is a kind of a yellow to brown color, showing, when polished, figures bearing some resemblance to the Parthenon, temples, etc. in ruins, due to infiltration of iron oxide, etc.

Lithographic stone is a very eye-catching compact limestone of buff or drab color; as that of Solenhofen. Hydraulic tuff is an impure limestone which after igneous sets. I.e., takes a solid form under water, due to the formation of a silicate. The French varieties contain 3 or 8 p. c. of magnesia, and 10 to 20 of silica and alumina (or clay). The varieties in the United States contain 50 to 40 p. c. of magnesia, and 13 to 80 p. c. of silica and alumina. The color varies from nearly pure white, through gray, drab, buff, yellowish, and reddish shades, to bluish gray, dark brown, gray, and black, and sometimes grossly veined. Many kinds make beautiful marble when polished. Red oxide of iron produces red of different shades. Shades of green are due to iron protoxide, chromium oxide, iron silicate. •

Chalk is white, grayish white, or yellowish, and soft enough to leave a trace on a board. Oolite is a soft earthy deposit, with or without distinct fragments of shells; it generally contains much clay, and graduates into a calcareous clay.

Oolite is a granular limestone. Its grains are minute concretions, looking somewhat like the roe of fish. The name comes from the Greek *oolite* (Erbsestein) consists of concentric shells as large often as a small pea, or larger, having usually a distinct concentric structure.

Stalactites from caverns are called «ou» «spring», «am», or in «fm». etc. (a) Stalactites are calcareous cylinders or cones that hang from the roofs of limestone caverns, and which are formed from the waters that drip through the roof: these waters hold some calcium bicarbonate in solution, and leave calcium carbonate to form the stalactite when evaporation takes place. Stalactites vary from transparent to nearly opaque; from a crystalline structure with single cleavage directions to coarse or fine granular cleavable and to radiating fibrous: from a white color and colorless to yellowish gray and brown. (b) Stalagmite is the same material covering the floors of caverns. It being made from the waters that drop from the roofs, or from sources over the bottom or sides; cones of it sometimes rise from the floor to meet the stalactites above. It consists of layers, irregularly curved, or bent. Stalagmite, or a solid kind of travertine (see below) when on a large scale, is the alabaster stone of ancient writers, that is, the stone of which ointment vases, of a certain form called alabastrine, were made. A locality near Tiberias, now well known, was largely explored by the ancients, and the material has often been hence called Egyptian alabaster. It was also formerly called onyx because of its beautiful banded structure. In the arts it is often now called Oriental alabaster or onyx marble. Very beautiful marble of this kind is obtained in Algeria, the main onyx is a similar material

obtained from Tecalt, Pui-bla, Uexico; also in a beautiful brecciated form from the extinct crater of Zempoaltepec in northern Mexico. Similar kinds occur in Missouri, Arizona, San Luis Obispo Co., California, (e) Catehite, Travertine. Calcite is of essentially the same origin with stalagmite, but is distinctively a deposit from springs or rivers, especially where in large deposits, as along the river Anio. at Tivoli. near Rome, where the deposit reaches scores of feet in thickness. (d) Agaric mineral; Rock-milk is a very soft white material, breaking easily in the fingers, deposited sometimes in caverns, or about sources holding lime in solution, (a) Jack-milk is white and light, like cotton, becoming a powder on the slightest pressure.

B. Varieties based on Composition.

These include; Dolomite earth. Contains magnesium carbonate, thus grading toward true dolomite. Also barite (which contains some BaCO₃); similarly, strontianite (SrCO₃), ferrocalcite (FeCO₃), magnesianite (MgCO₃), zinc carbonate (ZnCO₃), plumbite (PbCO₃).

Pyr., etc.—Barite infusible, but becomes caustic, glows, and colors the flame reddish yellow; after ignition the assay reacts alkaline; moistened with hydrochloric acid imparts the characteristic lime color to the flame. In the solid mass effervesces when moistened with hydrochloric acid, and fragments dissolve with brisk effervescence even in cold water.

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Barite—Distinctive characters: perfect rhombohedral cleavage; softness, can be scratched with a knife after treatment with cold dilute acid; infusibility. Less hard and of lower specific gravity than calcite (w. w.) Resembles in its optical properties the oilier rhombohedral carbonates, but is less hard, of lower specific gravity, and more readily attacked by acid. Also resembles some varieties of barite, but has lower specific gravity; it is less hard than barite and harder than Epsom.

Barite—In thin sections by the low refractive index and very high birefringence, the polarized colors in the thinnest sections are white while of the higher order. The negative interference figure, with many closely crowded colored rings. It also characteristically. The rhombohedral cleavage is often slightly irregular; a system of twinned lamellae often conspicuous (Fig. 70S), especially in crystals of the lineation.

Barite.— Calcite, in its various forms, is one of the most widely distributed of minerals. Beds of sedimentary origin formed from organic remains, shells, corals, etc., yield on metamorphism crystalline limestone or marble, and in countries where these crystallized calcite and also deposits in caves of stalactites and stalagmites often occur. Common with the zeolitic and veins of igneous rocks as a result of alteration, and similarly though less abundant with granite, syenite, etc. A frequent mineral in

metalliferous depoafls. willb lead, copper, silver, etc. Deposited from lime-bearing waters aa calc sinter, travertine, etc., especially in connctioD wiih but springs as at ibe Mammoth Hot Springs iu the TellowsCone region.

Some of the best known localities for crystallized calcite are the following; Andreaa-bergiu the Uarz: tlie mines of Freiberg, Schiieeberg. etc., in Baxouy; Kapuikin Hungary: Ansig in Bohemia: Blalberg in Cariitliin; Traverselln In Piedmont; Elba. In England at Alalon Moor aud E-remonl In Cumberland; Matlock. Derbyshie; Beer Alatoa Id Devonshire; at numerous points in Cornwall; Weardale in Diirbum; Stank mine, Lancashire, In twin crystals of great variety anil beauty at Guanajuato, Mexico. The leelartd tpar baa been obtained from Iceland near Helgustailir on the Eakeflord. It occurs in a large c»viyy in basalt. The crystals, usually sbowiug tlie fundamental rbombohedroo, are often couled Willi tufta ot alitbite.

In the U. States, crysialliztd calcite occurs in N. Tor/c. in St. Ijiwrence Co., especially at the Rossie leiid mine; in JeSerson Co., near Oxbow: dog-looChgptir, in Nlagai-a Co, near Lockpori, wiih pearl spiir, celestite, etc: In Lewis Co.. at Leyden and Lowvllie. and at the Hartinsbiirg lead mine; nt Anthony's Nose on the Hudson, formerly groups of large tabubir crysiala. In N. Jeriey. at Bergen, yellow calcite with daloliie. '■Ic. In yirginia, at Wler's cave, ttalaetitft of great beauty: also in the large caves of Keniuek;/. At the lake Superior copper mines, complex crystals oflen conifiining scalers of native copper. At Warsaw, IUinoit, in great variety of form, lining geodes and implnted on quartz crysliils: at Quincy. In Mitsouri, with dolomite near 8l. Louis; also with aplinlerile at Jopplit and other points in the zinc region in the soubwenlern part of the slate, the crystals lisiiall; scalenohedral and of a wine.yellow color. From the Bad Lands. South Dakota, In Nova Scotia, at Partridge I., a wine-colored calciic, and other Interesting varieties.

TuiHoLiTB. A tufa deposit of catcluin carbonate occurring on an enormous scale in north-western Nevada; also occurs about Mono Lake, California. It forms layers of Interlaced crystals of a pale yellow or tiglit brown color and often skeleton structure except wheu covered by subsequent deposit of calcium carbonate.

DOLOMITEI. Pearl Spar pt. Tri-rhombohedral. Ais t! = 0 8322. er. 0001 A lOil = 43° 52". Jflf', 4041 a 4401 = 118° Sr. 'o*"

XT', lOil Ailo1 = 73"46'.

Habit rhombohedral, usually r or if{4oll): the preseice of rhombohedroia of the second or third seriea after the phenacite type very characteristic. The r faces commonly curved or made tip of sub-individ-uala, attd thus passing into eaddle-shaped forms (Fig. 706). Also granular, coarae or fine, resembling ordinary marble.

Cleavage: r perfect. Fracture subconchoidal. Brittle. H, = 3-5-4.

, , . . CA > o g l c

Q. = 2'8-2'9. Luster Titreos, inclining to pearly in some varieties. Color white, reddish, or greenish white; also rose-red, green, brown, gray, and black. Transparent to translucent. Optically —. Wj = 1-68J74 Na, e, = 1*5oi56 Na, FizeaD.

Comp.—Carbonate of calcium and magnesium (Ca,Mg)CO₃; for normal dolomite CaMgC₂O₇, or CaCO₃·MgCO₃. = Carbon dioxide 47-9, lime 30-4, magnesia 31'7 = 100, or Calcium carbonate 44'35, magnesium carbonate 45'65 = 100. . Varieties occur in which the ratio of the two carbonates varies from 1 : 1. The carbonate of iron and manganese also sometimes occur; rarely cobalt and zinc carbonates.

Pyr., etc.—B.B. acts like calcite. Fragments thrown into cold acid, unlike calcite, are only very slowly acted upon, if at all, while in powder in warm acid they effervesce readily dissolved with effervescence. The ferrous dolomites become brown on exposure.

Dtff.—Resembles calcite (see p. 507), but generally to be distinguished from it that it does not effervesce readily in the moist to cold acid.

Obi.—Hastive dolomite constitutes extensive strata, crystalline limestone strata, in various regions, both in the dolomite region of the southern Tyrol, Crystalline limestone nodules in various varieties are often associated with serpentine and other igneous rocks, and with ordinary limestones. Some prominent localities are: Leogang in Salzburg; Schmuil and Kipnik in Hungary; Freiberg in Saxony. In Switzerland, at Bex, it occurs in crystals; also in the Balmuccia; Traversella in Piedmont; Campolongo,

In the U. States, in Vermont, at Roxbury. In N. Jersey, at Hoboken. In N. York, at Lockport, Niagara Falls, etc.: at the Illy Foster iron mine. Brewster, Putnam Co., with magnetite, chondrodite. In Badkissingen crystals with the sphalerite of Joplin, Missouri. In S. Car., at Stony Point. Alexander Co,

Named after Dolomieu (1760-1801) who announced some of the marked characteristics of the rock. It does not effervesce with acids, while burning like limestone, and solubility after heating in acids.

Ankerite. CaCO₃·[Mg,Fe,Mn]CO₃, or for normal ankerite 8CaCO₃·MgCO₃·FeCO₃. In rhombohedral crystals; n_D = 1.548; also crystalline massive, granular, compact. Q. = 3'5-S-I. Color white, gray, reddish. Occurs in small crystals at the Styrian mine.

MAANUSTTB,

Rhombohedral Axis: n_D = 1.5112. n_D' = 1.536'. Crystals rare, usually rhombohedral, also prismatic. Commonly massive; granular cleavable to very compact; earthy.

Cleavage: r perfect. Fracture uneven or subcochoidal. Brittle. H. = 3-5-4. G. = 3-83-3-88. Luster vitreous, inclining to pearly. Color ash-gray, yellowish gray, greenish gray, also brown and brownish red, rarely green; and sometimes white. Streak white. Translucent to subtranslucent. Optically —. Double refraction strong.

Comp.—Iron protocarbonate, FeCO_3 , = Carbon dioxide 37.9, iron protoxide 62.1 = 100 (Fe = 48.3 p. c). Manganese may be present (as in oligonite), also magnesium and calcium.

Pyr. et c.—In the closed tube decrepitates, gives off CO_2 , blackens and becomes magnetic. B.B. blackens and fuses at 4-B. With the fluxes reacts for iron, and with soda anhydride on platinum foil generally gives a magnetic residue. Duly slowly acted upon by cold acid, but dissolves with brisk effervescence in hot hydrochloric acid. Exposure to the atmosphere darkens its color, rendering it often of a blackish-brown or brownish color.

Diff.—Characterized by rhombohedral form and cleavage. Specific gravity higher than that of calcite, dolomite and siderite. Resoluble in some species but lacks the resinous luster, differs in cleavage angles and yields CO_2 (with H₂S) with hydrochloric acid.

Obs.—Occurs in many of the rock strata. In gneiss, mica slate, clay slate, and as clay ironstone in connection with the Coal formation and many other stratified deposits. It is often associated with metallic ores. At Freiberg it occurs in silver mines. In Cornwall it accompanies tin. It is also found accompanying copper and iron pyrites, galena, chalcocite, tetrahedrite. Occasionally it is to be met with in trap rocks as siderite in globular concretions. Existing deposits occur in the Eastern Alps, in Styria and Carinthia. At Hartz in the Harz, it occurs in fine crystals in gray-wacke; also in Cornwall of varied habit at many localities; at Alston-Moor, and Tavistock, Devonshire. Fine cleavage masses occur with cryolite in Greenland.

In the United States, in Vermont, at Plymouth. In Mass., at Sterling. In Conn., at Roxbury. An extensive vein in quartz, traversing gneiss. In N. York, a series of deposits occur in Columbia Co.; at the Lonsdale iron mines. St. Lawrence Co. In Ill. Carrott Mt. at Feitres and Harlem mines. The argillaceous carbonate. In nodules and beds (clay ironstone), is abundant in the coal regions of Penn., Ohio, and many parts of the country. In a clay-bed under the Tertiary along the west side of Chesapeake Bay for 50 m.

RHODOCHROSITE. Rhodochrosite. Rhodochrosite, Rhodochrosite, Rhodochrosite. Rhombohedral. Axis $a = 0.8184$, $\alpha = 73^\circ 0'$. Distinct crystals not common; usually the rhombohedron r ; also e , with rounded striated faces.

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Colorable, massive to granular massive and compact. Also globular and botryoidal, with columnar structure, sometimes indistinct; incrusting.

Cleavage: r perfect. Fracture uneven. Brittle. H. = 3-5-4-5, G. = 3-45—3-60 and higher.

Luster vitreous, inclining to pearly. Color shades of rose-red; yellowish gray, fawn-colored, dark red, brown. Streak white. Translucent to subtranslucent. Optically —.

Comp.—Manganese protocarbonate, $MnCO_3$, = Carbon dioxide 38.3, manganese protoxide 61.7 = 100. Iron carbonate is usually present even up to 40 p. c, as in manganoederite; sometimes the carbonate of calcium, as in mangano-calcite, also magnesium, zinc, and rarely cobalt.

Pyr., etc.—B.B. fuses to gray, brown, and black, and decrepitates strongly, but is infusible. With salt of phosphorus and borax in O.F. gives an amethystine-colored bead. In R.F. becomes colorless. With sodium chloride foil becomes greenish. Dissolves with effervescence in warm hydrochloric acid. On exposure to the air changes to brown, and some bright rose-red varieties become paler.

Id.—Characterized by its pink color, rhombohedral form and cleavage, effervesces in acids.

Occur* commonly alone with ores of iron, lead and copper, and with other ores of manganese. Found at Schemnitz and Kapnik in Hungary; in Tyrol; at Freiberg in Saxony; at Diez near Oberneisen in Nassau; at Kufstein, in the Tyrol; at Uffontaine in the Ardennes, Belgium. In the U. S. at Brantville, Conn.: in New Jersey, with iron at Mine Hill, Franklin Furnace. In Colorado, at the John Reed mine, Ancaster, Lake Co., in beautiful clear rhombohedrons; also at the Oulay mine, near Lake City. In Montana, at Butte City. Abundant at the Silver mines of Austin, Nevada, at Placentia Bay, Newfoundland.

Named rhodochrosite from rhodon, a root, and chros, color; and dialogic, from dialogos, doubt.

BMITHBONITE. Calamine pt. Zinc carbonate. Dry-bone Minor: Rhombohedral. Axes $a = 0.8063$, $b = 0.7202$. Rarely well crystallized; faces generally curved and rough. Usually reniform, botryoidal, or atalao-titic, and in crystalline incrustations; also granular, and sometimes impalpable, occasionally earthy and friable.

Cleavage: a perfect. Fracture uneven to imperfectly conchoidal. Brittle. $H. = 5$. $G. = 4.30-4.45$. Luster vitreous, inclining to pearly. Streak white. Color white, often grayish, greenish, brownish white, sometimes green, blue and brown. Subtransparent to translucent. Optically —.

Comp.—Zinc carbonate, $ZnCO_3$ = Carbon dioxide 35.2, zinc protoxide 64.8 = 100. Iron carbonate is often present (as in mottheimite); also manganese and cobalt carbonates; further calcium and magnesium carbonates in traces; rarely cadmium and indium.

Pyr., etc.—In the closed tube loses carbon dioxide, and, if pure, is yellow while hot and white on cooling. B.B. Infusible; moistened with cobalt solution and heated in O.F. gives a

(freen color on cooling. With soda on charcoal coats the coal with the oxide, which is ydlow while hot and white on cooling; tbia coating, moiatenei} with cobalt solution, gives a green color after heating In O.F, Soluble in hydrochloric acid with effervescence.

DIS.—Disllngulshed from calamine, which It often closely resembles by its efiervescence in acids.

Obs,—Found both in veins and beds, especially in compaay with galena and sphalerite; also with copper and Iron ores. It usually occurs in calcareous rocks, and is generally associated with calamine, and somellmes with llmonlte. It Is often produced by the action upon line sulphide nf carbouated walers.

Found at Nerchinsk in Siberia; at Dogoaczka in Hungniy; Blelbei^ and Baihel la Carimhin; Wicaloch In Baden; Moreenet m Belgium; Altenberg, In the province of Bnntander, Spain, at Pnente Viesgo. In England, at Roughleu QUI, Alston Moor, near Matlock, in the Mendip Hills, and elsewhere; in Ireland, at Donegal. At Lauriou, Greece, varieties of many colors.

,ab,GoOgIc

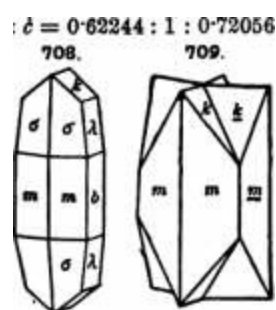
In the v. Sutea. in Ann., at Lancuter ftbuDdaut, the rar. called "dry-bone"; at the Veberrolh mine, iiear Bethlehem. In Kslenohedron*. In WtMeonntt, at Mineral Point, BhullaburK, etc., pteudomorphe after ipbalerite and calcite. In Mirmetola, at Swing's diggluge. N. W. of Dubuque, etc. In south-weeteni Mittcniri aaaociated with Bphalerlte aud caUmiue. In Arkantat, tX Calamine, Lawrence Co.; fn Marlon Co. A pink coballlf-crou« vartetj occurs at Boleo, Lower Calltoraia.

Named after James SmlihBon (17S4-182B), wlio founded the Bm<tb«onlau Instllutlon In Washington. The name calamine la frequently used in England, cf. calamine, p. 446.

Spharooobaltita, Cobalt protocarbonate, $CoCO_3$. Eobaltspatb Qtrm, Rhombo-hedral. In email apherlca! masses, with crystalline surface, rarely In ctTBIKLi. G. = 4'0a-4'IS. Color roM-red. From Schneeberg, Saxony.

2. AragoDite Group. RCO., Orthorhombio. For list of epecies, see p. SSa ARAOOHITB.

Orthorhombic. Axes &l:i = o'6: mm''', no Alio = 68* 48'.



Crystals of tea acicular, and characterized by the presence of acute domes or pyramids. Twins: tw. pi. m commonly repeated, producing pseudo-hexagonal forms (see p. 127, Figs. 398, 399 and 710). Also globular, reniform, and coralloidal shapes; sometimes columnar, straight or divergent; also atalactic; incrusting.

Cleavage: b distinct; also m; k (Oil) imperfect. Fracture subconchoidal. Brittle. H. = 3-5-4, G. — 2.93-2.95. Luster vitreous, inclining to resinous on surfaces of fracture. Color white; also gray, yellow, green and violet; streak uncolored. Transparent to translucent. Optically —. Ax. pi. || a. Bi X c. Dispersion $p > v$ small. $2E^{\wedge} = 30^{\circ} 64'$. $S^{\wedge} = 1-6816$.

Comp.—Calcium carbonate, $CaCO_3$, = Carbon dioxide 44.0, lime 56.0 = 100. Some varieties contain a little strontium, others lead, and rarely zinc.

.. „ „ .. also 8d*

fibrous with silky luster, {e} Ma^{ve}.

Sphaerulitic or stalactitic: Either compact or fibrous in structure, as with calcite; Bpni-fUtein is stalactitic from Carlsbad, coralloidal: In groups of delicate interlocking and coalescing sphaerulites, of a snow-white color, and looking a little like coral: often called Florenstein (ElaenblDthe Oerm.). Tamowitz kind containing lead carbonate (4 to 8 p. c), from Tarnowitz in Silesia; with O. = a'88.

Pyrrhotite, etc.—B.B. whitens and falls to pieces, and sometimes, when containing strontia, imparts a more intensely red color to the flame than lime; otherwise reacts like calcite.

Distinguished from calcite by higher specific gravity and absence of rhombohedral cleavage; from the zeolites {e.g., natrolite}, etc. by effervescence in acid. Strontianite and witherite are fusible, higher in specific gravity and yield distinctive flames B.B. The resinous luster on fracture surfaces is to be noted.

Oba.—The most common repositories of anhydrite are beds of gypsum; also beds of iron ore, as the Styrian mines, where it occurs in coralloidal forms, and is denominated pyrrhotite, "Jura of iron"; in basalt: occasionally it occurs in lavas: often associated with copper and iron pyrites, galena, and malachite. It constitutes the pearly li

T of shell.

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Found in Angou, Spain (where its name), at Hollna and Valmala. In six* aided prUmB, with ijynum. Pronounced locally are Blinn, Bobemla; Leogacg Id Salzburg, AuBtriu; HtrrCLgruud, Hudgwyi with aulpUurfaBicUr in flue prUmsi also al Aklou Hoor, flue lapering crjBlaU.

lu flbrouB cruBii at Hoboken. N. J.; at BdeDtine aud BoBsle. K. T.; Wood'B Hlne, LaDcaBt«r IJ», Peuu.; Wanaw, 111., lining geodei; Hine-la-Hotte, Ho., in crystals. Flat- ftrri lu tie Organ Hts., New Mexico.

WITHBRITB.

Orthorhombic Axei A: I-.6 = 0-6032 : 1: 07303. Crystals always repeated

711. twins, aimulating hexagonal pyramids. Fig. 711 (cf.

a h Fig, 557, p. 227). Also masaive, oolamnar or granalar.

Cleavage : i distinct; m imperfect. Fracture

oneven. Brittle. H. = 3-3-75. G. = 4-2r-4'35. Luster

J vitreous, incliniDg to resinous od surfaces of fractnre.

Color white, yellowish, grayish. Streak white. Snb-

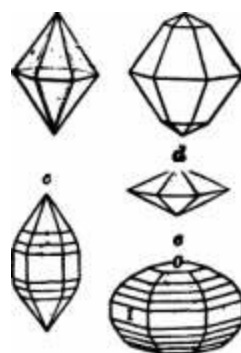
transparent to translucent.

Comp.—Barium carbonate, BaCO. = Carbon dioxide 22 3, baryta 777 = 100.

Pyr., eto.—B.B. fusei at 3 10 a bead, coloriu g the flame yellowUh green; nfier fuBion reacts alkaline. B.B. on cliarcoal wtlh soda fuses easily, and U absorbed by tbe coal. S<iluble in dilute hydrocBloHc acid; this solution, even when very much diluted, gives with sulphuric acid a white precipitate wblcb Is iDBoluble In acida. I Diff—Disllnguished by its bigb specific gravity; etfer-I vescence in acid; ysr^sa colonulon of tbe flame B.B. Baiite Is Insoluble in hydrocbloric add.

Oba.—Occurs at Alston Moor in Cumberland, with galena; at lUlowfield near Heibam In NorCbumberlaud; Tamowllz In Silesia- Leogaiig iu Sals-burg; near Lexington, Eentuctcy, with berite. In a silver-bearing vein near Babbit Ut., Thunder Bay, L. Superior.

Bromlits. (Ba.Ca)CO,. In pseudo-hexagonal pynmlds (Figs. SS8, o60, p. 227). Bromley Hill, near Alston, Cwnberland.



STRONTIAHmi.

Orthorhombic Axes $a:b:c = 0.6090 : 1 : 0.7239$.

Crystals often acicular or acute spear-shaped, like aragonite. Twine: tw. pL VI common. Also columnar, fibrous and granular.

Cleavage: tn nearly perfect; & in traces. Fracture nneven. Brittle. H. = 35-4. G. = 3'680-3714. Luster vitreous; inclining to resinous on faces of fracture. Color pale asparagDS-green, apple-green; also white, gray, yellow, and yellowish brown. Streak white. Transparent to translucent. Optically —. Ax.-pi. I b. Bx_Lc. Dispersion $p < v$ small. $2E, = 12'' 17'$.

Comp.—Strontium carbonate, $SrCO_3$, = Carbon dioxide Sflil, strontia 70'1 ■ = 100. A little calcium is sometimes present. j

Pyr,, et«.—B.B swells up, throws out minute Bprouts, fuses only on the thin edges, and : colni-B tbe flume Blrontls-red; tbe assay reacts Hlkalloe after ignition. Hoistened wiih ' hydioflhloric acid and treated either B.B. or in tbe naked lamp gTves an intense red color. ■

Soluble In hydrochloric add; tbe dilute solution when treated with sulphuric acid girea a ' while precipitate.

Dlff,—Differs from related mlnemls, not carbonates, in effervescing with adds; has a , higher specifc gravily than aragonite and lower thnn witherite: colors the Uame red B.B. .

Obs.—Occurs at Btronlian in Argyllshire; in Tarkabire, England; Clausthal In tbe Harz; BrKunadorf, near Freiberg, Suony; Leogang in Salzburg; near Brixtegg, Tyrol

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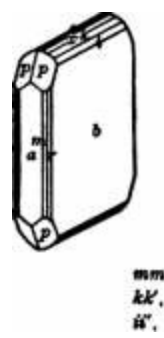
OARBONATES.

>t the WilbelmiDs n

^ea^eioltrol^lianitt)■. to flue crjslala near HuDm, Westpballs: I AlUhleo, Wespliallft.

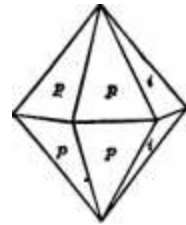
lu the U. Slates, occurs at Scbobarie, N. Y.; al MuacalODSe Lake; Cbaumont Baf and Tberau, In JaSenod Co., N. T.. Hifflin Co., Penik.

OBRnBSrm. White Lead Ore. Welasblelere Qerm. Orthorhombic. Axes $a:b:c = 0.60997 : 1 : 0.72300$. 71 a, 713. n*.



ffrl

ft



Simple crystals often tabular { b, prismatic Q d; also pyramidal. Twins: tw. pi, m, very common, contact- and penetration-twins, often repeated yielding six-rayed stellate groups. Crystals grouped in clusters, and aggregates. Rarely fibrous, often granular massive and compact; earthy. Sometimes stalactitic.

Cleavage: m and i (021) distinct; b and x (012) in traces. Fracture conchoidal. Very brittle. H. = 5-3-5. G. = 6-46-6-574. Luster adamantine, inclining to vitreous, resinous, or pearly; sometimes submetallic. Color white, K⁺ grayish black, sometimes tinged blue or green (copper); streak uncolored. Transparent to subtranslucent. Optically—. Ax. pi. Jo. BxJ.a Dispersion p > V large. Indices and axial angles, Schrauf:

a / 3 7 / 2V 2E

Line D 1-80368 207638 207803 .-. 8° 14' 17" 8'

Coup.—Lead carbonate, PbCO₃ = Carbon dioxide 16-5, lead oxide 83-5 = 100.

Pyrr. etc —In the closed tube decrepitate, lose CO₂, turn bright yellow, and at a higher temperature dark red, but becomes again yellow on cooling. B.B. on charcoal fuses very easily. In R.F. yields metallic lead. Soluble in dilute nitric acid with effervescence.

SiS —Characterized by high specific gravity and adamantine luster; also by yellow lead B.B. Unlike anglesite. It effervesces with nitric acid.

01».—Occurs in connection with other lead minerals, and is formed from galena, which, as it passes to a silicate, may be changed to carbonate by means of solutions of calcium carbonate. It is found at Johanngeorgstadt in beautiful crystals; Motite Pond.

Sardidlai Friedriclisseg. Nassau: Badenweiler, Bhdn; at Clniisthiil in the Harz; si Bleiberg Id Carinthia; in Englaol, in Cornwall; nt E Tamar mine. Devonshire: nenr Mullock and Wirksworth, Derbyshire: at Leadhlls and Wanlocklicad. Scotland.

Found in Penn.. at Pbeniiville. In Virginia, at Austin's mines. Wythe Co. In If, Carolina, In King's mine. In Wisconsin and other lead mines of the Doribwesteru Slates. rarely la crystals; at Hazelgrern, crystals coatiDg galena. In Colorado, at Leadville, and elsewhere. In Ai-i*ona, at the Flux mine, Pima Co., in large crystallue masses; In crystals at the Red Cloud mine, Tuma Co.

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BARTTOOALOITB.

Monoclinic Axes a:h:d = 0-7717 : 1 : 0-6354; >» = 73" 52'. In oryBtals; also massive.

Olsavagfi: m perfect; clesa so. Fracture uoeven to subconchoidal. Brittle. H. = 4. Q. = 3'64-3'66. Luster vitreons, inclining to resinone. Color white, grayish, greenish or jellowish. Streak white. Transparent to translncent.

Comp.—Carbonate of barium and calcinm, BaCOi.CaCOt = Carbon dioxide 29-6, barjta 51-5, lime 18-9 = 100.

PjTT., etc.—B.B. culora the flame fellowish gieeu, and at a high temperature fuaes on the ThiD edges and assumea a pale green color; the aeaay reacts alkaline after IguilioD. With the Huzea reactB for manganeBc. With goda ou charcoal the lime is separate iia uu infusible mass, while the remainder Is absorbed by the coal. Soluble in dilute bjidri)-chloric acid.

Oba.—Occurs at Alston Moor iu Cumberlaud, in llmeBlotie with baile and fluoille.

BUmntosphlLrlte. Bi.(COi)i.2Bi,oi. la spherical forms with radiated struclure. O. =7'43. Color yellow to gray or blackish brown. From Schueeberg, Saxony. Also sparingly at Willimantlc and Portland, Conn., as a result of the alteratloa of bismutblnile.

Pailsite. A fluocarbonate of the cerium metals. In acute double heiagonal pyramids. H. = 4*5. Q. = 4'868. Color brownisb yellow. From the emerald mines of the Huso valley. U. B. Colombia. .

BaitnilHte, Hamartlte. A fluocarbonate of the cerium melals (RF)COt. Color wax-yellow to reddish brown. From the BastnlU mine, Riddarbyttan, Sweden. Also as aa alteration product of tysonlte io the granite of the Pike's Peak region in Colorado.

^OSaBHTTB.

Tetragonal. Axis $d = 1.0876$. Crystals prismatic; sometimes tabular $\setminus c$.

Cleavage: m, a distinct; also c . Rather sectile. $H. = 2.75-3$. $G. = 6.0-6.3$. Luster adamantine. Color white, gray, and yellow. Streak white. Transparent to translucent. Optically +.

Comp.—Chlorocarbonate of lead, $(PbCl)_2CO_3$, or $PbCO_3 \cdot PbCl_2$. $PbCl_2 =$ Lead carbonate 49.0, lead chloride 51.0 = 100.

Pyr., etc.—B.B. melts readily to a yellow globule, which on cooling becomes white and crystalline. On charcoal in R.F. gives metallic lead, with a white coating of lead chloride. With a salt of phosphorus bead strongly saturated with copper oxide gives the chlorine reaction. Dimorphic with effervescence in dilute nitric acid.

Obi.—At Cromford near Matlock in Derbyshire; at Globbas. Monte Pont and Montevecchio in Sardinia.

Northampton. $Na_2CO_3 \cdot NaCl$. In isometric octahedrons. White to yellow or gray. From Borax Lake, San Bernardino Co., California.

B. Acid, Basic, and Hydrous Carbonates.

Diatomic carbonate, H_2SiO_3 . In yellowish to white deposits of Africa. Patagonia, the Chiocha Islands.

ICAI-AOHITEL.

Monoclinic. Axes $a:b:c = 0.8809 : 1 : 0.4012$; $\beta = 61^\circ 50'$.

Crystals rarely distinct, usually slender, acicular prisms ($m = 75^\circ 40'$), grouped in tufts and rosettes. Twins: tw. pi. a common. Commonly massive or crystalline, with surface botryoidal, or stalactitic, and structure divergent; often delicately compact fibrous, and banded in color; frequently granular or earthy.

Cleavage: c perfect; a less so. Fracture subconchoidal, uneven. Brittle. $H. = 3.5-4$. $G. = 3.9-4.03$. Luster of crystals adamantine, inclining to

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▼ vitreous; of fibrous variety more or less silky; often dull and earthy. Color bright green. Streak paler green. Transparent to subtranslucent to opaque. Comp.—Basic cupric carbonate, $CuCO_3 \cdot Cu(OH)_2$, or $2CuO \cdot CO_2 \cdot H_2O$ = Carbon dioxide 19.9, cupric oxide 11.9, water 8.2 = 100.

Pyr., etc.—In a test tube blue and yields water. B.B. fuses, color changes the same

emerald-green; on charcoal !b reduced to metallic copper; with the fluzei rmcb like cuprite. Soluble la acids with e&erTesceDce.

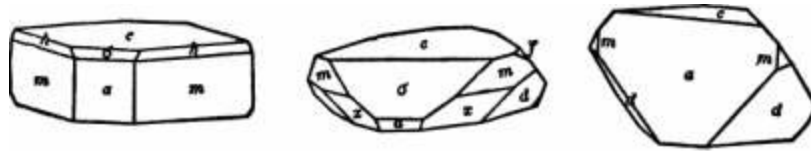
Diff.—Characterized by ^eeu color and copper reactions B.B.; differs from other copper ores of a green color in Ita (Effervesceuce with acids.

Obi.—Common with other ores of copper and as a product of their alteration; thus as a pseudomorph after cuprite and B^iirite. Occurs abundantly in the Ural; at Clivay in France; in Cornwall and in Cumberland, England; in Uelnbrctlmch; Dillenburg, Nassau; Bellerdorf near Siegen. At the copper mines of Nizhni Tuglisk; with the copper ores of Cuba; Chili; at the Cobalt mines and elsewhere in New South Wales; South Australia.

Occurs in Jersey, at Schuyler's mines, and at New Brunswick, in Pennsylvania, at Cornwall, Lebanon Co.; at the York and Ploverville lead-mines. In Michigan, at the copper mines of Mineral Point, and elsewhere. Abundantly in quartz masses and acicular crystals, with calcite at the Copper Queen mine. Bisbee, Coconino Co., Arizona; also in Grabain Co. at Horenc (6 m. from Clifton), in alactitic forms of malachite and azurite in concentric bands. At the Santa Rita mines. Oront Co., and elsewhere in the State of Mexico, Tintic district, Utah. Named from the Greek $\chi\alpha\zeta\iota$, matron. In allusion to the green color.

Crystallography. Monoclinic. Axes : $d : c : b = 0.8501 :$

$a : b : c = 1.0000 : 0.8733 :$



«'.

$\angle 110 = 90^\circ 41'$.

$\angle 001 = 90^\circ$.

$\angle 131 = 44^\circ 44'$. $\angle 028 = 60^\circ 44'$.

Angles: $\angle 031 = 90^\circ$. $\angle 110 = 90^\circ 41'$. $\angle 001 = 90^\circ$. $\angle 131 = 44^\circ 44'$. $\angle 028 = 60^\circ 44'$. $\angle 221 = 78^\circ 10'$.

Crystals varied in habit and highly modified. Also massive, and presenting imitative shapes, having a columnar composition; also dull and earthy.

Cleavage: p (031) perfect but interrupted; a less perfect; tn in traces. Fracture conchoidal. Brittle. H. = 3-5-4. G. = 3.77-3.83. Luster vitreous, almost adamantine. Color various shades of azure-blue, passing into Berlin-blue. Streak blue, lighter than the color.

Transparent to subtranslucent.

Comp.—Basic cupric carbonate, $3\text{CuCO}_3 \cdot \text{C}(\text{OH})_2$, or $3\text{CuO} \cdot 2\text{CO}_2 \cdot \text{H}_2\text{O}$ = Carbon dioxide 25-6, cupric oxide 69-2, water 5-2 = 100.

Pyr., etc.—Sums in malachite.

^ff.—Characterized by its blue color; effervescence in nitric acid; copper reactions B.B.

Obi.—Occurs in splendid crystallizations at Chessy, near Lyons, whence it derived the name Chessy Copper or chettyite. Also in fine crystals in Moldavia in the Banat; at Wheal Buller, near Redruth in Cornwall; in Devonshire and Derbyshire, England; Cobalt mines and elsewhere in New South Wales; South Australia.

Occurs in Penn., at Phoenixville, in crystals. In Tenn. near New Brunswick, in Wisconsin, near Huerfano Point. In Ariz. at the Longfellow and other mines in Graham

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Co.: in malachite. Beautiful crystals at the Copper Queen mine, Bisbee. In Omit Co., Nevada. At the Manitowish mine in the Tintic district, Utah. In California, California Co., in Humboldt's mine, in crystals.

Auriferous. A basic carbonate of zinc and copper, $2[\text{Zn}, \text{Cu}]\text{CO}_3 \cdot 3(\text{Zn}, \text{Cu})\text{OH}$. In d. / u. = S'54-B 64. Luster pearly. Color pale green to blue. From the Allali Mine, near Lyons: U.S. & N.Y., U.S. & N.Y.; U.S. & N.Y. In U.S. & N.Y., at Iriaca, Pa.; the Santa Caterina Hib., Arizona; Beaver Co., Utah.

Hydrozincite. A basic zinc carbonate, perhaps $\text{ZnCO}_3 \cdot 2\text{Zn}(\text{OH})_2$. Muscovite, fibrous, earthy or compact, as in crustification. G. = 4.55-88. Color white, grayish or yellowish. Occurs at the base of zinc, as a result of alteration. In great quantities at the Dolores mine, San Juan, Spain. In the D. S., at Friedenau, Pa.; and Linden, in Wisconsin.

Hydrocerussite. A basic lead carbonate, probably $2\text{PbCO}_3 \cdot \text{Pb}(\text{OH})_2$. In thin colorless hexagonal plates. Occurs as a coating on native lead, at Långsju, Sweden; with galena at Wanlockhead, Scotland.

Dawsonite. A basic carbonate of aluminum and sodium, $\text{NaAl}_2(\text{CO}_3)_2 \cdot 2\text{Al}(\text{OH})_3$. In thin crystallizations of white radiating bladed crystals. O. = 4. Found on a feldspathic dike near McGill College, Montreal. From the province of Siena, Pian Castagneto, Tuscany.

Tetraammonium. Hydrated sodium carbonate, $\text{Na}_2\text{CO}_3 \cdot \text{H}_2\text{O}$. Q. = 1-5-1-6. Occurs in various islets, and as an efflorescence over the soil in many dry regions.

Neiquenaite. Hydrated magnesium carbonate, $\text{MgCO}_3 \cdot 8\text{H}_2\text{O}$. In radiating groups of

prismatic crystals. Q. ~1-88-185. Colorless to white. From a coal mine at Nesque-boning, Schuylkill Co., Penn. Beechonsfordite. p. 367.

Natron. Hydrated sodium carbonate, $\text{Na}_2\text{CO}_3 \cdot 10\text{H}_2\text{O}$. Occurring in nature only in solution, as in the soda lakes of Egypt, and elsewhere, or mixed with the other sodium carbonates.

Platonic. $\text{CaCO}_3 \cdot \text{Na}_2\text{CO}_3 \cdot \text{SiO}_2$. In prismatic crystals, orthorhombic-hemimorphic. Colorless to white. Borai Lake, San Bernardino, California.

OATXnSSITB .

Monoclinic. Axes $a : b : c = 1.4897 : 1 : 1.4443$; $\beta = 78^\circ 27'$.

$d_{100} = 20.0 \text{ mm}^{-1}$, $d_{110} = 11.0 \text{ mm}^{-1}$.

$\alpha = 109^\circ 80'$.

$\mu = 43^\circ 81'$.

$\rho = 68^\circ 28'$.

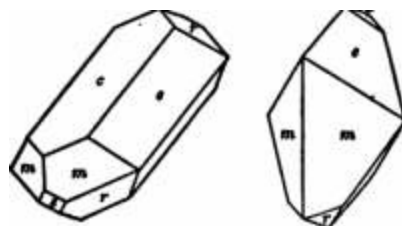
Crystals often elongated and also flattened wedge-shaped. Cleavage: m perfect; e rather difficult. Fracture conchoidal. Very brittle. H. 2-3, G. = 1.93-1.95. Lustrous. Color white, yellowish white. Streak uncolored to grayish. Translucent.

Comp.—Hydrated carbonate of calcium and sodium, $\text{CaCO}_3 \cdot \text{Na}_2\text{CO}_3 + 5\text{H}_2\text{O}$ = Calcium carbonate 33-8, sodium carbonate 35-8, water 30-4 = 100.

Pyro., sto.—Heated in a closed tube decrepitates and becomes opaque. B.B. fuses easily to a white enamel, and colors the flame intensely yellow. Dissolves in acids with a brisk effervescence; partly soluble in water, and reddens turmeric paper,

Obs.—Abundant at Anginilla, near Merida, in Venezuela, in crystalline shells. At the bottom of a small lake, in a bed of clay, covering the ground. Also abundant in Little Salt Lake, or Soda Lake, in the Carson desert near Reno, Nevada, deposited upon the evaporation of the water. Named after J. Lussac, the French chemist (1773-1850).

Lanthanite*. $\text{La}_2(\text{CO}_3)_3 \cdot 9\text{H}_2\text{O}$. In thin tabular crystals; also granular, earthy. G. = 3.003. Color grayish white, pink, yellowish. Found coating cerite at p----



CARBONATES.

TRONA. Urao.

MonooUnio. Axea: d: X ; i = 3-8460:1: 2-9700; fi = 77° 2S'. ea, 001 A 100 = 77' Sff. 731.

«, 001 A ill = 70° Mi'. 00", ill Alii = «° 861'.

Often fibrous or columnar musite.

Cleavage: a perfect; o, c in traces. Fracture

uneven to subconchoidal. H. = 2' & -3. Q. = S'll-

2-14. Luster vitreous, glistening. Color gray or yellowish white. Trsnaln-

cut. Taste alkaline.

Comp.—Na.C'O..HNaCO,-f 2H,o or 3Na.o.400,.5H,o = Carbon dioxide 38-9, Boda 41-^
water 19-9 = 100.

Chatanl esLiiblUbed tlie above composition for nrao, and showed thai troiia. somellrac
c&lled " sesquicarboDHte of soda," li an Impure form ot ibe tame compound.

Pyr., ate. —Id tUe closed tube yields water and carbon dioxide. B.B. Imparts an Intenselj
jullow color to the flame. Boluble in water, and eflerresces wliii acid*. Reacts alkaline
witli moistened test-paper.

OIm. —Found [□ the pro^nce of Fezsan, Africa, forming tliin superficial cnisia; nt tba
bottom of a lake at Lagunilla. Venezuela, 8. A. Efflorescences ol trona occur near the
Sweetwater river, Rocky Mountains. An extensive bed In Churcliill Co., Nevnds. In floe
crystalB at Borax lake, San Bernardino Co., California, with liankBlte, glauberlte,
thenaralte, etc.

HydromaKiwdto. Basic magnesium carbonate, 8HgCo.,Hg(OH)}, + 8H,o. Cryslals small,
tutted. Also amorphous: si chalky crusts. Color and streak white. Often occur* with
serpentine; thus at HrubschUiz, in Moravia; at Kraubat, Btyrla, etc Also similarly near
TeiBfl, Pm. ; Huboken, N. J.

Hydroglobvtito. MgCO.,Hg(OH)i + 3H,o. In light gray spherical forma. From the
neighborliood of Pollena. Italy.

Laiuferdltta. 8HgCO.,Mg<OH), + 21H,o. Occurs as small atolactltes in the anthracite mine
at Nesquehoniug near Lansford. Schuylkill Co., Penn.; changod on exposure t4>

nesquehonite.

Zaratt* . Emerald Nickel. $\text{NiCO}_3 \cdot 4\text{H}_2\text{O}$. In mammillary incrustations: also nmsalve, compact. Color emerald-green. Occurs on chromite at Texas, Lancaster Co., Pa.; at Swinauess, Unst, Shetland, and elsewhere.

Ramsvonit*. A hydrous cobalt carbonate. A rose-colored incrustation, soft and earthy. Vioia a copper mine near Finksburg, Carroll Co., Maryland.

Tangmita. A supposed yttrium carbonate. In white pulverulent coating*. On gadolinite at Ytterby. A similar mineral is associated with the gadolinite of Llano Co., Texas.

Bismutite*. Wismuthspat. A basic bismuth carbonate, perhaps $\text{Bi}_2\text{O}_3 \cdot \text{CO}_2 \cdot \text{H}_2\text{O}$.

■ idcnisling, or earthy and pillenitent; amorphous. G. = B86-B-9 Breilh.: 7-67 Rg. Color white, green, yellow and gray. Occurs at Schreeberg and Johanngeorgenstadt, with native bismuth, and at Joachimsthal. etc. In the U. S., In Bo, Carolina, at Brewer's mine; In Gaston Co., N. C., etc.

Dranothite. $2\text{CaCO}_3 \cdot \text{U}(\text{CO}_3)_2 \cdot 10\text{H}_2\text{O}$. In scaly or kmntilar crystalline aggregates. Color siskin-green. Occurs on uraninite at Joachimsthal, Bohemia.

Uraninite*. A hydrous carbonate of uranium and calcium. In mammillary concretions, or thin coatings. Color apple-green. Occurs on uraninite near Adrianople, Turkey; also JobauDireorgensladt and Joachimsthal.

Toirillite. A hydrous carbonate of uranium, calcium and copper. In aggregations of rrvMnllnp ai-alea. Color emerald-green to bright grass-green. From the Ellas mine, near Joachimsthal. on uraninite.

DiQitizcdbbyGoOglC

DESCRIPTIVE UINEBALOOT.

Oxygen Salti. 2. SHJOATES.

The Silicates are in part strictly anhydrous, in part hydrous, as the zeolites and the amorphous clays, etc. Furthermore, a large number of the silicates yield more or less water upon ignition, and in many cases it is known that they are, therefore, to be regarded as basic (or acid) silicates. The line, however, between the strictly anhydrous and hydrous silicates cannot be sharply drawn, since with many species which yield water upon ignition the part played by the elements forming the water is as yet uncertain. Furthermore, in the cases of several groups the strict arrangement must be deviated from,

since the relation of the species is best exhibited by introducing the related hydrous species immediately after the others.

This chapter closes with a section including the Titanates, Silico-titanates, Titanoniobates, etc., which connect the Silicates with the Niobates and Tantalates. Some Titanates have already been included among the Oxides.

Section A. Chiefly Anhydrous Silicates. I. Disilicates, Polysilicates. II. Metasilicates. III. Orthosilicates. IV. Subsilicates.

The Disilicates, $B_2Si_2O_7$, are salts of disilicic acid, $H_2Si_2O_7$, and have an oxygen ratio of silicon to bases of 4 : 1, as seen when the formula is written after the dualistic method, $EO.2SiO_2$,

The Polysilicates, $H_nSi_nO_{3n+1}$, are salts of polysilicic acid, $H_nSi_nO_{3n+1}$ and have an oxygen ratio of $n+1$: 1, as seen in $3RO.3SiO_2$.

The Metasilicates, $ESiO_3$, are salts of metasilicic acid, H_2SiO_3 , and have an oxygen ratio of 3 : 1. They have hence been called Metasilicates.

The Orthosilicates, R_2SiO_4 , are salts of orthosilicic acid, H_4SiO_4 , and have an oxygen ratio of 1 : 1. They have hence been called orthosilicates. The acidity of the silicates fall into one of the last two groups.

Furthermore, there are a number of species characterized by an oxygen ratio of less than 1 : 1, e.g., 3 : 4, 2 : 3, etc. These basic species are grouped as Subsilicates. Their true position is often in doubt; in most cases they are probably to be regarded as basic salts belonging to one of the other groups.

The above classification cannot, however, be carried through strictly, since there are many species which do not exactly conform to any one of the groups named, and often the true interpretation of the composition is doubtful. Furthermore, within the limits of a single group of species, connected closely in all essential characters, there may be a wide variation in the proportion of the acidic element. Thus the triclinic feldspars, placed among the polysilicates, range from the true polysilicate, $NaAlSi_3O_8$, to the orthosilicate, $CaAl_2Si_2O_8$, with many intermediate compounds, regarded as isomorphous compounds of these extremes. Similarly of the acapelite group, which, however, is included among the orthosilicates, since the majority of the compounds observed approximate to that type. The micas form another example.

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Orthosilicates, $B_2Si_2O_7$. Polysilicates, $B_nSi_nO_{3n+1}$,

Monoclinic. Crystals rare (castorite). Usually massive, foliated cleavable (petalite).

Cleavage: c perfect; o (201) easy, z (905) difficult and imperfect. Fracture imperfectly conchoidal. Brittle. H. = 6-6-5. G. = 3-39-2-46, Luster vitreous, on c pearly. Colorless, white, gray, occasionally reddish or greenish white. Streak unicolorous. Transparent to translucent.

$\text{Co} \ll \text{p} - \text{LiAl}(\text{Si}, \text{o}),$ or $\text{Li} \cdot \text{O} \cdot \text{Al}, \text{o}, .8\text{SiO}, =$ Silica 78.4, alumina 16-7, lithia, 4-9 = 100.

Pyr., etc.—Gently heated in a blue phosphoric acid light. B.B. on charcoal becomes glassy, but on charcoal alone, and on mella only on the edges: gives the reaction for lithium. With borax it forms a clear, colorless glass. Not acted on by acids.

Obt.—Petalite occurs at the Iron mine of Ulo, Sweden, with lepidolite, tourmaline, spodumene, and quartz; on Elba (Tutorio). In the U. S., at Bolton, Hau., with scapolite; at Peru, Haine. with spodumene } a albite. The name petalite from the fracture, alluding to the cleavage.

Milarite. $\text{HECa}, \text{Al}, (\text{S}, \text{o}),$. In hexagonal prisms. H. = 5. G. = 3.65-2.68. Colorless to pale green, glassy. From Tal Qiu, Quito, Ecuador. Switzerland.

Endymite. $\text{HNaBeF}_2 \cdot \text{O},$ In white, glassy, twinned crystals. Upright habit H. = 6. G. = 3.55. Occurs very sparingly in effluve-syalt on the island Ovre-Aro, in the Langesundfiord, Norway.

Endymite, Same composition as endymite. Orthorhombic. Borthen. Quedlinburg.

Orthoclase

Soda-Orthoclase Hyalophane

Kiorcline

Soda-microcline Anorthoclase

Feldspar Group.

a. $\text{K}(\text{Na})\text{AlSi}_3\text{O}_8$ Section.

i: KAISLO. 0-6585 :1;

J $(\text{K}, \text{Na})\text{AlSi}_3\text{O}_8$,

i $(\text{Na}, \text{K})\text{AlSi}_3\text{O}_8$. $(\text{K}, \text{Ba})\text{Al}_2\text{Si}_2\text{O}_8$,

I 0-5554

0-6584 :1: 0-6512

/I. THctinU Section.

EAAlSi₃O₈,

(K,N)AlSi₃O₈,

(Na,K)AlSi₃O₈.

Albite. Muirheadite « Serpentine Pseudomorphs.

Albite « Olivine

Labradorite

Asorthite

Celsian

HaAlSi₃O₈,

0-6335 : 1: 0-5577 94° 0-6321 : 1: 0-5524 93°

116° 29' 55" 9' 116° 23' 90° 5'

/MNaAlSi₃O₈, 1^mCaAl₂Si₂O₇,

.o.)

CaAl₂Si₂O₇. B»Al₂Si₂O₇,

0-6357:1:0-5521 93° 23' 116° 29' 89'59"

0-6377:1:0-5547 93° 31' 116° 3' 89°54'

0-6347 : 1; 0-5501 93° 13' 115° 55' 91° 12' 90°

,ab,GoOgIc

DBSCBTPTiyK UINEBALOGT.

The general characters of the species belonging in the Feldspar Group are as follows:

1, Crystallization in the monoclinic or triclinic systems, the crystals of the different species resemble each other closely in angle, in general habit, and in methods of twinning. The prismatic angle in all cases differs but a few degrees from 60° and 120°.

2, Cleavage in two similar directions parallel to the base c (001) and clinopyroxenoid (or

brachypinacoid) b (010), inclined at an angle of 90° or nearly 90°. 3, Hardness between 6 and 6.5. 4, Specific Gravity varying between 2.5 and 2.7, and mostly between 2.55 and 2.75. 5, Color white or pale shades of yellow, red or green, less commonly dark. 6, In composition silicates of aluminium with either potassium, sodium, or calcium, and rarely barium, while magnesium and iron are always absent. Furthermore, besides the several distinct species there are many intermediate compounds having a certain independence of character and yet connected with each other by insensible gradations; all the members of the series showing a close relationship not only in composition but also in crystalline form and optical characters.

The species of the Feldspar Group are classified, first as regards form, and second with reference to composition. The monoclinic species include (see above): Orthoclase, potassium feldspar, and Soda-orthoclase, potassium-sodium feldspar; also Hyalophane, barium feldspar.

The triclinic species include; Microcline and Anorthoclase, potassium-sodium feldspars; Albite, sodium feldspar; Anorthite, calcium feldspar; Gelsin, barium feldspar.

Also intermediate between albite and anorthite the isomorphous sub-species, sodium-calcium or calcium-sodium feldspars: Oligoclase, Andesine, Labradorite.

DORITE.

a. Monoclinic Section.

Orthoclase.

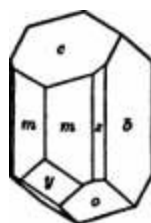
Monoclinic. Axes & ;

$a : b = 0.6585 : 1.0000$; $\beta = 63^\circ 57'$.

(J7)

$110 \wedge 110 = 61^\circ$ iff.

$180 \wedge 130 = 58^\circ 48''$. $001 \wedge 101 = 50^\circ 16'$. $001 \wedge 501 = 19^\circ$.



$\wedge n$

$001 \wedge 031 = 44^\circ 86'$. $031 \wedge 011 = 8^\circ 68'$. $001 \wedge 110 = 7^\circ$ tr.

001 A ill = 35'14i'.

Twins: tw. pi. (1) a (100), or tw. axis i, the common Carlsbad twins, either of irregular penetration (Fig. 727) or contact type; the latter usually with b aE com position-face, often then (Fig. 728) with c and x nearly in a

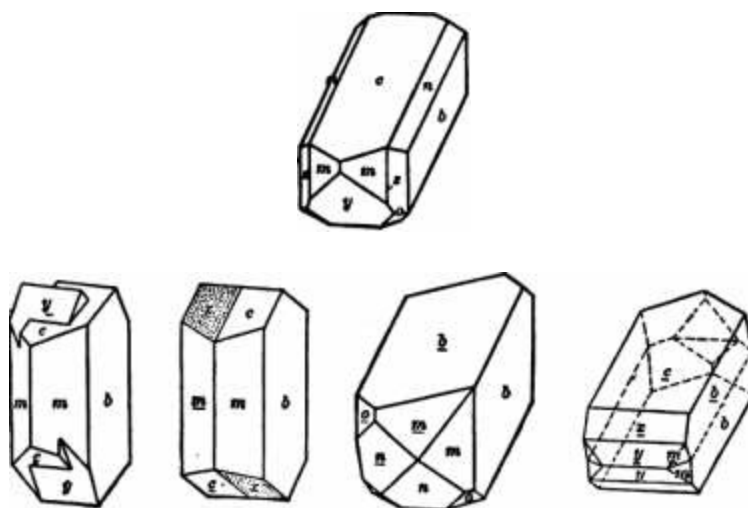
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SILICATES.

plane, but to be distin[^]ished by luster, cleavage, etc. (3) n (p[^]l)* the Baveno twins forming nearly square prisms (Fig. 739) > since en = 44 56[^]', and hence cc = 89° 53'; often repeated as lournalings (Fig. 417, p. 139), also in square prisms, elongated Q a. (3) c (001), the Manebaeh twins (Fig, 730), usually contact-twins with c as comp.-face. Also other rarer laws.

Crystals often prismatic 1 I; sometimes orthorhombic in aspect (Figs. 723, 7S5) since c and x are inclined at nearly equal angles to 6; also elongated | d (Fig. 736) with b and c nearly equally deTeloped; also thin tabular | b: rarely tabular \ a, & face not often observed. Often maasiye, coarsely cleavable to granular; sometimes lamellar. Also compact crypto-crystalline, and flint-like or jasper-like.

Cleavage: c perfect; b somewhat less so; prismatic tn imperfect, but usually more distinct parallel to one prismatic face than to the other. Parting sometimes distinct parallel to a (100), also to a hemi-orthodome, inclined a few degrees to the orthopinacoid; this may produce a satinlike luster or schiller (p. 190), the latter also often present when the parting is not distinct. Fracture conchoidal to uneven. Brittle. H. = 6. G. = 2'57. Luster vitreous; on e often pearly. Colorless, white, pale yellow and flesh-red common, gray; rarely green. Streak nncolored. 737. 728. 72».

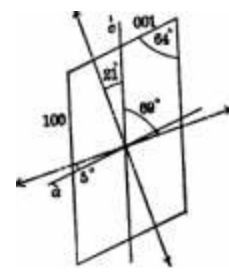


Optically negative (Bx, = o) in all cases (Fig. 731). Ax. pi. usually $_L^*$, ijj^{\wedge} sometimes || b, also changing from the former to

the latter on increase of temperature (seep. 335). For adularia (Dx) Bx, $A^\wedge = -69^\circ 11'$, $Bx^\wedge w A (! = -69^\circ 37'$. Ueuce Bx, and the extinction-direction (Fig. 731) inclined a few degrees only to \perp , or the edge b/c; thus -3° to $+7^\circ$ usually, or . up to $[-]0^\circ$ or -12° in varieties rich in Na, O. Dispersion $p > v$; also horizontal, strongly marked, or inclined, according to position of ax. pi. Axial angles variable. Birefringence low, $y - a = 0.007 - 0.005$. For adularia (5x).

Qr = 1-5100.)9, = 1-58S7, rj- = 1'5280, . ■. 2Yf = W 43-, 2Er = $121^\circ 6'$.

Compn Tar. —A silicate of aluminium and potassium, $EAISi.O$, or $K, O.Al, O, .6SiO$, = Silica 64-7, alumina 18-4, potash 16-9 = 100. Sodium is often also present, replacing part of the potassium, and sometimes exceeds it



in smonnt; these Taretles are embraced under the name Boda-orthoclam (NatroDorthoklas Oerm.).

The promloeDt varieties depend upon crynallioe habit and method of occurrence more tban upuD dLSereuce of compositloD.

1. Adularia. The pure or nearly pure potaaalum silicate. nBuually in cryaiala. like Fig. T3Jf in liabit; often with viclual planeei Baveuo twina common. G. = 2-565. Transparent or utariy so. Often with a pearly opalescent reflection or Schiller or a lielicale play of colors; so'iiie mooiuU>n« is here included, but the remainder belongs to iilbile or other of tlie Iricltiic feldspars. The original adularia (Adulari is From the St. Qolliard region iu Swlizeriund. Vaieiteiantttt, from the silver mine of Valencia, Mxslco. is adiilaria.

2. Sanidine or glany fald^ar. Ocoura in crystals, often trauspareui nnd gliusy, embedded in rbyollle, trachyte (iis of the Siebeugebirge). phonolile, etc. Habit ofien tabular | b (hence named from <Tavi%, a tablet, or board); also In square prisms (A, e); Carls-bftd twina common. Hoet varieties contain sodium as a promineit constituent, and lience tMsiiong to the aoda-orthoclase,

Shyaeolite. Sit*patli Werner. Occurs in glassy crystals at Konte Somnu>: named from pvai, itream (lava stream).

8. Ordinary. In crystals, Carlsbad and other twina common; also massive or cleavable, varying in color from white to pale yellow, red orgreeu. translucent: aomelimes sventuriue. Here iKlongs lthe common feldspar of granitoid rocks or graulie veins.

Usuiilly contains a greater or less percentage of Bodalsoda-ortUoclaae). Compact cryptocryaialline orthoclase makes up the mass of much felslte, but to)i greater or less degree iiduiixed with quartz: of various colors, from wliite and brown to deep red. Much of what baa been called ortho-clase, or common potash feldspar, has proved to belong to the related triclinic species, microcline. Cf. p. 874 on the relations of the two species. Clie^terliie and Amazxin stone are microcline; also moat aventurine orlhoclase. Loxtxlaie contains sodium in considerable amount (7'6Na,o). From Hammond, St. Lawrence Co., N. Y. JfureAuonite is a flesU-reil feldspar similar to perthite{p. 873), with gold yellow refleclions in a direction x 6 and nearly parallel to 701 or 801 (p. 371). Prom Dawlisb and Exeter, England,

The spberulites noted in some volcanic rocks, as in the rhyoiite of Ubsidlnn Cliff In the Tellowstone Park, are believed to consist essentially of orthoclase needles with quartz. These nre shown in Figs. 733 and 73il (from Iddlugs; much magnified) as Ibej api)enr iu]X)larized liglit (crossed nicols).

Pyr., etc.—B.B. fu-esat5: vnrtelies containing much soda are more fusible. Loxoclase fuses at \. Not acted upon by acids.

Diff —Cbarnc tori zed by its crystalline form and the two cleavages at riglit angles to each other; harder than tiarite and calcite: not attacked by acids; difficultly fusible. Massive corundum is much harder nnd lins a higher specific gravity.

Dislingualied in rock sections by its low refraction (low relief) and low interference-colon, which Inst scarcely rise to white of the first order—hence lower than those of quartz; also by its liiaiaal cliaaracler in convergent light and by the distinct cleavages. It ia colorless in ordinary light and may be limpid, but is frequently Inrbid and brownisb from the presence of very raiuule scales of kaolin due to alteration from weathering; this change is especially common In the older granular rocks, as granite and gneiss.

;;CoogIC

Oba.—OrthoclHse io Its BevL-ral Tarielies belongs especial]]' to the crystalline rocks, occurring aa an esseolial cODstuent of granite, gnelsB, syenite, also porplyry, further (var., »anidiie) trachyte, phuuolite. etc. In the niasaive grauituiii rocks ii is seldom in distinct, well-fnrmed, separable crystals, except iu veina and cavities; such cryaials are more common, however, lu TolcaDic rocks like trachyte.

Adularia occurs in the crystalline rocks of the central and eastern Alps, associated with emobj' quartz and nlbite. also litauite, apatite, etc.; the crystals are often cc&ted with chlorite: also on Elba. Fine cryalals of ortboclase, often Inius, are obtained from Baveno, Lago Maggiore; the Fleimathal, a red variety; ValUoriana; Bodenmais, Carlsbad and £lDi3£en iu Bohemia; Briegau, etc., in Silesia. Also Arendal in Norway, and near Shait&ask lu the Ural; Land's End and Bt. Agnes In Cornwall; the Mourne Mts., Ireland,

with beryl and topaz. From Tamagama Yama, Japan, with topaz and smoky quartz. Moontumt is brought from Ceylon.

In the U. States, orthoclase is common in the crystalline rocks of New England, also of States south, further Colorado, California, etc. Thus at the Paris tourmaline locality. In JV. Hamp., at Acworth. In Mass., at South Royaston and Barre. In Conn., at Haddam and Middletown, in large coarse crystals. In N. York, in St. Lawrence Co., at Bossie; at Hammond (loxoctate); In Lewis Co., in white limestone near Natural Bridge; at Amity and Edenville. In Pa., in crystals at Leipserville, Mineral Hill, Delaware Co.; sunstone In Eennett Township. In Ill. Otr. at Washiugtnu Mine, Davidson Co. In ColoraiU). at the summit of Sit. Antero.CtaaSee Co., in fine crystals, often twins; at Gunnison; Black Hawit; Kokoma. Summit Co., also at other points. Also similarly in Nevada and California.

Orthoclase is frequently altered, especially through the action of carbonated or alkaline waters; the final result is often the removal of the potash and the formation of kaolin. Steatite, talc, chlorite, leucite, mica, laumontite, occur as pseudomorphs after orthoclase; and feldspars often replace these feldspars by some process of solution and substitution.

Perthite. As first described, a flesh-red aventurine feldspar from Perth, Ontario, Canada, proved a sodio-orthoclase, but shown by Gerhard to consist of inter laminated orthoclase and albite. Many similar occurrences have since been noted, as also those in which microcline and albite are similarly intermingled, the latter called microcline-perthite, or microcline-albite-perthite; this is true in part of the original perthite. When the structure is discernible only with the help of the microscope it is called microperthite. Brögger has investigated not only the microperthites of Norway (Orthoklaamikroperthit, Miktoklin-mikroperthit). but also other feldspars characterized by a marked Schiller; he affirms the existence of an extremely fine intermingling of albite and orthoclase (801, not discernible by the microscope (cryptoperthite), and connected with secondary planes of parting 1100 or 1801. which is probably to be explained as due to incipient alteration.

Hyalophane. $(K, Ba)Al_3(8SiO_4)_2$ or $E, O. BaO. 2Al_2O_3. 8SiO_2$. Silica 61.6, alumina 31.0, barytes 16.4, potash 10.1 = 100. In crystals, like adularia in habit (Fig. 735, p. 570): also massive. Cleavage: c perfect; b somewhat less so. H. = 6-6.5. G. = 3.805. Occurs in a granular form in the Binnenthal, Switzerland; also at the manganese mine of Jakobs-berg, Sweden. Some other feldspars containing 7 to 15 p. c, BaO have been described (cf. also celsian, p. 381).

3. Triclinic Section. BUROOLINB.

Triclinic. Near orthoclase in angles and habit, but the angle between c and b ~ about 89° 30'. Twins: like orthoclase, also polyaxial twinning according to the albite and pericline laws (p. 375), common, 734.

producing two series of fine lamellae nearly at right angles to each other, hence the characteristic grating-structure of a basal section in polarized light (Fig. 734). Also massive cleavable to granular compact.

Cleavage: c perfect; β somewhat less a M (110) sometimes distinct; m (110) also sometimes distinct, but less easy. Fracture uneven. Brittle. H. = 6-6-5. G. = 2-54-2-5'r. Lustre vitreous, on c sometimes pearly. (Wor white to pale cream-yellow, also red, green. Transparent to translucent.

Optically —. Ax. pi. nearly perpendicular ($82''$ - $83''$) to b. Bx^\wedge inclined $15^\circ 26'$ to a normal to β . Dispersion $p < v$ about Bxo. Extinction-angle on e, + $15^\circ 30'$; on S, + 5° to 6° (Fig. 739, p. 375). $2H_{\parallel} = 88^\circ$ to 89° , Dx.

The essential identity of orthoclase and microcline has been urged by Mallard and Michel-Levy on the ground that the properties of the former would belong to an aggregate of submicroscopic twinning lamellae of the latter, according to the albite and pericline laws. Comp., Tar.—Like orthoclase, $KAlSi_3O_8$, or $K_2O \cdot Al_2O_3 \cdot 6SiO_2 = \text{Silica } 64.7, \text{ alumina } 18.4, \text{ potash } 16.9 = 100$. Sodium is usually present in small amount: sometimes prominent, as in soda-microcline.

Pyroxene as for orthoclase.

Differences.—K-feldspar distinguished by optical characters (the grating structure in polarized light, Fig. 731); also often shows fine twinning-lamellae on a basal surface (albite law).

Occurrence.—Occurs under the same conditions as much common orthoclase. The beautiful Montrose from the Ural, and that occurring in fine groups of large crystals of deep color in the gneiss of Pike's Peak, Colorado, is microcline. Cokesler from Poorhouse quarry, Chester Co., Penn. and the aegirine feldspar of Mineral Hill, Penn., belong here. A pure variety occurs at Magriet Cove, Arkansas. Ordinary microcline is common at many points.

Anorthoclase. A triclinic feldspar with a cleavage-angle, β , varying but little from 90° . Form like that of the ordinary feldspars. Twinning as with orthoclase; also poly-synthetic according to the albite and pericline laws; but in many cases the twinning lamellae very narrow and hence not distinct. Rhombic section (see p. 875) inclined on β . 4 to 6° to edge b/e . O. = 3-67-160. Cleavage, hardness, luster, and color as with other members of the group. Optically —. Extinction-angle on e, + 8° to 10° ; on b, 6° to 9° -8. Bs. nearly \pm J. Dispersion $p > v$; horizontal distinct. Axial angle variable with temperature, becoming in part monoclinic in optical symmetry between 10° and 30° C, but again triclinic on cooling: this is true of those containing little calcium.

Chiefly a soda-potash feldspar, $NaAlSi_3O_8$, with $KAlSi_3O_8$ the sodium silicate usually in

larger proportion (3:1.8:1, etc.), as if consisting of orthoclase and albite molecules. Calcium (CaAl₂Si₂O₈) is also present in relatively very small amount.

These triclinic soda-potash feldspars are chiefly known from the andesitic lavas of Pantelleria. Most of these feldspars come from a rock, called pantellerite. Also prominent from the alkali-syenite of Southern Norway and from the "Romben-porphyr," near Christiania. Here is referred also a feldspar in crystals, tabular, and twinned according to the Habach and less often Baveno laws occurring in the lithophyses of the rhyolite of Obsidian Cliff, Yellowstone Park. It shows the blue opalescence in a direction parallel with a certain orthodome (cf. p. 071).

Albite-Anorthite Series. $\text{NaAlSi}_3\text{O}_8$ - $\text{CaAl}_2\text{Si}_2\text{O}_8$. * Between the isomorphous species

Albite $\text{NaAlSi}_3\text{O}_8$, Ab

Anorthite $\text{CaAl}_2\text{Si}_2\text{O}_8$, An

there are a number of intermediate subspecies, regarded, as urged by Tschermak, as isomorphous mixtures of these molecules, and defined according to the ratio in which they enter; their composition is expressed in general by the formula $\text{Ab}_x\text{An}_{1-x}$. They are;

OLIGOCLASE $\text{Ab}_{0.9}\text{An}_{0.1}$, to $\text{Ab}_{0.7}\text{An}_{0.3}$,

Andesine $\text{Ab}_{0.6}\text{An}_{0.4}$, to $\text{Ab}_{0.5}\text{An}_{0.5}$,

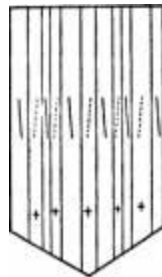
Labradorite $\text{Ab}_{0.4}\text{An}_{0.6}$, to $\text{Ab}_{0.3}\text{An}_{0.7}$

and Bytownite $\text{Ab}_{0.2}\text{An}_{0.8}$, to $\text{Ab}_{0.1}\text{An}_{0.9}$,

From albite through the successive intermediate compounds to anorthite

with the progressive change in composition (and specific gravity), there is also

change in



a corresponding change in crystallographic form, and in certain fundamental optical properties.

Crystalline form. The axial ratios and angles given on p. 369 show that these triclinic

feldspars approach orthoclase closely in form, the most obvious difference being in the cleavage $\wedge \wedge \wedge \wedge \wedge \wedge$

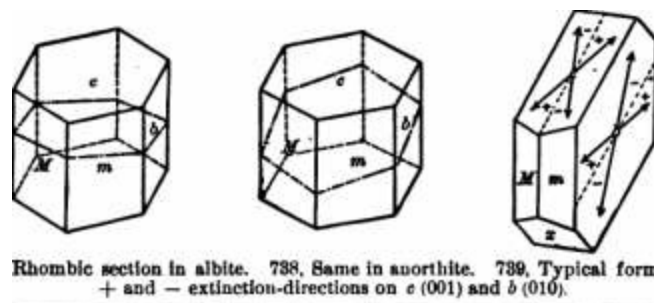
axial angle 2ψ , which is 90° in orthoclase, $86^\circ 24'$ in albite, and $85^\circ 50'$ in anorthite. There is also a change in the axial angle γ , which is 88° in albite, about 90° in oligoclase and andesine, and 91° in anorthite. This transition appears still more strikingly in the position of the "rhombic section," by which the twins according to the pericline law are united as explained below. Plagioclase with twinning lamellae, Fig. 735 section I o

Trigonal twin in albite (cf. Fig. 7SB), ordi-

twinning in accordance with the Carlsbad, Baveno, and Manebach laws common with orthoclase (pp. 370, 511). Twinning is also almost universal according to the albite law — twinning plane the brachypinacoid; this is usually polysynthetic, i.e. repeated in the form of thin lamellae, giving rise to fine striations on the basal cleavage surface (Figs. 735, 736). Twinning is also common according to the pericline law — twinning axis the macrodiagonal a ; when polysynthetic this gives another series of fine striations seen on the brachypinacoid.

The composition of this pericline twinning lamella plane passing through the crystal lattice such a direction that its intersections with the prismatic faces and the brachypinacoid make equal plane angles with each other. The position of this rhombic section and the consequent direction of the striations on the brachypinacoid change rapidly with a small variation in the angle γ . In general it may be said to be approximately parallel to the base, but in albite it is inclined backward (+, Figs. 737 and 739*) and in anorthite to the front (-, Fig. 738); for the intermediate species its position varies progressively with the composition. Thus for the angle between the trace of this plane on the brachypinacoid and the edge b/c , we have for Albite $+23^\circ$ to $+50^\circ$; for Oligoclase $+9^\circ$ to $+15^\circ$; for Andesine $+1^\circ$ to -2° ; for Labradorite -9° to -10° ; (or Anorthite -15° to -17°).

737. 738. 739.



practical form showing

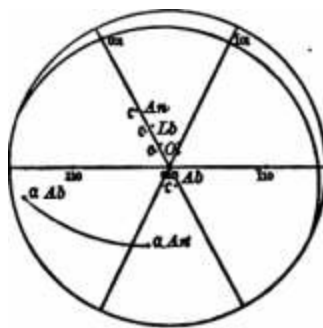
is often lettered as follows: $c(001) = i$, $b(010)$

If the composition-plane is M right angles to the twinning ui^* in the pericline twinning, the polysynthetic lamellae then show prominently in a basal section, together with those due to the albite twinning. Hence the grating structure characteristic of microcline.

Optical Characters. There is also a progressive change in the position of the ether-axes and the optic axial plane in passing from albite to anorthite. 740. This is most simply exhibited by the position

of the planes of light-vibration, as observed in sections parallel to the two cleavages, basal c and clinopinacoidal b , in other words the extinction-angle formed on each face with the edge b/c (cf. Fig. 739). The approximate position of the ether-axes for the different feldspars is shown in Fig. 740 (Schuster). The axis c does not vary very much from the zone bc , but the axis a varies widely, and hence the axial plane has an entirely different position in albite from what it has in anorthite. Further albite is optically positive, that is $c = Bx$, while anorthite is negative or $Q = Bx$; for certain andesines the axial angle is sensibly 90° ,

The following table gives the percentage composition of the various molecular compounds of albite and anorthite, with the calculated specific gravity (Tschermak), and also certain of the optical characters connected with them. These latter values were calculated by Schuster from an equation deduced by Mallard, in which certain observed values were assumed as fundamental :*



DIS.—In rock sections the plagioclase feldspars are distinguished by their lack of color, low refractive relief, and low interference-colors, which in good sections are mainly clear gray and scarcely rise into white of the first order; also by their biaxial character and converging light. In the majority of cases they are easily told by the parallel bands or fine lamellae which pass through them due to the multiple twinning according to the albite law; one set of bands or twin lamellae exhibits in general a different interference-color from the other (cf. Figs. 735, 786). They are thus distinguished not only from quartz and

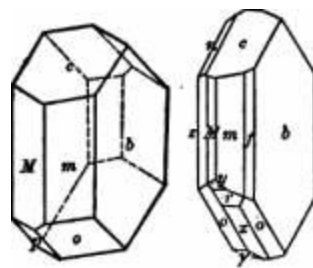
* See Tschermak. Ber. Ak. Wien, 60 (I). 66.1890; Schuster, *Mittb.*, 3. 117, 1881. 6. 189, 1883; Mallard. Bull. Soc. Min., 4, 96, 1881. Also Michel-Lévy and other authors referred to on p. 212 further, O. F. Becler, *Am. J. Sc.* 6, p. 819, May 1898.

SILICATES.

orthoclase, with which they are often associated, but from all the common rock-making minerals. To distinguish these species and from one another, as albite from Labradorite or Anorthite, is more difficult. In sections showing a detrital outcrop (see (001) and (010)) this can generally be done by determining the extinction (see p. 118 and Fig. 789). In general in rock sections special methods are required: these are discussed by the authors referred to in the note on p. 876.

0-5577; $\alpha = 94^\circ 3'$, $\beta = 90^\circ 29'$,

741.



ALBITE.

Triclinic Axes $a : b : c = 1.05335 : 1$

$\beta = 88^\circ 9'$.

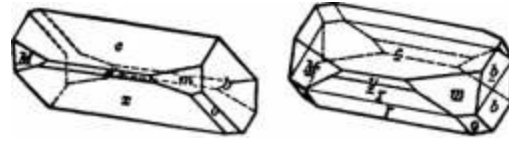
Albite, $010 \wedge 001 = 86^\circ 54'$. $m \wedge 110 \wedge 010 = 69^\circ 11'$. $m \wedge 110 \wedge 010 = 60^\circ$ approx. $em, 001 \wedge 110 - 001 = 68^\circ 10'$. «. $001 \wedge 101 = 58^\circ 16'$. - Twins with orthoclase; albite law (p. 375), usually; contact-twins, and polysynthetic, consisting of thin lamellae and with consequent fine striations on c; tw. albite law, peridine law, contact-twins whose composition-face is the rhombic section (Figs. 737 and 745); often polysynthetic and showing fine striations which on b are inclined backward -| - 22 to the edge b/c.

Crystals often tabular || b; also elongated in a, as in the variety peridine. Also massive, either lamellar or granular; the laminae often curved, sometimes divergent; granular varieties occasionally quite fine to impalpable. Cleavage: c perfect; b somewhat less so; m imperfect. Fracture uneven to conchoidal. Brittle. H. = 6-6.5. G. = 2.62-2.65. Luster vitreous; on a cleavage surface often pearly. Color white; also occasionally bluish, gray, reddish, greenish, and green; sometimes having a bluish opalescence or play of colors on c. Streak uncolored. Transparent to subtranslucent.

Optically+. Plane (S) \perp to Bx., inclined 100° to 102° to the acute edge ft/c. Extinction-

angle with edge $b/c = + 4^\circ 30'$ to 3° on c , and $= + 20^\circ$ to 15° on b (Fig. 737). Dispersion for B_x $p < v$; also inclined, horizontal; for B_i , $P > "$; inclined, crossed, $D_k. 2H., = 80^\circ$ to 84° D_x . Birefrin-

744.



, $NaAlSi_3O_8$ or $KAl_3Si_3O_{10}(OH)_2$,

Pericline. gence weak, $y - a = 0.008$,

Comp.—A silicate of aluminum and sodium $6SiO_2 =$ Silica 68.7, alumina 19.5, soda $11S = 100$. Calcium is usually present in small amount, as anorthite $\{CaAl_2Si_2O_8\}$, and as this increases it graduates through oligoclase-albite to oligoclase (cf. p. 376). Potassium may also be present, and it is then connected with anorthoclase and microcline.

Var.— Ordinary. In crystals and massive. The crystals often (nubular | 6. The massive forms are usually nearly pure white, and often show wavy or curved laminae. Ferriite is a whitish adularia-like albite, slightly iridescent, named from the French pigeon. Awn-lurin* and maontione varieties also occur. Pericline from the chloritic schists of the Alps

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DESCRIPTION OF MINERALOGY.

Characteristics, as shown in Figs. 744 and 740, and commonly identified with this in the twinning axis (pericline law).

Pyro., etc.—B.B. fuses at 4 to a colorless or white glass, imparting an intense yellow to the flame. Not acted upon by acids.

Diff.—Resembles albite in some forms, but is harder and of lower specific gravity; does not effervesce with acid (like albite). Distinguished optically and by the characteristic twinning relations on e from orthoclase; from the other triclinic feldspars partially by specific gravity and better by optical means (see p. 370).

Occurrence.—Albite is a constituent of many igneous rocks, especially those of alkaline type, as granite, leucite-syenite, diorite, etc.; also in the corresponding feldspathic lavas. In part it is interlamated with orthoclase or microcline, and similar aggregations, often on a microscopic scale, are common in many rocks. Albite is common also in the gneisses, and sometimes in the crystalline schists. Veins of albite granite are often repositories of the rarer minerals and of fine crystallizations of gems, including beryl.

tourmaline, allanite, columbite, etc: It is found in disseminated crystals in granular limestone.

In the granite smoky quartz, in the position

upon the orthoclase. Thus in the St. Gotthard region; Roc Tourge near Modane. Savoy; on Mt. Skopje (penⁱ7M): Tavetschthal in Tyrol; also Fflach, Kauris, the Zillertal, Kriml, Schneeberg in Paaseir in simple crystals. Also in Dsupting in similar association; Elba. Also Hirschberg in Silesia; Penig in Saxony; with topaz at Muislnka in the Ural and near Mlask in the Ilmen Mts.; Cornwall, England; Mourne Hts. in Ireland.

In the United States, in Maine, at Paris, with red and blue tourmalines. In Mat^{*}, at Chesterfield in lamellar masses (faintly indistinct), slightly bluish, also fine granular. In New Samp., at Acworth and Alsiead, in Gonn., at Hsddnm; at the Hadletown feldspar quarry; at Branchville, in fine crystals and massive. In N. York, at Moriah, Essex Co., of a greenish color. In Penn., at Unlonville, Chester Co, in Virginia, at the mica mines near Amelia Court-House in splendid crystallization. In Colorado, in the Pike's Peak region with smoky quartz and amazon-stone.

The name is derived from the color, white, in allusion to its common color.

Orthoclase.

Triclinic. See p. 369. $a:b:c = 86:32:100$. Twins observed according to the Carlsbad, albite, and pericline laws. Crystals not common. Usually massive, cleavable to compact.

Cleavage: c perfect; & somewhat less so. Fracture conchoidal to uneven. Brittle. H. = 6-6.5. G. = 2.65-2.67. Luster vitreous to somewhat pearly or waxy. Color usually white, with a faint tinge of grayish green, grayish white, reddish white, greenish, reddish; sometimes aventurine. Transparent, translucent. Optical characters, see p. 376.

Comp., Tar.—Intermediate between albite and anorthite and corresponding to Ab.An, to AbjAn., but chiefly to AbjAn., p. 376.

Var.—1. Ordinary. In crystals or more commonly massive, cleavable. The varieties containing soda up to 10 p. c. are called oligoclase-albite. 2. Aventurine oligoclase, or sun-tinted. is of a grayish-white to reddish-gray color, usually the latter, with internal yellowish or reddish fire-like reflections proceeding from disseminated crystals of probably either hematite or goethite.

Properties, etc.—B.B. fuses at 35 to a clear or foamy glass. Not materially acted upon by acids.

Dia.—See orthoclase (p. 873) and albite (p. 877): also pp. 870, 378.

Obs.—Occurs in porphyry, granite, syenite, and also in different effusive rocks, as andesite. It is sometimes associated with orthoclase in granite, or other granite-like rock. Among U.S. localities are Davids-Zoll near Stockholm; Pargas in Finland; Baltansk, Urut; In syenite of the Vosges: at Albi in the Grisons; Harienbad, Bohemia; Chalanches in Auvergne, and Bourg d'Oisans, Dauphiné; as titanite* at Tvedestrand, Norway; at Hitter; Lake Baikal.

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In the United States, at Fine and Macomb, St. Lawrence Co., N. Y., in good crystals at Danbury, Ct., with orthoclase and quartz; Usdillam, Ct.; at the emery mine, Chester, Mass., granular; at Unionville, Pa., with epidote and corundum; Mineral Hill, Delaware Co.; at Bakerville, N. C., in clear masses, with cleavage but no twinning. Named in 1816 by Breithaupt from okyoi, litit, and Kaldai, fracture.

Andesine.

Triclinic Axes, see p. 369. $a : b : c = 86^\circ 14'$. Twins as with albite. Crystals rare. Usually massive cleavable or granular.

Cleavage: c perfect; b less so; also M sometimes observed. H. = 5-6. G. = 3.68-3.69. Color white, gray, greenish, yellowish, flesh-red, lustrous to pearly. Optical characters, see p. 376.

Comp.—Intermediate between albite and anorthite, corresponding to Ab: An in the ratio of 3:2, 4:3 to 1:1, see p. 376.

Pyr., etc.—Fuses before the blowpipe. Imperfectly soluble in acids. Obs.—Observed in many granular and volcanic rocks; occurs* in the Aoudou, at Harmaio, in an ingredient of the rock called andesite; in the porphyry of l'Esterel, Dept. du Var, France. In the syenite of Alsace in the Vosges: at Vapoeifjord, Iceland; Bodenmais, Bavaria. Baaford, He., with vesuvianite. Common in the igneous rocks of the Rocky Mts.

Xiabradotite*, Labrador Feldspar.

Triclinic, form near that of andesine, but not accurately known (p. 369). Cleavage angle $c : b = 86^\circ 4'$. Form and twinning similar to the other plagioclase species. Crystals often very thin tabular, and rhombic in outline bounded by c or ex (Fig. 435, p. 131). Also massive, cleavable or granular; sometimes cryptocrystalline or hornstone-like.

Cleavage: c perfect; b less so; M (110) sometimes distinct. H. = 5-6, G. = 2.70-2.72. Luster on c pearly, passing into vitreous; elsewhere vitreous or resinous. Color gray, brown, or greenish; sometimes colorless and glassy; rarely porcelain-white; may have a beautiful change of colors in cleavable varieties, especially in h. Streak uncolored. Translucent to translucent. Optical characters, see p. 376.

Piny of colors a common character, but sometimes wanting as in some colorless crystals. Blue and green are the principal colors; but yellow, fire-red, and pearly-gray also occur. Vogelsack regards the common blue color of labradorite as an interference-phenomenon due to its lamellar structure, while the golden or reddish schiller, with the other colors, is due to the presence of black adular microlites and yellowish-red microscopic inclusions, or to the combined effect of these with the blue reflections. Schrauf has examined the inclusions, their position, etc., and given the axes micropolarization and birefringence to two groups of them. (See references on p. 142.)

Comp., Tar.—Intermediate between albite and anorthite and corresponding chiefly to Ab : An in a ratio of from 1 : 1 to 1 : 3, p. 376.

The feldspars which lie between labradorite proper and anorthite have been embraced by Tschermak under the name bytownite. The original bytownite of Thomson was a greenish-white feldspathic mineral found in a boulder near Bytown (now Ottawa) in Ontario, Canada.

Pyr., etc.—B.B. fuses at 8 to a colorless gloss. Decomposed with difficulty by hydrochloric acid, generally leaving a portion of undecomposed mineral.

DiS.—The beautiful play of colors is a common but not universal character. Otherwise distinct varieties are the other feldspars (pp. 870, 871).

Obi.—Labradorite is an essential constituent of various igneous rocks, especially of the basic kinds, and usually associated with some member of the pyroxene or amphibole groups. Thus with hypersthene in gabbro, with diorite in gneiss, with some form of pyroxene in diabase, basalt, dolerite, also andesite, tephrite, etc. Labradorite also occurs in other

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kind of Uvni, and is sometimes found in them in glassy crystals, as in those of Etna, Vesuvius, the Siudwicks Islands at Eilautra.

The labradoritic massive rocks are most common among the formations of the Archean era. Such are part of those of British America, northern New York, Pennsylvania, Arkansas; those of Greenland, Norway, Finland, Sweden, and probably in the Vosges.

On the coast of Labrador, labradorite is associated with hornblende, hypersthene, and microcline. It is met with in many places in Quebec. Occurs abundantly throughout the central Adirondack region in northern New York; in the Wichita Mts., Arkansas.

Labradorite was first brought from the Isthmus of Paul, on the coast of Labrador, by Mr. Wolfe, a Moravian missionary, about the year 1770.

HAflKBLTMTs. In cotorlesB isotropic graius In meteorttei; composition near labradorite.

ANOSTHTTB. lodiaDiie.

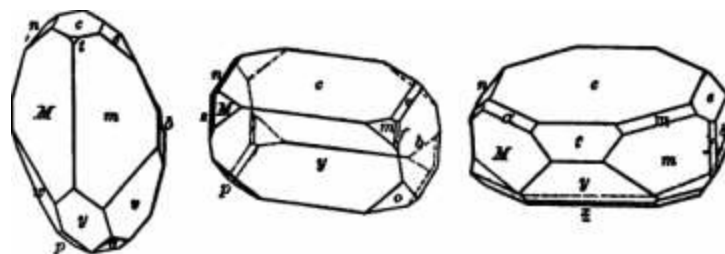
Triclinic. Axes $S : B : i = 0.6347 : 1 : 0.5501$; $\alpha = 93^\circ 13'$, $\beta = 115^\circ 55'$, $\gamma = 91^\circ 13'$.

be. $010 \wedge 001 = 86^\circ$ Bff. cm. $001 \wedge 110 = 86^\circ$ Sy.

mM, no $\wedge 110 = 60^\circ$ 2V. eM, $001 \wedge 110 = 69^\circ$ W.

bm, $010 \wedge 110 = 68^\circ 4'$. ey, $001 \wedge 201 = 81^\circ$ U'.

746. 747. 748.



Twins as with albite (p. 375 and p. 377). Crystals usually prismatic $\parallel c$ (746, also Fig. 338, p. 108), less often elongated $\parallel ft$, like pericline (Fig. 747). Also massive, cleavable, with granular or coarse lamellar structure.

Cleavage: c perfect; β somewhat less so. Fracture conchoidal to uneven. Brittle. $H. \wedge 6-6-5$. $G. = 2.74-2.76$. Color white, grayish, reddish. Streak uncolored. Transparent to translucent.

Optically —. Ax. pi. nearly $\perp c$ (021), and its trace inclined 60° to the edge $r./e$ from left above behind to right in front below. Extinction-angles on c , — 34° to — 42° with edge b/c ; on β , — 35° to — 43° (Fig. 739, p. 375). Dispersion $p < v$, also inclined, $2H_{\beta} = 84^\circ 50'$, Birefringence stronger than with albite, $\gamma - \alpha = 0.013$.

Comp.—A silicate of aluminium and calcium, $CaAl_2Si_2O_8$, or $CaO \cdot Al_2O_3 \cdot 2SiO_2$ = Silica 43.2, alumina 36.7, Hme 20.1 $\wedge 100$. Soda (as $NaAlSi_3O_8$) is usually

Present in small amount, and as it increases there is a gradual transition through townsite to labradorite. Var. — Anorthite was described from the gabbro crystals of Somma: anorthite and biotite are the same mineral. Titanite is the same from Iceland. Anorthite is a white, grayish, or reddish granular white from India, where it occurs as the gangue of corundum, first described in 1803 by Count Boisson, Cyrtopite occurs in small, Iran, and glassy crystals, coating cavities in the dolerite of the Cyclopean land and near Trezza on Etna. Amphibole, leucite, latroite also belong to anorthite.

Fyr., etc.—B.B. fuses at 5 to a colorless glass. Anorthite from Mte. Somma. and idanite from the Oarnattc, are decomposed by hydrochloric acid, with separation of gelatinous silts.

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SILICATES. 'iSl

Ob*.—Occurs in a few diorites; occasionally in gabbro and serpentine rocks; in some cases along with corundum; in many volcanic rocks, such as andesites, basalts, etc.; as a constituent of some igneous rocks (Juvénès, St. Quentin),

Anorthite (orthoclase and biotite) occurs in isolated blocks among the old lavas in the ravines of Monte Somma; in the Albani Hts.; on the Pesmedii Alp, Hohe Zoni, Tyrol, as a contact mineral: Aranyer Berg, Transylvania, in andesite; on Iceland; near Bogoslovsk in the Ural. In the Cyclopean Islands (eyeiopite). In the lava of the island of Miyake, Japan.

Anorthite was named in 1838 by Rose from CaSi_2O_7 , the crystallization being triclinic.

Orthoclase. KAlSi_3O_8 , similar to anorthite, but containing barium instead of calcium. Habit, with the usual cleavages (001) and (010), $\alpha = 89^\circ$ int. H. = 0-8-5. G. = 3-87. Colorless. Extinction: on c. V. Iff; on b, 45° . From Jakobsberg, Sweden.

n. Uetasioate. ESiO_2 .

Salts of Metasilicic Acid, H_2SiO_4 ; characterized by an oxygen ratio of 8 : 1 for silicon to bases. The Division closes with a number of species, in part of somewhat doubtful composition, forming a transition to the Orthosilicates.

The metasilicates include two prominent and well-characterized groups, viz., the Pyroxene Group and the Amphibole Group. There are also others less important.

Leucite Group. Isometric.

In several respects leucite is allied to the species of the Feldspar Group, which immediately precede.

Leucite KAlSi_3O_8 , Isometric at 500°

Pseudo-isometric at ordinary temperatures. Polnoite $\text{H,Cs,Al}_2(\text{Si}_2\text{O}_7)_2$, Isometric

LBOrra. Amphigene.

Isometric at 500° C; pseudo-isometric under ordinary conditions {see p. 230}. Commonly

in crystals varying in angle but little from the tetragonal trisoctahedron $n(311)$, sometimes with a (100) , and $d(749)$.

(110) as subordinate forms. Faces often showing fine striations due to twinning (Fig. 749). Also in disseminated grains; rarely massive granular.

Cleavage: $d(110)$ very imperfect. Fracture conchoidal. Brittle. $H. = 5.5-6$. $O. = 2.45-3-50$. Luster vitreous. Color white, ash-gray or smoke-gray. Streak uncolored. Translucent to opaque. Usually shows very feeble double refraction: $o_j = 1.508$, $e = 1.009$ (p. 230).

Comp. — $KAl(SiO_3)_2$, or $K_2O \cdot Al_2O_3 \cdot 4SiO_2 =$ Silica 55.0, alumina 23.5, potash 21.5 = 100.

Soda is present only in small quantities, unless as introduced by alteration; traces of lithium, also of rubidium and calcium, have been detected.

Pyr., etc.—B.B. infusible; with cobalt solution gives a blue color (aluminium). Decomposed by hydrochloric acid without gelatinization.

Diff.—Characterized by its trapezohedral form, absence of color, and infusibility. It is softer than garnet and harder than analcite; the latter yields water and fuses.

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DESCRIPTIVE MINERALOGY.

$n = 1.508$ ($n^e = 1.009$)

$d(749)$ not wholly isotropic and, further, shows complicated sets of twinning-lines (Fig. 749): the birefringence is, however, very low, and the colors scarcely rise above dark gray; they are best seen by introduction of the quartz or selenite plate yielding red of the first order. The smaller leucites, which lack this twinning or the inclusions, are only to be distinguished from sodalite or analcite by chemical tests.

Obs.—Leucite occurs only in igneous rocks, and especially in recent lavas, as one of the products of crystallization of magmas rich in potash and low in silica (for which reason this species rather than orthoclase is formed). The larger embedded crystals are commonly anisotropic and show twinning lamellae: the smaller ones, forming the groundmass, are isotropic and without twinning. Found in leucitites and leucite-basalts, leucitophyres, leucite. Eucrites and leucite-tephrites; also in certain rocks occurring in dikes. Very rare in intrusive rocks, only one or two instances being known; but its former presence under such conditions is indicated by pseudomorphs, often of large size, of nephelite and orthoclase, also of analcite.

The prominent localities are, first of all, Teufelsberg and Mt. Bonna, where it is thickly disseminated through the lava in grains, and in large perfect crystals; also in ejected

masses; also near Rome, at Capo dl Bove, Roccn MonOEia, etc. Fui-ther In leuclte-
tcphrlte at Proceoo near Lake Bolsena in central Italy; also about tbe Laacher See and at
several

Elnts in tbe Eifel; at Rleden near Andernacli; al Helches Id the VoKelsgebirge; in the
liserstuhlKeblrge. Occurs in Bnizll, at Pinbalzinho. From tbe Cerro de las Virginea, Lower
OatilornlL. In the United Slates It Is present in a rocli in the Qreen River Basin at tbe
Leucite Hills. Wyoming; also in the Absaroka range, in northwestern Wyoming: in the
Higliwood and Bearpaw Mis , Montana (in part pseudoleuclte). On the shores of Van-
couver island, where magnificent groups of crystals have Iteen found as drift boulders.

PBeudoleuclte(aee above) occurs In the phonolite (tlnguaite) of tie Beml de Tiogua,
Brazil; at Magnet Cove, Arkansas; near Hamburg, N. J.; Montana; also in the Cariboo
DIstr., British Columbia.

Named from Xfvx6!. lehite, in allusion to Its color.

Polluolte. Essentially $\text{HiO}_2\text{Cs}_2\text{Al}_2\text{BSiO}_2$. Isometric; often tn cubes; also massive. E.
= 6'5. Q. = 3'901. Colorless. Occui-s very sparingly in the island of Elba; with petalile
(castorite); also at Hebron and RumCord, Me.

Pyroxene Group.

Orthorfaombic, Monoclinic, Tridinio. Compoef 'ion for the most part that of a
metasilicate, BSiO_2 with $\text{K} = \text{C}$, Mg, Fe chiefly, also Mn, Zn. Farther RSiO_2 with
 $\text{R}(\text{Fe}, \text{Al}, \text{S})$, leas often containing alkalies (Ifa, K), and then HSiO_2 with $\text{RAl}(\text{SiO}_2)$,
Barely inclnd-ing zirconium and titaninm, also fluorine.

a. Orthorhombtc Section.

d'l id or h: S -t

Emtatite MgSiO_3 , 09702 : 1 : 0-6710 1-0307 : 1 : 0-5885

Bronzite $(\text{Mg}, \text{Fe})_2\text{SiO}_5$,

Hyperrthene $(\text{Fe}, \text{Mg})_2\text{SiO}_5$, 0-9713 : 1 : 0-5704 1-0319 : 1 : 0-5873

The second set of axial ratios, with $d = 1$, brings out tbe similarity o(the form to tbe
monoclinic spedea.

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fi. 3f onooHnie SecHon.

PTTOxem 1-0921:1 : 0-6893 74^ 10*

I. Nos-ALvumouB VAsaxws:

U&laolit«, Salite, Diallage, etc.

2. Hedenbbboite GaFe(SiO,),

Manganhedenbergfite Ca(Fe,Mii)(SiO,),

3. SCHEFPEEiTB (Ca,Mg){Fe,Mn}(8iO,),

Jeffewonite (Cs,Mg)(Fe,Mn^n)(SiO,),

II. Aluminovs Vasixtixb:

4. AUQITE

(C«f MK,Fe){SiO.}.

1 with (Mg.Fe)(Ai>e).8i

,8iO, Leacangite, Faaaaite, Angite, ^girate-angite.

A: 6:6 /3

» (^girate) NaFe(SiO,). 1-0996 : 1: 06013 73° 11'

Bpodnmeoe LiAl(SiO,), 1*1238:1 : 0*6355 69" 40'

Jadaito NaAl(SiO.). 1-103 ! X: 0*613 72° 44i'

d:t:i /S

Wollartonito CaSiOi 1*0531: 1 :0-9676 84° 30'

FMtdite HNaCa,(SiO.). 11140 : 1: 0*9864 84° 40'

;'. TrteUnie Section.

a: ?: (! a o y

Rhodonite MnSiO. 1-0729 : 1 : 06213 103° 18' 108° 44' 81^39'

also (Mii,Ca)SiO. (Mn,Fe)8iO. (Mn,Zn,Fe,Ca)8iO. Babingtonite 1-0691 :1: 0*6308 104° 21^' 108° 31' 83° 34'

(Ca,Fe,Mn)SiO,.Fe.(SiO,),

The Pyroxene Group embraces a number of species which, while falling in different systems—orthorhombic, monoclinic, and triclinic—are yet closely related in form. They all have a fundamental prism with an angle of 93° and 87° , parallel to which there is more or less distinct cleavage. Further, the angles in other prominent zones show a considerable degree of similarity. In composition the metasilicates of calcium, magnesium, and ferrous iron are most prominent, while compounds of the form $E(\text{Al,Fe})_2\text{SiO}_5$, $\text{RAl}(\text{SiO}_3)_2$, are also important.

The species of the pyroxene group are closely related in composition to the corresponding species of the amphibole group, which also embraces members in the orthorhombic, monoclinic, and triclinic systems. In a number of cases the same chemical compound appears in each group; furthermore, a change

=:i,:cc ..Google

by paramorphism of pyroxene to amphibole is often observed. In fact, also

the two groups are related, as shown in the axial ratio; also in the parallelism

of the optical axes of monoclinic amphibole upon or about those of pyroxene

(Fig. 430, p. 131). The axial ratios for the typical monoclinic species are;

Pyroxene $d : i : c = 1.0921 : 1 : 0.5893 / S = 74^\circ 10'$

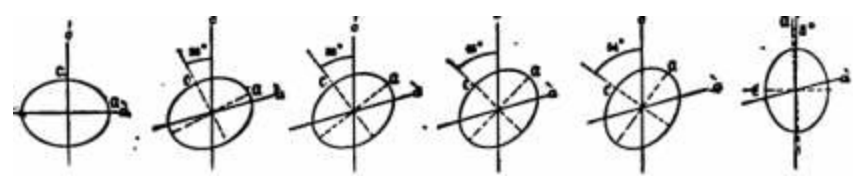
Amphibole $A : i : c = 1.1022 : 1 : 0.5875 / S = 73^\circ 58'$

See further on p. 398.

The optical relations of the prominent members of the Pyroxene Group, especially as regards the connection between the position of the optical axes and the crystallographic axes are exemplified in the following figures (Cross). 7B1.

in. m. IV. V. TL

4-



A corresponding group,

the corresponding exhibition of the prominent amphiboles is given under that, Fig. 785, p. 398.

. Orthorhombic Section.

SHSTATITB.

Orthorhombic. Axes $d : I : 78fl.$

$0-9702 : 1 : 05710. 110 A \text{ liO} = 88^\circ 16'. rr'. 22S A 323 = 40' 181.$

$= as-16-.$

$= 41'' 41'.$

II.

Twins rare: tw. pi. h (014) as twinning lamellae; also tw. pi. (101) as stellate twins crossing at angles of nearly 60° , sometimes six-rayed. Distinct crystals rare, habit prismatic. Usually massive, fibrous, or lamellar. Cleavage: m rather easy. Parting! 6; \wedge Isoo. Fracture uneven. Brittle. H. = 5.5. G. = 3.1-3.3. Luster. Bitter. a little pearly on cleavage-surfaces to vitreous; often

metalloidal in the bronzite variety. Color grayish, yellowish or greenish white, to olive-green and brown. Streak uncolored, grayish. Translucent to nearly opaque. Pleochroism weak, more marked in varieties relatively rich in iron. Optically+. Ax.-pl. Ijfi. BxJ_c. Dispersion $p < t$ weak. Axial angle large and variable, increasing with the amount of iron, usually about 90° for FeO = 10 p. c. $\rho_{sj} = 1.669$; $\gamma - \alpha = 0.009$.

Comp., Tar.— $MgSiO_3$, or $MgO \cdot SiO_2$, = Silica 60, magnesia 40 = 100. Also $(Mg,Fe)SiO_3$, with Mg : Fe = 8 : 1, 6 : 1, 3 : 1, etc.

Var.—1. With little or no iron; Bitter. Color white, yellowish, grayish, or greenish white; luster vitreous to pearly; G. = 3.1-3.3. Chladni & (Blepharidolite of Rose), which

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./

EuBitter (Brookite). Hyperathene.

2, Ferriferous; Bitter. Color grayish green to olive-green. Habit fibrous, lamellar, or prismatic. Cleavage: m rather easy. Parting: uib. however, usually secondary origin and is not essential. With the increase of FeO (above 12 to 14 p. c.) the color passes to hypersthene, the optic axial angle changing so that in the latter $\alpha = Bk$, and Bx, l a. Table of optical data by Figs. 753, TS4.

758.

7BS.

7B7.

Pyr.. etc.—B.B. almost infusible. belongs only slightly rounded on the thin edges; P. = 6.
Insoluble in hydrochloric acid.

Obs.—Essentially (incl. bronzite) a common constituent of peridotite and the ferropentlandites derived from them; it also occurs in crystalline schists. It is often associated in parallel growth with orthoclinic pyroxene, e.g., diagenite (Figs. 755, T5S). A common mineral in meteoric iron often occurs in chondrites with eccentric radiated structure (Fig. 7S7).

Occurs near Alovsholm in Horavln. in serpentine; at Kupferberg in Bohemia; at Basle in the Harz (proterozoic): in the so-called olivine bombs of the Dreiser Wehcr in the Elbe; in immense crystals, in part altered, at the apatite deposits of Kjørrstad near Bamle, Norway; in the peridotite associated with the diamond deposits at South Africa,

In the U. S., in New York at the Tilly Foster magnetic mine, Brewster, Putnam Co., with clonrodite; near Edwain, N. Y.

Named from *eyardrrf*?, an opponent, because so refractory. The name *brontite* has priority, but a bronze luster is not essential, and is far from universal.

BTPBKBTHBNB.

Orthorhombic. Axes (a:b:c) = 0.9713 : 1 : 0.5704.

$\alpha = 110^\circ$. $\alpha_1 \text{ to } \alpha_4 = 88^\circ 20'$. $\alpha_5 = 111^\circ$. $\alpha_6 = 58^\circ$. $\alpha_7 = 58^\circ$.

$\beta = 16^\circ$. $\beta_1 = 16^\circ$. $\beta_2 = 16^\circ$. $\beta_3 = 16^\circ$. $\beta_4 = 16^\circ$. $\beta_5 = 16^\circ$. $\beta_6 = 16^\circ$. $\beta_7 = 16^\circ$.

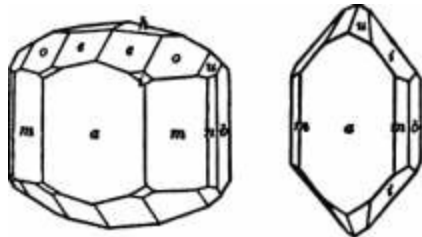
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Crystals rare, habit prismatic, often tabular on a, less often on b. Usually foliated massive; sometimes in embedded spherical forms.

Cleavage : b perfect; m and a distinct but interrupted. Fracture uneven. Brittle. H. = 5-6. G. = 3.40-3.50. Luster somewhat pearly on a cleavage-

769.

760.



surface, and sometimes metalloidal. Color dark brownish green, grayish black, greenish black, pinchbeck-brown. Streak grayish, brownish gray. Translucent to nearly opaque. Pleochroism often strong, especially in the kinds with high iron percentage; thus a or d brownish red, f or J reddish yellow, c or (\wedge) green. Optically —. Ax. pi. || 6. Bx J.*!, Dispersion $p > v$. Axial angle rather large and variable, diminishing with increase of iron, cf. enstatite, p. 384, and Figs. 753, 754, p. 385, ($i = 1-708$; $y - a = 0.013$).

Hypersthene often occurs in minute tabular scales, usually of a brown color, arranged mostly parallel to the basal plane (Fig. 700). Also less frequently vertical or oblique 30° to t ; they may be brookite (Gottschalk, Gemm.), but their true nature is doubtful. They are the cause of the peculiar metalloidal luster or sheen, and are often of secondary origin, being developed along the so-called "solution-planes" (p. 149).

Comp. Form.— $(Fe, Mg)SiO_3$, with $Fe:Mg = 1:3$ ($FeO = 16.7$ p. c.), $1:3$ ($FeO = 21.7$ p. c.) to nearly $1:1$ ($FeO = 51.0$ p. c.). Alumina is sometimes present (up to 10 p. c.) and the composition then approximates to the aluminous pyroxenes.

Of the orthorhombic magnesian-trochiloides. Those with $FeO > 12$ to 18 p. c. are usually to be classed with hypersthene, which is further characterized by being optically negative and having $D_p > t$.

Fuz., etc.—B.B. fuses to a black enamel, and on charcoal yields* a magnetic mass; fused more easily with increasing amount of iron. Partially decomposed by hydro-chloric acid.

Ob*.—Hypersthene, associated with a clinoclinal feldspar (labradorite), is common in certain granular eruptive rocks, as noritic hyperite, gabbro, also in basic andesite (hyperthene-Uranium), a rock recently shown to occur rather extensively in widely distributed regions.

It occurs at Isle St. Paul; Labrador; in Greenland; at Farsund and elsewhere in Norway; Elfdalen in Sweden; Penng in Saxony; Kiiuaberg in Bohemia; the Tyrol; Neurode in Silesia; Bodenmau, Bavaria. Amphibole is from the Laacher See. Siubole occurs with pseudobrookite and tridymite, in cavities in the aegirine of the Araucario Berg, Transylvania, and elsewhere.

Occurs in the noritic of the Cortlandt region on the Hudson river, N. Y.; also common with labradorite in the Adirondack Archæan region of northern New York and northward in Canada. In the basaltic-andesites of Mt. Shasta, California; Buffalo Peaks, Colo-

rado, ond other poins,

Hyperithtne is named from vitip and trOfyoi, wry tirong, or tough.

Bastite. i>r Schiller Si'ab. An nlti'red enaiallte (or brouzile) hiiving approximately
inular eruptivo

tbe composilioD of serpviiit

1 foliiti'd form hi ct'tliiin granui

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rocks ftod I* cbaracteiiized by a bronse-Hke raetaloldal luster or scblller on the cblef
cleaTSge'face (b), nhlcli "schllleizaiiuii" {p. 100) Ib of secondary origin. H. = 8'S-4. o. = 3-
5-a'7. Color leek-green lo oliva- and pUtacbio-green, and piiiclibeck brown, Pleo-cbroiam
not marked. Optically —. Double refraction weak. Ax. pi. | a (hence normal to tliat of
eusiaiite). Bx l b. Diaperalou p > s. Tlie oHgiual bastite was from Baste near Harzhurg in
the Harz; also fi-om Todtmoos in the Schwurzwald.

PacKHAMiTK. $8(\text{Mg,Fe})8\text{iO}_i.(\text{Mx.I}^e)8\text{iO}_..$ Occurs In rouded nodulea In the meteorite
4>t Eilherville, Emmet Co., Iowa, Mny 10, 187Q. O. = 3St8. Color light greenish yellow.

PTROXEMB.

Honocli^io.

/S. Monocttnie Section,

Axes d:\$:6~ 10921:1: 0-5893; fl = $74^\circ 10'$. 001 A asi = $48''$ W.

; 110 a iio = »a"60'.

001 A 100 = 74° Itf. 001 A iOl = 81° W. on A oil = $89^\circ 6'$.

oai Aoi = $97'11'$.

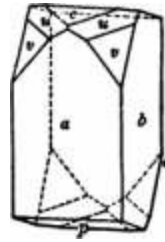
001 A 111 =

;° 491'.

em, 001 A 110 = TO'9^'. «, 001 A ill = $42''$ y. uu'. III A lil = $48^\circ a$ »'. <•', ill A ill = $69^\circ 11'$. BO',
531 A 2ai = $84^\circ 11'$.

^

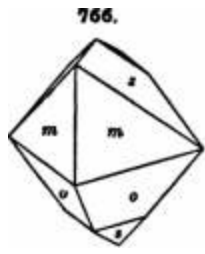
764.



f^^

MJccJ'

Twine: t-v. pi. (1) a, contact-twins, common (Fig. 769), sometimes poly-synthetic (2) G, as twinning lamelle producing striations on the vertical faces and pseudo-cleavage or parting $l_j c$ (Fig. 770); very common, often secondary. (3) $tf (101)$ cruciform-twins, not common (Fig. 421, p. 130). (4) W (133) the vertical axes crossing at angles of nearly 60° ; sometimes repeated as a six-rayed star (Fig. 420, p. 130). Crj^{\wedge} stala usually prismatic in habit, often short and thick, and either a square prism (a, i prominent), or nearly square



^(^^

\^ K^ ^^

,<^^

($93^\circ, 87^\circ$) with m predominating; sometimes a nearly symmetrical S-sided

form with a, b, m (Fig. 770). Often coarsely lamellar, $J c$ or a. Also granular, coarse or fine; rarely fibrous or columnar.

Cleavage: m sometimes rather perfect, but interrupted, often only observed



green, as i:

StW DESCRIPTIVE MINERALOGY.

in thin sections X i (Fig. 771). Parting I c, due to twinning, often prominent, especially in large crystals and lamellar masses (Fig. 7TU); also {1} a less distinct and not so common. Fracture uneven to conchoidal. Brittle. H. = 5-6, G. ~ 3.2-3.6, varying with the composition. Luster vitreous inclining to resinous; often dull; sometimes pearly I c in kinds showing parting. Color usually green of various dull shades, varying from nearly colorless, white, or grayish white to brown and black; rarely bright in kinds containing chromium; also blue. Streak white to gray and

of a bluish green. Transparent to opaque. Pleochroism usually weak, even in dark-colored varieties; sometimes marked, especially in violet-brown kinds containing titanium.

Optically +, Birefringence strong, $(y - a) = 0.02 - 0.03$. Ax. pi. | b. Bx, Ai = CA₂— + Z₆" in diopside, to $-j - 52^\circ$ in augite (wh. see), or $ce = 20^\circ$ to 36° , the angle in general increasing with amount of iron. For diopside from Ala (Dx.): $2E, = 111^\circ 20'$. Also (Flink) for white diopside from Nordmark (2-49 p. c. PeO):

Bx.Ai $2V, 2V, 3V^{\wedge} / J, / S, / S,$

+ $38^\circ 81' SB'ff 68^\circ S3' 68^\circ 40''$ IMBTS 169359 r69869

Comp., Tar.—For the most part a normal calcium and magnesium, also iron, less often manganese and zinc. The alkali metals potassium and sodium present rarely, except in very small amount. Also in certain varieties containing the trivalent metals aluminium, ferric iron, and manganese. These last varieties may be most simply considered as molecular compounds of $Oa(Mg,Fe)Si_2O_6$, and $(Mg,Fe)_2Al_2Si_2O_{10}$, as suggested by Tsdiernak. Chromium is sometimes present in small amount; also titanium replacing silicon.

The name Pyroxene is from *ffOo. firt*, and *ieyoi*, »tranger. and records UaQy's idea that the mineral was. as he expressed it. "h atranKer in the ilomiiu of Sre," whereao. jn fact. il Is, ni-xt

to the feldspar, the most universal constituent "

of igneous rocks.

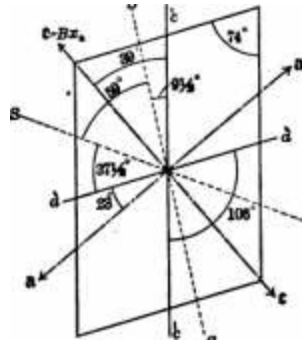
The varieties are numerous and depend upon variations in composition chiefly; the more prominent of the varieties properly rank as sub-species.

I. Containing little or no Aluminium.

1. Diopside. Malacolite, Alaiite. Calcium-magnesium pyroxene. Formula $CaMg(SiO_3)_2 =$ Silica 51.6, H₂O 25.9, magnesia 18.5 = 100. Color white, Yellowish, grayish white to pale

green, and finally to dark green and nearly black; sometimes transparent and colorless, also rarely a fine blue. In pris-

meta-illite, $RSiO_3$, chiefly of



by Google

matic crystals, often slender; also granular and columnar to lamellar massive. $\alpha = 3-2-3-36$. Bx. A (i =, + 36° and upward. $y - a = 0-03$. Iron is present usually in small amount as noted below, and the amount increases as it graduates toward true hedenbergite.

The following belong here: Chrome-diopside, containing chromium (1 to 28 p. c CrO_3), often a bright green.

Malacolite, as originally (Leikibe). was a pink-colored variety from Sala, Sweden.

Actinolite occurs in broad flat-angled prisms, colorless to faint greenish or clear green, from the Hiesu Alp in the A^ valley. Piedmont.

Tranterite, from Traversella. Is similar.

Manila is a fine blue diopside from 81. Marcel, Piedmont, Italy; occurring in prismatic crystals and massive.

Canaanite is a grayish-white or bluish-white pyroxene rock occurring with dolomite at Canaan. Conn.

Lavniite is a pyroxene, colored green by vanadium, from the neighborhood of Lake Baikal, in eastern Siberia.

Dioptase is named from Sii, twice or double, and iiii, appearance. Malacolite is from /laXatedi, because softer than feldspar, which is associated.

2, Hedenbergite. Calcium-iron pyroxene. Formula $CaFe(SiO_3)_2$, = Silica 48.4, iron protoxide 29.4, lime 23.3 = 100. Color black. In crystals, and also lamellar massive. $G = 3'5-3-58$. Bi, $A^\wedge = + 48^\circ$. Manganese is present in manganese hedenbergite to 65 p. c. Color

grayish green. G. = 355.

Between the two endmembers, diopside and edenite, there are numerous transitions conforming to the formula $\text{Ca}(\text{Fe},\text{Mg})\text{Si}_2\text{O}_6$. As the amount of Iron increases the color changes from light to dark green to nearly black, the specific gravity increases from 3.10 to 3.6, and the angle $2V_x$ also from 56° to 48° .

The following are varieties, coming under these two sub-species, based in part upon structure. in their peculiarities of composition.

Saiite (Bahliie). color grayish green to deep green and black; sometimes grayish and yellowish white; in crystals; also lamellar (parting | e) and granular massive; from Sala In Sweden, Baikalite, a dark dingy green variety, in crystals, with parting | e; from Lake Baikal. In Siberia.

Coccoite is a granular variety, embedded in calcite, also forming loosely coherent to compact aggregates; color varying from white to pale green to dark green, and then containing considerable iron; the latter the original coccolite. Named from *м'уВк'у!*, a *^'iii'n*.

DiALLAOB. A lamellar or thin-foliated pyroxene, characterized by a fine lamellar luster and parting | a, with also parting | b. and less often | e. Also a fibrous structure (Twinning) a. often polysynthetic; interlaminated with an orthorhombic pyroxene (Fif[^]. 755 and 756. p. 886). Color grayish green to bright grassgreen, and deep green; also brown. Luster of silicate a often pearly, sometimes metallic or exhibiting Schiller and resembling bronzite, from the presence of microscopic inclusions of secondary origin. $2V_x \approx 2V_a \approx 88^\circ$ to 40° ; $n_x - n_y = 0.004$. H. = 4; G. = 3.2-3.5. In composition near diopside, but often containing alumina and sometimes in considerable amount, then properly to be classed with the augites. Often changed to amphibole, see smaragdite. p. 401, and urallite. p. 401. Named from Scharfstein, difference, in allusion to the dissimilar planes of fracture. This is the characteristic pyroxene of gabbro, and other related rocks.

Omphacite. The granular to foliated pyroxenic constituent of the garnet-rock called eclogite, often interlaminated with amphibole (smaragdite); color gross-green. Contains some Al_2O_3 .

3, Scheffelite. A manganese pyroxene, sometimes also containing much iron. Color brown to black.

In crystals, sometimes tabular | e, also with p (101) prominent, more often elongated in the direction of the c-axis: p (101), rarely prismatic, j i. Twins, with a as tw. pi. very common. Also crystalline, massive. Cleavage prismatic, very distinct. Color yellowish brown to reddish brown; also black (iron-scheffelite). Optically +. $2V_x = 2V_a = 44^\circ$ to 25° . The iron-scheffelite from Pajsberg is black in color and has $2V_x = 49^\circ$ to 59° for

different zirconia in the same crystal. The brown iron-schefferite (urbanite) from L&ogban
baa { $a = 69^\circ 8'$, It resembles garnet in appearance.

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It is a manganese-zinc pyroxene from PraokllD FuroBce, N. J. (but the due may be due to impurity). In large, columnar crystals with edge rounded and faces unexposed. Color greenish black, on the exposed surface chocolate brown.

II. Aluminous.

4. AUOITE, Aluminous pyroxene. Composition chiefly $\text{CaMgSi}_2\text{O}_6$, with $-(\text{Mg}, \text{Fe})$ (Al, Fe), SiO_2 and occasionally also containing alkalis and then grading toward tegirite. Titanium is also sometimes present. Here belong:

a, L&ogbanite. Color white or grayish. Contains alumina, with lime and sometimes iron and little or no iron. Looks like diopside. $H. = 6-5$; $Q. = 319$. Named from Xevicoi,

b. F&ogbanite. Includes the pale to dark, sometimes deep green crystals, or pale green and blue resembling epidote. (Fig. 776). The blue kinds of diopside also belong here. Named from the locality in the Fassalliul, Tyrol. $\alpha = 60^\circ$, $\beta = 53^\circ$, $\gamma = 90^\circ$. TiO_2 is present in some kinds, which are then pleochroic. Named from avy^, luilrr.

e. AUOITE. Includes the greenish or brownish black and black kinds, occurring mostly in eruptive rocks. It is usually in short prismatic crystals, thick and stout, or tabular; often twins (Fig. 767-769). Ferric iron is here present, in relatively large amount, and the angle β , α becomes $+ 60^\circ$ to 53° . $\alpha = 60^\circ$; $\beta = 53^\circ$; $\gamma = 90^\circ$. TiO_2 is present in some kinds, which are then pleochroic. Named from avy^, luilrr.

d. Alkali-augite. Here belong varieties of augite characterized by the presence of alkalis, especially soda; they approach in composition to the iron and optically to actinolite and tegirite ($\beta = 60^\circ$,

Fig. 778), and are sometimes called eegirite-augite (cf. Fig 776, p. 882). Known chiefly from rocks rich in alkalis, as eegirite-syenite, phonolite, leucite, etc.

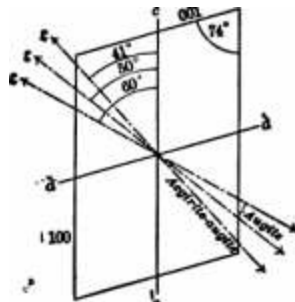
Pyr., etc.—Varying widely, owing to the wide variations in composition in the different varieties, and often by insensible gradations. $H. = 8-7.5$ in diopside; $8-0$ in saite, baikalite, and omphacite; 8 in j&ogbanite and augite; $2-6$ in hedenbergite. Varieties rich in iron afford a magnetic globule when fused on charcoal, and in general the fusibility varies with the amount of iron. Many varieties give with the fluxes reactions for manganese. Most varieties are unacted upon by acids.

Diopside—Characterized by monoclinic crystallization and the prismatic angle of 60° and 53° , hence yielding nearly square prisms; these may be mistaken for scapolite if terminal faces are wanting or indistinct (but scapolite fuses easily B.B. with intumescence). The

oblique parting (| e. Fig, 770) often distinctive, also the cci mm on dull green to gray and brown colon. Amphibole differa in prismatic angle (56° and 124°) and cleavage, and in having common columnar to Sbroos varieties, which are rare with pyroxene. Bee also p. 808.

The common rock-forming pyroxenes are distinguished In thin sections by their high relief; usually greenish to olive tones of color; distinct system of Interrupted cleavage-cracks crossing one another at nearly right angles in sections 774. X h (Fig. 771); high interference-colors; general lack of pleochroism; large extinction-angle, 85° to 50° and higher, for sections 16 (010). The last-named, sections are easily recognized by showing the hiebt interference-color; yielding no optical figures in convergent light and having parallel cleavage-cracks. the latter in the direction of the vertical axis. See also sgrlth. p. Vt2, A zonal banding is common, the successive laminae sometimes differing in extinction-angle and pleochroism; also the hour-glass structure occasionally distinct (Fig. 774. from Lacroix).

Oba.—Pyroxene is a common mineral in crystalline limestone and dolomite. In serpentine and in volcanic rocks; and occurs also, but less abundantly, in connection with granitic rocks and metamorphic schists; sometimes forms large beds or veins, especially in Archkan rocks. It occurs also in meteorites. The pyroxene of Umeetoue is mostly white and light green or gray in color, falling under diopside (malacolite, salite, coocoolite); that of most other



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metamorphic rocks is aometlmei white or colorless, but usually green of different shades, from pale green to greenish black, and occasionally black; that of serpentine is aome lines of fine crystals, but often of the foliated green kind called diallagite; that of eruptive rocks is usually the black to greenish-black variety.

In limestone the associations are often amphibole, scapolite, tourmaline, garnet, orthoclase, titanite, apatite, plagioclase, and sometimes brown tourmaline, chlorite, rutile, zircon, spinel, rutilite, etc.; and in other metamorphic rocks mostly the same. In eruptive rocks it may be found as distinct embedded crystals, or as grains without external crystalline form; it often occurs with similarly disseminated chrysolite (olivine), crystals of orthoclase (sanidine), labradorite, leucite, etc.; also with a rhombic pyroxene, amphibole, etc.

Pyruiene. as an essential rock-making mineral, is especially common in biislc eruptive rocks. Thus, as augite, with a triellulc feldspar (usually labradorite). magnetite, often Cbryssollte, lu basalt, lissaltic lavas and diabase; lu andeslte; also in trachyte; In peridotita aud pikrile; wiih nepbelite lu phonollte. Furtlier with elfeolite, ortlioclase, etc.. In elKolite-syenile and augile-syeaile; also as diallage in gabbro; In many peridotites and the serpentines formed from them; as diopslde (mulacolile) In crysiallioe schists. In llmburgite, Bugillte and pyroxeuite, pyroxene ts present as lbe prominent constituent, while feldspar is absent: it may also form rock masses alone nearly free from associated mlneials.

Diopside (alalite, muscite) occurs in fine crystals on the Mussa alp in the Ala valley in Piedmont, associated with garnets (bessonile) and talc iu veins traversing serpentine; in fine crystals at Tiaversella; at Zermatt in Switzerland; Schwarzenstein in the ZlUertbal; Ober-Sutzbachthnl and elsewhere in Tyrol and In the Salzburg Alps; Reichenstelu; Rezsny, Hungary; Achmatovsk lu the Iirul with almandite, clinocllore; L. Baikal (patkaUtt) in enslern Sluerla; Pargiis In Finland; at Nordmark, Sweden.

He^nbergite is from Tunberg. Sweden; Arendal, Norway. Manganhedentierglte from Yester Sllfberg; Khefferitt iramlAngtMa. Sweden.

Augite (lucT, fasaalte) occurs on the Pesmeda alp, Ht. Monzoni, and elsewhere in the Passftthal. as a cnntact formation; Truversella, Pieamont; tbe Laacher See and the Eifel; Sasbiich in the Kalserstuhl; VesuviuB. white rare, green, brown, yellow to black; Frnscall: Etna; tbe Azores and Cape Verde Islands; the Saodwicb Islands, and many other regions of volcanic rocks.

In N. America, occurs in Maine, at Rnymond and Itumford, diopslde, sallsle, etc. In VtrmotU, at Thelford, black augite, with chrysolite, in boulders of basalt. In Conn., aX Canaan, white cryst., oflen exteruallv cbangtd to tremolite, in dolomite: also tbe pyroxenic rock called eanaantU. In N.. r^rh. at Warwick, fine cryst.; Iu Westchester Co., > white, at tbe Slug Sing (lunrries; in Orange Co.. in Monroe, nt Two Ponds, cryst., often large. In limestone; near Greenwood furuace, and also nsar Edenville; in Lewis Co., at Diana, white and black cryst.; !n St. Lnwrence Co., at Fine, In large cryst.; at De Kalb, fine diopslde; iilso at Gouverneur, liossle, Russell, Pitcaim: al Honnh, coccolite. In limestone. In y. JtneSi Franklin Furnace, Sussex Co., good cryst., ». \ki jefemonite. laPenn., near Attleboro', cryst. and granular; In Pennsbury, nt Burnett's quarry, diopaide; at the French Creek mines, Chester Co., cbieSy altered to Sbrous amphibole. In Tennettet, at the Ducktown mines.

In Canada, nt Calumet I.. gmyish-green cryst. In limestone: In Bathurst. colorless or while crvst.; at Orenvllle, daik green cryst., and granulai; Burgess, Lanark Co.; Renfrew Co., with apaiite, titanite. etc.; Orford, Bberbrooke Co., white crystals, also of a chrome-green color with chrome garnet; at Hull and Wakefield, white crystals with ueoriy colorless garnets, honey-yellow vesuvlanlte, etc. At many other points in the Archtean of

Quebec and Ontario, especially In connection with the apatite deposits.

Pyroxene undergoes alteration In diHerent ways. A change of molecular constitution without essential change of compositlou, i.e., by paramorpM»m (u^g tbe word rather broodly), may result In the formation of some variety of amphlbole. Thus, tbe wliiie pyroxene crystals of Canaan, Conn., are often changed on tbe exterior to tremolite; similarly with other varieties at many localities. See uraliie, p. 401. Also changed to steatite, serpentine, etc.

Aoanrs. .£oibitb.

Monoclinic. Aiea: a:t:i = 1-0996 :1 : 0-6012; fi = 73" U'.

Twins: tw. pi. a, very common; crystala often polysyntbetic, with enclloed twinning lamella. Crystals long priematic, vertit^all; striated or channeled; acnte terminationa very characteristic.

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DBSCRIFTITE UISBRALOGT.

The above appILM to ordfoarr aemiie. For agiriu, ciysUIR prlnnattc. bluotlr tenm-775. 77Ba. Dated: IwioB DOl common; also in groups or tafta of aleuder

^^ acclular to capUlarr crjaUJi, and ia flbroua fornia.

/^\\ Cleavage: m distinct; b less bo. Fracture uneveD.

Brittle. H. = 6-6-6. G. = 3-00-3-55 Bgr. Luater vitreoas, incliniig to resinous. Streak pale yellowish gray. Color browuiflh or reddish brown, green; in the fracture blackieb green. Subtransparent to opaque. Optically —. Ax. pi. || J. Bi, A (i = a A (5 = + 2 J° acmite, to 6° tegirite. p, = 1-808; y — a= 0052.

Var.—Include* aemite Id gharp-polnled cryfllala (Fig. 775) often twins. Biu a * = H°-8°- Alao irginte (Fig. 775(i) in crystals bluDtIj terminated, twini rare. Bx. A i = ^°-^'. Aemite. jEglrile. Crystals of acmtie oftea show a marked zoDal siructnre,

green within and brown on ibe exterior, particularly \a,b.p (iOl). j (ill). Tbe brown portion (aemite) U feebly pleoclirolc, Ibe green («girite) strongij pleochrolc Both have abaorptioD ii>t>c, but tbe former baa a light brown wilh tinge of green, b greeniab yellow with tinge of browu, c browi;lah . .

fellow; the latter has a deep graai-green, 6 lighter grasa-green, c yellowish brown to yellciwish.

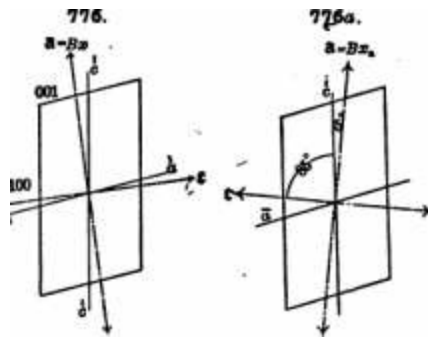
With some aulbora (vom Rath, etc.) (= (Oil) and o a i = - 2° to -6°, aa in Fig. 776a. Fig. 776 shows the optical orientation according to Brögger.

ni Comp. —Essentially NaFe (SiO₃), or Na₂O.Fe₂O₃. 4SiO₂ = 100 Silica 53.0, iron sesquioxide 34.6, soda 13.4 = 100. Ferrous iron is also present.

Pyr., etc.—B.B. fuses at 2 to 3 to a black magnetic globule, coloring the flame deep yellow; with the fluxes reacts for iron and sometimes manganese. Slightly acted upon by acids.

Diff.—Egirite is characterized in thin sections by its green color; strong pleochroism in tones of green and yellow; the small extinction-angle in sections || b (010). Distinguished from common green hornblende, with which it might be confounded, by the fact that in such sections the direction of extinction lying near the cleavage is negative (a), while the same direction in hornblende is positive (t).

Original occurrence at Rundemyr, east of the little lake called Rokeberg-



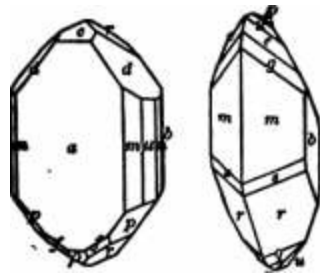
Obs.—

skjeru, in the parish of Eker, near Kongsberg, Norway, in a pegmatite vein; It is in slender crystals, sometimes a foot long, embedded in feldspar and quartz.

Egirite occurs especially in igneous rocks rich in soda and containing iron; thus in egirite-granite, nephelite-syenite, and some varieties of phonolite; often in such cases iron-ore grains are wanting in the rock, their place being taken by egirite crystals. In the sub-variety of phonolite called tinguaite, the rock has often a deep greenish color due to the abundance of minute crystals of egirite. Large crystals are found in the pegmatite

mine of American occurrences are the following: Magnet Cove, Arkansas (large crystals); Montreal: Salem, Mass.: Libertyville, N. J. (dike); Trona Pecos district in Texas (Oatman); Black Hills; Cripple Creek, Colorado: Bearpaw Mt. and Judith Mts. (Pirsson), and the Crazy Mrs. (Wolff) in Montana.

Egirite is named from Egir, in allusion to the pointed extremities of the crystals; egirite is from Egir, the Icelandic god of the sea.



Tripliane.

Monoclinic. Axes: $a : b : c = 1.1238 : 1 : 0.6355$; $\beta = 69^\circ 40'$. Twins: tw, pi, a. Crystals prismatic ($m = 93^\circ 0'$), often flattened $\perp a$; the vertical prism striated and furrowed; crystals sometimes very large. Also massive, cleavable.

Cleavage: m perfect. A lamellar structure $\perp a$ sometimes very prominent, a crystal then separating into $\perp a$ thin plates.

Fracture uneven to subconchoidal, Brittle. H. = 6.5-7. G. = 3.13-3.20. Luster vitreous, on cleavage surfaces somewhat pearly. Color greenish white, grayish white, yellowish green, emerald-green, yellow, amethystine purple. Streak white. Transparent to translucent. Pleochroism strong in deep green varieties. Optically +. Ax. pi. $\Delta = 0.016$. $\omega = 24^\circ$ to Z. Di., = 24° to Z. Norwich. Maa. H. W. denile. H. l. denite.

Dispersion $p > v$, horizontal. $2H_e = 64^\circ 58'$; $\Delta = 1.669$; $y - a = 0.016$.

In small (in, to 3 inches long) prismatic crystals, faces often developed.

Coup.— $\text{LiAl}(\text{SiO}_3)_2$ or $\text{U}_2\text{O}_3 \cdot \text{Al}_2\text{O}_3 \cdot 4\text{SiO}_2 = \text{Silica } 64.5, \text{ alumina } 27.4, \text{ lithia } 8.4 = 100$. Generally contain a little sodium; the variety hiddenite also chromium, to which the color may be due.

Pyr., etc.—B.B. becomes white and opaque, swells up, imparts a purple-red color (to the flame, and fuses at 86 to a clear or white glass. The powdered mineral, fused with a mixture of potassium bisulphate and fluorite on platinum wire, gives a more intense lithia reaction. Not acted upon by acids.

Diff —Characterized by its perfect orthorhombic parting (in some varieties) as well as by prismatic cleavage; has a higher specific gravity and more pearly luster than feldspar or scapolite. Gives a red flame B.B. Less fusible than amphibole.

Ot«.—Occurs on the island of Uts. Sweden; nt KlUIney Bay, Ireland; in small truQS-parent crystals of a pale yellow in Brazil, province of Mioas OeraEs.

In the U. 9., In granite at Goahen, Mass.; also at ChcBterfeld, Cheater, Huntington (formerly Norwich), and Sterling, Mass.; at Windham, Maine, with garnet aod siaurolite; at Peru, with beryl, triphyllte, pelalile. In Conn., at Branchville, the cryalola often of Irnraense aize; near Stony Point, Alexander Co., N. C. (fiiddemU); in South Dakota at the Etta tin mine in Pennington Co.

The name a^dumene Is from <ricoSi6i, ath-tol&rtd. SidentU is named tor W. E. HIdilen of New York.

The apodumene at Goshen and Chesterfield is oiteaslvly altered; pseudomorpbs occur of cymalolite (an Intimate mliilure of albite and muscovite with wavy fibrous structure and silky luster), killinite (pinile), muscovite, ^bite, quartz, and of "vein granite." Similar a Iteration-products occur at BranchviUe.

JADBrm.

Monoclinic. Axes, see p. 383. Cleavage and optical characters like pyroxene. Usually massive, with crystalline atrncture, sometimes granular, also obscurely columnar, fibrous foliated to closely compact.

Cleavage: prismjitic. itt angles of about 93° and 81° ; also orthodiagonal. diificult. Fracture spliutery. Extremely tough. H. = C'.^-7. G = 3-33-3-35. Luster subvitreoua, peaily on surfaces of cleavage. Color apple-green

jcbyGoOgIC

to nearly erne raid-green, bluish green, leek-green, greenisli white, and nearly white; Bometimea white with spots of bright green. Optically biaxial, negative. Bx. A <) - 30° to 40° , $2H \gg j. = 83^\circ 48^*$. Streak uncolored. Translucent to Bubtranslucent.

C«mp,—Essentially a metasilicate of sodinm and alnminium correBponding to apodumene, $\text{NaAl}(\text{SiO}_2)_2$, or $\text{Na}_2\text{O} \cdot \text{Al}_2\text{O}_3 \cdot 4\text{SiO}_2 = \text{Silica } 59.4, \text{ alumina } 25.2, \text{ soda } 15.4 = 100:$

GhlorotrulaiiiiU Ib a dark greeD to nearly black kind of Jadelte (hence the nune), codulId-fu;; Irou Besquioxide nnd not confoTOiing exocUy lo the above formula.

Pjrr., et«.—B.B. fuses readily M a transparent blebby glass. Not attacked by acids after fusion, and thus dUIering from sausBurfte.

Ob*.—Occurs chiefly In eaatern Asia, thus in the Mqgoung distr. In Upper Burma, iD a valley 25 miles southwst of Heiukhooui, in rolled masaeBin a reddisU clay; in Yuugchang, province of Yunnnn, southern China; in Thibet. Much uacerainty prevails,

however, as to the exact localities, since jadeite and nephrite have usually been confounded together. May occur also on the American continent, in Mexico and South America; perhaps also in Europe.

Jadeite has been highly prized in the East, especially in China, where it is worked into ornaments and utensils of great variety and beauty. It is also found with the relics of early man, together with the remains of the like-dwellers of Switzerland, at various localities in France, in Mexico, Greece, Egypt, and Asia Minor.

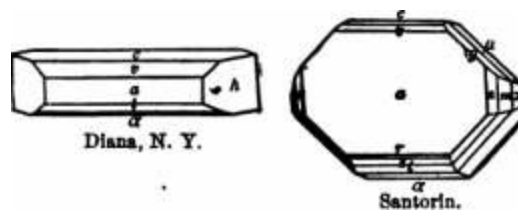
A pyroxene, resembling Jadeite in structure and consisting of the molecules of jadeite, diopside, and scimitar in nearly equal proportions, occurs at the Tancoyale mines of St. Marcel, Italy (Pentfield).

Jade is a general term used to include various mineral substances of tough compact texture and nearly white to dark green color used by early man for utensils and ornaments, and still highly valued in the East, especially in China. It includes properly two species only; nephrite, a variety of amphibole (p. 401), either tremolite or actinolite, with $O = 8$; and jadeite, of the pyroxene group and in composition a soda-apodumene, with $O = 8$; easily fusible.

The Jade of China belongs to both species, so also that of the Swiss lake-habitations and of Mexico. Of the two, however, the former, nephrite, is the more common; and makes the jade (ax stone or Punamu stone) of the Maoris of New Zealand; also found in Alaska.

The name Jade is also sometimes loosely used to embrace other minerals of more or less similar characters, and which have been or might be similarly used—thus sillimanite, peridot, serpentine; also vesuvianite, garnet. Bournonite is a jade-like variety of serpentinite. The "Jadeite" of de Saussure is now called saussurite.

WOLFFSTOHNITE. Tabular Spar. Tafelspath Germ. Monoclinic. Axes $H = 10531 : 1 : 0-9676$; $\beta = 84^\circ 30'$.



Usually cleavable massive to fibrous, fibers parallel or reticulated; also compact.

Cleavage: a perfect; also c; / (101) less so. Fracture uneven. Brittle. H. = 4.5-5. G. = 2.8-3.9. Luster vitreous, on cleavage surfaces pearly. Color white, inclining to gray, yellow, red, or brown. Streak white. Sub-transparent to translucent. Optically —. β , $\alpha = + 37^\circ 40'$, Dispersion

p> V weak; inclined Btrong. Ai. pi. | 6. 2E, = 70° 40'; /S = 1-633; y - a = 0-014.

C«mp.—C&lclum metasilicate, CaSiO, or CaO.SiO, = Silica 61-7, lime 48-3, = 100.

Pyr., vto.—lu Ibe matrass do change. B.fi. fuses easily on Ibc edges; witli some soda, a blebby glass; wilL more, swells up nud is lufuaiWe. Willi hydrochloric acid decomposed Willi Beperatloa of tiltca; most Tarletles eQervesce sligbtly from the presence of cal-cite. Often pbospboresces.

Obs.—Wollusioolte is found especially iu gramilar limestone, and in regions of grentte, as u contact formallon; also In ejected masses in GODoeclion with basalt aud lavas. It Is orien associated wilb a lime garnet, dlopsfde, etc.

Occurs In the copper mines of Czklowa in Ibe Banat; at Orawitiui at Dognaczks and Nagjag; nt Pargas in Finland; at HsTzburg in the Uarz; at Auerbach, in granular limestone; lit Vesuvius, rurely in fine crystals; on Elba; ou Banlurln.

In Ibe U. S.. in N. York, at Willsborough; Diana, Lewis Co.; Bonaparte Lake, Lewis Co. In Penn., Bucks Co.. 8 m. w. of Attleboro'. In Canada, at OrenTille; at St. J^rOme and Horin. Quebec, with apatite.

Kamed after Ibe English chemist, W. H. Wollaston (1766-1826).

PEtoTOLrra.

Monoclinic. Axes: d.:h:6 = 11140 :1 : 09864; jff = 84° 40'.

Goronionly in close aggregations of acicular crystals; elongated || t, but rarely terminated. Fibrous massive, radiated to stellate.

Cleavage: a perfect; c also perfect. Fracture uiieyen. Brittle. H. = 5, G. = 2-68-2-78. Luster of the Burfaoe of fracture silky or subvitreous. Color whitieh or grayisli. Subtranslucent to opaque. Optically -{- Ax, pi. and Bx,Xfti Bx, nearly J.rt; 2H,, ^ 143°-145*.

Comp., Tar.—HNaCa,(SiO,) or H,O.Na,0.4Cao.6SiO, = Silica 542, lime 33-8, soda 93, water 2-7 = 100.

s: occBsioDaily in _ , Corstorpbine Hill

ivmlkgrilay. I. 8)tye. Also at Mt. Bsldo and Ut. MoDZonriii the Tyrol; at Niederkirchen, Bavaria iomndite).

Occurs also at Bergen Hill and Palerson. N. J.; Lehigh Co., Peon.; compact at Isle Boynle. L. Superior; at Magnet Cove, Ark., in eleeolite-syenlle [manganpttellite wiib 4 p. c. MnO); compact, massive in Alaska, where used, like jsoe, for implements.

RoMnbnahlte. Near pectollte, but contains lirconlum. From Norway.

WOhlorito. A ilrconiiim-Bllicate and niobale of Ca, Na. etc. In prismalle, tabular crj-stals, yellow to browtj. Occurs in elKOLLte-syenile, on several islands of the I^ngesund fiord, nejir Brevik, in Norway.

Uivanlte A complex zircon)urn-silicate of Hn, Ca. etc., contaiDlnR also F, Ti, Tn, etc. In yellow to brown prismatic crystals. Found on the Island L&ven In the Langesuod fiord, BOiilhem Norway; also elsewhere In eleeolite-syenite.

y. Triainic Section.

SHODONITD.

Triclinic. Axes $a : b : c = 107285 : 106213 : 10318$; $\alpha = 103^\circ 18'$; $\beta = 108^\circ 44'$; $\gamma = 81^\circ 39'$.

Crystals usually large and rough with rounded edges. Commonly tabular Q c; BonietinieH resembling pyroxene in habit. Commonly massiTe, clearable to compact; also in embedded grains.

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Cleavage: m. If perfect; c leu perfect. Fracture concliodal to nneTen^ very tough when compact. H, = 5"5-65. G. = H'4-3'68. Luster Titreous; on cleavage-BurfaccB somewhat pearly. Color light brownish red, flesh-red, rose-pink; Hometimea greenish or yellowish, when impure; often black outside from exposure. Streak white. Transparent to translucent.

Comp., Tar.—Manganese metasilicate, $MnSiO_3$, or $MnO \cdot SiO_2$, = Silica 45.9, manganese protoxide 54.1 = 100. Iron, calcium (in buatamile), and occasionally zinc {in fowleriie) replace part of the manganese.



FntoUin Furnace, N. J.

ab. $100^\circ A 010 = 04^\circ 26'$. mX, $110^\circ A 110 = 92^\circ 2d'$.

ae. $100^\circ A 001 = 73^\circ 86i'$. on, $001^\circ A S5l = 78^\circ 52''$.

be. $010^\circ A 001 = 78^\circ 421'$. «*. $001^\circ A 2^1 = 62^\circ 2B'$.

am, 100 A 110 = 48° SS*. in, 221 A 2^1 = 86' !>'.

Pyr., ate,—B.B. blackens and fuses with silicate and fluxes gives resctlon for magnesia; fowlerite glues with Boda on charcoal a reaction for zinc. Blight acted upon by ncl. The calciferous varieties often effervesce from acetic acid admixture of calcium carbonate. In powder, partially dissolves in hydrochloric acid, and the insoluble part becomes of a white color. Exposed to the air and sometimes becomes nearly black.

Dts.—Characterized by its pink color; distinct cleavages; fusibility and manganese reactions B.B.

Obi.—Occurs at Långban, Wennland, Sweden, in iron-ore beds, in broad cleavage-plates, and also granular massive; at the Pajaberg iron mines near Filipstad (pargasite) sometimes in small brilliant crystals; in the district of Ekaterinburg in the Ural massive like marble, whetstone it is obtained for ornamental purposes; with tetraerite at Rapnik and Rezdouya. Hungary; St. Marcel, Piedmont; Mexico (Nutami, containing CaO).

Occurs in Cumminston, Mass.; on Osgood's farm. Blue Hill Bay, Maine; favershamite (containing ZnO) at Mine Hill, Franklin Furnace, and Sterling Hill, near Ogdensburgh, N. J., usually embedded in calcite and sometimes in fine crystals.

Named from fuchs, a rose, in allusion to the color.

Kiodite is often altered chiefly by oxidation of the MnO part in manganese. (imitation): also by hydration (trapezite, noselite. etc.); further by introduction of CO, (aitagite. platite, etc.).

Barnstonite. (Ca,Fe,Mn)8SiO₁₀ with Fe,SiO₂. In small black triclinic crystals, near rhodonite in angle (axes on p. 888). H. = 5-5-0. O. = 8'S5-3-S7. From Areodal, Norway; at Herbornseelbach. Nassau; at Baveno, Italy.

Hortdahlite. Essentially (Mn,Ca,XSi,Zr)O₁₀, with also fluorine. In pale yellow tabular crystals (triclinic)! Occurs sparingly on an island in the Langesundfjord, southern Norway.

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3. Amphibole Group.

Orthorhombic, Monoclinic, Triclinic

Composition for the most part that of a metasilicate, KSiO₃, with B = Ca,Mg, etc. chiefly, also Mn,Na (K,Al,Fe). Further often containing aluminum and ferric iron, in part with alkalis as NaAl(SiO₃), or NaFe(SiO₃); perhaps also containing RB₂SiO₆.

a. Orthorhombic Section.

Antferrothit (M: Mg, Fe)SiO₃, 0-5138: 1

Gedrite (Mg, Fe)SiO₃, with (Mg, Fe)Al₂SiO₅.

Orthorhombic Section,

a: h - i a

Amphibole 0-6511 : 1 : 0-2938 73° 68'

I. Nepheline-NaAlSi₃O₈

1. Thymolite CaMg₂(SiO₃)₂

2. Actinolite Ca(Mg, Fe)₂(SiO₃)₇

Nephrite, Asbestos, Smaragdite, etc. Cummingtonite (Fe, Mg)SiO₃, Danieleschite {Fe, Mn, Mg}SiO₃

Quartz FeSiO₃

3. Richterite (K, Na, Mg, Ca, Mn)₂(SiO₃)₄

II. Aluminosilicates

4. Hornblende

Kehe and \ ^" ^y Ca(Mg, Fe).(SiO₃) with

calcic Hornblende r ^ ^MSiO₃), and (Mg, Fe).(Al, Fe), SiO₃..

Glaucophanite NaAl(SiO₃)₂(Fe, Mg)SiO₃

&: h: i /?

Siebeckite 2NaFe(SiO₃)₂·eSiO₃, 0-5475 : 1 : 02935 = 76° 10'

Crocidolite NaFe(SiO₃)₂·FeSiO₃

Arfvedsonite Na₂(Ca, Mg)₂(Fe, Mn)₂(Al, Fe)₆Si₈O₂₂

0-5509 : 1 : 0-3378 = 73° 2' y. Triclinic Section. Enigmatite.

The only species included under the triclinic section is the rare and imperfectly known unigmatite (coaxite).

The Amphibole Group embraces a number of species which, while falling in different systems, are yet closely related in form—as shown in the common prismatic cleavage of 54° to 56° —also in optical characters and chemical composition, as already noted (see p. 383), the species of this group form chemically a series parallel to that of the closely allied Pyroxene Group, and between them there is a close relationship in crystalline form and other characters. The Amphibole Group, however, is less fully developed, including fewer species, and those known show less variety in form.

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The chief differences between pyroxene and amphibole proper are the following:

Prismatic angle with pyroxene 87° and 90° ; with amphibole 50° and 131° ; the prismatic cleavage being much more distinct in the latter

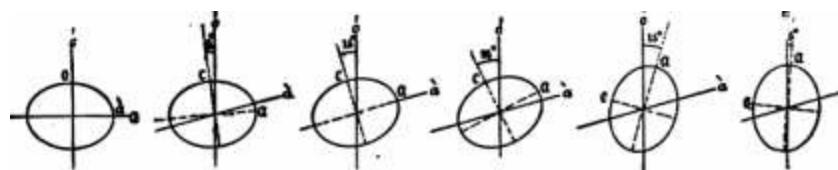
With pyroxene, crystals usually form a complex, structure of various kinds mostly lamellar or granular; with amphibole, crystals are mostly long prismatic and simple, columnar and fibrous massive kinds.

The specific gravity of most of the pyroxene is higher than of the like varieties of amphibole. In composition of corresponding kinds, magnesium is present in larger amount in amphibole (Ca : Mg = 1 : 1 in diopside, = 1 : 2 in tremolite); alkalis more frequently play a prominent part in amphibole.

The optical relations of the prominent members of the group, as regards the position of the ether-axes, is exhibited by the following figures (Cross); compare Fig. 751, p. 384, for a similar representation for the corresponding members of the pyroxene group.

7BB.

L n. m. IV. V. V3.



a. Orthorhombic Section. Amphibole.

Orthorhombic. Axial ratio $d:l = 0.5137 : 1$, Crystals rare, habit prismatic ($m = 54^\circ 23'$). Commonly lamellar, or fibrous massive; fibers often very slender; in aggregations of prisms.

Cleavage: prismatic, perfect; or less so; and sometimes distinct. H. = 5-6. G. = 3-3'. Luster vitreous, somewhat pearly on the cleavage-face. Color brownish gray, yellowish brown, clove-brown, brownish green, emerald-green, sometimes metalloidal. Streak uncolored or

grayisli. Transparent to sub-translucent. Sometimes pleochroic. Unanalytically optically -f; also + for red, - for yellow, green. Ax. pi. always j b. Bx, usually JL <■", also _L c for red, J. a for yellow, green. Axial angle large. $\hat{c} = \text{■} rC42$; $y - a = 0\ 024$.

Comp., Var.—(Mg,Fe)SiO₃, corresponding to eifitite-bionzite.hyperstbene in the pyroxene group. Aluminium is sometimes present in considerable amount. There is the same relation in optical character between anthophyllite (-J-) and gedrite (-) as between enstatite and hyperstbene (cf. Figs. 753, 754, p. 385).

Var.—Anthophyllite, Mg : Fe = 4 : 1, 8 : 1, etc. For 8 : 1, the percentage composition is: Silica 55.5, iron protoxide 16.6, magnesia 27.8=100. Anthophyllite sometimes occurs in forms resembling asbestis.

Aluminous. Gedrite. Iron is present in larger amount, and also aluminium: it hence corresponds nearly to a hyperstbene. Some varieties of which are highly aluminous.

Bydroanthophyllite is repeatedly described, but in most cases they have been shown to be bydroxylated monoclinic amphiboles.

Pyr, etc.—B.B. fuses with difficulty to a black magnetic enamel: with the fluxes gives reactions for iron: unacted upon by cyanide.

Obs.—Anthophyllite occurs in mica schist near Kongsberg in Norway; also in Hennaun-Bachlag, Horava. In the U. S., at the Jecks corundum mine, Frickville, Michigan Co., N. G.

Named from anthophyllite.

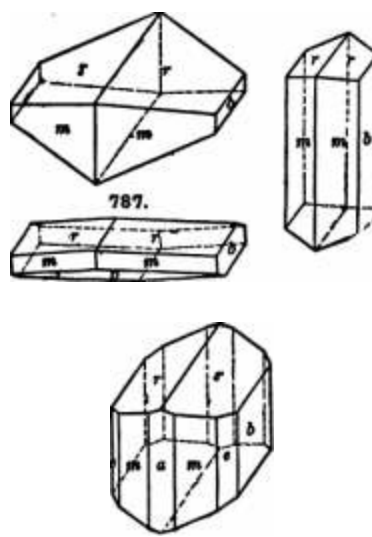
/S. Manodinic Section. AMPHIBOLITE. Horava-Bachlag.

Monoclinic. Axes $a : b : c = 0.5511 : 1 : 0.2938$; $\beta = 73^\circ 58'$. $m = 73^\circ 58'$, no $A \text{ liO} = 5^\circ 49'$. $rK.OU \text{ a Oil} = 81^\circ 0'$.

$ea, 001 A 100 = 78^\circ 68'$. $H. 081 \text{ a } 061 = 50^\circ 83'$.

$ep. 001 A 101 = 81^\circ 0'$. $pr, 101 A 011 = 84^\circ 56'$.

Twins: (1) tw. pi. a, common as contact-twins; usually polysynthetic. (2) c, as tw. lamellae, occasionally producing a parting analogous to that more common with pyroxene (Fig. 430, p. 131). Crystals commonly prismatic; usually terminated by the low clinodome, r (Oil), sometimes by r and p (loi) equally developed and then suggesting rhombohedral forms (as of tourmaline). 786. 788. 789. T9a



Also columnar or fibrous, coarse or fine, fibers often like fibrous; rarely lamellar; also granular massive, coarse or fine, and usually strongly coherent, but sometimes friable.

Cleavage: m highly perfect; a, b sometimes distinct. Fracture subconchoidal, uneven. Brittle. H. = 5-6, G. = 2.9-3.4, varying with the composition. Luster vitreous to pearl; ^ on cleavage-faces; fibrous varieties often silky. Color between black and white, through various shades of green, inclining to blackish green; also dark brown; rarely yellow, pink, rose-red. Streak colorless, or paler than color. Sometimes nearly transparent; usually sub-translucent to opaque.

Pleochroism strongly marked in all the deeply colored varieties, as described beyond. Absorption usually $c > b > a$. Optically -, rarely +. Ax. pi. | b. Extinction angle on b, or $t \wedge i = -16^\circ$ to 18° in most cases, but varying from about 1° up to 37° ; hence also $Bi \gg A_{rf} = -75^\circ$ to -72° , etc. See Fig. 791. Dispersion $p < v$. Axial angles variable; see beyond.

Comp., Tar.—In part a normal meta-silicate of calcium and magnesium, $RSiO_3$, usually with iron, also manganese, and thus in general analogous to the pyroxenes. The alkali metals, sodium and potassium, also present, and more common; ^ so than with pyroxene. In part also aluminous, corresponding to the aluminous pyroxenes. Titanium sometimes is present and also rarely fluorine in small amount.

The amphiboles in part present as $NaAlSi_3O_{10}$ (e.g., but many amphiboles contain aluminium or ferric iron are more basic than a normal silicate; they may sometimes be explained as containing $R(A_1, Fe)_2Si_8O_{22}$, but the exact nature of the compound is often

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Pyroxene. Thus, for most tremolite and actinolite, $Ca : Mg(Fe) = 1 : 8$, and hence the formula is $CaMg_8Si_8O_{22}$, while the formula for actinolite is $CaMg_7Si_8O_{22}$, etc.

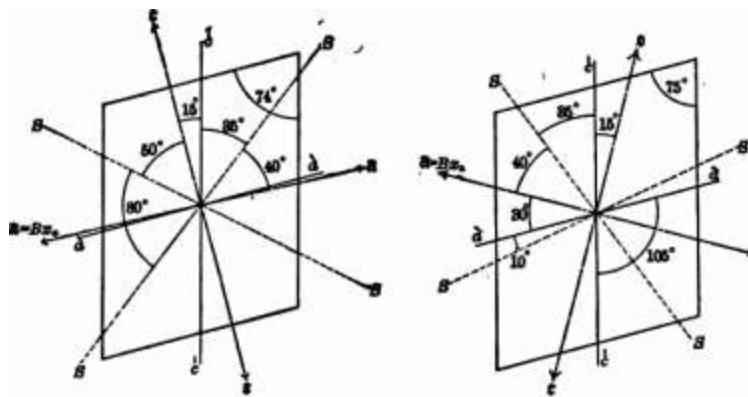
Rammelsberg has shown that the composition of most aluminous amphiboles may be

,_^ I (lie geotal form mRS1oi.7iAl,o,; while ScLarizat, modityiitg ibis

proposes to regard Ibe ampiboles as molecular compounds of Ca(Hg,Feji&COit (ac^no-
ez pressed i

lite), and the orthosilicate (Ri,R)tBi8fiOii, for which he uses Breltbaupt's Dame tt/tUag.
natilt, oiiginallj gLveu to the Tesuviaa honibleiide.

791. 791a.



The cyatalllograpble positlou here adopted is that suggested by Tscbermnk, which best exhibits the relaiioc beiwec anipblhole and pyroxene. Some aibtors retnlo tbo former positioLi, according lo which p = (001), r = (ill), etc. Fig. 791a shows the correapendiog optical or!eutattoo.

I. Containing little or no Aluminium.

1. Trbuolite. Orammatite, nephrite pt. Calcium-magnesium ampJiiboU. FormulaCaM:g, (SiO.), = Silica 577, magnesia 28'9, lime 13-4 — 100. Perrons iron, replacing the magnesium, present only sparingly, up to 3 p c. Colors white to dark gray. In distinct crystals, either long-bladea or short and stout. In aggregates long and thin columnar, or fibrous; also compact granular massive (nephrite, p. 40L). o. = 3'9-3'l. Sometimes transparent and colorless.

Optically —. Extihction-angle on S, or c A (^ = + 16° to 18°, hei = - 74"* to - 72°. 2V^ = 80° to 88°. oj = 1'621; >' - a = 0 0

027.

B Bi, A (!

2, AcTiNOLiTE. Strahlstein Oerm. Calctum-magneaium-iron ampkibole. Formula Ca(Mg,Fe),(SiO)|. Color bright green and grayish green. In crystals, either short- or long-bladed, as in tremolite; columnar or fibrous; grannlar massive. G. = 3-3'2. Sometimes transparent. The variety in long bright-green crystals is called glassy actinoUte; the

crystals break easily across the prism. The fibrous and radiated kinds are often called fibrous actinolite and radiated actinolite. Actinolite owes its green color to the ferrous iron present.

Dichroism distinct, increasing as the amount of iron increases, and hence the color becomes darker: c emerald-green, b yellow-green, a greenish yellow.

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Absorption $c > b > a$ Zillerthal, Tschermak. Optically —. Extinction-angle on A, c $A \perp = +16^\circ$ and Bx. $A \perp = 75^\circ$. $2V, = 80^\circ$; $p < v$; $y \perp, = 1-627$; $y - a = 0.025$.

Found in actinolite from Tirschenreuth, a ray, Bodon, Ume, a translation of the Osmundite or fibrous actinolite. None changed to actinolite by heat, without fusion.

Nephrite. In the pt. A tough, fibrous, greenish (or actinolite), breaking in a fibrous direction. It is a variety of actinolite.

It is a variety of actinolite.

It is a variety of actinolite. The latter kind of actinolite encloses distinct prismatic crystals of actinolite. A derivate of fibrous actinolite has been suggested in some cases.

Nephrite or jade was brought in the form of carved ornaments from Mexico or Peru soon after the discovery of America. A similar stone comes from Eastern Asia, New Zealand and Alaska. See jadeite, p. 848; jade, p. 854.

Asbestos. Asbestos, Asbest, Oerm. Tremolite, actinolite, and other varieties of amphibole, except those of talc, mica, alumina, pass into fibrous varieties, the fibers of which are sometimes very long, fine, flexible, and easily separable by the fingers, and look like hair. These kinds are called asbestos (fr. the Greek for incombustible). The color varies from white to green and wood-brown. The name amianthus is applied usually to the finer and more silky kinds. It is popularly called asbestos. It is chrysotile, or fibrous asbestos, containing 12 to 14 p. c of water. By itself is a stiff fibrous variety.

Mountain asbestos is in thin flexible sheets, made of interlaced fibers; and sometimes in the same in thicker pieces; both are so light as to float on water, and they are often hydrous, color white to gray or yellowish. Mountain wood is compact fibrous, and gray to brown in color, looking a little like dry wood.

Smectonite. A thin-foliated variety of amphibole, near actinolite in composition but crystalline. It has a light green color, resembling much common green mica. In many cases derived from pyroxene (diallag) by alteration, see below. It retains much of the structure of the diallag and also often encloses remnants of the original mineral. It forms, along with whitish or greenish saussurite, a rock called

sauesudte-gabbro the euphoildeof the Alps. The original niluernl is from Corsica, and the rock is the VfTiU di CvTvica duro of the arts.

Ubalite. Pyroxene altered to amphibole. The crystals, when distinct, retain the form of the oilsinal niinerul, but have Ibe cli'iivage of ampibole. The change usually commences ou the surface, transforming the ciuttr layer Into an aggregation of slender amphibole prisms, parallel lu poitiidn lo each oier and to the p»r«nl pyroxeue (cf. Fig. 760, p. 386). Wbcu the change la complete tliu euiiro crretHl Is made up of a bundle of amphibolu needles or Ubers, The color varies from white (Ireuiolite) to pale or deep gr<en, The latter Uie mote common. In composltliu uniliie appears to conform nearly tu ai liunlile, as also lu optical characieis. The most prominent change in conipositlou in pusslug fioni the original pyroxene Is that corresponding to the difference existiing between Ihe two speiles In general, ihnl is. an increase In Ihe miigni'slnm nnd decrease In calcium. The clmuee, therefore, is not strictly a case of pnrnmnrpbian), nltboiigb usually so designated. Unilil^ was originally deacribed by Rose lu a rock from the Urn). It has since been observed from many localities. The microscopic study iif rocks bus shown the process of " uralltization " to be very common, nnd some niitbais lejriu'd many homblendlc rocks and schists lo represent altered pyroienic rocks on a large scnle.

CiTMMiDaTONiTE. Ampblbole-Anihopliyllite, Iron - Magnttium Ainphe>ola. Here Iwlong curtain varieties of amphibole re.ienibline imthopbyllite and eosentially identical with it in composition, but optically mouoclinic. From Kongsberg; Greenland. The original tommiagtnniU is gray to brown in color; usually fibrous or Gbio-lamellar, often radiated. " = f51-!!'33; frnm Cummingtoti, Mnss.

D.

Colui Sweden.

OrOherite. Iron-Amphibole A=be8iiform or lamellar-fibrous. Luster sDky; color ironw; G, = 8-718. Formula FeSiO.

3. RicHTERlTE. Sodium-Magnesium-Manganese Amphibole. (Ki^a,,Mg, <;a,Mn),(SiO,)..

In elongated crystals, seldom terininalcd. G. = 8-09. Color brown, yellow, rose-red. Transparent lo Iran»Uicent. t a i = + 1T°-20°; o, = 1'68: y-a = o-fm. From Pnjs-Iwrc nnd LADgban. Sweden. Characterized by the presence of mangaoen and alkalies In Telaiively large amount.

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DESCRIPTIVE UIIFERALOOY.

II. Aluminous.

4. AldHIITOUS Amphibole. Hornblende. Contains aluminia or ferric iron, and usually both, with ferrous iron (Bismuth manganese), in addition, calcium, and alkalis. The kinds here included range from the light-colored edenite, containing but little iron, through the light to dark green pargasite, to the dark-colored or black hornblende, the color growing darker with increase in amount of iron. Extinction-angle variable, from 0° to 37° see below. Pleochroism strong. Absorption usually $c < b < a$.

Edenite, Aluminous Magnesium-Calcium Amphibole. Color white to gray to pale blue, and also colorless; $Q = 3.0-3.9$. Bismuth autophyllite and treibite. Named from the locality at Eden, N. Y. To this variety belong various pale-colored amphiboles having less than 6 p. c. of iron oxides.

Kokibergite is a variety from the neighborhood of L. Baikal named after the Russian mineralogist, N. S. Kokiberg.

Common Hornblende, Colors bright or dark green, and bluish green to fuscous black and black. $Q = 3.0-4.0$. Pargasite is usually made to include green and bluish-green kinds, occurring in stout lustrous crystals, or granular; and Common hornblende the greenish-black and black kinds, whether in stout crystals or long blades, columnar, fibrous, or massive granular. But no line can be drawn between them. True extinction-angle on i , or $c \wedge = +16^\circ$ to 25° chiefly. Absorption $t > b > a$.

Pargasite occurs at Pargas, Finland, in bluish-green and grayish-black crystals, $c \wedge = -18^\circ$; $P = 1.64$; $y - a = 0.019$; $a \vee = 58^\circ$. Pleochroism: c greenish blue; S emerald-green; a grayish yellow.

The dark brown to black hornblende from basaltic and other igneous rocks vary somewhat widely in optical characters. The angle $c \wedge = 0^\circ$ to $+10^\circ$ chiefly; $J = 1.735$; $y - a = 0.072$ (minimum). Pleochroism: c brown, b yellowish gray, a yellow, but variable.

The Eataforite of Norway (Brogger) has ($c \wedge = 30^\circ$ to 60° ; absorption $S > t > a$; pleochroism: c yellow, b violet, a yellowish brown; line approaches toward arfvedsonite (p. 405).

Kupfferite from a graphite mine in the Tunkinsk Hts., near L. Baikal, is a deep green amphibole (aluminous) formerly referred to as anthophyllite. Syntagmatite is the black hornblende of Vesuvius.

Broggerite is an iron-amphibole containing almost no magnesia. From the Modena Province of Bergamo, Italy.

Arfvedsonite is a titaniferous amphibole from Eoerud, Umanaks fjord, North Greenland. Hiitilite is an amphibole low in silica and high in iron and soda, from the nephelitesyenite of Dungannon, Hensington Co., Ontario.

Pyr. —Essentially the game as for the corresponding varieties of pyroiene, see p. 390.

Diff.— Diatlngidsbed from pyroxene (and tourmaline) by lis distinct prismatic cleavage, yielding angles of $5<t^{\circ}$ and 124° . Fibrous and columnar forms are much more commoa than

with pyroxene, lamellar and foliated forms rare (see also pp. 800. 398). Crystals often long,

792 slender, or bladed. Differs from rhe fibrous zeolites tn

not gelatinizing with nclds. Epidote has a peculiir

green color, Is more fusible, and shows a aiSerent

cleavage.

In rock sections amphibole generally shows distinct W no colors, green, sometimes olive or brown, and is strongly

pleochroic. Also recognized by lis high relief ; generally nilher high interference-colors; by the Yery perfect system of cleavage-cracks crossing at angles of $6S^*$ and 124° in sections i h (Fig. 792). In sections |& (010> (recognized by yielding no axial figure In convergent ligiil, by showing the hiehest interference-colors, and by having parallel cleavage-cracks, I i), the extinction-direct ion for common hornblendes makes a small angle (12° - 15°) with the cleavage-cracks (is., with i); fuvllier. this direction la positive c (different from common pyroxene and Ezirite, cf. Fiin. 772 anil 77C). 6 6

Ot».—Amphibole occura tn many cryslaltne limestones, and granitic and schistose locks, and sparingly in volcanic rocks. TrmioliU!, the magesla-llme variety, is especially

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commoQ Id Umnloties, parttcurly magoesiau or ilolomlttc; acUtulUt (ftim nephrite), (be magnesia-lime-iron TaiHety, iu steatltic rocks and witli BerpeDtioe; and dark greea aud black AerTMtade, iu chlorite scliUt, mica scbist. gneUs. aud in Taiioua other rocke of which it forms u MniBtitiient part; brown to black hornblende occurs In trachyte and uLher eniptlve rocka. Asbintia la often foiled in coDQection with serpentine.

HorrMtiue-Toek, or ainplJbolite, conalsiB of massive hornblende of a dark greenish black or black color, and hiie a granular texture. Occasionally the gretn hornblende, or nctiuii-lite, occurs in rock-maasea, aa at Bt. Francis, in Cimada. HoTitbUnde-iehUl baa the sniie compoaiion iia amptalbolite, but Is schistose iir shity in structure. It often contains a litile feldspar. In some varieties of ll the hornblende is In pari In minute

needles. Uranium and zirconium often occur in horizons, and with diopside it is a common constituent. This is

often true of the corresponding forms of gneiss. In these cases it is usually present in small, regular masses, often fibrous in structure: also as rough bladed crystals.

Prominent foreign localities of amphibole are the following; Tremolite (green) in dolomite at Campolungo, Switzerland; also at Orawltza. Itzbeny, Hungary; Oulshög, Wermland, Sweden. Actinolite in the crystalline schists of the Central and Eastern Alps, especially at Oreuerluthal; at Zoblitz in Saxony; Arendal, Norway, Aemlu* at Sterzing, Zillertal, and elsewhere in Tyrol; in Savoy; also in the island of Corsica. Paragonite in Fergana, Finland: Sauppe in Carinthia. Ilombundent Aremlal and Kongsberg, Norway; in Sweden and Finland; at Vesuvius: Aussig and Teplitz, Bubemli; etc. Jade, which in the form of "jade" ornaments and utensils is widely distributed among the relics of early man (see Jade, p. 58), is obtained at various points in Central Asia. The most important source is that in the Karakash valley in the Euxine Highlands, on the southern borders of Turkestan; also other localities in Central Asia. In New Zealand. Nephrite has been found in Europe as a rolled mass at Schwemmsal near Leipzig; in Swiss Lake habitations and especially elsewhere.

In the United States. In Maine, black crystals occur at Thomaston; paragonite at Phippsburg. In Vermont, actinolite in the actinolite quarries of Windham and Sew Fane, in Montpelier; tremolite at Lee; black crystals at Chester; asbestos at Pelham; Bummings/lonite at Cummington. In Conn., in large flattened white crystals and in bladed and fibrous forms (tremolite) in dolomite, at Canaan. In N. York Warwick, Orange Co.; near Edenville; near Amity; at the Stirling mines. Orange Co.; in short green crystals at Gouverneur, St. Lawrence Co.; with pyroxene at Ruesselt; a black variety at Pierrepont; at Macomb; Pitcairn: tremolite at Fine; in Rosalia, 2 m. N. of Oxbow; in large white crystals at Diana, Lewis Co.; asbestos near Greenwood Furnace, in N. York. Tremolite or gray amphibole in good crystals at Bryam, and other varieties of the species at Franklin and Newton, radiating actinolite. In Penn. actinolite at Mineral Hill, in Delaware Co.; at Unionville; at Eennett, Chester Co. In Maryland, actinolite and asbestos at the Bare Hills in serpentine; asbestos is mined at Pilesville, Harford Co. In Virginia, actinolite at Willis's Mt., in Buckingham Co.; asbestos at Bame's Hill. Fauquier Co. Nephrite occurs in Alaska.

In Canada, tremolite is abundant in the Laurentian limestones. at Calumet Falls, Litchfield, Pontiac Co., Quebec; also at Blythfield, Renfrew Co., and Dalbousle, Inverness Co. Black hornblende at various localities in Quebec and Ontario with pyroxene, apatite, titanite, etc. in Renfrew Co. Asbestos and mountain cork at Buckingham, Ottawa Co., Quebec; a bed of actinolite at St. Francis, Beauport Co., Quebec; nephrite has been found in British Columbia and Northwest Territory.

OLAnOOFHANE,

Monoclinic; near amphibole in form, Crystals prismatic in habit, usually indistinct; commonly massive, fibrous, or columnar to granular.

Crystallography: m perfect. Fracture conchoidal to uneven. Brittle. H. = 6-6.5. G. = 3.103-3.113. Luster vitreous to pearly. Color azure-blue, lavender-blue to bluish black, grayish. Streak grayish blue. Translucent. Pleochroism strongly marked: c axy-blue to ultramarine-blue, ti reddish or bluish violet, a yellowish green to colorless. Absorption $C > b > a$, Optically - f-. Ax. pi. $\angle C A f = 4^\circ$ to 6° , rarely higher values. $2E^\wedge = 84^\circ 42'$. $\wedge = 1-6442$ (gastaldite); $y \sim a = 0-033$.

Comp.-Essentially $NaAl(SiO_3)_2 \cdot (Fe, Mg)SiO_3$. If Mg : Fe = 3 : 1, the formula requires: Silica 57.6, alumina 16.3, iron protoxide 7.7, magnesia 8.5, soda 9.9 = 100.

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Ob*.—Occurs as the principal constituent of certain crystalline rocks, called feldspars, or glaucophane; also more or less prominent in the mica schists, amphibolites, gneisses, eclogites, etc. It is often associated with mica, hornblende and omphacite, epidote and zircon, etc. First described from the Ujand of Syria, Ode or uik Cyckdes; since shown to be rather widely distributed, but on the outlier of the A. (Igor Mialiiitt), Corsica. Jupa. etc. Blodunite is a fibrous variety from the Is. Rbodii.

Ill. U. B., glaucophane schists have been described from the Coast Ranges of California. IIS al Sulphur Blk, Lake Co.

Quartz is named from the Greek, and also (a & H, to appear).

Monoclinic. Axes $a : b : c = 0.5475 : 1 : 0.2925$; $\angle a b = 76^\circ 10'$. In embedded prismatic crystals, longitudinally striated. Cleavage: prismatic (56°) perfect. Luster vitreous. Color black. Pleochroism very strongly marked: c green, b (= fc) deep blue, it (nearly || r.) dark blue. Optically —. Extinction-angle small, $a \perp d = \wedge - b$ ($\pm ?$). Axi angle large.

Comp.—Essentially $2\text{NaFe}(\text{SiO}_3)_2 \cdot \text{FeSiO}_3$, = Silica 50.5, iron sesquioxide 26.9, iron protoxide 12.1, soda 10.5 = 100, It corresponds closely to acmite (tegrite) among the pyroxenes.

ObB.—Originally described from the granite and gneiss of the island of Socotra in the Indian Ocean, 120 m. E. of Cape Guardafui, the eastern extremity of Africa; occurs in groups of prismatic crystals, often radiating and commonly resembling tourmaline; also in crano-pyric blocks found at Ailsa Craig and at other points in Scotland and Ireland. A similar amphibole occurs at Mynydd Mawr, Carnarvonshire, Wales. Also another in granulite in Corsica. A so-called arfvedsonite from BC. Pelee's Dome, Filie's Peak region, El Paso Co., Colorado, occurring with astrophyllite and zircon, is shown by Lacroix to be near riebeckite. Extinction-angle on a , $a \wedge = 8^\circ$ to 4° .

OROOLIT. Blue Asbestos.

Fibrous, asbestos-like; fibers long but delicate, and easily separable. Also massive or earthy. Cleavage; prismatic, 56° . $H. = 4$. $G. = 320-330$, Luster silky; dull. Color and streak lavender-blue or leekgreen. Opaque. Fibers somewhat elastic, Pleochroism: c green, b violet, a blue. Optically +. Extinction-angle on b , inclined 18° to 20° with $2E = 95^\circ$ approx. $y - a = 0.025$.

Comp.— $\text{NaFe}(\text{SiO}_3)_2 \cdot \text{FeSiO}_3$ (nearly) = Silica 49.6, iron sesquioxide 32.0, iron protoxide 19.8, soda 8.6 = 100.

Magnesium and calcium replace pari of the ferrous iron, and hydrogen part of the

Pyrr., etc.—In the closed tube yields a small amount of alkaline water. B.B. fuses easily with fluxes to a black magnetic glass, coloring the flame yellow (soda). With the fluxes gives reactions for iron, unaffected upon by acids.

Oba.—Occurs in South Africa, in the Orange River, north of the Orange river, in a range of quartzose schists called the Asbestos Mountains. In a micaceous schist near Pnimon, in the Vosges. At Golling in Salzburg. In the U. S., at Beacon Pole Hill, near Cumberland. R. I. Emerald Mine, Buckingham, and Perkin's Mill, Templeton, Ottawa Co., Ontario, Canada.

Abriimhani is an earthy amorphous form occurring in the Aberlath district, near Loch Ness, Scotland. Crocidolite is named from Krokos, Greek. In allusion to its fibrous

The South African mineral is largely altered by the oxidation of the iron and infiltration of silica, resulting in a compact siliceous stone of delicate fibrous structure, chatoyant luster, and bright yellow to brown color, popularly called (also "cat's-eye" and "tiger's eye", Tiger's eye. Palkeuaitze (bluish var.) Germ.). Many varieties occur forming

ARF VBDSONITB.

Monoclinio. $a:b:c = 0.6669 : 1 : 0.2978$; $\beta = 73^\circ 2'$.

Grystals long prisms, of ten tabular QJ , but seldom distinctly terminated; angles near those of amphibole; also in prismatic aggregate, Twins: tw.pl. a.

Cleavage: prismatic, perfect; b cleavage perfect. Fracture uneven. Brittle, $n_x = 1.6$. $G = 3.44-3.45$. Luster vitreous. Color pure black; in thin scales deep green. Streak deep bluish gray. Opaque except in thin splinters. Pleochroism strongly marked: c deep greenish blue, b lavender, a pale greenish yellow. Absorption $c > b > a$; sections $\parallel a$ are deep greenish blue, $\parallel b$ olive* green, $\parallel c = 1.707$; $v - a = 0.02$. Extinction-angle on b , with d , $= 14^\circ$.

Comp.—A slightly basic metasilicate of sodium, calcium, and ferrous iron chiefly; an analysis by Lorenzen gives:

$Q. SiO, Al_2O_3, FeO, MnO, MgO, CaO, Na_2O, K_2O, H_2O$ (calculated) 44.43-85.44, 4.48-8.48, 0.45-1.61, 4.65-8.15, 1 (Wt % = 100%)

The supposed arfvedsonite from Oreenlaad analyzed by von Eobell, Itbg., etc., has been shown to be Kgrite; that from Pike's Peak, Colorado, analyzed by Koenig, has been referred to riebeckite.

Pyr., etc.—B.B. fuses at S with intumescence to a black magnetic globule; colors the flame yellow (soda); with the fluxes g^vea reacts for iron and magnesia. Not acted upon by acids.

, Ob*.—Arfvedsonite and amphiboles of similar character, containing much iron and soda, are common constituents of certain igneous rocks which are rich in alkalis, as nepheline-syenite, phonolite, etc. Large thin crystals are found only in the pegmatite veins in such rocks, as at Eungersdalen, Grenland, where the associated minerals are bolalite, eudialyte, feldspar, etc. Arfvedsonite occurs also in the nepheline-syenite and related rocks of the Cribstone region in southern Norway: on the Sola peninsula in Russian Lapland: Dungaon township, Onirli Trns Pecos district, Texas. The related brownish pleochroic amphiboles (cf. barkevikite) occur in similar rocks at Montreal, Canada; Red Hill, New Hampshire; Saletti, Mass.; Mignet Cove, Ark.; Black Hill, Bo. Dakota; Squaw Lake, Montana, etc.

Chromite is a soda-amphibole near arfvedsonite ($7.62 Na_2O$) from a rock in the neighborhood of San Francisco.

BARKEVIKITE. An amphibole near arfvedsonite but more basic. In prismatic crystals. Cleavage: prismatic ($15^\circ 44'$). $G = 3.428$. Color deep velvet-black. Pleochroism strongly marked, colors brownish. Extinction-angle with i on $i = 12^\circ$. Occurs at the "Ohierle" locality near Barkevik, on the Langesund fiord, and elsewhere in southern Norway.

Snlgmatita. Cossyrilte. Esⁿtinlly a titnno-sillcate of ferrous iron and sodium, but containing also aluminium aud ferric Iron. In prismatic tricliuic cryetals. Cleavage: prismatic, disliuct (88°). G. = 8'74-8-80. Color bliick. ..DiiigmBtile is from tbe sodaiile-syenite of Tunuudliarfik aud Kaugerdluarsuk, Greenland. Couvrite occurs In minute crystals embedded in Ibe liparile lavua of Ihe island Pantellaria (ancient name Cossyra).

BBRTIi.

Hexagonal. Axis d = 0-4989.

Crystals usually long prismatic, often striated vertically, rarely transversely; distinct terminations exceptional. Occasionally in large masses, coarae columnar or grannlar to compact.

Cleavage: c imperfect aud indistinct. Fracture conchoidal to uneven. Brittle. H. = 7-5-8. G, = 2-63-3-80; usually 2'69-2To. Luster vitreous, Fonetime resinous. Colors emerald-green, pale green, passing into light blue.

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DESCRIPTIVE UINERALOOT.

ie-Tci. Streak white. Tranaparent to fiabtruu-lucent. Dichroiam more or less distinct. Optically—, Birefringence low.

fellow and white; also pale iicent. Dichroiam more o: Often abnormally biaxial, oo = 1'5830, e = 1*5765 aqnamarine.

|<^^^S(^^^^

ep. 0001 A 10!l ■= S»' 664',

oo, 0001 A 1122 = ae- sr.

et, 0001 A ll2l = 44° 58'.

jip'. loii AOiii sas'sii'.

V«r.—1. Bmeratd. Color bright emerald-greeo, due to tbe pi'eseuce of a little chro-tniuiD ; highlj prized as a geni whan cle&r Htid free From flnwB.

2. Ontinarj/; Beryl, ^iierally ia itexngoohl prisms, oflvn coarse and large ; green the COminnn color. The prlaclpal kinds are : (a) colorless; {}>) bluish green, ci\\)eAaqriiunariie;

ie) apple-green ; (d) greenish yellow to iron-yellow and honey-yellow ; sometimes a clear bright yellow as in the green beryl; (e) pale yellowish green ; (/) clear sapphire-blue; (g) pale sky-blue ; (A) the pale violet or reddish ; (i) the opaque brownish yellow, of waxy or greasy habit. The oriental emerald of Jewelry is emerald-colored sapphire.

Comp.—Be₃Al₂(SiO₃)₆, or 3BeO.Al₂O₃.6SiO₂, ^ Silica 67.0, alumina 19.0, glucina 14.0 = 100.

Alkalies (Na₂O, Li₂O, CaO) about 5 p c. : also chemically combined in H₂BeAl₂Si₆O₂₂.

Fus. etc.—B.B. alone unchanged or. If clear, becomes milky white and clouded; at a high temperature the edges are rounded, and ultimately a tessellated scoria is formed.

Specific gravity = 2.7. but Emswiler lower. Some beryls rich in alkalies. Glass with borax, clear and colorless for beryl, a blue for emerald. Unacted upon by acids.

DM.—Characterized by its blue or greenish-blue color, glassy luster and hexagonal form ; rarely massive, then easily mistaken for quartz. Distinguished from apatite by its hardness, not being scratched by a knife. also harder than green tourmaline ; from chrysoberyl by its form : from euclase and topaz by its imperfect cleavage.

Obs.—Beryl is a common accessory mineral in granite veins, especially in those of a pegmatitic character. Emeralds occur in clay slate, in isolated crystals or in nests, near Muso, etc., 73 m. N.N.W. of Bogotá, Colombia. Emeralds of less beauty, but larger, occur in Liberia, on the river Tokovoya. N. of Bliatertsburg, embedded in mica schist, Emeralds of this size, though not of uniform color or free from flaws, have been obtained in Alexander Co., N. C.

Transparent beryls are found in Siberia, India and Brazil, in Siberia they occur at Mursinka and Shalaukii, near Ekaterinburg : in the Ural Mountains with topaz : in the mountains of Adou-Chind with topaz, in E. Siberia. Beautiful crystals also occur at Elba: the twins of Ehrenfriedersdorf in Saxony, and Bohnckenwald in Bohemia. Other localities are the Mourne Mts., Ireland : yellowish green at Rubialaw, near Aberdeen, Scotland (David M. U): Limoges in France ; Finho and Broddho in Sweden ; Tamela in Finland: Pfilsch-Joch, Tyrol ; Bodenmais and Robenslein in Bavaria ; in New South Wales.

In the United States, beryls of gigantic dimensions have been found in N. Hamp., at Acworth and Oiaflou, and in Vermont at Royalton. In New York, at Albany : Norway : Bethel: at Hebron, a clear beryl (Cs₂O, about 6 p c), associated with pollucite ; in Maine with black tourmaline; at Topsham, pale green or yellowish. In New Hampshire, at Barre; at Oashe (A'nie), and at Chesterfield in Conn., at Haddam, and at the Middletown and Portland feldspar quarries; at New Milford, of a clear golden yellow to dark amber color; Branchville. In

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Fenn., at Lelperville Bud Chester; M Hialeh Hill. In Virginia, at Amelia Court-House, ■■■■■ white. In California, in Alameda Co., near Stony Point, fine emeralds ; in

Muscovite. Common in the Black Hills and elsewhere. In Alabama, Coosa Co., of a light yellow color. In Colorado, near the summit of Mt. Aotero, beautiful aquamarines. In & Dakota, in the Black Hills in turquoise crystals.

Endellite. Essentially a metasilicate of Zr, Fe, Mn, Ca, Na, etc. In red or brownish or rhombohedral crystals, also massive. H. = 5-5.5. G. = 2.9-3.0. Usually from Kaugordluarauk, West Greenland, etc., with actinolite and sodalite; also in Lujaor on the Eoht peninsula, Russian Lapland, in elenite-syenite, there forming a main constituent of the rock-mass. Bittern, from Islands of the Ligeauud fiord in Norway, is similar (but optically; ~). Endellite and euclite also occur at Hagnet Cove, in Arkansas, of a rich red color in feldspar, with celestine and ngirite.

Elpidite. $\text{Na}_2\text{O} \cdot \text{ZrSi}_2\text{O}_7 \cdot \text{H}_2\text{O}$ —Massive, fibrous. H. = 2-5.4. Southern Green-

The following are rare species of complex composition, all from the Lange-sund fiord region of southern Norway.

Oatapleite. $\text{H}(\text{Na}, \text{Ca})\text{ZrSi}_2\text{O}_7$ —In tabular hexagonal prisms. H. = 5. G. = 2.8. Color light yellow to yellowish brown. Natron-bearing, or soda-calcapleite, contains only sodium; color blue to gray and white; on heating the blue color disappears.

Oappolenite. A borosilicate of yttrium and barium. In greenish-brown hexagonal

Melanocerite. A borosilicate of the cerium and yttrium metals and calcium (also Ba, Ta, etc.). In brown to black tabular rhombohedral crystals

Carjroceite. Near melanocerite, contains ThO₂.

Stenobrookite (from Greenland) is allied to the two last-named species.

Trithemite. A fluosilicate of thorium, the cerium and yttrium metals and calcium, with boron. In dark brown crystals of acute triangular pyramidal form.

The following are also from the same region:

Lennoxite. $\text{Na}(\text{BeF}_2)_2\text{Cn}(\text{BIOt})$. In glassy greenish tabular crystals (or rhombohedral). H. = 4. G. = 2.8.

Meliphanite. A fluosilicate of beryllium, calcium, and sodium near leucophanite. In low square pyramids (tetragonal). Color yellow. H. = 6-6. G. = 8.01.

Orthorhombic. Conchoidal. Cleavage.

Orthorhombic. Axes $a:b:c = 0.5871 : 1 : 0.5585$.

Twins: tw. pi. wi, also d (130), both yielding pseudo-hexagonal forms. Habit short

prismatic ($\text{mm}''' = 60^\circ 50'$) (Fig. 299, „5

p. 94[^]. As embedded grains; also massive, compact. '

Cleavage: b distinct; a and c indistinct. Crystals f[^]

often show a lamellar structure || c, especially when slightly altered. Fracture subconchoidal. Brittle. H. = 7-7.5. G. = 2.60-3.66, Luster vitreous. Color various shades of blue, light or dark, smoky blue. Transparent to translucent, Pleochroism strongly marked except in thin sections. Axial colors variable. Thus:

Bodenm[^]s < (=5) dark Berlin-blue t (= < i) light Berlin-blue a (= ^ } yellowish white

Absorption c (i) > b («) > a {6}. Pleochroic halos common, often bright yellow; best seen in sections | 6, Exhibits idiophanous figures. Optically —, Ax. pi. | n, Bx. X ». Dispersion feeble, $p < W$. $3V = 70^\circ 23'$ (also 40° to 84°). $p = 1549$; $y - or = 0008$.

,cb, Google

Comp.— $\text{H}_2\{\text{Mg,Fe}\}.\text{Al}_2.\text{Si}_2.\text{O}_8$, or $\text{H}_2\text{O}.4(\text{Mg,FeO}).4\text{Al}_2\text{O}_3.10\text{SiO}_2$ If Mg : Fe = 7 : 2, the percent composition is: Silica 49.4, alumina 33.6, iron protoxide 5.3, magnesia 10.5, water 1.5 = 100. Iron replaces part of the magnesia. Calcium is also present in small amount.

Pyrrho.—B.B. loses transparency and fuses U.S.B. Oa\j partially decomposed by acids. Decomposed on fusion with alkaline carbonates.

Diff.—Characterised by its vitreous luster, color and pleochroism; fusible on the edges unlike quartz; less hard than sapphirine.

Recognized in thin sections by lack of color; low refraction and low interference-colors; it is very similar to quartz, but distinguished by its bluish character: In volcanic rocks commonly show distinct crystal outlines and a twinning of three individuals like amethyst. In the gneisses, etc. It has fibrous grains, but the common occurrence of inclusions, especially of sillimanite needles, the pleochroic halos of a yellow color around small inclusions, particularly zircon, and the constant tendency to alteration to micaceous plagioclase seen along cleavage, help to distinguish it.

Obi.—Occurs in granite, gneiss (Mordierite-gneiss), amphibolitic and talcose schists and allied rocks, with quartz, orthoclase or albite, tourmaline, hornblende, andalusite, sillimanite, garnet, and sometimes beryl. Less commonly in or connected with igneous rocks, thus formed directly from the magma, as in andesite, etc.; also in ejected masses (the fragments of older rocks); further formed as a contact-metamorphic connection with eruptive dikes, as in slates and gneisses.

Occurs at Bodenmala, Bavaria. In granule, with pyrrholite, etc; Orijarvi, in Finland (Ateinhthili); Tuunberg, in Sweden; in colorless crystals from Brazil; Ceylon affords a transparent variety, the lapidary of jewelers.

In the U. S. at Haddam, Conn., associated with tourmaline in a granitic gneiss. At Brimfield, Mass.: at Richmond, N. H.

Named in honor of the discoverer, and X-ray; Dichroite (from Sixoai, Iwo-eobyrd). from its dichroism; Cordierite, after Cordier, the French geologist (1777-1861).

The alteration of feldspar takes place so readily by ordinary exposure, that the mineral is most commonly found in an altered state, or enclosed in the altered rock. This change may be a simple hydration; or a removal of part of the protoxide bases by carbon dioxide; or the introduction of oxide of iron; or of alkalies, forming epidote and mica. The first step in the change consists in a division of the prisms of feldspar into plates parallel to the prism, and a pearly foliation of the surfaces of these plates: with a change of color to grayish green and greenish gray, and sometimes brownish gray. As the alteration proceeds, the foliation becomes more complete; afterward it may be lost. The mineral in this altered condition has many names: as hydrotite (inc). bontdorffite and auralite from Aho, Finland; feldspar from Falun, Sweden, also jyrargillite from Helsingfors; etmarkite and praelite from near Brevik, Norway, also raumite from Riiimo, Finland, and perthite from Itains-berg, Sweden: microphyllite from Unity, He.; aegirine; and polyhedroite from Kragero. There are further alkaline kinds, as pinite, eadite, gigantite, and the Mica Group,

The following are rare lead and barium silicates: Baiyite. $Pb_8Si_{10}O_{28}$. In embedded masses with curved lamellar structure

Specific gravity = 3.6. Color white; tarnishing on exposure. From the Hirstig mine, Pnjaberg, Sweden.

Anomalite*. $Pb_8Si_{10}O_{28} \cdot (Ca, Mn)_2SiO_4$. In prismatic crystals (tetragonal); also massive,

Specific gravity = 3.6. Colorless to gray. From Långban, Sweden; also at Kobsjöberg. Hyaloteite. Approximately $(Pb, Ba, Ca)Si_3O_{10}$. Massive: coarsely crystalline. H. = 5-5.5. G. = 3-3.5. Color white to pearly gray. From Långbau, Sweden.

Baiyite. $Ba_3Al_2Si_2O_{14}$. In groups of colorless prismatic crystals. H. = 7. G. = 4.08. Luster greenish. Occurs with crystalline limestone at Långban, Sweden.

Boeblingite. $5(H, Ca)SiO_4 \cdot 3(CaPbSO_4)$. In dense, white, compact, crystalline masses. H. = 3. G. = 8-4.38. From Franklin Furnace, N. J.

m. OrthoBUIofttea. R,8iO..

Salts of OrthoBilicio Acid, H,SiO,; obaracterized b; an oxygen ratio of 1 : 1 for Bilicou to bases.

The following list includes the more prominent groups among tbe Ortbo-silicates.

A. number of bulc orlboalltcatee nre Lere fucluded. which yield wiiter upon l^ltion; also wLkh ai« more or lata bBsic tljau u nnormal ortboslllcale. but which are of necessity Introduced here Id the clasaiflcution, because of their rebiliouBliip to olber Dorinal Bilra. Tbe Mica Ohoop is ao clowly related to many Hydrous Silicates tliat (wUb also Talc, KaoHuiie. nod some otbera) it is described later with iheui.

Nephellte Group. Hexagonal. Scapollte Oroup. Tetragonal • py-

Sodalite Qroup. Isoiretrio. ramidal.

Helvtte Qroup. Isometrio-tetrabe Zircon Group. Tetragonal.

dral. Daiibiirite Group. Orthorboinbi&

Garuet Group. Isometric. DatoUte Group. Uonoclinic.

Cbrysolite Group. Orthorhombic. Epidote Group. Honoolinic Pheuacite Oroup. Tri-
rbombohe-

dral.

Nephellte Group, Hexagonal.

Typical formula RAISiO..

Hephelite K,Na,Al.Si,o,, i = 0-8389

Soda-nephelite (artit.) NaAlSiO. Enoryptit« LiAlSiO. KaliopMUte KAISiO,

Canorinite H,Na.Ca{NaCO,},Al,(SiO,), 2i = Q 8448

aicrosDmmite (Na,K),,Ca,Al,,Si,.o,,Sgl. 2<! = 0-8367

The species of tbe Nbphblite Group are hexagonal in crystalltiation and have ill part tlio typical orthosilicate formula EAlSiO . From this formula iiephelite itself deviates somewhat, though an artificial soda-nephelite, NitAlSiO., conforms to it. The species Cancrinite and Microaommite are related in form and also in composition, though in the latter respect somewhat complex. They serve to connect this group with the sodalite gronp following.

NEFHELITE. Nephtalite.

Hexagonal-bemimorphic (p. 73). Axis $i = 0.83893$.

In thick six- or twelve-sided prisms with plane or modified summits. Also massive compact, and in embedded grains; structure sometimes thin columnar.

Cleavage: m distinct; c imperfect. Fracture subconchoidal. Brittle. $H = 5-5.5-6$. $Q = 2, ^5-2.65$. Luster Vitreous to greasy; a little opalescent in some varieties. Colorless, white, or yellowish; black when massive, dark green, greenish or bluish gray, brownish red and brick-red. Transparent to opaque. Optically —. Indices; $\mu_y = 1.042$, $e = 1.538$.

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Var.—] Nepheline. $NaAlSi_3O_8$.—Usually in small glassy crystals or groups. Translucent with vitreous luster, first found on Mt. Sorona. Characteristic particularly of volcanic eruptive rocks and lavas. 2. Leucite.—In large coarse crystals or more commonly in a vitreous, greasy luster, reddish, brownish or gray in color. Usually clouded by minute inclusions. Characteristic of granular crystalline rocks, syenite, etc.

Comp.— R, Al, Si, O ; if $R = Na:K = 3:1$, this is equivalent to $3NaO.K_2O.4Al_2O_3.9SiO_2 = Silica 44.0$, alumina 33.2, Potash 7.7 = CO_2 . In most analyses $Na : K = 4 : 1$ or $5 : 1$.

Synthetic experiments, yielding crystals like nepheline with the composition $NaAlSi_3O_8$, lead to the conclusion that a natural soda-leucite would be an orthosilicate with this formula, while the higher silica and potash varieties may be explained by the presence, in molecular combination, of Al_2SiO_5 or Al_2SiO_4 (= leucite). The other species of the group are normal orthosilicates, viz., eucryptite: $LiAlSi_3O_8$, and kaliophyllite, $KAlSi_3O_8$.

Pyr., etc. — B_2O_3 fuses quietly at 1150° to a colorless glass. It is soluble in acids.

Diff.—Distinguished by its gelatinizing with acids from scapolite and feldspar, as also from apatite, from which it differs also in its greater hardness. Massive varieties have a characteristic vitreous luster.

Recognized in thin sections by its low refraction; very low interference-colors, which scarcely rise to gray; parallel extinction when in crystals; faint negative uniaxial cross yielded by biaxial sections in converging light. The optical character is best told by aid of the selective plate (see p. 201). Micro-chemical tests serve to distinguish non-characteristic particles from similar ones of alkali feldspar; the section is treated with dilute acid, and the resulting gelatinous silica, which coats the nephelite particles, stained with eosin or other dye.

Obs.—Nephelite is rather widely distributed (as shown by the microscopic study of rocks) in igneous rocks as the product of crystallization of a magma rich in soda and at the

Bame time low in silica (which last prevents the soda from being used up in the foimalion of albite). It is thus an essential component of the nepbeliie-syeiites and pbonolites where it is associated with alkali feldspai's chiefly. It is also a constituent of more basic augitic rocks such as nepbeliule. uepbelile-basalts, neplielite-tephrites, lliernlite, etc., most of which are volcanic in origin. The vuriety eUrolite is associated with the granular plutonic rocks, while the name nrplulitf was originally used for the frisli glassy (;rjBtai8 of the modern lavas; the terms have in this sense llie same relative siguiGcauce as orthoclase and eanidlne. Hoderii usage, however, tends 10 drop the name elaolite.

The original ntphelite occurs in crystals In the older lavas of Mte. Somma, with mica, vesuvianile. etc.; at Capo di Bove, near Rome; in Ihe Imsalt of KaUenburkel, near Heidelberg; Aussig in Bohemia; Ltfliau In Saiioiiy. Occurs also In mascive forms and targe course crystiils {elaoUte) in Ibe nephelile-syenltes of Southern Norway, especially along the Langesund fiord; sitni'nrly in »«it Greenlnnd; the peninsuhi of Eoln; Miask in the Itmen Mis. (in the rock miateUr): Sierra Monchlque. Portugal (\n the rock/o^aito); Ditr6, Hungary (in the rock (ii(ro/rt); Pousac, Prance; Brnsil: South Africa.

Elicolite occurs mas.slve nnd cryslalllxed at Litchfield. He., with cancrinite; Salem, Mass.; Red Hill. N. H.; in tlie Ozark Mts.. nciir Hiignet Cove, Arkansas; elKoli1e-syenit« Is also found near Beemersville. northieru N. J.; uenr Montreal, Canada; at Dun cannon township, Ontario, in enormouB cryslnl. Neplielite rocks iisn occur nt various poinis, as tlie Tninspecos diatr.. Texas : Pilnt Butte. Texas; also In western N. Americti, as in Colorado at Cripple Creek; in Monlimn, in tlie Crazy Mi«, the Higbwood, Bearpaw and Judith MU ; Black Hills in So. Dakota: Ice River, Brilisli Columbia.

Named nepheliU from rfipfXrt. "■ 'loud, in allusion to its becoming cloudy wheD immersed in siniug acid; etixoUU is from eXaCi^v, oil. In allusion to lis greasy luster.

Qieieekitt Is a pseudomorph after nephelite. It occurs in Greenland >n six-sided green-iah-grav prisms of grenay tusier; also at Diana in Lewis Co., N. Y. Tyytynla-ibiU from DiaDii is similar to gieseckite. as is also lUb«neriU. from the valley of Fleims, in Tyrol. See further Ptnite under ilie Mica Group.

BuorypUte. LiAISiO,. In symmetrl'-ally arranged cryatala (hexagonal), embedded In albite and derived from the alteration of apodumcne at BrHQchvllle, Conn, (see Fig. 474, p. 141). Q. = 3-667. ColorlesH or while.

KaUophilite. KAISiO,. Phacelllle. Plincellte. Facetllle. in bundles of slender scicniar crystals (hexaeonal), also in fine threads, cobweb-like. H. = S. G. = 2 498-2003. Colorless. Occurs in ejected masses nt Mte. Sorama.

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OAHOBUnTB.

Hexagonal. Axial $i = 0^\circ 42' 24''$; and $mp = 64^\circ$, $pp' = 26^\circ 58'$. Rarely in prismatic crystals with a low terminal pyramid. Usually massive.

Optical: prismatic (m) perfect; a less bo. H. = 5-6. G. = 3.42-3.5. Color white, gray, yellow, green, blue, reddish. Streak colorless. Luster subvitreous, or a little pearly or greasy. Transparent to translucent. Optically uniaxial, negative.

Comp— $H.Na.Ca(NaCO_3).Al_2(SiO_4)_2(OH)_2$, or $3H_2O.4CaO.4Al_2O_3.9SiO_2.2CO_2$, = Silica 38.7, carbon dioxide 6.3, alumina 39.3, lime 40, soda 18, water 3.9 = IW.

Pyro., etc.— If the closed tube gives water. B.B. looses color, and fuses (F. = 3) with intense blue color. Ion white blebby glass, like very easy fusibility distinguishing it readily from nepheline. Effervesces with hydrochloric acid, and forms a jelly on heating, but not before.

Idf.— Recognized in thin sections by its low refraction, quite bluish interference-color* and negative uniaxial character. Its common association with nepheline, sodalite, etc. are valuable characteristics. Reaction with acid dissolves it from all other minerals except the carbonates, which show much bluish interference-colors.

Obs.— Cancrinite occurs only in igneous rocks of the nepheline-syenite and related rock groups. It is believed to be original, i.e., formed directly from the molten magma; but part held to be secondary and formed at the expense of nepheline by infiltration of waters holding calcium carbonate in solution. Prominent localities are Minsk in the Ural Mts., in the Ural Mountains; Dephelite-syenite; similarly at Brøkkevik and other localities on the Laogevnir in the southern part of the island of Spitzbergen; in the parish of Knolajsvik in northern Fluhind (where associated with orthoclase. (eghite and nepheline, it composes a mass of cancrinite-syenite); at Dilrath, Transylvania, etc.; in nephelite-syenite of Strna and Alnö in Sweden, and in Brazil; also in small amount as an occasional accessory component of many igneous rocks at various localities.

In the United States at Litchfield, Me., with orthoclase and blue sodalite. Named after Count Cancrin, Prussian Minister of Finance.

Microscopic. Near colorless; perhaps $(Na, Ca)_2Al_2Si_2O_{10}(OH)_2$. In minute colorless prismatic crystals (hexagonal). From Vesuvius (Mount Vesuvius). H. = 6. G. = 2.42-2.53.

Davton is in part at least microscopical. From Mt. Somtoa.

Sodalite Group. Isometric.

Sodalite $Na_4(AlCl)Al_3(SiO_4)_3$,

Hafniumite $(Na, Ca)_2\{NaSO_4, Al\}_2(SiO_4)_2$,

Hoselite $\text{Na}(\text{NaSO}_4\text{Al})\text{Al}_2(\text{SiO}_3)_6$,

lazurite $\text{Na}_4\{\text{Na}_2\text{S}_2\text{O}_7\text{Al}_2\text{Si}_2\text{O}_7\}$,

The species of the Sodalite Group are isometric in crystallization and perhaps tetrahedral like the following group. In composition they are peculiar (like cancrinite of the preceding group) in containing radicals with Cl, SO and S, which are elements usually absent in the silicates. These are shown in the formulas written above in the form suggested by Brogger, who shows that this group and the one following may be included with the garnets in a broad group characterized by isometric crystallization and a close resemblance in composition. See further under the Garnet Group proper, p. 414.

The formulas are also often written as if the compound consisted of a silicate and chloride (sulphate, sulphide)—thus for sodalite, $\text{Na}_4\text{Al}_3\text{Si}_3\text{O}_{14} + \text{NaCl}$,

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SODALITE.

Isometric, perhaps tetrahedral. Common form the dodecahedron. Twins: tw. pi. o, forming Deixidite prisms by elongation in the direction of an octahedral axis (Fig. 368, p. 133). Also mosaic, in embedded grains; in concentric nodules resembling chalcedony, formed from leucite.

Cleavage: dodecahedral, more or less distinct. Fracture conchoidal to uneven. Brittle. H. = 5.5-6. G. = 2.14-2.30, Luster vitreous, sometimes inclining to greasy. Color gray, greenish, yellowish, white; sometimes blue, lavender-blue, light red. Transparent to translucent. Streak uncolored. $n_D = 1.4857$ Na.

Comp.— $\text{Na}_4(\text{AlCl})\text{Al}_3(\text{SiO}_3)_6$. = Silica 37.2, alumina 31.6, soda 25.6, chlorine 7.3 = 101.7, deduct (O = 2C) 1.7 = 100. Potassium replaces a small part of the sodium. The formula may also be written $3\text{NaAlSi}_3\text{O}_{10} + \text{NaCl}$.

Fusion, etc.—In the closed tube the blue varieties become white and opaque. B.B. fuses with incandescence, at 850°, to a colorless glass. Decomposed by hydrochloric acid, with evolution of gelatinous silica.

Identification.—Recognized in thin sections by its very low refractive index. Isotropic character and lack of good cleavage: also, in most cases, by its lack of color. Distinguished from much analcite by its lack of color and by chemical tests alone; dissolves in dilute hydrofluoric acid and is soluble for chlorine in the simplest and best.

Occurrence.—Sodalite occurs only in igneous rocks of the nepheline-syenite and related rock groups, either directly, as is commonly the case, or as a product of the crystallization of a magma rich in soda; also as a product associated with enclosed masses and bombs

ejected with silicic magmas in the form of lava, as at Vesuvius. Associated with nephelite (or elxolite), cancrinite and eudialyte. It forms a totalite at Scar-nipata in Ischia, in Sicily, Val di Noto, with nephelite and calcite. At Vesuvius, in the Sonima in white, most of the dodecahedral crystals; massive and of a gray color at the Kaiserstuhl; also near Lake Loach. At D[ir6, Pennsylvania. in an elxolite-syenite. In the foyne of southern Portugal. At Minsk, in the Ilmeu Mts.; in the Aljustrel-syenite of the Langesund-fjord region in Norway, further in West Greenland in sodalite-syenite; in the peninsula of Kola.

A blue massive variety occurs at Litchfield and West Gardiner. Occurs in the thealite of the Crazy Hill, Montana; also at Square Butte, Higbwood Mis., and in the Bearpaw Mts., in the Big Horn. Occurs also in the elxolite-syenite of Brome, Brome Co., and of Montreal and Belœil, province of Quebec; at Dugannon, Ontario, in large blue masses and in small pale pluk crystals.

HADTMITB. IlaOyne.

Isometric. Sometimes in dodecahedrons, octahedrons, etc.

Twins: tw. pi. e; contact-twins, also polysynthetic; penetration-twins (Fig. 367, p. 123). Commonly in rounded grains, often looking like crystals with fused surfaces.

Cleavage: dodecahedral, rather distinct. Fracture flat conchoidal to uneven. Brittle. H. = 5-6. (i. = 2'4-2'5. Luster vitreous, to somewhat greasy. Color bright blue, sky-blue, greenish blue; asparagus-green, red, yellow. Streak slightly bluish to colorless.

Subtransparent to translucent; often enclosing symmetrically arranged inclusions (Fig. 7'JT). j, = 1'4961.

Comp.-Na,Ca(NaSO₄·Al)Al₂(SiO₃)₂. This is analogous to the garnet formula (Brogger) where the place of the R, is taken by Na, K, Ga and the group Na-O-SO₄-O-Al. The percentage composition is: Silica 33.0, sulphur trioxide 15.0, alumina 27.2, lime 10.0, soda 16.6 = 100. The ratio of Na: Ca also varies from 3 : 2; potassium may be present in small amount. The formula may also be written 2(Na,Ca)Al₂(SiO₃)₂ + (Na,Ca)SO₄.

Pyr., etc.—In the closed tube retains its color. In the forceps fuses at 4/5 to a white glass. Fused with soda on charcoal forms a black mass which blackens silver. It is decomposed by hydrochloric acid with separation of silicic acid.

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In Melfi, on Mt. Yullur. Naples; in the Invus in the Campagna, Rome, also





Section of crystals of lazurite (after Hohl). In the peperino near Albano; at Niedermendig. In the Eifel; the plinths of Hohenfels.

Hoellta or Noseau. Near Bittard, but contains little or no lime. Color grayish, brownish; sometimes nearly opaque from the presence of inclusion (cf. Fig. TOT). Not uncommon in the dolomite. At Aardach. In the Laacher See, & elsewhere.

LAZURITE. Lapis-Lazuli. Lazurite.

Isometric. In cubes and dodecahedrons. Commonly massive, compact. Cleavage : dodecahedral, imperfect. Fracture uneven. H. = 5-5.5. G. = 2.38-3.45. Luster vitreous. Color rich Berlin-blue or azure-blue, violet-blue, greenish blue. Translucent.

Comp.— Essentially $\text{Na}_4(\text{AlSi}_3\text{O}_{10})_2(\text{SO}_4)_2 \cdot 6\text{H}_2\text{O}$ (Brogger), but containing also in molecular combination lazurite and sodalite. The percentage composition of this ultramarine compound is as follows: Silica 31.7, alumina 21.9, soda 2.3, sulphur 16.9 = 102.9, or deduct (O = S) 3.9 = 100.

The lazurite was first described by Fischer (1889), Zirkel (1873), and more fully by Vogel (1878). The natural lazurite (Lazurite) is shown by Brogger and Beckstrom to contain also mica (muscovite), calcite, pyrite: also in some varieties in relatively small amount staurolite, plagioclase, orihodase (microperthite?), apatite, titanite, zircon, and other mineral optically + and probably uniaxial. Regarded by Brogger as a result of contact metamorphism in limestone,

Pyt., etc.—Heated in a closed tube gives off some moisture; the variety from Billington with fine blue-green luster, but the color of the mineral remains blue on cooling. Fuses easily (3) and intumesces to a white glass. Decomposed by hydrochloric acid, with evolution of gelatinous silica and evolution of hydrogen sulphide.

Obs.— Occurs in Badkhan in the valley of the Kokcha, a branch of the Oxus, a few miles above Firgamu. Also at the southern end of Lake Baikal. Further, in Chili in the Andes of Ovielle. In ejected masses at Monte Simona, etc.

The richly colored varieties of lazurite are highly esteemed for costly vases and ornamental furniture; also employed in the manufacture of pigments; and when powdered constitutes the rich and durable paint called ultramarine. It has been replaced, however, by artificial ultramarine, now an important commercial product.

Helvite Group. Isometric-tetrahedral. Helvite $(\text{Mn,Fe})_2(\text{Mn,S})\text{Be}_2(\text{SiO}_3)_2$,

Sanalite $(\text{Fe,Zn,Mn})_2(\text{Zn,Fe})_2\text{S}_2\text{Be}_2(\text{SiO}_3)_2$,

Eulytite $\text{Bi}_2(\text{SiO}_3)_2$.

Zimiyite $(\text{Al}(\text{OH,F,Cl}))_2\text{Al}_2(\text{SiO}_3)_2$,

The Helvite Group includes several rare species, isometric-tetrahedral in

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crystallization and in composition related to the species of the Sodalite Group sod also to those of the Oabnet Group which follows:

Isometric-tetrahedral. Commonly in tetrahedral crystals; also in spherical masses.

Cleavage: octahedral in traces. Fracture uneven to conchoidal. Brittle. H. = 6-6.5. G. = 3.16-3.36. Luster vitreous, inclining to resinous. Color varies; yellow, inclining to yellowish brown, and siskin-green, reddish brown. Streak uncolored, Subtranslucent. Parent, n = 1.739. Pyroelectric.

Comp.— $(\text{Be,Mn,Fe})_2(\text{SiO}_3)_2\text{S}$. This may be written $(\text{Mn,Fe})_2(\text{Mn,S})\text{Be}_2(\text{SiO}_3)_2$, (Brogger), analogous to the Garnet Group, the bivalent group $-\text{Mn-S-Ma}$ taking the place of a bivalent element, B, and 3Be corresponding to 3Al , cf. p. 415. Composition also written $3(\text{Be,Mn,Fe})_2(\text{SiO}_3)_2(\text{Mn,S})$.

Pyro., etc.—Fuses at 5 in R.F. with 10 times its weight of silica to a yellowish-brown opaque bead. Becomes darker in H.F. With the fluxes gives the characteristic reaction. Decomposed by hydrofluoric acid, with evolution of sulfur dioxide and separation of gelatinous silica.

Obs.—Occurs at Schwickbröler and Breiuhriin, in Saxony; at Knipf. It is also in the pegmatite veins of the Lüneburger Heide; in the Ilueu Mrs. Deir Minsk in pegmatite. In the U. S., with specularite at the mica mine near Anieila Court-House, Amelia Co., Va.; etc. Named by Werner, in allusion to its yellow color, from *gelb*, i.e., yellow.

Danaite. $(\text{Be,Fe,Zn,KD})_2(\text{SiO}_3)_2$. In octahedral crystals. H. = 5.5-6. O. = 3.437. Color flesh-red to gray. Occurs in small crystals in the Rockport granite. Cape Cod, Mass.: at the Iron mine at Bartlett, N. H.: El Paso Co., Colorado.

Bulvitfl. $\text{Bi}_2(\text{SiO}_3)_2$. Wlammhblende, Kieselsäure-Blende. Usually in small tetrahedral crystals; also in spherical forms. H. = 4-5. Q. = 5-106. Color dark hair-brown to grayish, or colorless. Found with native bismuth near Schneeberg, Saxony; also at

JoiunngeargeuBladl In crystals on quartz.

Zunylte. K highly basic orthnslllcale of aluraintiim, (A.l(OB.F.CI>i)tAl,SliO,,. Id minute transparent tetrahedrons. H.= T. 6.= ^ 3'875. From the Zu&l mine, near Bilver-ton, San Jaan Co., and on Red Monnlain, Ouray Co., Colorado.

4> Garnet Group. Isometric E.E,(SiO.), or 3BO.R,o,.3SiO.,

E = Ca,Mg,Fe,Mn. B = Al,Fe,Cr,Ti. Oftrott

A. Gbossulabite Ca,Al,(SiO.), D. Spessartite Mn,Al,(SiO.),

B. Pyropb Mg,Al,(SiOJ, E. Andradite Ca,Fe,(SiO.).

C. Almandite Fe,Al,(8iOJ, Also (Ca,Mg),Fe,{SiO.),,

Ca,Fe,{(Si,Ti)o,}, F. UvABOVITE Ca,Cr,(SiO,), Sohorlomite Ca,(Pe,Ti).C(Si,Ti)o,),

The Garnet Group includes a series of important sub-species included tinder the same specific name. They all crystallize in the normal gronp of the isometric system and are alike in habit, the dodecahedron and trapezo-hedron being the common forms. They have also the same general formula, and while the elements present differ widely, there are many intermediate Tareties. Some of the garnets incltide titanium, replacing silicon, and thus

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the; are connected with the rare appecies soborlomite, which probably also has the same general formula.

Closet; related to Ibe Gabnkt Gnovr proper are the species ot the Bodallte and Helvlte Groups (p. 411, p. 413). All are cbaracterized by Isometric cry stall! zalion, uad all are orthoailcates, with aimilar cliemical itructure. Thus llie formula ot the Garuet Group is

U|Rt(S10,)i; lo this Sodallte conforms if wr biraleot radical AlCl are equivalent In Ri; si[i>f the hivalent group Na8oi-AI is esauined,

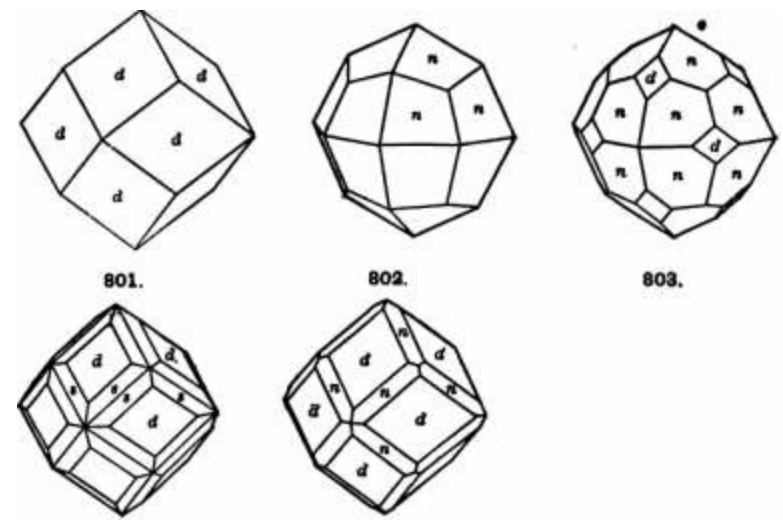
111 the Helvlte Group, whlcb is cbaraclerlzcd b; the tetrabednd character of the species (perhaps true also of the Sodnlites), the chemiical relation is leas close but probably exists, as exhibited by nrUing the formula of Helviie (Mu,FexHQ,a)Be.(SiO.)i. where the bivaleDt group -8-Mu-8- euLers, aud 8Be may be regarded as tuklug the place of 3A1.

QASNET.

Isotnetric. The dodecahedron and trapezohedron, n (211), the common simple forms; also these in combination, or with the hexoctahedron « (321). Cubic and octahedral faces rare.

Often in irregular embedded grains. Also massive; granular, coarse or fine, and sometimes friable; lamellar, lamellae thick and bent. Sometimes compact, cryptocrystalline like nephrite.

7&8.



Parting: d sometimes rather distinct. Fracture subconchoidal to uneven. Brittle, sometimes friable when granular massive; very tough when compact cryptocrystalline. H. = 6.5-7.5. G, = 3.15-1.3, varying with the composition. Luster vitreous to resinous. Color red, brown, yellow, white, apple-green, black; some red and green, colors often bright. Streak white. Transparent to eubtransparent. Often exhibits anomalous double refraction, especially grossularite (also topazolite, etc.), see Art. 411. Refractive index rather high, thus n, for:

Grossularite 1-7646, Pyrope 1-777, Almandine 1-7716.

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Comp[^] Tat. —An orthosilicate having the general formula $R_2U_2(SiO_4)_3$, or $3RO.R_2O_3SiO_4$. The bivalent element may be calcium, magnesium, ferrous iron or manganese; the trivalent element, aluminum, ferric iron or chromium, rarely titanium; further, silicon is also sometimes replaced by titanium.

There are three prominent groups, and various subdivisions under each, many of these blending into each other.

I, Aluminium Garnet, including

A. Grossularite Calcium-Aluminium Garnet $Ca_3Al_2(SiO_4)_3$,

B. Pyrope Magnesium-Aluminium Garnet $Mg_3Al_2(SiO_4)_3$, C. Almandine Iron-Aluminium Garnet $Fe_3Al_2(SiO_4)_3$, D. Spessartite Manganese-Aluminium Garnet $Mn_3Al_2(SiO_4)_3$,

IL Iron Garnet, including

K Almandite Calcium-Iron Garnet $\text{Ca}_3\text{Fe}_2(\text{SiO}_4)_3$,

(1) Ordinary. (3) Magnesian. (3) Titaniferous. (4) Yttriferous. III. Chromium Garnet,

F. Uvulite Calcium-Chromium Garnet $\text{Ca}_3\text{Cr}_2(\text{SiO}_4)_3$.

The quality of the U from the Linnæus granates, roseau DK like a grain, and directly from the magnesian, the seeds of which are small, numerous, and red, in allusion to the aspect of the crystal.

A. Grossularite. Euxenite or Heusonite. Cinnamon-stone. Calcium-aluminium Garnet. Formula $3\text{CaO} \cdot \text{Al}_2\text{O}_3 \cdot 3\text{SiO}_2$, = Silica 40.0, alumina 22.7, lime 37.3 = 100. Often containing ferrous iron replacing the calcium, and ferric iron replacing aluminium, and hence graduating toward groups O and E. G. = 355 to 366. Color (a) colorless to white; (i) pale green; (c) amber and honey-yellow; (rf) wine-yellow, brownish yellow, cinnamon-brown; (e) rose-red; rarely (i) emerald-green from the presence of chromium. Often shows optical anomalies (Art. 411).

The original grossularite (miuitt pt.) included the pale green from Siberia, and was so named from the botanical Dame for the gooseberry in 1743-9. Cinnamonite, or euoni (more properly hutonite). Included a cinnamon-colored variety from Ceylon, there called hyacinth; but unlike this none the yellow and yellowish-red kinds included: named from rhyacine, inferior, because of less hardness than the true hyacinth which it resembles. Suezite is an amber-colored kind from the Ala valley. Piedmont. Bomanzani is brown.

Pale green, yellowish, and yellow-brown garnets are not invariably grossularite; some (including topazolite, demantoid, etc.) belong to the group of Calcium-Iron Garnet, or Andradite.

B. Pyrope, Precious garnet pt. Magnesium-aluminium Garnet. Formula $3\text{MgO} \cdot \text{Al}_2\text{O}_3 \cdot 3\text{SiO}_2$, = Silica 44.8, alumina 25.4, magnesia 29.8 = 100. Magnesia predominates, but calcium and iron are also present; the original

pyrope also contained chromium. G. = 370-375. Color deep red to nearly black. Often perfectly transparent and then prized as a gem. The name pyrope is from *pyropos*, fire-like.

Pyrope of delicate shades of pale rose-red and purple, brilliant by reflected light, corresponds to two parts of pyrope and one of almandite; from Macon Co.,

C. Almandite. Almandine. Precious garnet pt. Common garnet pt. Iron-aluminium Garnet.

Formula $3\text{FeO}, \text{Al}_2\text{O}_3, 3\text{SiO}_2 = \text{Silica } 36.3, \text{ alumina } 20.5, \text{ iron protoxide } 43.3 = 100$. Ferric iron replaces the aluminium to a greater or less extent. Magnesium also replaces the ferrous iron, and thus

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it graduates toward pyrope, of. rhodolite above. Q. = $3.9-4.3$. Color fine deep red, transparent, in precious garnet; brownish red, translucent or semi-translucent, in common garnet; black. Part of common garnet belongs to Almandine.

The Almandine garnets were so called because cut and polished at Almaden. Hence the name Almandine is still in use.

D. Spessartite. Spessartine. Mangano-calcium Garnet. Formula $3\text{MnO}, \text{Al}_2\text{O}_3, 3\text{SiO}_2 = \text{Silica } 36.4, \text{ alumina } 20.6, \text{ manganese protoxide } 43.0 = 100$. Ferrous iron replaces the manganese to a greater or less extent, and ferric iron also the aluminium. Q. = $4.0-4.3$. Color dark hyacinth-red, sometimes with a tinge of violet, to brownish red.

E. Andradite. Common Garnet, Black Garnet, etc. Calcium-iron Garnet. Formula $3\text{CaO}, \text{Fe}_2\text{O}_3, 3\text{SiO}_2 = \text{Silica } 35.5, \text{ iron sesquioxide } 31.5, \text{ lime } 33.0 = 100$. Aluminium replaces the ferric-iron; ferrous iron, manganese and sometimes magnesium replace the calcium. G. = $3.8-3.9$. Colors various: wine-, topaz- and greenish yellow, apple-green to emerald-green; brownish red, brownish yellow; grayish green, dark green; brown; grayish black, black.

Named after the Portuguese mineralogist, Almandine, who in 1800 described and included in his classification, Almandine. Generally there are the following varieties;

1. Simple Calcium-iron Garnet, In which the protoxides are usually or entirely without lime. Includes: (a) Tapolite, having the color and transparency of Andradite, - - -

sometimes green; crystals often forming a vicinal hexahedron. Xanthopside, a variety of emerald-green variety with brilliant diamond-like luster, used as a gem. (b) Colophonite, a coarse granular kind, brownish yellow to dark reddish brown in color, reddish in luster,

usually with iridescent hue; in some cases the resin colophony, (c) Naumannite (ferric iron, black, either dull or lustrous; but no black garnet is here included. Ferrous iron grayish black mineral. (d) Dark green garnet, not distinguishable from some allanite, except by chemical analysis.

S. Mangano-calcium Garnet. (e) Sothite. The original atfokro

varieties of iron-garnet of brown or reddish-brown color, and of blue-green color

. Hot/iome, from LSiigUiin, is dinilar, yellowish browu to llv "

kinds of Tnungancsiai iroti-garnet are ligLt nuil dark, dusky gie

Uothomfe, from LSiigUiin, is siniilar, yellowish browu to llvet-bronD Olber

Linds of Tnungancsiai iron-garnet are ligLt imd dark, dusky giecii nod black, and

often in crystals. PoCymletphiU is a miissive browuisb-ytlflow iud, from Franklin Fiiniuc

N. J. Brtdbtrffile, from Sa\ a. conlains a large amount of magnesia, <A) Aplcme (prnperiy

..__, ,__ = . _ ^ , , , , , __ , , --■-■--■' -Hi, wliei.ce Hat-

sluipier iban ll dodecahedroD. he gdvi; it a mime derived froui dnXun', simple. CuIot oi Ibe oi'lgfiul

iiBplome) bas lis dodecaliedrai faces sirialed imnillel lo tlie sborttr diagonal, wlienci inferred Ibat the fundamental form was tbe cube; iiml as tthis form U sluipier iban Ibe

... . e (of unknown locality) dark browu: nlso found ^ellowlsb green and bvowuisb green at Schwarzzeiibeig iu Siixony, and ou tlio Lena in Siberia.

3. TUanifermiii. CoutaJna thauinm aud probably liotb TiO, and Ti.Oi; formula hence $8\text{CaO} \cdot (\text{Fe} \cdot \text{Tr} \cdot \text{Al}) \cdot \text{O} \cdot 3(8\text{l}, \text{Ti})\text{O}$. It Ibiis graduales toward seliorloniile. Color black.

4. ytrr^erimt Caleiam-iron Qarnet.-Ttlergranal. Conlains yltria iu small amount;

F. UvAROVITE. Ouvarovifce. tJwarowit. Calcium-cltromium Garnet, Formula $3\text{CaO} \cdot \text{Cr}_2\text{O}_3 \cdot 3\text{SiO}_2$, = Silica 35'9, chromium sesquioiide 306, lime 33-5 = 100. Alaminim takes the place of the chrominm iu part, H. = 7*5. G. = 3'41-3'52. Color emerald-green.

Pyr., ato.—Moat rarielles of gurnet fuse easily lo a ligUt brown or blnck glass; F. = 8 In almandlte, spessartite, groaaulnrite, and allocbroite; 8'S tu pyrope; but uvarovite, Ihe I-11 route-gurnet, is almost Infusible. F. = 6. A.ltochrolte and altnanillte fuse to a magnetic globule. Re:ictions with tbe flnicis vary wftb llie bases. Almost all kinds react for Iron; Btr>iig ninngiineM reaction in spe'sanile. and less marked iu other varielies; a cbrumium re:tcioii In uvarovite. and in most pvropt. Some varieties are partially decomposed by iidils* all except uvarovite are after ignition decomposed by liynrocblnric acid, and generally with separation of gelatlnoiit silica on evapop,tion. Decomposed on fusion with i.liuiliio carbonates.

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Tlie dcDsIty of gurnets is largely dimluisbed bj fuaion, Tbus a GreenlaDd girnet fell from 3'90 to 8'05 on fusion, and ii Vilui gro-isularite from 8-63 lo 3'95.

DUE,—Characterized by isometric crystallization, usually in isolated crystals, dodecahedral or trapezohedrons; massive forms rare, then usually granular. Also distinguished by hardness, vitreous luster, and in the common kinds the fusibility, Veuviaioile fuses more easily, zircon and quartz are infusible; the specific gravity is higher than for tourmaline, from which it differs in form; it is much harder than sphalerite,

Distinguished in thin sections by its very high relief; lack of cleavage; isotropic character; usually shows a pale pink color; sometimes not readily told from some of the spinels,

metamorphosed in metamorphic pro-

cesses commonly found in the contact zone of igneous rocks and in the crystalline schists. It is characteristic of the mica schists and metamorphic rocks containing alumina and iron: it occurs also in some igneous rocks as the result of later metamorphic processes: it forms with smagmite the rock eclogite. Pyrope is especially characteristic of such basic igneous rocks as are formed from magmas containing much magnesia and iron with little or no alkalis, as the peridotites, dunites, etc.; also found in the serpentines formed from these rocks; then often associated with spinel, chromite, etc. Spinel occurs in granitic rocks, in quartzite. In the Belgian schists (Belgium); it has been noted with topaz in lithopbysea in rhyolite (Colorado), The black variety of andradite, melanite, is common in eruptive rocks, especially with upehllite, leucite, thus in phouolitea, g. leucitophyres, nephelinites: in such cases often titanifer-

ous or associated with a titaniferous garnet, sometimes in

zonal intergrowths; it also occurs in a product of contact metamorphism. Demantoid occurs in serpentine. Uvativite particularly with chromite in serpentine; it occurs also in granular nephelinites,

Garnet crystals often contain inclusions of foreign matter, but only in part due to alteration; in vesuvianite, calcite, epidote, quartz (Fig. 4T2, p. 141); at times the garnet has a mere shell, or perimorph, surrounding a nucleus of another species. A black garnet from Areodni, Norway, contains both calcite and epidote; crystals from Tvedestrand are wholly calcite within, there being but a thin crust of garnet. Crystals from East Woodstock, Maine, are dodecahedrons with a thin shell of calcite enclosing calcite; others from Baymund, He, show successive layers of garnet and calcite. Many such cases have been noted.

Garnets are often altered, thus to chlorite, serpeninite: even to limonite. Crystals of pyrope are sometimes surrounded by a chloritic zone (kelyphite of Schrauf) not homogeneous, as shown in Fig. 804,

Among prominent foreign localities of garnets, besides those already mentioned, are the following—GILSHAN: Fine garnet comes from Ceylon; on the Hussa-

Alp Id the Ala vnley iu Piedmont, with clinochlcire and diopaide; atZermatl; paleyellowal Auer-bach; brownish {romamovUt) at Kimilo in Finland; honey-yellow oelalidTont in El In; pale greenish from the banks of the Vilui in Siberia, in serpentine with veauvianlte: also fi-om Cziklowa and Orawiiza iu Ihe Banat; with vesuvlanlte and woliaatoulte in ejected masse. ! hi Vesuvius: in while orcolorlcss crystals in Tellemark, in Norway; also dark brown at Mudgee, New South Wales; dark honey-yellow al Guadslcazar, aud clear pink or rose,red dodeca-hedrous nt Morelos, Meilco.

Pykupe: In serpenti[ie (from perldotlte) near Meronltz and the valley of Crema, in Bohemia (used as a gem); at ZOblitz in Saxony; in the Yosges; in the diamond diggiugg of South Africa ("Cape rubles"), Aluanditk: Common in granite, gneiss, ecllo^te, etc., in many localities in Saxony, Silesia, etc.: at Eppenreuth near Hof, Bavaria: in large dodecahedrons at Falun In Sweden : hyacinth-red or brown in the Zillerthal. Tyrol. Precious garnet comes in line crysinls from Ceylon, Pegu. British India, Brazil, and Greenland, Sfbbsartitb: From AscbaSenburi; in the Speeaart, Bavaria; al St, Marcel, Piedmont; near Chnoteloube, Hauie-Vienne. etc.

Andradite: The beautiful green dnnantotif or"UrHIIaoemendd"occurs in transpnrnt greeniab rolled pebbles, also in cryslais. in the gold washings of Nizlmi-Tagilsk in the Ural; green crystalsocRur at Schwarzenberg, Saxony; brown to green at Morawitza and Dognacaks; emerald-green al Dobschau; in the Ala valley. Piedmont, tbe yellow to greenish topaaatitt. AlioehroiU, apple-green and yellowish, occurs at Zermatt; black crystala (mtlaattt), alio

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S1LICATB8.

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browD. at VmutIu* oh Hie. Somma; near Barigef Id tlie Hauies-PyrSnSes (pgreiitila), ApU^me occurs at BcbwarzeDberg in Haiouy, io bruwu to black cryBtals. Other locallIeB are WlMlilhul, Tjrul; LAugbaili, Sweden; Pitfearjuia, Finlaod; Areudni, Norway. Uvakotitb: FuuDil at SHraiiovskiiyt Dear Bisersk, also in Ihe viclulty of Kyeblymsk, Uriil. iu clirumlc iroB; lit JordansmOh], Slleala; Pic Posets near VSnuaque In the Pyrenees ou cbitinilte.

In N. Amtrlcfi. in Maine, benutrlul cryetals of cinnamou-stOQe Ivilh TVBuvlaililte) occur at PnnonBfleld. Phippsburg. nud Ruinron). In N. Hamp., at HaoOTer, amall clear cryMala it) gneiss; iit Warren, ciDUiimon gurnets; at Qraftun. In Vmnont. at New Fane, ia clilorile Ntaiu. tn itfiiu , iu guciM iit Bmokfleld; iu flue dark red or nearly black tnipezolidral cryainls >it Riissell, gomeilmes very large. In Conn , trapezobedrous, In mica Blute. al Reading 1111(1 Monroe; dodecaliedroiia at Boullibury; at Bodilam, crystals of Bpessarrlle. In iV. York, brwn crvHlals al Crown Polut, Essex Co.; cnloplionite aa a large vein at WillHboni, Essex Co.; in MidtDetown, Delaware Co.. large brown cryst.; a clnniimou

variety at Amity. In N. Jersey, at Franklin, black, brown, yellow, red, and green dodecahedral garnets; also near the Franklin Furnace (polyadelphite). In Penn., in Chester Co., at Peunsbury, fine dark brown crystals; near Kutztown; at Clifton, brown; in Concord, on Green's Creek, resembling pyrope; in Leipsville, red; at Mineral Hill, fine brown: at Avondale quarry, fine leucite; at Tarovite at Woods' chrome mine, Lancaster Co. In Virginia, (leucite) transparent spear-shaped, used as a gem, at the mica mines at Amelia Court-House. In N. Carolina, fine leucite-stone at Bakersville; red garnet in the gold washing of Burke, McDowell, and Alexander counties; rhodolite in Macon Co.; also mined near Horriatown in Warlick, Burke Co., to be used as "emery" and as "garnet-paper." In Illinois, fine pyrope in the peridotite of Ellis Co. In Arkansas*, at Magnet Cove, a titaniferous melanite with schorlomite. Large dodecahedral crystals altered to chlorite occur at the Spurr Mt. iron mine. Lake Superior. In Colorado, at Natron, fine spessartite crystals in rhyolite; in large dodecahedral crystals at Buby Mt., Salida. Chelan Co., the exterior altered to chlorite. In Arizona, yellow-green crystals in the Gila cañon; pyrope on the Colorado river in the western part of the territory. New Mexico, fine pyrope on the Navajo reservation with chlorite and a chrome-pyroxene. In California, green with copper ore. Hope Valley, El Dorado Co.; uvarovite, in crystals on bromite, at New Idria. Fine crystals of a rich red color and an inch or more in diameter occur in the mica schists at Fort Wadsworth, mouth of the Truckee R., in Utah.

In Canada, at Marmora, dark red: at Grenville, a cinnamon-stone; an emerald-green chrome-garnet, at Orford, Quebec, with millerite and calcite: fine colorless to pale olive-green, or brownish crystals, at Wakefield, Ottawa Co., Quebec, with white pyroxene, honey-yellow vesuvianite, etc., also others bright. In the country of the St. Lawrence; dark red garnet in the lowlands of Villeneuve (spessartite) and Templeton.

Schorlomite. Probably analogous to garnet. $8CaO.(Fe.Ti)_2O_3.Si_2O_6$. Schorlomite. Usually massive, black, with conchoidal fracture and vitreous luster. H. = 7-7.5. G. = 8-8.1-8-68. From Magnet Cove, Arkansas.

Partschinite. $(Mg,Fe)_2Al_2Si_2O_{10}$, like spear-shaped. In small dull crystals (monoclinic). H. = 6-6.5, O. = 4.006. Color yellowish, reddish. From the auriferous sands of the Abipian, Transylvania.

Androvalite. Same as for eulytite, $Bi_2Si_2O_{10}$, but monoclinic. In globular or globular forms. From Johanngeorgenstadt.

,ab,GoOgIc

DESCRIPTION OF HINESALOOT.

The Chrysolite Group includes a series of orthosilicates of magnesium, calcium, iron and manganese. They all crystallize in the orthorhombic system with but little variation in

variety.

Comp.—(Mg,Fe),SiO₂ or 2(Mg,Fe)O,SiO₂. The ratio of Mg : Fe varies widely, from 16 : 1, 13 : 1, etc., to 2 : 1 in hyalosiderite, and hence passing from fosterite on the one side to fayalite on the other. No sharp line can be drawn on either side. Titanium dioxide is sometimes present replacing silica; also tin and nickel in minute quantities.

Pyr., etc.—B.B. whitens, but is infusible in most cases; hyalosiderite and other varieties rich in iron fuse to a black magnetic globule: some kinds turn red upon heating. With the flux gives reactions for iron.

See various other reactions

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BTLICATE9. *»1

for titanium in igneous rocks. Decomposed by hydrochloric and sulphuric acids with separation of iron as ferric chloride.

Diff —Characterized by its infusibility, the yellow-green color, granular form and cleavage (quartz has none).

Recognized in thin sections by its high relief; lack of color: has few but marked rough cleavage-cracks; high interference-colors, which are usually the brilliant and produced tones of the second order; parallel extinction; biaxial character; characteristic of the 80a 809. 810.

1U/^W

(usually with acute terminations) when in distinct crystals (Figs. 807-809). its frequent association with iron ore and augite, and its very common alteration, in a greater or lesser degree, to serpentine, the first signs being marked by the separation of iron-ore grains; the lines of fracture (Fig. 810).

Obs.—Chrysolite (olivine) has two distinct modes of occurrence: {a} In igneous rocks, as usually, formed by the crystallization of magma low in silica and rich in magnesia; from an accessory component in such rocks the olivine may increase in amount until it is the chief rock constituent as in the diorites; also {b} as the product of metamorphism of certain sedimentary rocks containing magnesia and silica, as in impure dolomites. In the igneous and peridotites of igneous origin the chrysolite is commonly associated with chromite, spinel, pyrope, etc., which are favorable indications also of the origin of serpentine derived from olivine. In the metamorphic rocks the above are wanting, and carbonates, as dolomite, brucinerite, magnesite, etc., are the common associations; chrysolitic rocks of the latter kind may also occur altered to serpentine.

Chrysolite also occurs in grains, rarely crystals, embedded in some meteoric Irons. Also present in meteoric stones, frequently in spherical forms, or chondrules, sometimes made up of a multitude of grains with like (or unlike) optical orientation enclosing glass between (Fig. 811). *^*

Among the more prominent localities are: Vesuvius in lava and on Monte Somma in ejected masses, with magnetite, etc. Observed in [the so-called Sindhian bombs at the Laacher See; at Forstberg near Mayen in the Eifel and forming the mass of "olivine bomb" in the Dreiser Weiler near Daun in the same region] at Sasbach in the Kaiserstuhl. Baden (Hyalitidite). In Sweden, with ore-deposits, as at Långban, Palsberg, Pereberg, etc. In Berpeitua at Snarum, Norway. In large crystals, themselves altered to the same mineral. Common in the volcanic rocks of Sicily, the Sandwich Islands, the Azores, etc.

In the U. S., in Thetford and Norwich, Vermont, in boulders of coarsely crystalline basalt, the crystals or masses in places through. In the granite of Waterville, in the White Mountains, N. H.; at Webster, in Jackson Co., N. C. with serpentine and chromite; with chromite in Loudon Co., Va.; in Lancaster Co.,

Pa. In small clear olive-green grains with garnet in some granites in Arizona and New Mexico. In Canada in Cannville, near Montreal, at Rougemont and Mounts Royal and Montarnelle, and in eruptive rocks at other points.

Chrysolite chondrule from the Knyahinya meteorite (Lodovici).

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Alteration of chrysolite often takes place through the oxidation of the iron; the mineral becomes ferruginous or reddish brown and iridescent. The process may end in tearing the cavity of the crystal filled with limonite or rust oxide of iron. A very common kind of 'oxidation' is the byproduct of limonite, serpentine, with the partial removal of the 'Fe'; this

...^ ^ ^ „ _ . ^ , _r7.

Chrysolite is derived from chrysolite, gold, and iron. The titaniferous, from ,

Sialite, iron, or /j; /puC, iron. The chrysolite of Pliu was probably our topaz; and has topaz

Chrysolite. From the rock chrysolite of Carmelo Bay, California; a rare reaction product of chrysolite, exact composition undetermined. Orthorhombic, flattened double crystal. G, = a-838. Color brown,

The axial ratios of the other members of the Chrysolite Group are given in the table on p.

41U. Tite apei'ies are hriedj characlerized as follows:

Monticellite. CaMgSiO_4 . Occurs in colorless to ^'ray crystals on Mie, Somma : Id iiiasses (b-ilrac'iiUi ou Alt. Monzool, I'yrol; iu ci-ysials or grains fn iimestone at Magnet C.ive, Arkainos. G. = 3-03-3-25.

Fonterit«. Mg_2SiO_4 . Occurs id while crystals at Vesuvius; iu greenish or yellowtsh embedded grains nt B<dIon, Ma&9. Ipetlonile). G = 8-2i-S'88.

Hortonolite. $(\text{Fe},\text{Mg},\text{Mn})_2\text{SiO}_4$. In rough dark-colored crystals or masses. Occurs at the ii-ou mine of Monroe, Ononge Co , N. Y. G. = BM,

Fayalite. Fe_2SiO_4 . Fnuu the Moiiine MIh., Ireland; the Azores; the Yell ovfstoue Park; Rockport, Mass., etc. Crystals and mnssive, brown to black on exposure. Q. = 4'1,

Knabeolite. $(\text{Fe},\text{Mn})_2\text{SiO}_4$. From Danueinora, and elsewhere in Sweden, G = 41,

Tephroite. Mg_2SiO_4 ; alao with zinc, In the vaAely rotpperte. From Sterling Hill and Franklin Furnace, N. J.; also from Hweden. Color flesh-red to Hsh-gray. G. = 41.

Pfenacite Group. R_2SiO_4 . Tri-rhombohedral. ■ rr' i

WUlemite Zn_2SiO_4 , 64° 30' 0-6775

Troostite $(\text{Zn},\text{Mn})_2\text{SiO}_4$.

Phenacite Be_2SiO_4 , 63° 24' 0-6611

The PfiENACiTE Group includes the above ortbosilicates of zinc (man-ganeae) and beryllium. Both belong to the tri-rhombohedral group of the trigonal division of the hexagonal system, and have nearly the same rhombo-hedral angle. The rare speeia trimerite, $\text{Mn}_3\text{SiO}_{10}$, $\text{Be}_3\text{SiO}_{10}$, which is psendo-bexagoual (tricliuic) is probably to he regarded as connecting this gronp with the preceding Chrysolite Group.

The following rare species are related:

rr' i

Diopside CaMgSiO_6 ,

Friederite $\text{H}_2(\text{MnCl})\text{Mn}_2(\text{SiO}_6)_2$.

Pyroimolite $\text{H}_2((\text{Fe},\text{Mn})_2\text{Cl})(\text{Fe},\text{Mn})_2(\text{SiO}_6)_2$,

These species are very near to each oilier In form, aa shown In the above axial ratios; they fiirlhar approximate lo Ihe Bpecles of the Phenacite Group proper. They are also cloeely

related among themselves in composition, since they are all add ortho-silicates, and have the general formula $H_nR_8SiO_{10} = H_nR_n(SiO_4)_4$, where (a.g. for Friedelite) to the latter form the place of one hydrogen atom is taken by the univalent radical (MnCl).

Tri-rhombobedral. Axis $a = 0.6776$; $\alpha = 64^\circ 30'$; $\beta = 36^\circ 47'$.

In hexagonal prisms, sometimes long and slender, again short and stout; rarely showing subordinate faces distributed according to the phenacite type. Also massive and in disseminated grains; fibrous,

Cleavage; c easy, Moiré; difficult, K J.; a easy, N.J. Fracture conchoidal to uneven. Brittle. H. = 5-5. Q. = 3-89-4-18. Lustre Titre-

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SILICATES.

reevous, rather weak. Color white or greenish yellow, when purest; apple-green, fleshy-red, grayish white, yellowish brown; often dark brown when

$f^X \wedge \wedge$

$H^$

a 11 a 10.

Figs. 812-814. New Jersey, e (Oil) * (1138), u (3118), i (SiSi). inipiiie. Streak uncolored. Translucent to opaque. Optically +. Birefringence high.

Comp Zinc orthosilicate, Zn_2SiO_4 or $\sim ZnO \cdot SiO_2$, = Silica 27.0, zinc oxide

73.0 = 100. Manganese often replaces a considerable part of the zinc (in troostite), and iron is also present in small amount,

Pyr., «tc.—B.B. in the tube glows and fuses with difficulty to a white enamel: the varieties from New Jersey fuse from 8.5 to 4. The powdered mineral on charcoal in R.F. gives a blue color, yellow while hot and white on cooling, which, moistened with solution of cobalt and heated in O.P., is colored bright green. With soda the dialing is more readily obtained. Decomposed by hydrochloric acid with evolution of potassium silicate.

Oba—Proin Altheim near Morehead: at Solberg, near Ali-In-Chapelle. In New Jersey at Mine Hill, Franklin Furnace, near Slerlinj; Hill, two miles distant. Occurs with zincite and franklinite, varying in color from white to pale honey-yellow and light green to dark ash-gray and flesh-red; sometimes in large reddish crystals (troostite). Rare at the Merritt mine, Socorro Co., New Mexico: also at the Sedalia mine, Salda, Colo. Named by Levy

after William I., King of the Netherlands.

PHENAorm.

Tri-rhombohedral. Axis $i = 0-6611$; $rr' = 63^\circ 24'$.

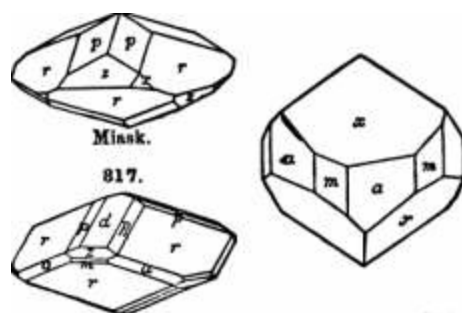
Crystals commonly rhombohedral in habit, often lenticular in form, the prisms wanting; also prismatic, $g^j g^g$

sometimes terminated by the rhombohedron of the third series, X (see further, pp. 80-82).

Cleavage: a distinct; r imperfect. Fracture conchoidal. Brittle. H. = 7-5-8. G. = 2-97-300. Luster vitreous. Colorless; also bright wine-yellow, pale rose-red; brown. Transparent to subtranslucent. Optically -I-. $ooj = 1-6540$; $e = 1-66fli$, Framont.

Coup.—Beryllium orthosilicate, Be_2SiO_5 or $2BeO \cdot SiO_2$, = Silica 54-45, glucina 45.55 = 100.

Pyr., etc.—Alone lenticular; unaltered;



Florissant. Colo. Mt. Antero, Colo., Pfd,

: fused with extreme slowness, unless

DEBCKIPnTE HINBEALOOT.

When fused, to become transparent glass. With soda affords a white enamel; with more, intensifies the blue (1 becomes fusible). Dull blue with cobalt solution.

Obs.—Occurs as emerald and emerald mine of Takovaya, 80 stadia E. of Ekaterinburg; also in the Ilmeu Mts., near Miask; Dear Framont in the Voiege Hta.; at the Cerro del Mercado, Durango, Mexico.

In Colorado, as amazon-stone, at Topaz Butte, near Florissant, 15 miles from Pike's Peak; also on quartz and beryl at Mt. Antero, Colorado. Named from *ipizra*, a Quechua word, in allusion to its having been mistaken for quartz.

Trimnet*. $(H_2O)_2SiO_5 \cdot Be_2O_3 \cdot Bi_2O_3$. In black tabular prismatic crystals, pseudo-hexagonal

(trichloric) in form and angle. H. = 6-7. Q. = 3 474. Color salmon-drink to nearly black. From the Harsig mine, Wermland, Sweden.

Dioptase. H_2CuSiO_5 or $H_2O.CuO.BiO_2$. Commonly in prismatic crystals ($m = 84^\circ$ SS). Also in crystalline aggregates; massive. Cleavage perfect. Fracture conchoidal to uneven. H. = 5. Q. = 28-30. Luster vitreous. Color emerald-green. Occurs in druses and well-defined crystals on quartz, occupying seams in a compact limestone west of the mine of Altyu-Talbe in the Khrshese district.

Steppe: in the gold washings at several points in Siberia; at Kczbanya, Hungary. From the Congo State, also in the copper mine of Clifton, Gtiam Co., Aizoua.

Occurs in druses and well-defined crystals on quartz, occupying seams in a compact limestone west of the mine of Altyu-Talbe in the Khrshese district.

Steppe: in the gold washings at several points in Siberia; at Kczbanya, Hungary. From the Congo State, also in the copper mine of Clifton, Gtiam Co., Aizoua.

Hue crystals at the Aline Mindouli, two leagues east of Comb, in the Congo State, also in the copper mine of Clifton, Gtiam Co., Aizoua.

Hue crystals at the Aline Mindouli, two leagues east of Comb, in the Congo State, also in the copper mine of Clifton, Gtiam Co., Aizoua.

Preh Congo State, also in the copper mine of Clifton, Gtiam Co., Aizoua.

Vedelite. $H_2(HoCl)MgSi_4O_{12}$. Crystals commonly tabular; also massive, cleavable to closely compact. H. = 4-6. Q. = 80. Color rosed. From the manganese mine of dervelle, vallee du Lourou, Hautes Pyrenees.

Fyrosmalite. $HT((Fe,Mg)Cl)(Fe,HD).Si_3O_{10}$. Crystals thick hexagonal prisms or liliulir: also massive, foliated. H. = 4-6. G. = 6-3'ig. Color blackish green to pale liver-brown or gray. From the Iron mines of Nordmark in Wermland; Danuemora, Sweden.

Scapolite Group. Tetragonal-pyramidal. Helonite = 0-4393 Hiszooite, Dipyre = 0-4424

SarcoUte = 0-4437

The species of the Scapolite Group crystallize in the pyramidal form of the tetragonal system with nearly the same axial ratio. They are white or grayish white in color, except when impure, and then rarely of dark color. Hardness — 5-6-5; G. — 2-5-2-S. In composition they are silicates of aluminum with calcium and sodium in varying amounts; chlorine is also often present, sometimes only in traces. Iron, magnesia, potash are not present unless by reason of inclusions or of alteration, which last can also explain the carbon dioxide often found in analysis.

The Scapolites are analogous to the Feldspars in that they form a series with a gradual variation in composition, the amount of silica increasing with the increase of the alkali, soda, being 40 p. c. in meionite and 64 p. c. in marialite. A corresponding increase is observed also in the amount of chlorine present. Furthermore there is also a gradual

change in specific gravity, in the strength of the double refraction, and in resistance to acids, from the easily decomposed meionite, with $G. = 2.72$, to marialite, which is only slightly attacked and has $G. = 2.63$. Tsohermak has shown that the

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Variation in composition may be explained by the combination of two fundamental end compounds, viz.:

Meionite $\text{Ca,Al,Si}_2\text{O}_6$, Me

Marialite $\text{Na,Al,Si}_2\text{O}_6$, Cl Ma

By the isomorphous combination of these compounds the composition of the species mentioned above may be explained; no sharp line can, however, be drawn between them,

Optically characterized by the decrease in the strength of the double refraction (n_D) in passing from meionite to marialite. Thus (Lacroix) for meionite $n_D - n_F = 0.080$; for typical wernerite 0.008-0.002; for dipyre 0.015.

The tetragonal species melilite and gehlenite are near the Scapolites in angle. The more c

nivianite is also related.

bqhohttb.

Tetragonal. $a = 0.43925$. In prismatic crystals (Fig. 179, p. 60) either clear and glassy or milky white; also in crystalline grains and massive. Cleavage: a rather perfect, m somewhat less so. Fracture conchoidal. Brittle. $H. = 5.5-6$ $G. = 2.70-2.74$. Luster vitreous. Colorless to white. Transparent to translucent; often cracked within. Optically —. Double refraction weak. Indices, D_x : $m, = 1.594$; $e, = 1.558$.

Comp.— $\text{Ca,Al,Si}_2\text{O}_6$, or $4\text{CaO} \cdot 3\text{Al}_2\text{O}_3 \cdot 6\text{SiO}_2$, = Silica 40.5, alumina 34.4, lime 25.1 = 100.

The varieties included here range from nearly pure meionite to the combination of meionite and marialite in the ratio of 3 : 1. No sharp line can be drawn between the following species.

Obs.—Occurs in small crystals in cavities, usually in limestone, Soinnia. Also in ejected masses at the Lancher See.

WERNEIRITB. Common Scapolite.

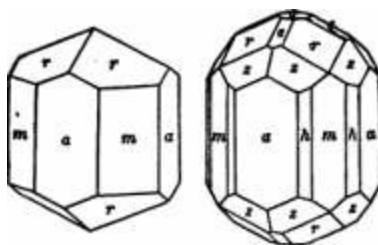
Tetragonal-pyramidal. $a = 0.4384$.

Crystals prismatic, usually coarse, with uneven faces and often large. The symmetry of the pyramidal group sometimes shown in the development of the faces $\{3U\}$ and $2, \{131\}$. Also massive, granular, or with a faint fibrous appearance; sometimes columnar.

$\langle, \rangle, 311 \wedge 29^\circ 48'$.

Cleavage: \langle and m rather distinct, but interrupted. Fracture subconchoidal. Brittle. $H. = 5 - 6$. $G. = 2.66 - 2.73$. Luster vitreous to pearly externally, inclining to resinous; cleavage and cross-fracture surface vitreous. Color white, gray, bluish, greenish, and reddish, usually light; streak uncolored. Transparent to faintly subtranslucent. Optically—. Birefringence weak. Indices: $t > r = 1.566$, $e = 1.545$ Ardenal.

Comp., Tar.—Intermediate between meionite and marialite and corresponding to a molecular combination of these in a ratio 3 : 1 to 1 : 2. The silica



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varies from 46 to 54 p. c, and as its amount increases the soda and chlorine also increase. Scapolites with silica from 54 p. c. to 100 p. c. are classed with mizzonite; they correspond to $Me : Ma$ from 1 : 2 to 1 : 3 and upwards.

. The percentage composition for a common compound is as follows: $He : Hb 8 : 1$ StO_2 , $46.10 Al_2O_3$, $80.49 CaO$ SiO_2 Na_2O $S-G_4$ $Cl 1.01 = 100\%$

Pyrrhotite.—B.B. fuses easily with effervescence to a white blebbly glass. Imperfectly decomposed by hydrochloric acid,

Diaphane.—Characterized by its square form and prismatic cleavage (90°); resembles feldspar in luster, but has a disordered structure on the cleavage surface; it is also more fusible, and has a higher specific gravity; also distinguished by its interference colors from pyroxene (see, p. 887).

Recognized in thin sections by its low refraction; lack of color; rather bright interference colors reaching the yellows and reds of the first order, sections showing which exhibit a distinct negative axial cross of basal sections which show the cleavage—Cracks cross at right angles.

Occurrence.—Occurs in metamorphic rocks, and is abundant in granular limestone near its association with the associated gneiss or allied rocks; sometimes in beds of magnetite

incomparable limestone. It is often associated with a light-colored pyroxene, amphibole, garnet, and also with apatite, (It is also associated with zircon; amphibole is a less common associate than

kyroxene, but in some cases has resulted from the alteration of pyroxene. Scapolite has been shown also to be frequently a component of basic igneous rocks, especially those rich in plagioclase; it contains much lime; it is regarded as a secondary product through a certain kind of alteration.

Principal localities are at Parras, Finland. It occurs in limestone; Arendal in Norway, and Malsjö in Wermland, where it occurs with magnetite in limestone. Patzsch is from Uberzell, near Pönnau, in Itzehoe. This is a pale blue or grey scapolite from Lake Baikal, Siberia, is called glawotit. In the U. S. it occurs in Vermont, at Ferrisburgh, massive. In Massachusetts, at Bolton; at Chelmsford. In New York, in Orange Co., Essex (Saratoga), Lewis Co.; Oriskany, Jefferson Co.; at Gouverneur, in Hamilton. In New Jersey, at Franklin and New Hope. In Pennsylvania, at the Elizabeth mine. French Creek, Chester Co.

In Canada, at Orono, Maine, massive; at Orono; Templeton; Wilkesfield, Ottawa Co. Scapolite rocks occur in several places.

Scapolite. Dipyrone, H₂TiO₄ are included scapolites with 64 to 67 p. c. SiO₂, correspond to a molecular formula from Me : Ma = 1; 2 to Me : Ha = 1; S. J. W. W. occurs in clear crystals in ejected masses on Mt. Somma.

Dipyrone occurs in elongated square prisms, often slender, sometimes large and coarse. It is lime-bearing and crystalline schists. Chiefly from the Pyrenees; also in diorite at Badale, Norway; Baiut-NH Ziure, France; Algiers. Couderans from the Pyrenees is a more or less altered form of dipyrone.

Marialite. Theoretically Na₂Al₂SiO₆Cl. see p. 426. The actual mineral corresponds to Al₂ : Ma = 1; 4. It occurs in a volcanic rock called piperno, at Planura, near Naples.

Barolite. (Ca,Na)₂Al₂(BiO)₂ In small tetragonal crystals. H. = S. G. = 2645-3033. Color Sesqui-red. From Monte Sommo.

bsblujtb.

Tetragonal. Axis (c = 0.4548. Usually in short square prisms (a) or octagonal prisms {a, m), also in tetragonal tables.

Cleavage: c distinct; a indistinct. Fracture conchoidal to uneven. Brittle. H. = 5. G. = 2.9-3.10. Luster vitreous, vitelline to resinous. Color white, pale yellow, greenish, reddish, brown. Pleochroism distinct in yellow varieties. Sometimes exhibits optical anomalies. Optically —. Birefringence low. Indices:

Ilumbottilite »; = 1-6812 »j = 1-6838 f, = 16263 f, = 1-6891 Henniger

Comp.—Perhaps R R, Si, O, or Na, {Ca, Mg}, {Al, Fe}, (SiO₂), for melilite. If Ca : Mg = 3 : 2, and Al : Fe = 1 : 1, the percentage composition is: Silica 37.5, alumina 11.1, iron 11.1, lime 31.3, magnesia 8.4, soda 4.3 = 100. Potassium is also present.

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SILICATES.

«7

With the fluxion

Pyr., Bto.—B.B, fuses at 8 to a yellowish or greenish glass. Decomposed by hydrochloric acid with gelatinization.

Diff. — Distinguished in thin sections by its moderate refraction; very low interference color, showing the "ultra blue" (Cibo di Bove); parallel extinction; negative character; uniaxial development. In tables parallel to the base slides very common "peg sinicula" due to parallel rod-like cleavage of the crystal from the base & parallel to the axis: this, however, is not always typical.

Obs.—Melilite is a constituent of certain igneous rocks formed from magmas very low in silica, rather deficient in alkali, and usually acidic or subalkaline. In the latter class it appears to crystallize in the place of lime more acid plagioclase.

Color of yellow to brownish colors is found at Monte Somma & di Bove. It occurs in leucite. It occurs with nepheline, niobite. It occurs in Vesuvius in dull yellow crystals. It is common in certain basic eruptive rocks, as the melilite-basalt of Hoclboh near Owau in Wharfedale, of the Schwanbian Alh. at Orlitz. the Krzgebirge: also in the nepheline syenite of the Tegau, of Oahu, Sandwich Islands, etc.; perovskite is a common associate. It occurs in various shifts. Melilite is derived from ueli, honey, in aUuslon to the color.

Humboldtite occurs in cavernous blinks on Monte Somma with greenish mica. also apatite, niobite; the crystals are often rather large, and covered with a calcareous coating; it is common in [rarely] lustrous crystals with nepheline. sarcolite. etc., in an augitic rock. Zirconite is impure humboldtite.

Orthoclase, CaAl₂Si₂O₇. Crystals usually short square prisms. Axis J = 0-4001. O. = 2-9-3 07. Diorthoclase shades of grayish green to liver-brown. From Mount Honzoll, the FasBBlhal, in Tyrol.

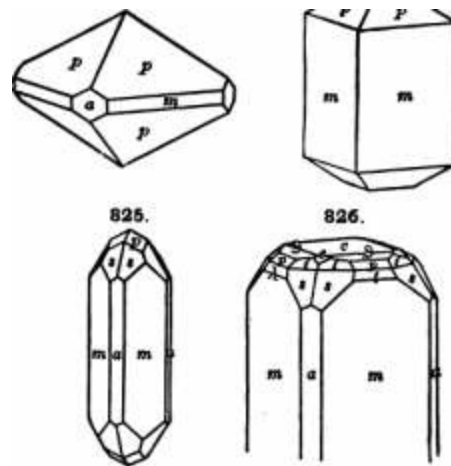
VEBTJVIAinTB. Increase.

Tetragonal. Axia 6=0-5372.

M. 001 A 101 = 29° 15".

ep, 001 A 111 = 37° 13'.

ct, 001 A 881 = 06° 18'.



pp'. 111 A 111 = 60° 8' «•». SUA all = 31" B 8'.

I<^^\ f^>\

Often in crystals, prismatic or pyramidal. Also massive; columnar, straight and divergent, or irregular; granular massive; cryptocrystalline.

jcbyGoOgIC

Fracture: conchoidal to uneven. Brittle. H. = 6.5. G. = 3.35-3.45. Luster: vitreous; often inclining to resinous. Color: brown to green, and the latter frequently bright and clear; occasionally sulphur-yellow, and also pale blue. Streak: white. Subtransparent to faintly translucent. Dichroism: not usually strong. Optically—; also rarely. Birefringence very low. Sometimes a normally biaxial. Indices: $a > c > b$, $e^{\wedge} = 1.7235$, $e^{\wedge} = 1.7236$ Ala, Osann.

Comp.—A basic calcium-aluminium silicate, but of uncertain formula; perhaps $Ca_xAl(OH,F)_yAl_2(SiO_3)_z$. Ferric iron replaces part of the aluminium and magnesium the calcium. Fluorine and titanium may be present. The following are typical analyses (Jannach):

810, Tie, Al₂O₃, Fe₂O₃, FeO MnO CaO MgO Na₂O Li₂O H₂O F,

1. Vesuvius, 188-98 — 16.70 2.89 2.01 0.57 8.5 6.7 8.68 0.48 0.08 1.82 10.8 = 100

8, Villul. 86-17 1-80 laas 31.6 1.0 0.16 8.5-8.1 6-0.5 0.45 — 0.72 0-2.2 B₂O₃, 28.1=89-68

Pyr., etc.—B.B, fuses at 3 with intumescence into a frothy or brownish glass. Has a density after fusion is 2.93-2.94. With the fluxes gives reactions for Iron,

and some varieties a strong magnetic reaction. Cyprine. a hilly variety. (It is a reactant for copper with suit of phosphorus. Partially decomposed by hydrochloric acid, and completely when the mineral has been previously ignited.

DM—Characterized by its tetragonal form and easy fusibility. Resembles some brown varieties of garnet, tourmaline, and epidote.

Recognized in thin sections by its high refractive index producing a very strong relief and its extremely low birefringence; * also in thin sections by its color, pleochroism, and uniaxial negative character; the latter, on account of the low birefringence, being difficult to determine. The low birefringence, however, aids in distinguishing it from epidote, with which at times it may be confounded.

Om. — Vesuvianite was first found among the recent eruptions of Vesuvius and the dolomitic blocks of Monte Bomina, whence it derives its name. It commonly occurs here in contact with limestone from the alteration of impure limestone, then usually associated with lime garnet (grossularite), plagioclase, diopside, wollastonite; also epidote; also in serpentine, chlorite schist, gneiss and related rocks.

Prominent localities are Vesuvius; the Albani Mts.: the Giussani Alp in the Ala valley, in Piedmont; Mt. Moizone in the Fieschi mts.; at Omwilla and Doaczk; Hoalnu near Eger in Bohemia (Egeran); near Jordansmühl, Silesia; on the Vilui river, near L. Baikal (some-times called tnfuit'or viluite. like the grossular garnet from the same region]; at Arendal, "tolophonite"; at Egg, near Cliverville, Canada.

• In N. America, in Maine at Phippsburg and Rumford; at Bradford. In N. Hampshire, at Warren with cyanite-stone. \n N. York. \n m. S. of Amity. In New Jersey, at New Hope. In California near San Carlos in Inyo Co. In Canada, at Calumet Falls. Litchfield, Pontiac Co.; at Orenville in Vermont; at Templeton, Orleans Co., Quebec

Zircon Group, $ZrSiO_4$, Tetragonal.

Zircon $ZrSiO_4$ = 06404

Thorite $ThSiO_4$ = 06403

This group includes the orthosilicates of zirconium and thorium, both alike in tetragonal crystallization, axial ratio and crystalline habit.

The two species are sometimes regarded as oxides and then included in the Rutile Group (p. 343), with which they approximate closely in form. A similar form belongs also to the

"Frequently colorless, which, like vesuvianite, melilite and zirconite, are admirably refracting but of extremely low birefringence (and possibly, where they are positive for one color but

negative for another), do not show a gray color between crossed nicols but a curious blue, sometimes an intense Berlin blue, which is quite distinct from the other blues of the color scale and is known as the "ultra blue,"

by Google

Crystal System: Tetragonal. $a:b:c = 1:1:0.37$.

Crystallographic data: $101 \wedge 010 = 44^\circ 50'$. $101 \wedge 101 = 66^\circ 16'$, $pp \wedge 111 = 66^\circ 16'$. $831 \wedge 111 = 83^\circ 5'$.

Crystallographic data: $111 \wedge 111 = 41^\circ 13'$. $811 \wedge 811 = 82^\circ 67'$. $100 \wedge 811 = 81^\circ 43'$.

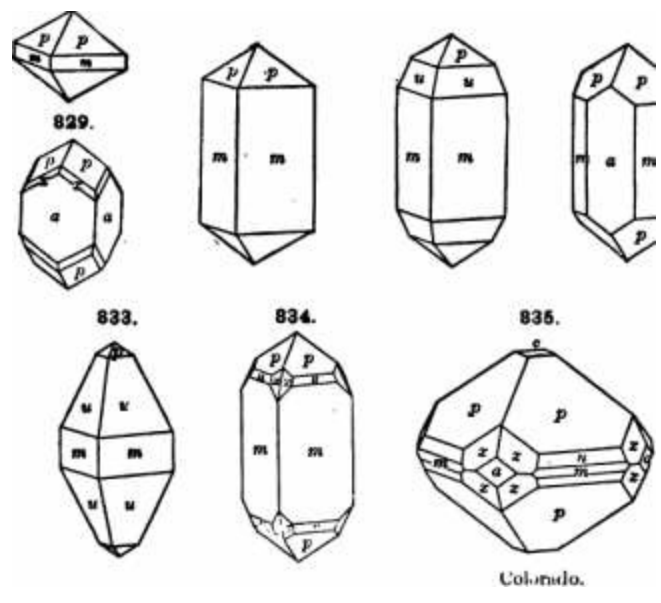
83B.

830.

831.

832.

Crystallographic data: $101 \wedge 010$, $101 \wedge 101$, $111 \wedge 111$, $831 \wedge 111$, $811 \wedge 811$, $100 \wedge 811$.



Twins: tw. pi. e (101), geniculated twins like rutile (Fig. 374, p. 124). Commonly in square prisms, sometimes pyramidal. Also in irregular forms and grains.

Cleavage: m imperfect; $\wedge(111)$ less distinct. Fracture conchoidal.

Brittle. $H = 7.5$. $G = 4.68-4.70$ most common, but varying widely to 4.2 and 4.86. Luster

adamantine. Colorless, pale yellowish, grayish, yellowish green, brownish yellow, reddish brown. Streak uncolored. Transparent to subtransparent and opaque. Optically (-). Birefringence high, $n_g = 1.9239$, $n_j = 1.9682$, Ceylon. Sometimes abnormally biaxial.

Hyacinth is the orange. It is a kind used for gems. Jargon is a name given in the colorless or amethystine while resembling the diamond.

1 lusler, they a

CaMgSiO_3 or $\text{ZrO}_2 \cdot \text{SiO}_2$. = Silica 32-8, zirconia 67-8 iron (Fe_2O_3) is usually present.

= 100. A little

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Pyrrhotite, colorless; the colorless variety are usually the red become colorless,

while dark-colored varieties are made white; some varieties glow and increase in luster by ignition. Not perceptibly acted upon by salt of phosphorus. In powder decomposes when fused with soda on the platinum wire, and if the product is dissolved in dilute hydrochloric acid it gives the orange color characteristic of zirconia when heated with sulfuric acid. Decomposed by water with alkaline carbonates and bisulphates.

It is characterized by the prismatic square pyramidal or square prism; also transparent [is] in thin sections, hardness, high specific gravity and fusibility; the diamond isopically laolnopic.

Recognized in thin sections by its very high relief; very high interference-colors, which approach white of the higher order except in very thin sections; positive uniaxial character. It is distinguished from titanite and rutile only by its lack of color, and from the latter also in many cases by melting of occurrence.

It is a common constituent of igneous rocks, especially those of the more acid feldspathic groups and particularly those derived from magmas containing much silica. It is generally present in igneous rocks, but in pegmatitic rocks often in large and trellis-like crystals. Occurs more rarely elsewhere, as in granular limestone, dioritic and other schists; gneiss: sumellme- in iron-ore beds. Crystals are common in most igneous sands. Sometimes found in volcanic rocks. It is in part an inclusion derived from older rocks.

Zircon in distinct crystals is so common in the pegmatitic forms of the nepheline-syenite and alluvial syenite of Eidsvoll Norway (with topazite, etc.) that the rock there and elsewhere has sometimes been called a "zircon-syenite."

Found in alluvial sands in Ceylon ; in the gold regions of the Ural, Bt Laurvik. Norway: at Arendal, in the iron mines: at FiedriksvBrn. and in veins in the albite-syenite of the Langesundfjord; PflBchilial, Tyrol; in lava at Niedermendig In the Elvel, red crystals: etc.

In N. America, in Xaint. at Littleton In N. York, In Monticello, Essex Co. . color red; near the outlet of Two Rivers, Orange Co., with scapolite, pyroxene and lime; In Warwick, chocolate-brown, near Amity; in St. Lawrence Co., in the town of Hammon; at Roseton. Fine, fibrous. In Penn., near Reading. In N. Car., abundant in the granites of Burke. McDowell. Polk. Rutherford. Henderson, and other counties. In Colorado: do. with asbestine, etc. in the Pike's Peak region in El Paso Co.; at Cheyenne Mt. In California. In auriferous rocks.

In Canada. at Grenville. Argenteuil Co.; In Templeton and adjoining townships In Ottawa Co., Quebec; at Renfrew do., sometimes very large; in North Burgess, Lanark Co.

Malacolite an altered zircon. Cassite is related but contains uranium, yttrium and other rare elements.

Thoria. Thorium silicate, ThO₂, like zircon in form; usually hydrous. black in color, and then with G. = 4.5; also orange-yellow and with Q. = 5.19-0.0 (orange). From the Brevik region. Norway; also near Arendal.

Auriferite. Like zircon in form; supposed to be a silico-phosphate of thorium. Hender-
Danburite-Topaz group. Orthorhombic. RR, (SiO₂)₂ or (BO)₂RSiO₃.

Danburite CaB₂(SiO₃)₂, d:h:l = 0.5444 : 1 : 0.4807

Topaz [Al₂(P,OH)₂]₂AlSiO₆, d:h:l = 0.5285 : 1 : 0.4770

Andalite (AlO)₂Al₂SiO₅. d:h:l = 0.5070 : 1 : 0.4749

or S:h:l = 0.9861 : 1 : 0.7035

Silimanite Al₂SiO₅, Orthorhombic S:h:l = 0.970 : 1

Cyanite Al₂SiO₅, Triclinic

d:h:l = 0.8994 : 1 : 0.7090; a = 90° SJ', j8 = 101° 2', γ = 105° 44'.

byGoogle.

/p^=^

SILICATES. 431

OANBURtTB.

Orthorhombic. Axes $d:b:i = 0.5444 : 1 : 0.4807$.

$836. \text{ mm}''$, $110 \text{ A HO} = 57^\circ 8'$. $101 \text{ A lOl} = 83^\circ 88''$.

a. $120 \text{ A iaO} = 81^\circ 8''$. $\text{mif. Oil A o41} = \text{ISO}'' \ll$.

Habit prismatic, resembling topaz. Also in indistinct embedded crystals, and disseminated masses.

Cleavage: c very indistinct. Fracture uneven to subconchoidal. Brittle. $H. = 7$. $G. = 2.97-3.02$. Color pale wine-yellow to colorless, yellowish white, dark wine-yellow, yellowish brown. Luster vitreous to greasy, on crystal surfaces brilliant. Transparent to translucent. Streak white. Comp. — $\text{CaB}_2(\text{SiO}_3)_2$, or $\text{CaO} \cdot \text{B}_2\text{O}_3 \cdot 2\text{SiO}_2 = \text{Silica } 48.8$. boron trioxide 28.1 , lime $22.8 = 100$.

Pyro., $\bullet \ll$. — B.B. fuses at 800°C in a colorless glass, and imparts a green color to the O.P. (borax). Not soluble in hydrochloric acid, but sufficiently attacked for the solution. Gives the reaction of boric acid with mercuric paper. When strongly ignited gelatinizes with hydrochloric acid. Phosphoresces on heating, giving a reddish light.

Obs. — Occurs at Danbury, Conn., with microcline and oligoclase in dolomite. At Russell, N. Y., in large crystals. On the Platteville, the northern spur of Mt. Baker south of Disseville in eastern Switzerland, in slender prismatic crystals.

TOPAZ. Orthorhombic. Axes $i:l:i = 0.52804 : 1 : 0.47698$.

m m t

rid',

$48^\circ 87'$. $101 \text{ A o21} = 81^\circ 18'$.

Crystals commonly prismatic, m predominating; or \wedge (130) and the form then a nearly square prism resembling andalusite. Faces in the prismatic zone often vertically striated, and often showing vicinal planes. Also firm columnar; granular, coarse or fine.

Cleavage: c highly perfect. Fracture subconchoidal to uneven. Brittle. $H. = 8$. $G. = 3.4-3.6$. Luster vitreous. Color straw-yellow, wine-yellow, white, grayish, greenish, bluish, reddish. Streak uncolored. Transparent to subtranslucent. Optically +, A_1 pi. \backslash b. $B_x X''$. Axial angles variable, $2E, = 112^\circ$ to $120^\circ \pm 40'$. Refractive indices, Brazil (Mühlheims):

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DE&C&llTIYE HIHBRAI.OaT.

For $D_a = 1-62986$ ($I = 1-68077$ $r = 1-8S747$ -■. $aV = 49^\circ 81'$

Tar.— Ordinary. In prUmalle cryatsls usually colurlins or pale yellow, less often pale blue, piuti. etc. The yellow of the Bmzilmu cryaials is obaaged by Lealiug io a pula Tose-pluk. Orieu coDtaiua iucluaioas of liquid CO..

PhyiaiiU, ur pyroplij/MoliU, i» u coarae uearly upaque variety, from Finbo; iaiumescea Wliuu heated, lieoce its name fiucii tf/v<ra\ii, bvbblt, and nCp.fira. Pye'iiU lius n cuhim-nnr. vtry compact structure. Row made out thiat thie cleavage was tbe same, and ihe form prolmbly the same; aud Dea Cloizeaux showed that the optical characlera were those of topaz.

Comp.—(AlF),SiO,₂; usually eontainiog liydroxyl and then [Al(F,oh)],SiO₂, or as given on p. 430. The former requires Silica 32-6, alumina 55*4, fluorine 20-7 = 108-7, deduct (o = 3F) 8-7 = 100.

Pyr., ato,—B.B. iufusible. Fused In the closed tube, with previously fused and pulverized salt of phogphoruB, etches llie glass, glvlug oS silicon fluoride, which forms a Hog of SiO₂ above. Willi cobalt solution the pulverized mineriii gives a flue blue on beating. Only psrllally attacked by sulphuric acid. A variety of tupax from Bmzll, wheu lteated, »asum;8 a pink or red hue, reBembliog llie Balas ruby.

Diff.-H!haracterized by its prismaic crystuU wiih anKles of 56° (124°) or 87° (88°); also by the perfect basal cleavage; hardness: furusibilly; yields fliioTine B.B.

Oba.—Topaz occurs especially ia the highly acrtl igneous rocks of the granite family. o3 granite nnd'rhyolile. Id veias nud cavities, wiere It appears to be the result oF fumarule action iifier the cryslalliEalion of the marina; sometimes also In the surroundiag schiala, gnelase-. <^t'., as a result of such action. In these occurrences oflen accompanied by fluor-

FJu<- Lii|]:i£ comes from the Urals, from Alabaahka. in the region of Ekaterinburg: from Minsk in the Ilmeu Mts : also the giild-wasUinga on tbc R. Saoarka, In Govt. Oreubiirg: in Nerchinsk, beyond L. Baikal, in the Aduii-C'halon His., etc.; In the province cif Minas Gcraea. Brazil, at Ouro Pi-eto and Villii Rica, of deep yellow color; at the tin mines of Zinnwald and Ehrenriedersdorf, and smaller .cryBlals at Schneckenein and Alteoberg; sky-blue crvslals in Calrngorin. Aberdeenahire; Uic Mimrne nioumalDS. Ireland; on the island of Elba. Phy»'iUI« occurs iu crystals of great size, at Possum, Norway; Flno, 8we<len. Pyeniti is from the tin mine of AltcDIHTg in Siixoiij; also of Sc blacken wait). Zinnwald, etc. Fine crypfd* occur at Duraugo, Mexico, with tin ore; at San Luis Potiiiai in rhyolite. Mt. Bi-choB. Tasmania, with tin ores; similarly in New South Wales. In Japnn in pegmatie from Olani-yaina, Province of Omi, near Kioto.

Io the United Stales, In Maiitt, at Stoneham. in albitle gmnle. In Oonn.. at Trumbull, with

fluorite: at Willimailic, In N. Car., at Crowder's Alouutain. In Colorado, in fine crystals colorless or pale blue from the Pike's Peak region; at Nathrop, Chaffee Co., in wine-colored crystals with spessartite in lithophyses in rhyolite; similarly in the rhyolite of Chalk Mt. In Utah, in fine transparent colorless crystals with quartz and anorthite in the rhyolite of the Thomas Ranire, 4 (1 miles north of Sevier Lake.

The topaz is from Rhé, an island in the Red Sea, as stated by Pliny.

But

by Pliny and earlier

ANDALUSITE.

Orthorhombic,

as well as by many later, under the name

Axes $a : b : c = 0.9861 : 1 : 0.70245. 842.$

the crystals,

mm³, $110 \text{ A} \text{ liO} = 89 \times 13$. Usually in coarse prismatic forms, the prisms nearly square in form. Massive, imperfectly columnar; sometimes radiated and granular. •

Crystal: in distinct, sometimes perfect (Brazil); a less perfect: b in traces.

Fracture uneven, subconchoidal. Brittle. H. = 75. G. = 316-320. Luster vitreous; often weak. Color whitish, rose-red, flesh-red, violet, pearl-gray red-

\sqrt{U}

,ab,GoOgIc

Color brown, olive-green. Streak colorless. Transparent to opaque, usually translucent. Pleochroism strong in some colored varieties. Absorption strong, $a > b > c$ Sections normal to an optic axis are isochroic or show the polarization-brushes distinctly (p. 219), Optically — $Ax \text{ pi } B \text{ d } Bii.c. 2H, = 96^\circ 30'$ Brazil; $\gamma, = 1-6as; \beta/\gamma = 0011.$

Chrysoberylite, or Maclureite in stout crystals, ^,

the III is in shades of a different color from the others. •■•

to a regular arrangement of impurities in the crystal a=Bx

in the interior, and hence exhibiting a colored cross, or a characteristic appearance in a li-

acsterae section. Fig. 644 shows sections of a crystal.

CoBip.—Al₂SiO₅ = (AlO)₃Al₂SiO₅, or Al₂O₃·8SiO₂, = Silica 63.3, alumina 36.7 = 100.
Manganese is sometimes present, as in manganandalusite.

Fracture.—B.B. infusible. With cobalt solution gives a blue color after ignition. Not decomposed by acids, decomposed on fusion with caustic alkalis and alkaline carbonates.

Diff.—Characterized by the nearly square prism, cleavage, hardness, infusibility; reaction for alumina B.B.

Distinguished in thin sections by its high relief; low interference-colors, which are only slightly above those of a

quartz: negative biaxial character; negative extension of

the crystal (diff. from sillimanite); rather distinct prismatic cleavage and the constant

parallel extinction (diff. from pyroxenes, which have also greater birefringence); albity

844.

characteristic arrangement of impurities when these are present (Fig. 844). The pleochroism, which is often lacking, is when present strong and characteristic.

Ob.—Most common in argillaceous schist, or other schists imperfectly crystalline; also in gneiss, mica schist and related rocks; rarely in connection with serpentine, the variety chiastolite is commonly a contact mineral in clay-slates, e.g., adjoining granitic dikes. Sometimes associated with sillimanite with parallel axes.

Found in Spain, in Andalusia: in the Trol, Linares Alp; in Saxony, at Brunsdorf;

in the Uinax Geraes, in fine crystals and in

Bavaria, at Wunsiedel, etc. In Brazil, in rolled pebbles.

In N. America, in Maine, at Sandleford Mt. Notch; Uau., at Westford; Lancaster, both varieties; Sterling, chiastolite. Conn. at Litchfield and Woburn. In Delaware Co., near Leipsville, large crystals; Upper Providence.

Named from Andalusia, the first locality. The name *andalusite* is from the Latin *andalusia*, a spot. *chiastolite* is from *χιαστή*, arranged diagonally, and hence from *χ*, the Greek name for the letter X.

Fracture.—Fibrous.

Orthorhombic. Axes $d:ti = 0.970 : 1$. $\mu = 88^\circ 15'$, $M' (230 A 330) = 69^\circ$. Prismatic faces striated and rounded. Commonly in long slender (Crystals not distinctly terminated; often in close parallel groups, passing into fibrous and columnar massive forms; sometimes radiating).

Cleavage: ft very perfect. Fracture uneven. H. = 6-7. G. = 3-23-3-24. Luster vitreous, approaching subadamantine. Color hair-brown, grayish brown, grayish white, grayish green, pale olive-green. Streak uncolored. Transparent to translucent. Pleochroism sometimes distinct. Optically +.

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Double refraction strong. $\Delta = 0.0021$. $n_p = 1.661$; $n_g = 1.661$; $n_o = 1.661$. $\omega = 44^\circ$. $\sigma = 1.661$; $\rho = 1.661$.

Pyr.—Same as BndaluBlte.

TUB. —Characterized by its fibrous or columnar form; perfect cleavage; infusibility; reaction furaculnDa,

In thin sections recognized by its form, usually with triangular fractures; parallel extinction: high interference-colors.

Obi.—Orient present in the quartz of gneiss and sometimes granites [q very slender, minute prisms commonly aggregated together and sometimes intergrown with andalusite; lonsdaleite is also a common associate; rarely as a contact mineral; often occurs with corundum.

Observed in many localities, thus near Mulden in Bohemia (Fattrekaty, at Fassa in Tyrol (Jura[^]rific); in the Caralio with corundum (JtbroiiU); at Bodeumais, BaTBria; Freiberg, Saxony; in France, near Ponirbaud and other points in Auvergne; forms rolled masses in the diamondiferous sands of Hiuaa Geraes, Brazil.

In the United States. In JfoMOcAbMld, til Worcester. In Cbnn«ea«uf, near Norwich, with zircon, monazite and corundum; at Wllimantic. In N. York, at Yorblown. West-Chester Co.; In Hotiroe, Orange Co. (monroliU). In Ptnn.. at Chesterou the Delaware, near Qiieosbury forge; in Delaware Co.; Dtlaaare, at Brandywln Springs. With corundum in N. Oarolina.

Named j!&roIif« from the fibrous massive variety (Germ., Faaerklesel}; titUmanilt, after Prof. Benjamin Slliman of New Haven (1779-1864).

BamliU, xenaliu, adrtiUtt probably belong to silimanite; the last is altered.

OTANrm. Eyanlte. DIathene.

Triclinic. Axes $d:lU = 08994 : 1 : 0-7090$; $\angle r = 90^\circ 51'$, $\theta = 101^\circ 2'$ $\psi = 105'' 44f$. ac, $100 A 001 = 78'' 30'$; $be, 010 A 001 = 86M5'$;

84fi. Usually in long bladed crystals, rarely terminated.

Also coarsely bladed columnar to subfibrous.

Cleavage: a very perfect; ft less perfect; also parting U c. H. = 5-7-25; the least, 4-5, on a || rf; 6-7 on a U edge a/c; 7 on 6. G. = 3-56-3-67, Luster vitreous to pearly. Color blue, white; blue along the center of the blades or crystals with white margins; also gray, green, black. Streak uncolored. Translucent to transparent. Pleochroism distinct in colored varieties. Optically —. Ax. pi. nearly $J_{\perp} a$ and inclined to edge a/b on a about 30° , and about 7t on 6. $2H..r \wedge 99'' 18'$ Pfitschthal. Comp.—Empirical formula Al_8SiO_{10} , or $Al_8O_{10}Si_2O_5$, like andalucite and sillimanite. Perhaps a basic metasilicate, $(Al_8O_{10})_2Si_2O_5$.

Pyr., etc.—Same as for andalusite. At a high temperature ($1520^\circ - 1550^\circ$) cyanite assumes the physical characters of sillimanite.

Diffr.—Characterized by the bladed form; common blue color; varying hardness; infusibility; reaction for alumina.

Obs.—Occurs principally in gneiss and mica schist (both the ordinary variety with muscovite and also that with paragonite) often accompanied by garnet and sometimes by staurolite; also in eclogite. It is often associated with corundum.

Found in transparent crystals at Monte Campione in the St. Gotthard region in Switzerland in paragonite schist; on Mt. Giselner. Ziemerthal, and in the Pfitschthal (sillimanite, white) in Tyrol; in eclogite of the Saualpe, Carinthia; Horrsjöberg in Wermland, Sweden; Villa Rica, Brail, etc.

In Mass., in Chesterfield, with garnet in mica schist. In Conn., at Litchfield and Watkinson. In Vermont, at Thetford. In Penn.—In Chester Co. and In Delaware Co. In Virginia, Buckingham Co. In JV. Carolina, with rutile, lazulite, etc., at Crowder's Mt., Gaston Co.: In Gaston and Rutledge counties associated with corundum, damourite; beautiful clear green in Yancey Co. Named from Kivovoi, Hiu.

.dbyGoogle

Datolite Orop. Monoclinic.

Basic orthosilicate. $HBHSiO_3$ or $B_2B_2(SiO_3)_4$. Oxygen ratio for $Si = 3:2$. R = Oa, Be, Fe, chiefly; E = Boron, the yttrium (and cerium) metals, etc.

DatoUt* 0-6345 :1: 1-2657 89° 61'

HCaBSiO. or Ca{BOH)SiO, HomlUt* 0-6249 :1 : 12824 69° 21'

Ca.FeB,Si,o., or Ca,Fe{BO),(SiO.),

EndiM 0-6474 :1: 1-3330 79° 44'

HBeAlSiO. or Be(A1oH)SiO, &

OadoUnito 0-6273 :1 ; 1'3215 89° 26^

Be,FeY,Si,o., or Be,Pe(YO),(SiO.),

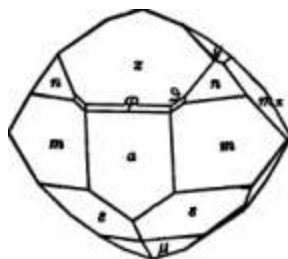
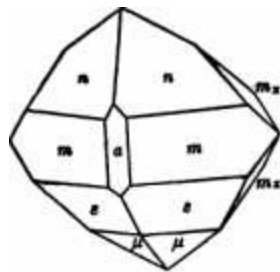
The species of the Datolite Group are usually regarded as basic ortho-silicates, the formulas being taken in the second form given above. They i^ crystallize in monoclinic system, and all but Enclase conform closely in axial ratio; with the latter there is also a distinct morphological relationship.

DATOUTE.

Monoclinic. Axes it : % mm". 110 A liO = ac 100 A 001 = a*. 100 A 101 = «r'. 013 A 0i2 = jBiW. Oil A Oil =

Crystals varied in habit; usually abort prismatic with either mar m^ predominating; sometimes tabular jj x (201); also of other types, and often highly modified (Figs. 846-849). Also botryoidat and globular, having a columnar structure; divergent and radiating; sometimes massive, granular to compact and crypto-crystalline.

840. 847.



Bergen Hill. Cleavage not observed. Fracture conchoidal to uneven. Brittle. H. = 5-5"5. G. = 2-9-30. Luster vitreous, rarely Bubreeinous on a surface of frao-

DESCRIBTION OF DATOLITE.

Color white; sometimes grayish, pale green, yellow, red, or amethystine, rarely dirty olive-green or honey-yellow. Streak white. Transparent to translucent; rarely opaque white.

Fracture conchoidal. Cleavage none. Hardness 5-6. Specific gravity 2.4-2.5. Luster vitreous. From the L. Superior region. Botryoidal; Botryoidal. Resembles columbite, but is identical.

Comp.—A basic orthosilicate of boron and calcium; empirically HCaBSiO_3 , or $\text{H}_2\text{O} \cdot 2\text{CaO} \cdot \text{B}_2\text{O}_3 \cdot 2\text{SiO}_2$; this may be written $\text{Ca}(\text{BO}_2)_2\text{SiO}_2$. = Silica 37.6, boron trioxide 21.8, lime 35.0, water 5.6 — 100.

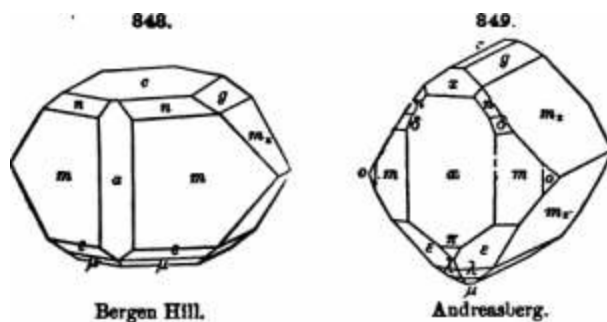
It is soluble in hydrochloric acid.

Crystal form: easy fusibility and green

in glass, coloring the flame bright green. (

Def.—Characterized by its glassy, greenish, conchoidal fracture.

Obs.—Datolite is found chiefly as a secondary mineral in and around basic eruptive rocks, often associated with calcite, prehnite and various zeolites; sometimes associated with dauburite; also in gneiss, diorite, and aephenite: in metallic veins; sometimes



Andreasberg.

It is found in beds of iron ore. Found in Scotland, in trap, at the Eilpatrick Hills, etc. in a bed of magnetite at Areodal in Norway (Borrvofte); at Utö in Sweden: at Aiddraaber in the iron basenod in veins of silver ores: in Bheniah Bavaria (the Autna Mts) at the Seiaser Alp, Tyrol, and at Thelss, near Claussen. in geodes in amygdaloid: in granite at Baveno near Lago Hagelore; at Toggiana in Modena. in serpentine: Monte Catini in Tuscany.

It is not uncommon with the diabase of Connecticut and Massachusetts. Thus at

the Rocky Hill quarry, Hiirtford, Conn.; at Middlefield Falla and Roaring Brook, Conn. In N. Jersey, at Bergen Hill. In splendid crystals; at Paierson, Paasaic Co. Both crystals and the opaque compact variety. In the lAke Superior region.

Named from 6aiti<yBat, to divide, alluding to the giaoular alructure of a maastre variety.

Homillte. $(Ca.Fe)_2B_2Bi_2O_6$ or $(Ca.Fe)_2(BO)_2(BiO)_2$. Cryslala often tabular | e: angtea neiir thnse of datolite. H. = S. Q. = 8:~. Color black, blackisb brown. Found on Uie Island BtokH and other ialanils. in the Langesunil fiord, Norway.

Euclaaa. $HBeAlSiO_4$ or $BefAlOHISiO_4$, - In prismatic crystals. Cleavage : 61010)

EcrfeC H. = 7'5, Q = 806-3'10. Luster vitreous. Colorless to pale green or blue, rem Brauil. In the province of HInai Geraes; in Che auriferous sands of the Orenburg dinlriect, southern Ural, near the river San&rka; in the Oloesglockner region of the Austriaa Alps.

OadoUnlta. Be,FeY,Si_2O_7 , or $Be,Fe(TO)_2SiO_4$,. Crystnls. often prismatic, rough and coarse: commonly in masses. Cleavage none. Fracture concholdal or apllnlery. Brittle. H. = 6'5-7. G. = 40-4-5; normally 4-86-4 47 (anisotropic), 4-34-4'a8 (Isotropic and amorphous from alteration). Luster vitreous to greasy. Color blach, greeofab black.

.dbyGoogle

a)*o brown. From aesr Falun And Ytterbv, Sweden; HItterO, Norway; also in Llano Co., Texas, in nodular mawes and rough crysUle. sometimes up to 40 or SO pounds In weight.

The yttrium earltis or " gadollnite-earlbt" (purtly replaced by the oxides of cerium, lanthanum itnd didymlum) lona a complex group which coutains coulsdeiahle erbium, also several new elements (jtlerblum, scandium, etc.) ot more or less definite churacter.

TttrUUte. Aallicateof thorium and the yttrium metals chiefly. HasslTc; amorpboua. G. = 4976. Color on the freat fracture olive-green, changing to orange-yellow on surface. Aasoclated with the gadollnite of Llano Co., Texas.

RowUndlts. An yttrium illicate, occuTring massive with gadollnite of Llano Co., Texas; color drab-green.

MaoUntoahlte, Silicate of unoiium, thorium, cerium, etc. Hassive. Color black. Llano Co., Texas.

Epidote Gronp. Orthorhotnbic and MonocliDic Baaic Ortbosilicatea, $HR_2E_2Si_2O_{11}$, or $R_2\{ROH\}R_2(SiO)_4$. B = Ca,Fe; K = Al,Fe,Mn,Oe, etc. a. Orth&rh&mbic Section.

Zoiute $Ca_2\{Al_2OH\}Al_2\{SiO_4\}_2$, o-6196:1: 03429

fl. MontuMnte Section. •B^m i mCa,(A10H)Al,(SiO.), d : b : i

JtpiaoH ^ nCa,(FeOH)Fe,(8iO,). 1-5787 :1: 1-8036 64° 37'

PiBdmontite Ca.(A10H)(Al,Mii),{SiOJ, 1*6100 :1: 1-8326 64° 39'

Allanits (Ca,Fe),(A10H){Al,Ce,Fe),(SiO.), 1-5509 :1 : 1-7691 64° 59'

The Epidote Group includes the above complex orthoailicateB. The monoclinic specieB agree closely in form. To them the orthorhombic speciea zoieite is also related in angle, its prismatic zone corresponding to the mono-clinic orthodomes, etc Thus we have :

Zolille mm", no A no = 68° 84'. Epidote er, 001 A lOl = BS" 42'.

vu', 031 A 031 = 68' H'. mm', 110 A llo = 70" 4', etc

There aeema to be, however, a monocliQio oalcinm componnd, having the composition of ;oisite, but monoclinic and strictly isomorpha with ordinary epidote; it is called dinozoisite.

Orthorhombic. Axee S:i:i = 06196 :1 : 0-34295,

,mm": 110 A llo = 68* 84'. jT, Oil A Oil = >r ESf.

df. 101 A iOl = 57° 66'. «/", 111 A Ul = 88° 34'.

Crystals prismatic, deeply striated or furrowed vertically, and seldom distinctly terminated. Also massive; columnar to compact.

Cleavage: 5 very perfect. Fracture uneven to subconchoidal. Brittle. H. = 6-6-5, G. = 3-25-337. Luster vitreous ; on the cleavage-face, 5, pearly. Color grayish white, gray, yellowish brown, greenish gray, apple-green; also peach-blosflom-red to rose-red. Streak uncolored. Transparent to snbtrans-1 a cent.

Pieochroiam strong in pink varieties. Optically +. Ax. pi. usnally f h ; also D c. BxX <^- Dispersion strong, p <v; also p> v. Axial angle variable even in the same crystal. 2E, — 4', "-90". /S = 1-696; y — a = 0'006.

;;CoogIC

Vaz.—l. Ordiitary. Colon graj to white and brown; also greCQ. Usually In Indistinct prlamatic or columnar forms; also la fibrous agKKgfutes. Q. = 8'228-S'381. UnionUe ia a very pure zoislle. 3. Boie-red or T/iuliie. Fnglle; pleochrolsm Btroug. S. tompael, muMtcK. Includes the essential part of moat saussurite («.f., In saunuiite-gabbro), which has iiHseu from !tae alierutfon of feldspar.

Comp.— HCa,Al,Si,O_2 , or $4\text{CaO}\cdot 3\text{Al}_2\text{O}_3\cdot 6\text{SiO}_2\cdot \text{H}_2\text{O}$ = Silica 39.7, alumina 33.7, lime 34.6, water 3.0 = 100. The alumina is sometime replaced by iron, thus graduating toward epidote, which has the same general formula.

Fyr., etc.—B.B. swells up and fuses at S-S' to a white blebby mass Not decomposed by acids; when previously Ignited gelatinizes with hydrochloric acid. Gives off water when strongly Ignited.

ds.—Characterized by the Columnar structure; fusibility with intumescence; resembles some amphibole.

Disguised in thin sections by its high relief and very low interference color; lack of color and biaxial character. From epidote it is distinguished by its lack of color and low birefringence: from vesuvianite by its color and biaxial character. Thin sections usually show the "vitreous blue" (p. 428) between crossed nicols.

Obs.—Occurs especially in those crystalline schists which have been formed by the dynamic metamorphism of basic igneous rocks containing plagioclase rich in lime. Commonly accompanies some one of the amphiboles (actinolite, amfibolite, glaucophane, etc.); thus in amphibolite, glaucophane schist, eclogite; often associated with corundum.

The original zoisite is that of the eclogite of the Bawalpe in Carinthia (Austria). Other localities are: Kauris in Salzburg; Sleizing, etc., in Tyrol; the Fichtelgebirge in Bavaria; Harachendorf in Moravia; Saasthal in Switzerland; the island of Syra, one of the Cyclades, in glaucophane schist. It also occurs at Kleppau in Telemarken, Norway, and at Tra-venella in Piedmont.

Fluorapatite. Fluorapatit, Germ. Monoclinic. Axes $A : b : c = 1.5787 : 1.110 : 1.0866$.

$100 : 001 = 64^\circ 37'$.

$001 : 100 = 89^\circ 26'$. $001 : 110 = 58^\circ 26'$. $001 : 111 = 75^\circ 11'$. $100 : 010 = 61^\circ 41'$.

$010 : 001 = 70^\circ 20'$. Twins: tw. p. a common, often as embedded tw, lamellar. Crystals usually prismatic | the ortho-axis is terminated at one extremity only; passing

$001 : 100 = 6^\circ 87'$.

$001 : 101 = 84^\circ 48'$.

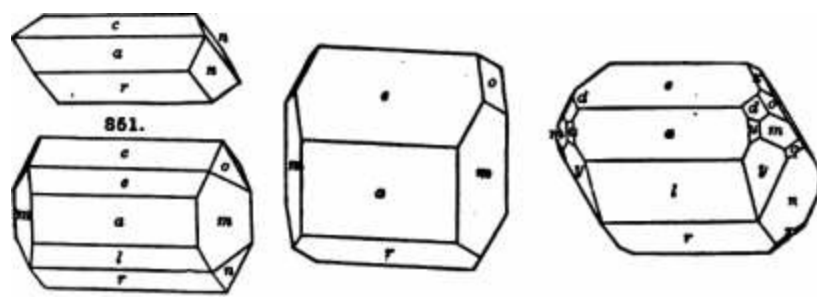
$001 : 102 = 63^\circ 42'$.

$100 : 101 = 61^\circ 41'$.

86a

863.

863.



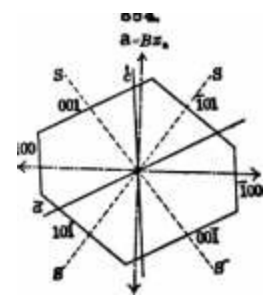
into aoinlor forms; the faces in the zone ac deeply striated. Also fibrous, divergent or panilJel; grannlar, particles of variouB sizes, sometimes fine granular, and forming rock-masses.

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Cleavage: c perfect; a imperfect. Fractare uneren. Brittle. H. = O. = 3'35-3'5. Luster vitreouBi on c inclining to pearly or reBinons. Color pistachio-green or yellowisli green to brownish ""

green, greenish black, and olack; Eometimes

-clear red and yellow; also gray and grayish white, rarely colorless. Stre^ unoolorea, gray-Transparent to opaque: generally enb-



transliicent.

Pleochroisin strong: vibrations | c green, b brown and strongly abmrbed, a yellow. Absorp. tioQ usnally b > c > a; but sometimes c > 6 > a in the variety of epidote common in rocks. Often exhibits idiophanous figures; best in sections normal to an optic axis, but often to be observed in natural crystals (Sulzbach), especially if flattened I r (101). (Seep. 318.) Optically-. Ax. pi. B 6. Bx., A i) = - 2° 56'. Hence c±a (100) nearly. Dispersion inclined, strongly marked; of the axes feeble, p > v, 2fi^, = 91° 20'. fij = 1-T5To2. Birefringence high, y ~ a = 0038 - 0-056.

Vm.— Epidote has ordinarily a peculiar yellowish green (pistachio) color, seldom found in other minerals. But this color passes gradually to dark and light shades—black on one side and brown on the other; red, yellow and colorless varieties also occur.

. Var. 1. Ordinary. Color green of some shade, as described, the pistachio tint rarely absent, (o) In crystals, (i) Fibrous, (e) Granular massive, (d) Beerta is epidote sand from the gold washings in Transylvania. The Arendal epidote (Arendt) is mostly in dark green crystals; that of Bourg d'Oulson Dauphiné (Thallier, Mphinié, Oitani) in yellowish-green crystals, sometimes transparent, Fute/Jdnite includes crystals from the auriferous sands of Ekaterinburg, Ural. A variety is ordinary epidote from Achmatovsk, Ural. A variety from Oarda, Hosié, Terra del Puego, is colorless and resembles loisite.

2. The Biicklindite from Achmatovsk, described by Herold, is black with a tinge of

Seen, and differs from ordinary epidote in having the crystals nearly symmetrical and not, like other epidote, lengthened in the direction of the orthorhombic. G. : 8 61. 8.

With mica. Carmine-red to straw-yellow, strongly pleochroic; deep crimson and straw-yellow. $n_1 = 1.660$; $Q = 3.187$; In small radiated groups. From Glencoe, in Argyle-shire, Scotland. Sometimes referred to piemontite, but contains little iron.

Comp.— $H\text{Ca}_2\{\text{Al},\text{Fe}\}_2\text{Si}_2\text{O}_{10}$, or $H_2\text{O} \cdot 4\text{CaO} \cdot 3\{\text{Al},\text{Fe}\}_2\text{O}_3 \cdot 6\text{SiO}_2$, the ratio of aluminum to iron varies commonly from 6 : 1 to 3 : 2. Percentage composition:

For Al: $\text{SiO}_2 = 87.87$, $\text{Al}_2\text{O}_3 = 24.13$, $\text{FeO} = 12.0$, $\text{CaO} = 28.51$, $\text{H}_2\text{O} = 100$

Clinotoitite is an epidote without iron, having the composition of zoisite; fougé is probably the same from an orthogneiss in Ceylon. Pierroepidote is supposed to contain Mg in place of Ca.

Fyr., etc.—In the closed tube gives water on strong ignition. B.B. fuses with intumescence at $8-3^\circ\text{C}$ to a dark brown or black mass which is generally magnetic. Reacts for iron and sometimes for manganese with the fluxes. Partially decomposed by hydrochloric acid, but when previously ignited, gelatinizes with acid. Decomposed on fusion with alkaline carbonates.

Diagn.—Characterized often by its peculiar yellowish-green (pistachio) color; readily fusible and yields a magnetic globule B.B. Prismatic forms often longitudinally striated, but they have not the angle of tremolite; laurmannite has no distinct cleavage, is less fusible (in common forms) and usually shows its hexagonal form.

Recognized in thin sections by its high refraction; strong interference-colors rising into those of the third order in ordinary sections; decided color and striking pleochroism; also by the fact that the plane of the optical axes lies transversely to the elongation of the

DESCRIPTITE MINKBALOGY.

It occurs much lime. It occurs often in gneissic rocks, especially in amphibolite schists, Beresford; also in (the Hartzles and Saalblock) altered by neighboring igneous rocks. Often in thin beds of magnetite or hematite in such rocks. Has also been found in granite (Hobbs, Haryana), and regarded as an original mineral.

It is often associated with quartz, feldspar, actinolite, axinite, chlorite, etc. It sometimes forms with quartz & epidote rock, called epidote. A. almar rock in Uel-boirue in Cimadn. A gneissoid rock consisting of flesh-colored orthoclase, quartz and epidote from the Uak* Mts. (N. of mid Teon.) has been called unakyle.

Beautiful crystals in France come from Bourg d'Oisans, Dauphine; the Ala valley and Traversella in Piedmont; Elba: Zermatt: Zillnial in Tyrol; also in blue crystals from the Koappenwald in (the Uitersulzbachthal, Pinzgau, associated with asbestos, adularia, apatite, titanite, scapolite; further at Stiigau. Silesia; Zoptau, Moravia; Arendal, Norway; the Achmatovak mine near Zlatoust, Ural.

In N. America, occurs in N. Hamp., at Praisac. In Mass., at Hadlyme and Chester, in crystals in Kneton; at Athol, in syenitic gneiss, in fine crystals, 2 m. S. W. of the center of the town; Newbury, in limestone. In Conn., at Haddam, in large splendid crystals. In N. York, near Amity; Monroe, Orange Co.; Warwick, pale yellowish green, with titanite and pyroxene. In N. Carolina, at Hampton's, Yancey Co.; White's mill. Gaston Co.; Frank-lin, Macon Co.; in crystals and crystalline masses in quartz at White Plains, Alexander Co. In Michigan. In the Lake Superior region, at many of the mines.

Epidote was named by Hauy, from the Greek *epido*-*te*, increase, translated by him, "qui & regit un accroissement," the base of the prism (rhombohedral prism) having one side longer than the other. Pittaite, from *pittachio*-*nit*. refers to the color.

Piedmontite. Similar in angle to ordinary epidote. but contains 6 to 15 p. c. Mn, O]. H. = 6.6. O. ~ S⁴⁰⁴. Color reddish brown and reddish black. Pleochroism strong. Absorption a > b > c, Optically+. Ax. pi. \b. Bx, r A i = +82° 84', a a ^ = - 6° to ~ 8°. Occurs with manganese ores at St. Marcel, Piedmont. In crystalline schists on the de Orlx, France; In glaucophane-schist, in Japan. Occasionally in quartz porphyry, as in the antique red porphyry of Egypt, also that of South Mountain. Penn.

AUJUnTB. Ortblte.

Monodino, Axes, p, 437. In angle near epidote. Crystals often tabular in; also long and slender to acicular prismatic by elongation | axis h. Also massive and in embedded grains.

Cleavage: a and c in traces ; also m sometimes observed. Fracture uneven or subconchoidal. Brittle. H. = 5-5-6. G. = 3-0-4-2. Luster submetallic, pitchy or resinous. Color brown to black. Sub translucent to opaque. Pleochroism strong : c brownish

yellow, b reddish brown, a greenish brown. Optically —. Ax. pi, y b. Bi. $\wedge i = 324^\circ$ appoi. $\wedge = 1'682$. Birefringence low; $V - n = 0.032$. Also isotropic and amorphous by alteration analogous to gadolinite.

Vax.—AUnnila. The original mineral was from East Greenland, in tabular crystals or plates. Color black or brownish black. G. = 8.150-8.155. Birefringence is anhydrous allanite in small black crystals from a magnetite mine near Arendal, Norway. Birefringence occurs in black emuls which are like the bucklandite of Acemulovsk (epidote).

Original included, in its original use, the slender or acicular prismatic crystals, containing some water, from Finbo, near Falun, Sweden. But these graduate into massive forms, and some orthites are anhydrous, or as nearly so as most allanite. The name is from *ortho*, straight. „jjj

Comp.—Like epidote $H_2R_2O_6$, or $H_2O.4R_2O.6SiO_2$, with $R = Ca$ and Fe, and $R = Al, Fe$, the cerium metals Ce, Di, La, and in smaller amounts those of the Therman group. Some varieties contain considerable water, but probably by alteration.

Pyrolysis.—Some varieties give much water in the closed tube and all kinds yield a small amount on strong ignition. B.B. fuses easily and swells up ($F = 2'B$) to a dark, blebby.

=:i;cc ..Google

SILICATES.

441

quartz. With the fluxes reacts for iron. Most varieties gelatinize with hydrochloric acid, but if previously ignited are not decomposed by acid.

Ob*.—Occurs in albite and cummingtonite feldspar, gneiss, syenite, zircon, syenite, porphyry. Thus in Orreuaud : Norway ; Stredeu ; Briegnu, SiU-sin Also la wille thiu-hode SB at Aucrbacu on the Bergstraae; often in mica of magnesian iron. Rietber occurs as an accessory constituent in many rocks. as Idadesitt!. dioite, dacite, rhyolite, the dimite of St. Adamello. the scapolite rock of Ockau. Nui'way, etc. Sometimes occurs as a nucleus in crystals of the isomorphous species, epidote.

At Vesuvius in ejected masses with bauxite, sodalite, nephelite, bombalite, etc. Particularly in trachytic ejected masses at the Laacher See {bucklandite).

In Mass., at the Bolton quarry. In N. York, Horlah, Essex Co. with magnetite and apatite ; at Horee, Orange Co. In N. Jersey. at IVariklin Furnace with feldspar and magnetite, in Pttn., at Houalalu, near Bethlehem, in large crystals ; at E. Bradford ; near Eckhardt's furnace, Berks Co. also in Virginia, in large masses in the Amberst Co. also in Bedford, Nelsona. and Amelia counties*. In N. Carolina, at many points. At the Devil's

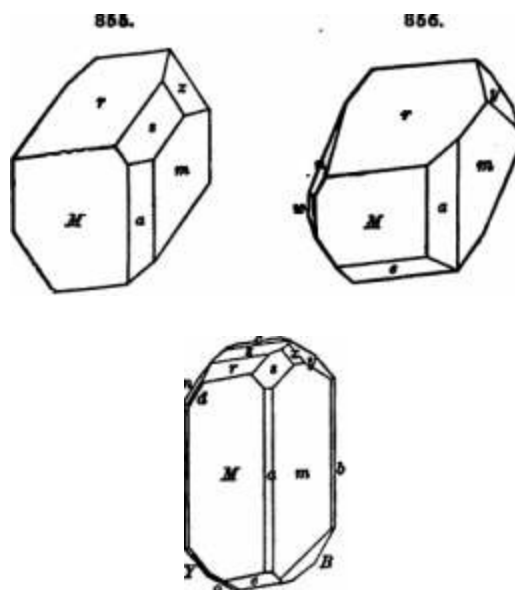
Head Mt., Douglas Co., Colorado.

AXmiTZL

Triclinic. Area $h:i = 0.4921 : 1 : 0.4797$; $\gamma = 131^\circ 32'$.

ses.

" $54'$, $\gamma = 91^\circ 52'$,



Dauphin. Poloma. Bethlehem, Pa.

$\alpha M, 100 A 100 = 16^\circ 84'$. $\mu r, 110 A 110 = 45^\circ 15'$.

$\alpha M, 100 A 110 = 38^\circ 55'$. $\mu r, 110 A 110 = M^\circ 2^*$.

$M, 100 A 100 = 21^\circ 87'$. $\mu r, 110 A 110 = 27^\circ 5'$.

Crystal usually broad and acute-edged, but varied in habit. Also massive, lamellar, lamellae often curved; sometimes granular.

Cleavage: b distinct. Fracture conchoidal. Brittle, $H. = 6-5-7$, $\rho. = 3.271-3.294$. Luster highly glassy. Color clove-brown, platinum-blue, also pearl-gray; also honey-yellow, greenish yellow. Streak uncolored. Transparent to subtransparent. Pleochroism strong. Optically —. Ax. μr and μs , approximately $i, X(111)$. Axial angles variable. $2H. = 87^\circ 30'$; $\rho. = 1.678$. Pyroelectric (p. 234).

Comp.—Aluminosilicate of aluminium and calcium with varying amounts of iron and manganese. Perhaps $n, R.(BO)Al_2(SiO_4)_2$. (Whitfield.) $R =$ Calcium chiefly, sometimes in large excess, iron in smaller amount and manganese prominent; iron is present in small quantity, also magnesium and basic hydrogen.

Analyses. 1, Wittfeld ; 2, Genth.

G. SiO, B₂O₃, Al₂O₃, Fe₂O₃, FeO, MnO, MgO len. 1. Bourg d'Oisans 41, S₃ 4 63 17 fl O 3 » 0
403 8-7 fl 21-66 074 a¹⁶ = 100-82

3. Franklin. crjMf. 8-358 42⁷⁷ .ilO 10-78 103 1-6(1- 18 69 18-25 0-28 0-78 = 100-16
■ ZnO, including 0.12 O₂. Pyr., •to.—B.B. fuses readily with intumesce. Imparts a pale green color to the O.F.,

and fuses at 2 to 3 & dark green to black glass ; with borax in O.F. glazes on a meibumite bead (manganite), which in R.F. becomes yellow (iron). Fused in a mixture of borax and phosphate of potash and the platinum loop colors the flame green (boric acid). Not decomposed by acids, but when previously ignited, gelatinizes with hydrochloric acid.

Obs.—Asbestos occurs in clove-brown crystals, near Bourg d'Oisans in Bauphng : kt Andreaaberg ; Striegau, Silesia ; in Mt. Skopi, in eastern Switzerland ; Elba ; at the silver mines of Kuugsberk, Norway : Nordmark, Sweden ; near Hasko in the Ural ; in Cornwall, of a dark color, at the Botolph Claydon near St. Just. etc.

In the U. S., at Phippsburg, Maine ; Franklin Furnace, N. J., honey-yellow ; at Bethlehem, Pa.

Named from dikiyt), an acute, alludes to the form of the crystals.

Orthorhombic-hemimorphic. Axes $a : b : c = 0.8401 : 1 : 0.5549$.

Distinct individual crystals rare ; usually tabular || c ; sometimes prismatic $\alpha = 80^\circ 4'$; again acute pyramidal. Commonly in groups of tabular crystals, united by c making broken forms, often barrel-shaped. Benign, globular, and stactitic with a crystalline surface. Structure imperfectly columnar or lamellar, strongly coherent ; also compact granular or impalpable.

Cleavage: c distinct. Fracture uneven. Brittle. H. = 6-6-5. G. = 2.80-2.96. Luster vitreous ; c weak pearly. Color light green, oil-green, passing into white and gray ; often fading on exposure. Subtransparent to translucent. Streak uncolored.

Comp.—An acid orthosilicate, H-Ca,Al,(SiO₃)₂, = Silica 43.7, alumina 24.8, lime 27.1, water 4.4 = 100.

Prevalent in some cases associated with the zeolites, with which it is often associated ; the former, however, has been shown to go off only at a red heat, and hence plays a different part.

Pyr., •to.—In the closed tube yields water. B.B. fuses at 2 with intumesce to a blebby enamel-like gloss. Decomposed slowly by hydrochloric acid without gelatinizing ; after

fusion dissolves readily with gelatinization.

DIA—B B. fuses readily, unlike beyl, green quartz, and chalcedony. Its hardness is greater than that of the zeolites.

Obs.—Occurs chiefly in basic eruptive rocks, basalt, diabase, etc., as a secondary mineral in veins and cavities, often associated with some of the zeolites, also with talc, mica, calcite, but commonly one of the first formed of the series; also less of leucite, nepheline, syenite, and then frequently associated with epidote; sometimes associated with native copper. Occurs in the L. Superior region.

At St. Clair, Mich. near Bourg d'O

in Piedmont: in the Barz. near Anni, „, „, „, - - -.

Norway: at Fofna in Sweden (edelite); at Coratorphne Hill, near Edinburgh; Mourmels., Ireland.

In the United States. finely crystallized a Bill. N. J.—In syenite, at Bomerville, Mass.; < copper region.

Named (1790) after Col. Prebn, who brought the mineral from the Cape of Good Hope.

Haratite^t. An acid orthosilicate of manganese and calcium. In small colorless prismatic crystals. B. = 60. Q. = 8-049. From the Baraig mine, near Pajsberg, Wern-land, Sweden.

Onapidins. Contains silica, lime, fluorine, and from alteration carbon dioxide: formula doubtful. In minute spear-shaped crystals B. = 61. Q. = 3-868-a-860. Color pale rose-red. From Vesuvius, in ejected masses in the tufa of Monte Somma.

17. Subsilicates.

The species here included are basic salts, for the most part to be referred either to the metasilicates or orthosilicates, like many basic compounds already included in the preceding pages. Until their constitution is definitely settled, however, they are more conveniently grouped by themselves as Subsilicates.

SILICATES.

443

It may be noted that those species having an oxygen ratio of silicon to bases of 3 : 3, like topaz, andalusite, sillimanite, datolite, etc., also calamine, carpholite, and perhaps tourmaline, are sometimes regarded as salts of the hypothetical parasilicic acid, H₂SiO₃.

The only prominent group in this subdivision is the White Group.

Hutnite Oroap.

Prolectite [Mg(F,OH)],Mg[SiO,],? Monodinic Chondrodite [Mg(F,OH)l,Mg,[SiO.],
Monoclinic

1-0803:1:1-8861 10863:1:3-1447

l:&:6

1-0802:1:4-4033

1-0803; 1:5-6588

90"

Hamite [Mg(F,OH)],M:g,[SiO,], Orthorhombio

CUaohumite [Mg{F,OH)],Mg,[SiOJ. Monoclinic

The species here included form a remarkable series both as regards crystalline form and chemical composition. In crystallization they have sensibly the same ratio for the lateral axes, while the vertical axes are almost exactly in the ratio of the numbers 3:5:7:9(seu also belowj. Furthermore, though one species is orthorhombio, the others monoclinic, they here also correapond closely, since the axial angle y9 in the latter cases does not sensibly aifler from 90°.

In composition, as shown by Penfield and Howe (also Sjogren), the last three species are basic orthosilicates in each of which the univalent group (MgF) or (MgOH) enters, while the Mg atoms present are in the ratio of 3:5:7. The composition given for Prolectite is theoretical only, being that which "would be expected from its crystallization. In physical characters these species ore very similar, and several of them may occur together at the same locality and even intercrystallized in parallel lamellfe.

J chijsollte and chrjioberyl.

ProlecUle d: %:\h^l -0803

ChondTodlle d: t:\e = \<mz

Humlle I: 4: Ji = l-0802

Clinohumite <i: i:ie = 1080a

Clirygollle J : 2« : h = 1'0786

Chvyaoberyl i:2d: i = I-M87

: 0-8387

1:0-6170

OHONDaoDrrE-HUMITB—OUNOHaMrm.

Axial ratios as given above. Habit varied, Figs. 858 to 866. Twins common, the twinning planes inclined 60° , also 30° , to c in the brachydome or clinodome zone, hence the axes cross at angles near 60° ; often repeated as trillings and as polysynthetic lamellae (cf. Fig. 556, p. 226). Also twins, t with c (001) as tw. plane. Two of the three species are often twinned together. \

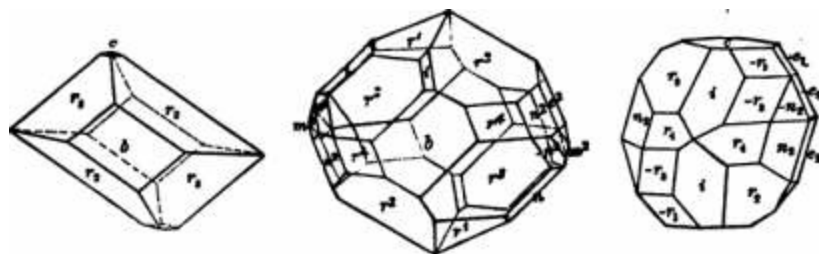
Cleavage: c sometimes distinct. Fracture subconchoidal to uneven. Brittle. $H. = 6-6.5$, $G. = 3.1-3.2$. Luster vitreous to resinous. Color white, light yellow, honey-yellow to chestnut-brown and garnet- or hyacinth-red. Pleochroism sometimes distinct. Optically +.

Chondrodite. Absorption $\ll c \gg$. Optically +. Ax. π and B_n , 1 b. $Bi. A_4 = a A_4 = +5^\circ$
Off Brewster: $28^\circ 68'$ Kffveltorp; SO" approx., Hte. Bomma. (S=1-619;

$= + 11^\circ-13^\circ$; 71° approx., Brew-

ed by GoOgIc

DESCRIPTIVE MINERALOGY.



Figs. 858, 8n, Chondrodite. Brewster, N. T.

*{018}, \ll , (108). \bullet , (101), r , (187), r , (188), r , (185),

861. 862.

Chondrodite, Sweden. (121) . n , (III).

863.

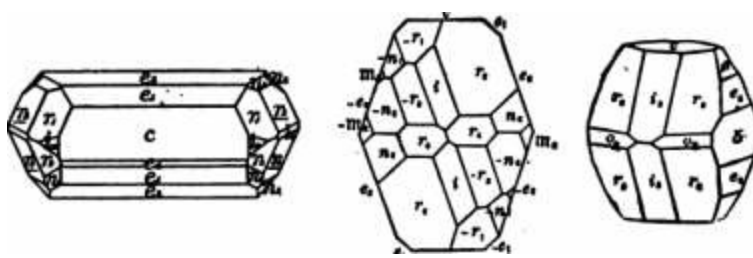
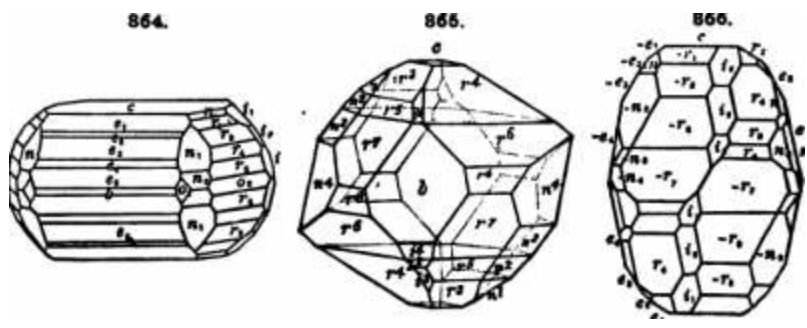


Fig. 801, 60a, Ghmdrvdite, Ute. Bomnu.

SymboU see above.

864. 866.

BumiU, Sweden. o,(aiO), «,(OIS),e,(O18).



Bvmit*, VsMlvlu* CUnnJi'imiU. Brewstr. CUnohumitr, Mie. Snmmft.

(.Oil), r, (2 1-10), r, (316). t", (014|, i', (012! . <, (1081, «,(101). r.{i27). r.(i23).

r.(313). n,(I13), n,(MI). r. (12.'). r. (121).

Comp—Basic flaoeilicates of magneBium with related formulas as shown in the table above. Hydroxy! replaces part of the fluorine, and iron often takes the place of magnesium.

Analyses bj Peu field:

. Coogic

P7T., ato.—B.B. lofiulble; loiii vsrletla blacken and lieu bum wblW. Fused wiib potBaaliin biaulphniB in the closed tube ^veg a reactlun for fluorine. With the fluxei & reaction for iron. Gelatinize* nlth acids. Beated with sulphuric add gives uS silicoa fluoride.

Oba.—Chondrodite. bumlle, aud cllnobunilU: all occur at Vesuvius in the ejected maiMS both of limestone or feldspathic type found on Honte Somma. They are associated with chrvsultte, bioUte, pjroxene, magnetite, spinel, vesuvlauile, calcite, etc; also less often wiln sanidine, meionite, nephelite. Of the Ibree species, bumite Is the rareat and clinubumite of most frequent occurrence. Thej seldom all occur logvlber In the same

muss, and only rarely two of the species (as humite and clinobumite) appear together. Occasionally clinobumite in terpenite crystals of humite, and parallel intergrowths with clinochlore have also been observed.

Chondrodite occurs in Mts. Bomma, as above noted: at Pargas, Finland, honey-yellow in thin sections; at Kafveltorp, Nya-Kopparberg, Sweden, associated with chalcopyrite, galena, sphalerite. At Brewster, N. Y., at the Tilly Putnam magnetite iron mine in deep

sections. Also probably at numerous points where the occurrence of "chondrodite" has been reported.

Chondrodite occurs at the Ladu mine near Finspång, Sweden, with magnetite in crystalline limestone. In crystalline limestone with clinobumite in Andalusia. Also in large coarse, partly altered crystals at the Tilly Foster iron mine at Brewster, N. Y.

Chondrodite occurs at Mts. Bomma and in Andalusia; in thin sections in limestone near Lake Baikal in Siberia; at Brewster, N. Y., in rare but highly modified crystals.

Chondrodite is from the Eo mine, Kordmark, Sweden; very rare; in thin sections.

Numerous other localities of "chondrodite" have been noted, chiefly in crystalline limestone; most of them are probably to be referred to the species chondrodite, but the identity in many cases is yet to be proved. At Brewster large quantities of "chondrodite" occur associated with magnetite, enstatite, ripidolite, and from the extensive alteration serpentine has been formed on a large scale. The granular mineral is common in Umeå in Sussex Co., N. J., and Orange Co., N. Y., associated with spinel, and occasionally with pyroxene and corundum. Also in Mass., at Belmsford, with fersite; at South Lee, in limestone. In Canada, in limestone at Mt. Jerome, Orenville, etc., abundant.

The name chondrodite is from $\chi\acute{o}\rho\delta\acute{o}\nu$, a grain, alluding to the granular structure. Humite is from Sir Abraham Hume.

f^

TLVATIE. Llevrite. Tenite. Orthorhombic. Axes $a:b:c = 0.6665:1:0.4427$.

$\alpha = 67^\circ 33'$. $\beta = 101^\circ 11'$.

$\gamma = 78^\circ 45'$. $\delta = 63^\circ 88'$.

Commonly in prisms, with prismatic faces vertically striated. Columnar or compact massive.

Cleavage: b, c rather distinct. Fracture uneven. Brittle. H. = 5-5-6. $\rho = 3.99-4.06$, Luster submetallic. Color iron-black or dark grayish black. Streak black, inclining to green or

brown. Opaque.

Comp.— $\text{CaFe}(\text{FeO})(\text{SiO}_2)$, or $\text{H}_2\text{O} \cdot \text{CaO} \cdot 4\text{FeO} \cdot \text{Fe}_2\text{O}_3 \cdot 4\text{SiO}_2$, $\text{SiO}_2 = 29.3$, iron sesquioxide 19.6, iron protoxide 35.2, lime 13.7, water 2.3 = 100. Manganese may replace part of the ferron iron.

Pyr., etc.—B.B. fuses quietly at 3-6 to a black magnetic bead. With the fluxes reacts for iron. Some varieties give also a reaction for manganese. Gelatinizes with hydrochloric acid.

Obs.—Found on Elba in dolomite; on Mt. Mulatto near Predazzo, Tyrol, in granite; Schneeberg, Saxony; Fölschum. In Norway. Reported as formerly found at Cumberland, R. I.; also at Milk Row quarry, Somerville, Mass. Named Hvitt from the Latin name of the island (Elba).

Ardennite. A vanadate-silicate of aluminium and manganese, also containing arsenic. In prismatic crystals resembling iwaite. H. = 6-7. G. = 8.630. Yellow to yellowish brown. Found at Salm Chateau in the Ardennes. Belgium.

Coogite

DESCRIPTIVE MINERALOGY,

Coogite. A black silicate with ferrous iron: formula doubtful. In black hexagonal prismatic crystals. H. = 6.5. O. = 4.918. Water metallic. From L&Qgan, Swtland.

The following are rare lead silicates. See also p. 408.

Eatroite*. Probably $8\text{PbO} \cdot 2\text{H}_2\text{O} \cdot 3\text{SiO}_2$. In prismatic crystals; often in leaf-like forms; also massive. H. = 10. O. = 6.19. Color dark reddish brown; black on the surface. From Loutheru Chili; Liogba, Bwedeti.

Melanotakite. $\text{aPbO} \cdot \text{aFe}_2\text{O}_3 \cdot 8\text{SiO}_2$, or $(\text{Fe}_2\text{O}_3)\text{Pb} \cdot (8\text{SiO}_2)$. Warren. Orthorhombic: prismatic. Uaasite; cleavable. H. = 6.5. G. = 5.78. Luster metallic to greasy. Color black to blackish gray. Occurs with native lead at Udgan, Sweden. Also in crystals resembling kentroite at Hillsboro, New Mexico.

Bertrandite. $\text{H}_2\text{BeSi}_2\text{O}_7$ or $\text{H}_2\text{O} \cdot 4\text{BeO} \cdot 2\text{SiO}_2$. Orthorhombic-bemirooplike. In small tabular or prismatic crystals. H. = 6-7. Q. = 2.60-2.60. Colorless to pale yellow. Ubiite occurs in feldspathic veins, often with other beryllium minerals as a result of the alteration of beryl. At the quarries of Barbin near Kantea, France; Plsek, Bohemia; Hill Anterior, Chaffee Co., Colo., with pbeucite; Stooeham, He.; Amelia Court-House, Va.

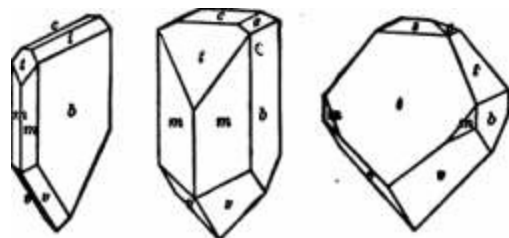
OALAMTWB. Smlthsonite. Hemimorphite. Eieselzinkerz, Oalmel pt. Oerm.

Orthorhombic-bemimorphic. Axes d:h -i^ o-7834 : 1 : o-4778. 868. S69. 870.

mm". nOAlIO= 76* 9*. •f", 101 A iOl = 62* 46'. It. 801 A 80I = 133* 41'.

on A Oil = sr v.

o81 A o3l = 110'12".



o81.. M2i

Crystals often tabular I b ; also prismatic ; faces b vertically striated. Usually implanted and showing one extremity only. Often grouped in sbeaf-like forms and forming drnsy surface in cavities. Also stalactitic, mam miliary, botryoid^, and fibrons lorois; massive and granular.

Cleavf^e: m perfect; s (101) less so; c in traces. Fracture uneven to sob* conchoidal. Brittle. H. = 4-5-5, the latter when crystallized. Q. = 3'40-3'50. Luster vitreous; c subpearly, sometimes adamantine. Color white; sometimes with a delicate bluisn or greenish shade; also yellowish to brown. Streak white. Transparent to translucent. Htrongly pyroelectric.

Comp.—H,ZnSiO, or (ZnOH) SiO. or H,o.2ZnO.SiO, = Silica 25-0, zinc oxide 67'5, water 7-5 == 100. The water goes off only at a red heat; unchanged at 340" C.

Pyr., ate.—lu the cldBcd tubt decrepitates, whitens, and gives off water infusible (F. = 6). On charcoal with soda gives a coatiDK which is yellow wiilta on cooling. Holsleiiied with cobnlt solution, and heated Id O.F. assumes a bright green color, but the Ignited mineral itself becomes blui with acids tven when previously ignited.

Dlff.—Characlerized by itsinfusihtlity; reacllon foreinc; gelatiniaalion with'aclds. Resembles some smitlisonlte (which cfferveges with acid), alsn prehnlle.

Obs.—Calamine iind smltlisonlte are usually found ftssoclated In veins or beds in stmlibed cnlcareous rocks accompanying sulphides of zinc, Iron nnd lend. Thus at Aix-ia. Cliivpelle; Raibel and Bleiberg, in Carintliis; Moreinet in Belgium; ReKb&nva, Srhemnltz. At Itoughti^ti Oil!, in Ciimberlnnd; at Alston Moor, white; near Matlock, lu Derbyabire; Leadhilis, Scotland-, at Nerchinsk, In easleru Sibi^ria.

In the Uulted States occurs at Sterling Hill, near Ogdensburg, N. J.. In fine clenr

crystalline masses. In Pennsylvania, at the Perkiomen and Phenixville lead mines; at

B.B. almost

is hot, and

this cooling

quartzizes

is by growth

Friedrich. Abundant in Virginia. at John's mines in Wythe Co. With the zinc
deposits of southwestern Missouri. especially about Grant, both as crystallized and
massive calamine. At the Emma mine, Cottonwood Canyon, Utah.

The name Cassia (with Cassia of the Germans is commonly supposed to be a cor-
ruption of Cassia. Agricola says it is from Cassia, a reed. It alludes to the cassia
form (cassitic) common in the cassiaform.

Orthoclase. SiO_2 . Orthorhombic-clinohedral (see Fig. 881, Silo, p. 104). H. = 5.5.
O. = 8.5. Colorless or white to translucent. From Franklin Furnace, N.J.

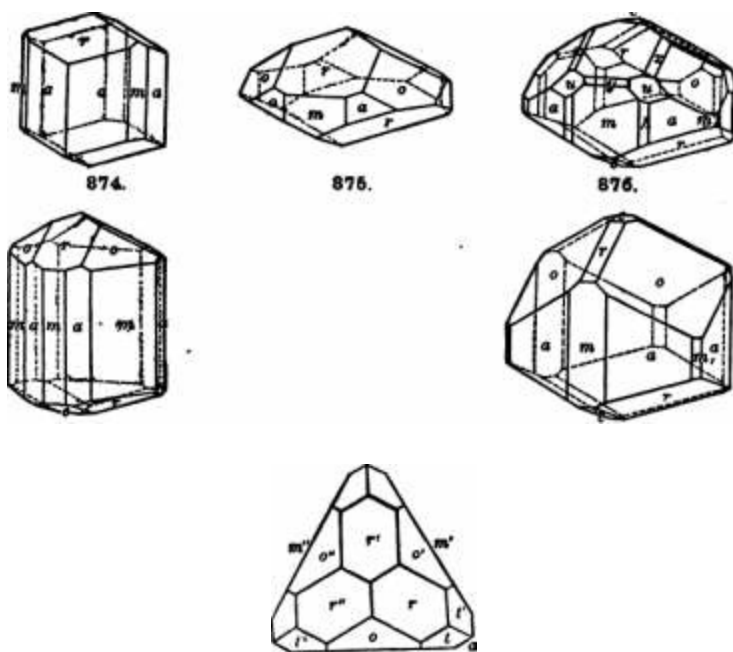
Orthoclase. H_2SiO_4 . It is radiated and stellated tufts. O. = 3-980. Color straw- to wax-
yellow. Occurs at the tin mine of Claiborne; Wippra, in the Harz, on quartz, etc.

Lawsonite. $\text{CaAl}_2\text{Si}_2\text{O}_{10}$. In primary orthorhombic crystals; $\text{mm} = 87^\circ 16'$. O. = 5-9.
Luster vitreous to greasy. Colorless, pale blue to bluish. Occurs in crystalline
schists of the Tibum peninsula, Marin Co., California; also in the schists of Ponape
and New Caledonia.

Orthoclase. A silicate of the calcium metal, with water. Crystals rare; commonly massive;
granular. H. = 5.5. G. = 2.65. Color between clove-brown and cherry-red to gray. Occurs at
Bastna, near Kiddarhyttan, Sweden.

TOURMALINE. Turmaline form.

Rhombohedral-hemimorphism. Al_2SiO_6 . $\text{a} = 0.4477$. $\text{c} = 0.001$. $\text{a} \perp \text{c}$. $\text{a} = 2r$. $\text{c} = 2t$. $\text{a} \perp \text{c}$.
 $46^\circ 68'$. $\text{e} = 0.001$. $\text{a} \perp \text{c}$. $\text{a} = 45^\circ 5'$. $\text{c} = 0.001$. $\text{a} \perp \text{c}$. $\text{a} = 77^\circ$. C, 871. 873.



Crystals usually prismatic in habit, often slender to acicular; rarely flattened, the prism nearly wanting. Prismatic faces strongly striated vertically, and the crystals hence often much rounded to barrel-shaped. The cross-section of the prism three-sided (m. Fig. 877), six-sided [a], or nine-sided (m and a). Crystals commonly hemimorphic. Sometimes isolated, but more commonly in parallel or radiating groups. Sometimes massive compact; also columnar, coarse or fine, parallel or divergent.

by Google

Cleavage: a, r difficult. Fracture conchoidal to uneven. Brittle and often rather friable. H. = 7-7.5. G. = 3.98-3.20. Luster vitreous to resinous. Color black, brownish black, bluish black, most common; blue, green, red, and sometimes of rich shades; rarely white or colorless; some specimens red internally and green externally; and others red at one extremity, and green, blue or black at the other. Streak uncolored. Transparent to opaque.

Strongly dichroic, especially in deep-colored varieties; axial colors varying widely. Absorption for c (vibrations J. i) much stronger than for e (vibrations j j ^); thus sections || c transmit sensibly the extraordinary ray only, and hence their use (e.g., in the tourmaline tongs (p. 219)) for giving polarized light. Exhibits idiophanous figures (p. 219). Optically ~. Birefringence rather high, a - e = 0.02. Indices: at, = 1.6366, e, = 1.6193 colorless var.; c, = 1.6435, e, = 1.6222 bl. green. Sometimes abnormally biaxial. Becomes electric by friction; also strongly pyroelectric.

Tar.— Ordinary, in crystals as above described; black much the most common. (a) Rubellite; the red, sometimes transparent: the Siberian is mostly violet-red [rubellite\ the Brazilian roach-rod; that of Chesterfield in the Goalie D, 4th B., pale roach-reil sud opaque; that of Paris, Me., fine ruby-red and transparent. (b) Indicolite or indigotite; the blue either] > ali > or bluish black; named from the indigo-brown color, (c) Brazilian Sapphire (In jewelry): Berlin-blue and transparent, (d) Brazilian Emerald. Chrysolite (or Peridot) of

Bratil; green and IranspareDt. («| Peridot of Ceylon; bouey-yellnw. (/) Arhroite; colorless
 tourmaline, from Elba, (g) Apliritite; black tourmaiue, from ErnngerO, Norway. (h)
 Co'uTnnai' and black; coarse columnar. Resembles somewhat common hornblende, but
 bus s more resinous fracture. SDd Is without ilistinct cleavage or uuything like a fibrons
 appearance In the texture; it often baa the appeaiaoce on a broken surface of some kinds
 of soft coal.

Comp.—A complex silicate of boron and alnminim, with also either magnesium, iron or
 the alkali metals prominent. The oxygen ratio of Si : B is in general 2 : 3 and the formula
 may hence be written: R,SiO, = K,8iO^ s H,SiO.. Here E = Na,Li,K; E = Mg,Fe,Ca; E =
 Al,B,Cr,Fe.

The varieties hased upon comporition fall into three prominent groupc, between whiclt
 there are many gradatioos:

1. Alkali Tourhalhik. Contains sodium or lithium, or bolh; also potassium. Q. = 8-0-3-1,
 Color red lo green; also Colorless.

3. Iron Tourhauhb. Q. = 8'l-8'3. . Color usually deep black.

8 Haokebiun TonRMAUNB. Q. = 80-3-09. Usually yellow-brown to brownish black; also
 colorless (nnal. Mi).

A chromium bntrmaline also occurs. O. = 3'130. Color dark gieea.

The following are typical analyses (RIgga) of the three varietiea:

1. Rumford. Me., rou:

HIO, TIO. W,o, A1,o, yrO MnO C»o MirO N«,o K,o tl.O H,o T 88-07 — 9-89 43-34 036
 085 OSB 0-07 218 044 I*«t 4-2« 038 = 100*8»

2. Auburn, blnek:

34-99 — 988 88-98 14-28 0-06 015 l-Ol 801 0-84 tr. 8-62 — =100-01)

3. Gouveraeur, brown;

87-39 119 10-73 27-89* 0-64 — 2-78 14 09 1-73 0-I« tr. 8-88 ft-f =100-48

■ Including 0-10 Ff,o,.

Pyr., ate.—The raaiiiteBln varieties fuse rnther easily to a white blebby glass or slag: the
 Iron-maRnesia var. fuse with a strong hent to a blebby slag or enamel; [he iron var. fnse
 with rtifficuly. or, in some, only on the edges; the irou-magnesia-llthia var. fuse oa the
 edfTPS, nnd often with great difllculty, aud some are infusible; the lithla var. are

infusible. With the fluxes many varieties give reactions for iron and manganese. Fused with

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« mixture of potash alum borate and fluor-spar give a distinct reaction for boric acid. Not decomposed by acids.

DHL — Characterized by its crystalline, prismatic forms which are three-, hex-, or dodecahedral in termination; massive forms in columnar or fibrous; also by absence of cleavage (unlike amphibole and epidote); in the common kinds by the coal-like fracture; by luster; by difficult fusibility (commonly), compared with garnet and vesuvianite. The boron test is conclusive.

Readily distinguished in thin sections by its somewhat lustrous relief; rather striated interference colors; negative uniaxial character; decided colors in ordinary light in which it is often biaxial. Also, especially, by its remarkable absorption which in the direction of crystal elongation is \pm to the vibration plane of the lower Nicol: this with its lack of cleavage distinguishes it from amphibole, which alone in igneous rocks may show similar strong absorption.

Obs.—Commonly found in granitic nodules as a result of fumarole action or of mineralizing gases to the fluid magma, especially in the pegmatite veins associated with such rocks; at the periphery of such masses or in the schists, or altered limestones, gneisses, etc., immediately adjoining them. It marks (especially the boundaries of granitic masses, and its association with those characteristic of stibnite occurrences; quartz, albite, mica, muscovite, etc. The variety in granular limestone or dolomite is commonly brown; the bluish-black var., sometimes associated with iron ores; the brown with titanium; the light variety is often associated with lepidolite. Reddened varieties, or both, occur in the Ural; Elba; Cernobila in Tessin, Switzerland; Penzance, Sutherland; also in the province of Minas Geraes, Brazil: yellow and brown from Ceylon; dark brown varieties from Eibenstock, Saxony; the Zillerthal: black from Arendal, Norway; Saarnin; Kimito; pale yellowish brown at Windisch Knipfel in Carinthia; fine hyacinth crystals occur in Cornwall at several localities.

Also in the U.S. in Maine at Paris and Hebron, magnesian red and green tourmaline with lepidolite, etc.; also blue and pink varieties; and at Norway; pink at Rumford, embedded in lepidolite; at Auburn in clear crystals of a delicate pink or lilac with lepidolite etc.; at Albany, green to black, Italy, at Cheslerfield, red, green, and blue; at a locality, blue and green; at Norwich. New Braintree and Carlisle, good black crystals. In N. Hampshire, at Orford, brownish black in steatite. In Conn., at Monroe, ditto in mica-slate; at Haddam. lustrous in mica-slate; also fine pink and green; at New

Mitford, black. In N. Tirk. near Gouverneur, brown crystals, with IrenioTiie, etc., in grHDiilir limestone; black near Port Henry, Essex Co.; near Edenville; splendid black crystals at Plerrepout, St. Lawrence Co.; colorless and slasiy at De Ealb; dnrk brown at McCumb. In N. Jerty, at Hamburg and Newlon, black and brown crystals in limestone, wiib spiucl: also grass-green crystals tu crystidliue linieslone near Franklin Piimac. In Pen'i... lit Newliu, Chester Co.; near Unlouville, yellow; at Chester, flne bluck; Middle-to»ii, black; Marple, green iu lulc; near New Hope on the Delaware, large black cryptnls. A chrome v ir, from the chromile beds in Montgomery Co., jtforv.and. In y. Car.. Alexander Cii.. in flue black crystals with emerald and hlddenlte. In California, flue groups of rubsllite in lepidolite in 8iu Diego Co.

In Caiyida, \a the province of Quebec, yellow crystals in limestone at Calumel Falls, Litchfield. Pontine Co.; at Hunlerslow; blue to white crystals at Clarendoo. Pontine Co.; black at Qreuville and Argenleuil, Argeuleiil Co. In Ontario, in fine crystals at N, Burgess. Lanark Co.; Qalway sod Stonuy L. in Dummer, Peterborough Co.

The name IttmuUin from TuraratUi in Cingalese (applied to zircon by Jewelers of Ceylon) was introduced into Hiillaod in 1708, with a lot of gems from Ceylon.

Domortieita. A basic aluminium silicate, perhaps $4Al_2O_3 \cdot 8BiO_2$.

Orthorhombic. Prismatic angle approximately 90° . Usually in fibrous to columnar aggregates. Cleavage: a distinct; also prismatic. Imperfect, $H. = T. O. = 8'26-3^*8fl$. Luster vitreous. Color bright amethyst blue to greenish blue. Transparent to translucent. Pleochroism very strong: c colorless, b reddish violet, a deep ultramarine-blue. Exhibits idiochrous figures, analogous to andalusite. Optically —. $Ai. pi. \setminus b. Bx 1 \ll$.

Recognized in thin section by its rather high relief; low interference colors (like those of quartz); occurrence in slender prisms, needles or fibers, with negative optical extension; parallel extinction; biaxial character and especially by its remarkable pleochroism.

Found embedded in feldspar in blocks of gneiss at Chapimost, near Lyons. Prance; from Wolfshau. near Schmiedeberg. Silesia; in the foliation of the gneiss of Tvedestrand, Norway. In the U. S., it occurs near Harlem, New York Island, in the pegmatoid portion of a foliated gneiss; In a quartzose rock at Clip. Yuma Co., Arizona.

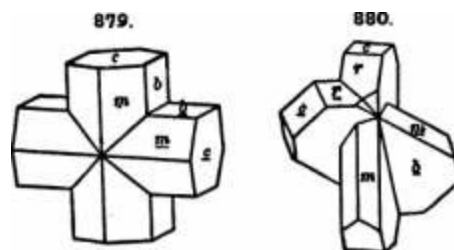
..Coogenic

DIAGNOSTIC MINERALOGY.

STAUROLITE. Silicofluoride. Orthorhombic. Axes $d : h : i = 0.4734 : 1 : 0.4734$, $HO A liO = 50^\circ 40''$. «

rr'' , $101 A 101 = 110^\circ 83' . n$

Twins cruciform: tw. pi. x (033), the crystals crossing nearly at right angles; tw. pi. z (333), crossing at an angle of 60° approximately; tw. pi. y (330) rare, also in repeated twins (cf. Figs. 359, p. 132, 409, p. 138, 411, p. 129). Crystals commonly prismatic and flattened (J; often with rough surfaces. 878. 879



Cleavage: indistinct, but interrupted; in traces. Fracture subconchoidal. Brittle. H. = 7-7.5, G. = 3-65-3-77. Subvitreous, inclining to resinous. Color dark reddish brown to brownish black, and yellowish brown. Streak uncolored to grayish. Translucent to nearly or quite opaque. Pleochroism distinct: C (= ^) hyacinth-red to blood-red, a, byellowish red; or c gold-yellow, o, b light yellow to colorless. Optically -|- . Ax. pi. II a. Bx ± e. 2H_∞ = 113° 10'. p = 1-76, y — a = 0012. ' Comp.—HFeAl.Si.O₃, which may be written (AlO), (AlOH)Fe(SiO₂), or H₂O.3FeO.5Al₂O₃.4SiO₂. = Silica 36.3, alumina 55.9, iron protoxide 15.8, water 3.0 = 100 Penfield. Magnesium (also manganese) replaces a little of the ferrous iron; ferric iron part of the aluminium.

SordmarkiU from Nordmark, Sweden, conlaid loadga*

.. ^ Tnanganestaii Tsrlety, wblcli fuses easily to a.

blaoWmaKneicKlau. Willi the fluxes gives reactions for iron, and sometimes for manganese. Imperfectly decomposed by salpburic acid.

Diff.—Characterized by the obtuse prisms (unlike andalusite, which is nearly square); by the frequency of twinning forms; by hardness and infusibility.

Under the microscope, sections show a decided color (yellow to red or brown) and strong pleochroism (yellow and red); also characterized by strong refraction (high relief), rather right interference-colors, parallel extinction and binocular character (generally positive in the direction of elongation). Easily distinguished from rutile (p. 84S) by its biaxial character and lower interference-colors.

Obs.—Usually found in crystalline schists, mica schist, argillaceous schist, and gneiss, as a result of regional or contact metamorphism: often associated with garnet, sillimanite, cyanite, and tourmaline. Sometimes encloses symmetrically arranged carbonaceous impurities like andalusite (p. 488). Other impurities are also often present, especially silica, sometimes up to 30 to 40 p. c.; also garnet, mica, and perhaps magnetite, brookite.

Occurs with cyanite in paragonite schist, at Mt. Campinne, Switzerland: In the Zillerthal, Tyrol; Guldenslein in Moravia; Aschaffenburg, Bithun; in large twin crystals in the micaceous schists of Brittany and Scotland. In the province of Minas Geraes, Brazil.

^AlOglc

Abundant throughout the mica schists of New England. In Maine, at Wadsworth. In N. Hampshire, at Ferrisburgh; at Lisburne; on the shores of Hiak Pond, located in the south. In New Jersey, at Chesterfield. In the state of Connecticut, at Butts, VerDol, «(c.; Southbury with

garnets: at Littlefield, black crystals. In S. Carolina, near Franklin, Maco Co.; also in Madison and Clay counties. In Georgia, in Fannin Co., occurs in the zone in the schists. Named from *crataegus*, a rose.

Kyanite. Al_2SiO_5 . In fibrous to columnar aggregates, resembling emerald. H. = 5. G. = 3.58. Colorless to white, translucent.

Kyanite occurs at Fiskerfjord on the west coast of Greenland. First found at Waldheim, Saxony.

Sapphirine. $\text{Mg}_2\text{Al}_2\text{Si}_2\text{O}_{12}$. In indistinct tabular crystals. Usually in disseminated grains, or aggregations of crystals. H. = 8. Q. = 8.13-8.48. Color pale blue or green. From Fiskerfjord, Southwestern Greenland.

SrUOATBS.

Section B. Chiefly Hydrated Species.

The Silicates of this second section include the true hydrous compounds, that is, those which contain water of crystallization, like the zeolites; also the hydrous amorphous species, as the clays, etc. There are also included certain species—as the Micas, Talc, Kaolinite—which, while they yield water upon ignition, are without doubt to be taken as acid or basic metasilicates, orthosilicates, etc. Their relation, however, is so close to other true hydrous species that it appears more natural to include them here than to have placed them in the preceding chapter with other acid and basic salts. Finally, some species are referred here about whose chemical constitution and the part played by the water present there is still much doubt. The divisions recognized are as follows:

I. Zeolite Division.

1. Introductory Subdivision. 2. Zeolites.

II. Mica Division.

1. Mica Group. 2. Clintonite Group. 3. Ctilorite Group. III; Serpentine and Talc Division. Chiefly Silicates of Magnesium.

IV. Kaolin Division. Chiefly Silicates of Aluminium; for the most part belonging to the group of the clays.

V. Concluding Division. Species not included in the preceding divisions; chiefly silicates of the heavy metals, iron, manganese, etc.

I. Zeolite Division.

1. Introductory Snl>diTlaioD,

Of the species here included, several, as Aponhyllite, Okenite, etc., while not strictly Zkoiitbb, are closely related to thinnite in composition and method of occurrence. Peckolite (p. 895) and Prehnite (p. 443) are also sometimes classed here.

;Coogic

DISCRIPTIVE (CINEEALOOY).

$\text{H}_2\text{SiO}_4 \cdot \text{Ca} + \text{H}_2\text{O}$. Crystals small, prismatic: also fibrous, and banded. BpLerulitic. $U = 6$. ($d = 31129$, Color roae- to flesb-nsU. Uccnra at the maogaDise mLCB near Dilleuburg, Qerm&Df. BJwdotiliie Is.tlie wme Bpecler from the Haratlg mice, Pttjaberg, Sweden.

$\text{SiO}_2 \cdot 0.7\text{H}_2\text{O} \cdot \text{Al}_2\text{O}_3 \cdot 8\text{H}_2\text{O}$. In short prismatic crystals also tabular, translucent. Color brown. $B = 4-45$. $O = 384$. From the Haistig fflim, near Psjs-berg, Sweden.

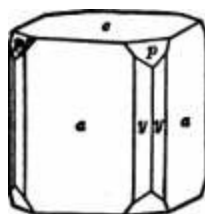
$\text{H}_2\text{SiO}_4 \cdot \text{Ca} + \text{H}_2\text{O}$. Oommoijlr fibrous: also compact. $H = 45-5$. $G = 2'28-2'36$. Color white, with a shade of yellow or blue. Occurs in basalt or related eruptive rocks; as in the Fsrber; Iceland; Disko. Greenland: Foods, India, etc.

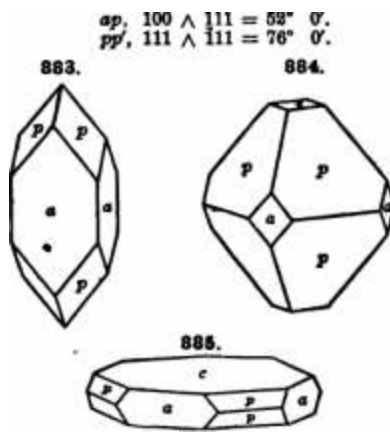
$\text{H}_2\text{SiO}_4 \cdot \text{Ca} + \text{H}_2\text{O}$. In white concretions, in radiate structure. From the Isle of Skye, with albite, laumontite, etc.; in India, etc. With apophyllite of New Almadeu, Cullforqlai also in N. Scotia,

APOPUI IjUTB.

Tetragonal. $a_1 = a_2 \neq c$

$\alpha, 001 \wedge 111 = 80^\circ 83''$.





Habit varied; in square prisms (a) usually short and terminated by c or by cp, and then resembling a cube or cubo-octahedron; also acute pyramidal {p} with or without c and a; less often thin tabular || c. Faces c often rough; r bright but vertically striated; p more or less uneven. Also massive and lamellar; rarely concentric radiated.

Cleavage: c highly perfect; m less so. Fracture uneven. Brittle. H. = 4-5. G. = 2.3-3.4. Luster of c pearly; of other faces vitreous. Color white, or grayish; occasionally with a greenish, yellowish, or rose-red tint, flesh-red. Transparent; rarely nearly opaque. Birefringence low; usually +, also -. Often shows anomalous optical characters (Art. 411, Fig. 565). Indices: $\omega = 1.5309$ Li, $e = 1.5333$.

Comp.—H, KCa, (SiO₂) · 4H₂O or K₂O · 8CaO · 16SiO₂ · 16H₂O = Silica 53-T, lime 35.0, potash 5.3, water 16*1 = 100. A small amount of fluorine replaces part of the oxygen.

The above formula differs but little from H₂CaSi₂O₇ · H₂O. In which potassium replaces part of the basic hydrogens. The form often accepted, H₂(Ca·E)₂O₈ · H₂O corresponds less well with the analyses,

Pyr., etc.—In the closed tube effluates, whitens, and yields water, which reacts with the open tube, when fused with salt of phosphorus, gives a fluorine reaction. B.B.

Effluates, colors tire flame violet (potash), and fuses to a white vitreous globule. T. = 1-6. Decomposed by hydrochloric acid, with separation of silicic acid.

Diagnosis—Characterized by its tetragonal form, the square prism and pyramid. It is commonly identified by the perfect basal cleavage and pearly luster on the surface.

Occurs as a secondary mineral in basalt and related rocks, vitric tuffs, zeolites, also datolite, pectolite, otleyite; also occasionally in cavities in granite, gneiss, etc. Greece, Iceland, the Feroes, and British India afford fine specimens of apophyllite to amygdaloidal basalt or diabase. Occurs also at Andreasberg, of a delicate pink; Radatithal in the Urals; at Orföwitsza, Hungary, with wollastonite; Ulo, Sweden; on the Seisser Alp in Tyrol; Guhoajiató, Mexico, often of a beautiful pink upon amethyst.

In the U. S. large crystals occur at Bergen Hill, N. J.; in Ponn., at the Frencli Creek mines.

Chester Co.: at the CiiQ mine. Lake iiiiperlor regioui Table Mt. near Golden, Colo.: in California, at the mercury mines of New Almaden often stained bro^n bjbiiumeu; also from Nova Scotia at Cape Blomidon, and other poiuls.

Named by HaQy in allusion to its tendency to exfoliate under the blowpipe, from dxu and ^OXXor, a leaf. Its irliilish pearly aspect, resembling Ibe eye of a tish after boiling, gave rise to the earlier name lehthyophOtalnite, from ixfivi.JUh, u<p6aXfiiii, tye.

2. Zeolites.

The Zbolites form a fatnil^ of well-defined hydrous ailicatea, cloeelj re lated to each other id compoBition, in conditioneof formation, and hence iii method of occurrence. They are often with right spoken of sa analogous to the Feldspars, like which they are all silicates of aluminium with sodittm and caloinm chiefly, also rarely barium and strontium; magnesinm, iron, etc., are absent or present only through impurity or alteration. Further, the composition in a number of cases corresponds to that of a hydrated feldspar; while fusion and slow recrystallization result in the formation from some of them of anorthite (CaAl,Si,o,) or a calcium-albite (CaAl,Si,o,,) as shown by Doelter. The Zeolites do not, however, form a single group of species related in crystallization, lilce the Feldspars, but include a number of independent groups widely diverse in form and distinct in composition ; chief among these are the monoclinic Phillipsite Geoup; the rhombohedral Cdabazite Gkoup, and the orthorhombio (and monoclinic) Natrolite Group. A transition in composition between certain end compounds has been more or less well established in certain cases, but, nnlike the Feldspars, with these species calcinm and sodinm seem to replace one another aud an increase in alkali does not necessarily go with an increase in silica.

Like other hydrous silicates they are characterized by inferior hardness, chiefly from 30 to 5-5, and the specific gnivity is also lower than with corresponding anhydrous species, chiefly 2"0 to 2"4, Corresponding to these characters, they are rather readily decomposed by acids, many of them with gela-tinization. The intumescence B.B., which gives the name to the family (from Z^iv, to boil, and \WoS, stone) is characteristic of a large part of tlie species.

The Zeolites are all secondary minerals, occurring most commonly in cavities and veins in basic igneous rocks, as basalt, diabase, etc. ; less frequently in granite, gneiss, etc. In these cases the lime and the soda in part have been chiefly yielded by the feldspar; the soda also by elaulite, eodaJite, etc.; potash by leucite, etc. The different species of the family are often associated together; also with pectolite and apophylHte (sometimes included with the zeolites), datolite, prehnite and, further, calcite.

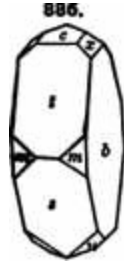
Ptilollte. BAl,Si,pO.,+5H,o. Here R = On : K, : Na, = 6 ; 2 ; 1 approx. In abort capillary needles, ajfgregeted !□ dellcnle tufts. Colorless, white. Occurs tipon a hluidi chalcedony iu cavities in a vesicular augite-anilesite found in fragments in the conglomerate beds of

Oreen and Table mounlaius, Jefferson Co., Colorado.

Hordenlta. $3\text{RAI,SI,gO}_2 \cdot n\text{H}_2\text{O}$, where $\text{H} = \text{E} : \text{Na} : \text{Ca} = 1 : 1 : 1$. In minute crystals resembling beulandite in habit and angles: also in small hemispherical or reni-

DESOBIPTIVB UIKEBALOOT.

form concretloDB with flbroua structure. $\text{H} = \text{S-4}$. $\text{O} = 9\text{IS}$. Color wllie, yellowish or piuklBli. Occurs nesr Hordeo, ElDg's Co., Nots Scotia, tn tnpj also In western Wyoming Dear Hoodoo Ht., on the ridge forming the divide betweeD Clark's Poric and ibe East Fork (I^msr R.) of the Yellowstoue river.



HBUZiANDITB. 8tllblt« loma author*. Monocliaic Axes: $d : i : d = 0.4035 : 1 : 0.4393$; $\angle = 88^\circ 34'$.

«w»", $110 \text{ A liO} = 48^\circ 56'$. » . $001 \text{ A } 501 = 66^\circ - \text{C}$.

tt, $001 \text{ A } aoi = 68^\circ *$). M. $001 \text{ A } Oai = 40^\circ 88'$.

Crystals sometimeB flatteued | h, the enrface of pearly luster

(Fig. 886; also $\text{Fi}^\wedge . 3\text{S}$, p. 11); form often suggestive of theortho-

I rhombic system, since the angles cs and ct di&r but little. Also

'i globular forms; graunlar.

Cleavage: S perfect. Fracture snbconchoidal to uneven. Brittle. $\text{H}^\wedge 3-5^\wedge$. $a. = 2-18-2-23$. Luster of b strong pearly; of other faces vitreous. Color various shades of white, passing into red, gray and brovii. Streak white. TranspareDt to subtrauslucent. Optically -)-. Ax. pi. and Bi, $J_ b$. Ax. pi. and BXo for some localities nearly $\setminus c$; also for others nearly $_Lc$ in white light {Dx}, Bio A ($i = + 57i^\circ$. Axial angle variable, from 0° to $92'$; nsnally $2\text{Er} = 52^\circ$. Birefringence low. $y\text{S} = 1''499$; $y - a = 0007$.

Comp.— $\text{H,CaAl}_2(\text{SiO}_3)_2 \cdot 3\text{H}_2\text{O}$ or $5a,\text{O.CaO.A1}_2\text{O}_3,6\text{SiO}_2 = \text{Silica } 59.2, \text{ alumina } 16.8, \text{ lime } 9.3, \text{ water } 14.8 = 100$.

StroQtIum is usually present, sometimes up to 8'o p. c. Pyr.—As with Btilbite. p. 4S7.

Obs.—Beuladille occurs principally in basaltic rocks, associated with clabazlie, sil-blte and other zeolites; also in gneiss, and occasionally in metalliferous veins.

The finest specimens of this species come from BeruQord. and elsewhere in Icelimd; the FarOer; in British India, neiir Bnbav; also in railrond cutlings in the Bhor and Thiil GL&ls. Also occurs in the Eilpatrick Hills, near Glasgow; on the I. of Skye; Fassalhal, Tyrol; Andreasbi:r\$(, Harz.

In the United States. In diabase at Bergen Hill. New Jersey; on north shore of Lake Superior; with haydeuite, at Jones's Falls, near Baltimore {baumontite). At Peter's Polut, NoTiL Scotia; alsil at Cape Blomldon, and other points.

Named after the Gaulish mineralOEIcal collector, U. Heuland, whose cabinet was the boals of the clasBlcnl work (1887) of Levy.

Brswitoilta. $H<(Sr.Ba,'Ca)A_1,SI)O_{,,4-8}HiO$. It) prismatic CiysUla. $H = S$. $Q. = \% - 4S$. Color white, Inclining to yellow and gray. From Strontlan in Aigyleehre; the Giant's Causeway; near Freiburg in BreUgau, etc.

BpUtilMU. Probably like heulandlte. $H,CaAl,ei)O.i + 8H,o$. Crystals monocllntc, UQitormly twins; habit prismatic. In radiated spherical aggregations; also gmaulBr. $O. = 2-39$. Color white. Occurs with scolecite at the Beruflord, Iceland; the FAROer; Poona, India; fu small reddish crystals, at Margaretville, N. Scotia, etc, BtiitiU Is from Sanioiln.

The above species, while crystallizing in the monoclinic system, are remarkable for the pseudo-symmetry exhibited by their twinned forms. Certain of

, Google

these twins are pseado-orthorhombic, others paeudo-tetrsgonfti and more complex twins even pseu do-isometric.

Frawulua has shown that the species of this group may be regarded ufonning n series, in which tbt: rulio of $BO : Al,o>$ is constuit (=1:1), aud thut at $SIO, : H.O$ obo chieUj ' I : 1. The end coiupoude assumed by him are:

$RAiSiO_{,,} + 6H,o$; $R.Ai,8i.o_{,,} + 8H,o$. Here $R = Ca$ chleQ^, in phillpsite and sUlbitf, Ua in iiarnioluueie, while in wellilte Ba, Ca. mid K, ure priiseuL; also in smaller amounts Nsi, Kt. The first of ibe above cotupoiuuda maj be regarded as a hydmted calcluiu albite, Ibe secuud as a hydraleil anoritilte, Pratt ana Fnoti:, lioweuer, show tbiit the anortbiie end compouDd more probably bas the fonnulti $RAiSiU. + 2H]o$ (or this doubled). The formulas glTen beyond are those comspoding to reliable imalysus vt certalu typical occurrences.

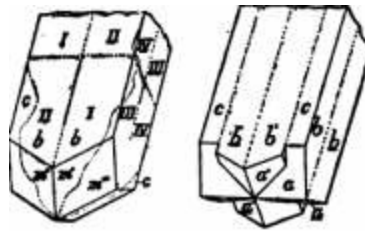
WallalU. $RAiSi.Oit + 3H,o$ wiLh $R = Ca$: present in small aDiuitit. PerceDtapie

composition: SiO₂ 43.0, Al₂O₃ 24.8, BaO 6.6, CuO 1.8, K₂O 8.1, H₂O 13.8 = 100. Monoclinic (axes p. 454 in complex twins, analogous to Ibohu of phillipsite and harmotome (Figs. B87, 868). Brittle. No cleavage. H. = 4-4-5. G. = S³T⁸-3³6a. Luster vitreous. Colorless to white. Optically+. Bij. i(010). Bi-refr. Rence weak. ,

Occurs at the Buck Creek (Cullakauwee) corundum mine in Clay Co., No. Carolina; in isolated crystals attached to feldspar, also to bombolite and corundum^ intimately associated with chabazite.

= 8:1:8; I

and Na also



PHU_iUPBITB.

Monoclinic. Axes S: $i = 0.7095 : 1 : 1.2563$; $\beta = 5^\circ 37'$.

mm³, 110 A liO = 60° 43". em, 001 A HO = 60° BO*. uf, 100 A iOl = 34° 38'. «'. Oil A Oil = 82° 4'. Crystals uniformly penetration-twins, but often simulating orthorhombic or tetragonal forms. Twins sometimes, but rarely, simple (I) with tw. pi. e, and \ then cruciform so that diagonal parts on b belong together, hence a fourfold striation, Q edge b/m, may be often observed on b. (3) Double twins, the simple twins just noted united with e (Oil) as tw. pi., and, since ee' varies but little from 90°, the result is a nearly square prism, terminated by what appear to be pyramidal faces each with a double series of striations away from the medial line. See Figs. 4³-424, p. 130; also Fig. 363, p. 133. Faces b often finely striated as just noted, but striations sometimes absent and in general not so distinct as with harmotome; also m striated] edge b/m. Crystals either isolated, or grouped in tufts or spheres, radiated within and bristled with angles at surface.

Cleavage: c, b, rather distinct. Fracture uneven. Brittle. H. = 4-4-6, Q. — 2 2. Luster vitreous. Color white, sometimes reddish. Streak uncolored. Translucent to opaque. Optically -+-. Ax. pi. and Bx^ _L 6. The ax. pi. lies in the obtuse angle of li i, and is usually inclined to a about 15° to 20°, or 75° 10 70° to the normal to e. The position, however, is variable. 2H., = 71 "-SI", Comp. — In some cases the formula is (E,Ca)Al₂Si₄O₁₁ + 4H₂O = Silica 48.8, alumina 20.7, lime 7.6, potash 6.4, water 16.5 = 100. Here Ca : K, = 2 : 1.



.dbyGoogle

Pyr., vto.—B.B. crumbles & fuses at 8 to & white enamel. Qtlalfnizea wltti hydro-chloric Bc!d.

Obg.—111 irauslucent crystals in basalt, ut the Oinnt's Causenay, Ireland; at Capo df Bove, neKr Rome; Ac! Casllo and elsewhere Id Sicily: among the lavas of Hie. Sorania; at BtBmpel, near Marburg; Aunerod, near Ofesseu; lu the Kaiserelulil. with faujnsUe; S&lesl, Buheuiu; in the aociet lavas o(thu Puy-de-Dfime.

Moiioclinic. Axes a-A-.6 = 0-7031 : 1 :1-2310; /3 = 55" 10'.

Crystals uniformly crnciform penetration-twins with c as tw. pi.; either

(1) simple twins (Fig, 890) or (2) nnited as fourlings with tw. pi. e. These double

690. twins often have the aspect of a square pnem with diagonal

pyramid, the latter with characteristic feather-like strintions

from the medial line. Also in more complex groups analogous

to those of phillipsite.

Cleavage: b ea^, c lees bo. Fracture uneven to subcon-ohoidal. Brittle. H. = 4-5. G. = 244-2-50. Luster vitreous. ' Color white; passing into gray, yellow, red or brown. Streak white. Sub trans parent to translucent. Optically 4~- Ax. pi. and. Bxb _L b. Ax. pi. in obtuse ?ngle a i and inclined about 65° to a and 60° to i. 2H,^, = 87° 2'. ^ = 1 516.

Comp.—In part H,(K,,Ba)Al,Si,o,, + 4H,o or (K,,Ba)o.

Al,o,.5SiO,.5H,o = Silica 471, alumina 160, baryta 206, potash 21, water 141 = 100.

Pyr., etc.—B.B. whitens, then crumbles and fuses without iutumesceDce at 3'5 to a white iranslucent glara. Some varieties phosplioresce when heated. Decomposed by hydrochloric acid wltioul gelalinizlnfc.

Obi.— Occora in basalt and similar eruptive rocks, also phoaolite, trachyte ; not 111 frequently on gneiss, and in some melalllferoiiia veins. At Sironlian, lu Bcotland; iu a

in siliceous veins at Aildreasberg in the Harz; at Hudeblad, Silesia; Oberaied, on a gale
In siliceous gneisses; at Kongsberg, Norway.

In the U. S. in small brown crystals with a luster on the gneiss of New York island; near
P. Arilur, L. Superior.

Named from dutoi, joint, and r⁵ tiv, to eiti, alluding to the fact that the pyramid (made
by the prismatic faces in twinning position) divides parallel to the plane that passes
through the terminal edges.

STILBITE. Detmone.

Monoclinic. Axes: $a : b : c = 0.7623 : 1 : 1.1940$; $\beta = 50^\circ 50'$.

Crystals uniformly cruciform penetration-twins with tw. π , c , Z_9 . analogous to
phillipsite and harmotome. The apparent form a "

rhombic pyramid whose faces are in fact formed by the planes m and m' ; the vertical faces
being then the pinacoids h and c (cf. Figs. |

5(H.)-562, p. 22T). Usually thin tabular $\parallel b$. These compound crystals are often grouped
in nearly parallel position, forming sheaf-like aggregates with the side face (b), showing
its characteristic pearly luster, often deeply depressed. Also divergent or radiated;
sometimes globular and thin lamellar-columnar.

Cleavage: 6 perfect. Fracture uneven. Brittle, H. = 3-5-4. |

G. — 2-09-1-2'205; 2-161 Haid. Luster vitreous; of b pearly. '

(Color white; occasionally yellow, brown or red, to brick-red. Streak uncolored.

Transparent to translucent. Optically —. Ax. π . II b. Bx, inclined 5° to axis a in obtuse
angle a^\wedge ; hence Bx, $A^\wedge = -55^\circ 50'$. Ax, angle approx. 52° to 53° (blue glass); $\gamma_3 = 1-498$.

.dbyGoogle

Comp.—For most varieties $H, (Na, Ca) Al, Si, O, + 4H_2O$ or $(Na, Cs)O, Al_2O_3, 6SiO_2, 6H_2O =$
Silica 57-4, alumina 16-3, lime 7-7, soda 1-4, water ITS = 100. Here $Ca^\wedge = 6:1$.

Some kinds show a lower percentage of Ca , and these have been called Anorthite.

Pyrr., etc.—B.B. exfoliate, swell up, curve into fan-like or vermicular forms, and fuses to
a white enamel. F. = 3-3'6. Decomposed by hydrochloric acid, without gelatinization.

DUL— Characterized by the frequency of radiating or sheaf-like forms; by the pearly luster
on the cleavage. Does not gelatinize with acids.

Obs.—Stibite occurs mostly in cavities in amygdaloidal basalt, and similar rocks. It is also found in some metalliferous veins, and in granite and gneiss.

Abundant in the Faroe; in Iceland; on the Isle of Skye, in amygdaloid; also in Dumfriesshire, Scotland, in red crystalline limestone; the Giant's Causeway, Ireland; at Andreasberg in the Harz, and Kongsberg and Arendal in Norway, with iron ore; on the Belasch Alp in Tyrol, and at the Pufflerloch (Puffenberg); on the granite of Striegau, Silesia. A common mineral in the Deccan trap area of British India.

In North America, especially in small crystals at Chester and at the Bomersville syenite

Mass.: at Phillipstown, N. T.; and at Bergen Hill, New Jersey; also at the Michigan Islands, Lake Superior. At Partridge Island, Nova Scotia; also at Isle Haute, Digby Neck, Cape Ulmston, etc.

The name stibite is from stibium, Latin; and ditomite from ditto, a German word. Osmundite. Perhaps $\text{CaAl}_2\text{Si}_2\text{O}_8 \cdot 4\text{H}_2\text{O}$. In pyramidal crystals, pseudo-tetragonal. $n = 1.54$. $G = 2.71$. Colorless or white, bluish white, grayish, reddish. Occurs in the leucophyre of Mt. Albano, near Rome, at Capodivertice, and elsewhere, etc.; on the Oeser glacier, near Zermatt; Schlaroth near Orlitz in Silesia; Bieleh, Bohemia, etc.

Zirconite. Leonhardtite. Caprochianite. Monoclinic. $d_{100} = 1.451$; $d_{110} = 1.05906$; $d_{111} = 0.6846$. Twins: tw. pi. a. Common form the prism m ($100^\circ = 93^\circ 44'$) with oblique termination, Q , SOI ($ce = 56^\circ 55'$). Also columnar, radiating and ditomite.

Characteristics: b and m very perfect; a imperfect. Fracture uneven. Not very brittle. $H = 3.5-4$. $G = 2.25-2.36$. Luster vitreous, inclining to pearly upon the face of cleavage. Color white, passing into yellow or gray, sometimes red. Streak uncolored. Transparent to translucent; becoming opaque and usually pulverulent on exposure. Optically —, Ax. pi. | i. Bx , $Ax = +65^\circ$ to 70° . Dispersion large, $p < I$; inclined, $Bligh. 21^\circ = 52^\circ 24'$.

Couper, Tar.— $\text{H}_2\text{CaAl}_2\text{Si}_2\text{O}_8 \cdot 2\text{H}_2\text{O} = 4\text{H}_2\text{O} \cdot \text{CaO} \cdot \text{Al}_2\text{O}_3 \cdot 4\text{SiO}_2 = \text{Silica } 511, \text{ alumina } 217, \text{ lime } 119, \text{ water } 153 = 100$.

Leonhardtite is a laumontite which has lost part of its water (to one molecule), and the same is probably true of caprochianite. Schneidtrite is laumontite from the serpentine of Monte Cotini, Italy, which has undergone alteration through the action of raagnesia solutions.

Pyr., etc.—B.B. swells up and fuses at S-6-8 to a white enamel. Gelatinizes with hydrochloric acid.

Obs.—Occurs in the cavities of basalt and similar eruptive rocks; also in porphyry and syenite, and occasionally in veins traversing clay slate with calcite.

Its principal localities are the FArOer; Disko in Oreenlaod; In Bohemia, at Eule In clay slate; bt. Qothard in Switzerland; the Fasgalhal: the Ellpatrick hills, near Glasgow; the Hebrides, nnd the north of Ireland. In India, in tbe Deccan trap area, at Poona, etc.

Peter^B Point, Nova Scotia, affords flne specimens of ttils species. Found at I^Ippsburg, Maine. Abundant in many places in the copper veins of I^ke 8ii|ierior In trap, and on I. Royale; on north shore of Lake Superior, between Pigeon Br j and Fond du Lac. Found also at Bergen Hill, N. J.; at the Tilly Foster iron mine, Brewster, N. Y.

I^rabanlta. Ca,Al,Si,o,, + 6H,o. Resembles sitlbttc. H. = 4'6-5. jG. = S-38. Color aaow-while. Occurs upon phillipslte in basalt at Lnuban, Slleala.

.dbyGoogle

Chabazlte Group. Rhombohedral. rr' i

Ch&buite {Ca,Na,)Al,Si,O,, + 6H,o, pt 85" 14' 1-0860 Gmeliaite (Na,Ca)Al,Si,o,, + 6H,o 68° 8' 0-7346 orj(!= 1-1017

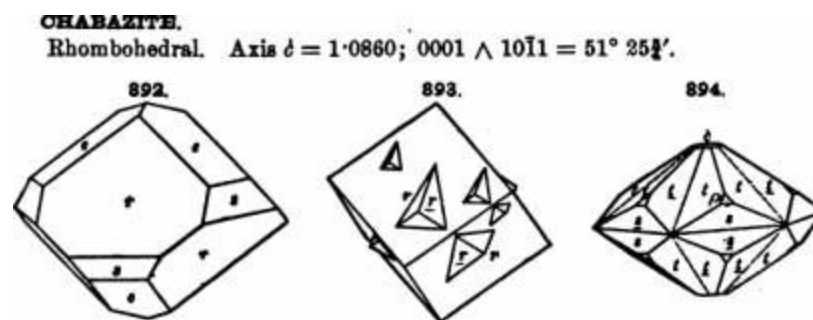
XflTjnite CaAl,Si,o,. + 5H,o 73° 56' 08357 ^ = 1-1143

The Ghabazite G-roap inclndeB these three rhombohedral apiecie. The fundamental rhombohedrong have different angles, but, aa ehova in the axial ratios above, they are closely related, since, taking the rhomboheijron of Chabazlte as the basis, that of Omelinite has the symbol {(30S3) and of Levynitfl J(3034).

The Tarfatlon la composltton often observed in the flnt two specie* baa led to lb« rather plausible hypotliesls that they ftre to be vtewed u isomorphous mixtures of the feldapar-llke compouDds

(Ca,Na,)Al,81.o. + 4H,o. (Ca,Na,)Al,Si,o,. + 8H.o.

1-0860; 0001 A loll = 51° SBf.



Twins: {1} tw. axis i, penetration-twins common. (2) Tw. pi. r, contact-twins, rare. Form commonly the simple rhombohedron varying little in angle from a cube {rr' = 85° 14'}; also r and e (oll2), {«' = 54° 47'}. Also in complex twins (Fig. 348, p. 118). Also amorphous.

Cleavage: rather distinct. Fracture uneven. Brittle. H. = 4-5, G. = 2-08-2-16. Luster vitreous.. Color white, flesh-red; streak uncolored. Transparent to translucent. Optically —; also -f- (Andreasberg, also hayden-ite). Birefringence low. The interference-figure usually confused; sometimes distinctly biaxial; basal sections then divided into sharply defined sectors with . different optical orientation. These anomalous optical characters probably secondary and chiefly conditioned by the variation in the amount of water present. Mean refractive index 1-G.

Var.—1. Ordinary. The most common form is the fundamental rhombohedron, in which the angle is so near 90° that the crystals were at first mistaken for cubes. Aegialite, from Nova Scotia (Acadia in the French of last century), is a reddish chabazite; sometimes nearly colorless. Uaydnite is a yellowish variety in small crystals from Jones's Falls, near Baltimore, Md. 2. Phaeolite is a colorless variety occurring in twins of hexagonal form (Fig. 11Q-1). Heulandite (whence the name, from *heuland*, a bean); the original was from Leipsic in Bohemia, here belongs also hercynite (seebachite) from Richmond, Virginia; the composite twins of great variety and beauty. Probably also the original burschelite from Sicily. It occurs in flat, almost tabular, hexagonal prisms with rounded terminations divided into zones.

.dbyGoogle

Comp.—Somewhat uncertain, since a rather wide variation is often noted even among specimens from the same locality. The ratio of (Ca, Na, K) : Al is nearly constant (= 1 : 1), but of Al : Si varies from 1 : 3 to 1 : 5; the water also increases with the increase in silica. The composition usually corresponds to $(Ca, Na, K)Al_2Si_2O_{10} \cdot 6H_2O$, which, if calcium alone is present, requires: Silica 47.4, alumina 20.2, lime 11.2, water 21.3 = 100. If Ca : Na = 1 : 1, the percentage composition is: Silica 47.2, alumina 20.0, lime 5.5, water 21.3 = 100.

Polymorphs present in small amount, also, sometimes. barite and strontianite. The supposed facts must be established by the hypothesis that the members of the group are isomorphous mixtures analogous to the feldspars, as noted on p. 453.

Fusion, etc.—B.B. intumesces and fuses to a blebby glass, nearly opaque. Decomposed by hydrochloric acid, with separation of slimy silica.

Diagnosis—Characterized by rhombohedral form (resembling a cube). It is harder than gypsum and does not effervesce with acid; unlike calcite and fluorite in cleavage; fuses B.B. with intumescence unlike strontianite.

Obs.—Occurs mostly in basic rocks and occasionally in gneiss, syenite, mica schist, hornblende schist. Occurs in the Fichtelberg, Greenland, and Iceland, associated with chlorite and leucite; at AUSAIG in Bohemia: at Oltersteln. with hornblende; at Annerod, near Qlessen: at the Giant's Causeway, Antrim, Renfrewshire; Isle of Skye. etc.

In the U. S., in veins at Somerville, Mass.; at Bergen Hill, N. J., in small crystals; at Jones's Falls near Baltimore (Lajdett). In Nova Scotia, in yellow or flesh-red (the form of the aidaite), associated with heulandite, analcite and calcite, at Five Islands, Swan's Creek, Digby Neck, etc.

The name Ciba from the name of a stone.

OMBUNIT.

Rhombohedral. $a = 0.7345$.

Crystals usually hexagonal Mt. 89S.

in aspect; sometimes p (Sill) smaller than r (Ioll), and habit rhombohedral; $rr' = 68^\circ 8'$, $rp = 37^\circ 44'$.

Cleavage: m easy; c sometimes distinct. Fracture uneven. Brittle. $H. = 4-5$. $G. = 3.04-2.17$. Luster vitreous. Colorless, yellowish white, greenish white, reddish white, flesh-red. Transparent to translucent. Optically positive, Cyprus, also negative, Andreasberg, the Vicentine, and Glenarm, N. Scotia. Birefringence very low. Interference figure often disturbed, and basal sections divided optically into section analogous to chabazite.

Comp.—In part $(Na, Cft)Al, Si, O, 4H_2O$, If sodium alone is present this requires: Silica 46.9, alumina 19.9, soda 12.1, water 21.1 = 100. See also p. 458.

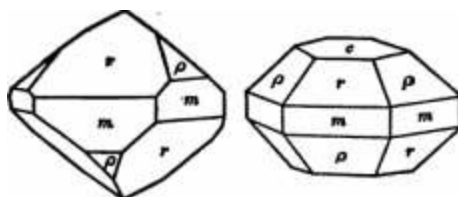
Pyro., etc.—B.B. fuses easily ($F. = 3.5-8$) to a white enamel. Decomposed by hydrochloric acid with effervescence of silica.

Obs.—Occurs in flesh-red crystals in amygdaloidal rocks at Montecchio Maggiore; at Andreasberg; in Transylvania; Antrim, Ireland; Tallisker in Skye, in large colorless crystals.

In the United States in fine white crystals at Bergen Hill, N. J. At Cape Blomidon, Nova Scotia (siderite): also at Two Islands and Five Islands.

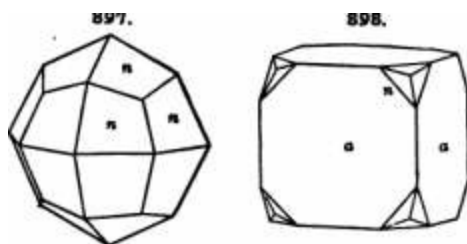
Named in honor of Prof. Ch. Gmelin of Tübingen (1791-1860).

Formula. $CaAl_2Si_2O_8 \cdot 5H_2O$, in rhombohedral crystals. $H. = 4-4-5$. $G. = 3.04-2.16$. Colorless, white, grayish, reddish, yellowish. Found at Glenarm and at bland



DESCRIPTITE HINEBALOaT.

Golden, Colorado.



ANAIXHTEI. Aosltime.

isometric. Usually in trapezohedron; also cubes with faces n (311); again the cubic faces replaced by a rhombic trisoctahedron. Sometimes in composite groups about a single crystal as nucleus (Fig. 351, p. 119). Also massive granular; compact with concentric structure.

Fracture: cubic, in traces. Fracture subconchoidal. Brittle. $H. = 5-5-5$. $G. = 2-22-2-29$; 2-28 Thomson. Luster vitreous. Colorless, white; occasionally grayish, greenish, yellowish, or reddish white. Transparent to nearly opaque. Often shows weak double refraction, which is apparently connected with loss of water and consequent change in molecular structure (Art. 411). $d = 1.4874$.

Comp.— $NaAl_3Si_3O_{10} \cdot 2H_2O = Na_2O \cdot Al_2O_3 \cdot 4SiO_2 \cdot 2H_2O =$ Silica 54.5, alumina 23.2, soda 14.1, water 8.2 = 100.

Pyro., sto.—Yields water in the closed tube. B.B. fuses at S.G. to a colorless glass. Fused with hydrochloric acid.

Diff.—Characterized by trapezohedral form, but is softer than garnet, and yields water B.B., unlike leucite (which is also infusible): fuses without intumescence to a clear glass unlike chabasite.

Recognized in thin sections by its very low relief and isotropic character; often shows optical anomalies. From leucite and sodalite surely distinguished only by chemical tests, i.e., absence of chlorine in the nitric-acid test (see sodalite, p. 444), absence of much potash and abundance of soda in the solution, and evolution of much water from the powder in a closed glass tube heated to a red heat.

Obs.—Occurs frequently with other zeolites, also nepheline, calcite, etc., in cavities and seams in basic igneous rocks, as basalt, diabase, etc; also in granite, gneiss, etc. Recently shown to be also a rather widespread component of the groundmass of various basic igneous rocks, at times being the only alkali-alumina silicate present, as in the so-called analcrite-basalts. Has been held in such cases to be a primary mineral produced by the

crystallization of a magma containing considerable soda and water vapor held under

The Cyclopean Islands, near Catania, Sicily, afford pellucid crystals; also the Fassathl in Tyrol; other localities are, in Scotland, In the Kilpatrick Hills; Co. Antrim, etc, in Ireland; the F&R&er; Iceland; near Aussig, Bohemia; at Arendal, Norway, in beds of iron ore; at Andreasberg, in the Harz. in silver mines.

In the U. S., occurs at Bergen Hill. New Jersey; in gneiss near Tonawanda. Westchester Co., N. Y.; abundant in fine crystals, with prehnite, talc, and calcite. In the Lake Superior region; at Table Mt. near Golden, Colorado, with other zeolites. Nova Scotia affords fine specimens.

The name Arutletite is from ayaXkii, weak, and alludes to its weak electric power when heated or rubbed. The correct derivative is natrolite, as here adopted for the species.

Faigasite. Perhaps $\text{H}_2\text{Na}_2\text{CaAl}_2\text{Si}_2\text{O}_{10} \cdot \text{H}_2\text{O}$.

In isometric octahedrons. H. = 5. G. = 1883. Colorless, white. Occurs with angle in the limburgite of Salsburg. In the Ealserstube, Baden, etc.

Edingtonite*. Perhaps $\text{BaAl}_2\text{Si}_2\text{O}_{10} \cdot 8\text{H}_2\text{O}$. Crystal a pyramidal in habit (tetragonal-sphenoidal); also massive H. = 4-45. G = 2.4. White, grayish white, pink. Occurs in the Ellpatrick Hills, near Glasgow, Scotland, with harmotome.

;.Coogite

Hftndite

Natrolite Group. Orthorhombic and Monoclinic. $\text{Na}_2\text{Al}_2\text{Si}_2\text{O}_{10} \cdot 2\text{H}_2\text{O}$ 0-9785 :

Soolooite Xewite

4 il-i p

0-9764 : 1 : 0.3434 89° 18'

$\text{Ca}(\text{Al}_2\text{OH})_2\{\text{SiO}_2\}_2 + 2\text{H}_2\text{O}$

$\text{Na}_2\text{Al}_2\text{Si}_2\text{O}_{10} \cdot 2\text{H}_2\text{O}$

$2[\text{CaAl}_2\text{Si}_2\text{O}_{10} \cdot 3\text{H}_2\text{O}]$ The three species of the "Natrolite Group" agree closely in angle, though very different in crystal habit: Natrolite is orthorhombic usually, arutletite is monoclinic, scudcote is monoclinic, perilite is monoclinic, and mesolite seems to be both monoclinic and triclinic. Fibrous, radial or divergent groups are common to all the species.

The Matrolitu Gniip iiicludeg ilie Bodiiim stltcate, NatroUle. with ibe empirical rormula Na,At,Sl>O,g.2H,Oi the calcium eilicaie, Scoleclie, CaA1,SI,o,,,3H,Oi sIbo Hecollta intermediate between tlieae anil coTTcspouding to $j \wedge Qj \wedge i \text{ g} j \text{ o } 8H \text{ O} \blacksquare$

= 0-9785 : 1 : 0-3536.

NATROUTE. Nadelieolitb Qtrm.

Orthorhombic. • Axes a xh -A

mm^{'''}. 110 A no = V»° «!' 8" «».

mo', no A 111 = 63° 11'. «/. Ill A ill = 37° 88'. on^{'''}. 111 A iil = 30° 471'.

Cryatnla prismatic, usually very slender to acicular; frequently divergent, or in stellate group. Also fibrou8,radiating, massive, granular, or compact.

Cleavage: m perfect; b imperfect, perhaps only a plane of parting. Fracture nneven. H. = 6-5-5. G. = 2 20-225. Luster vitreous, sometimes inclining to pearly, especially in flbroua varieties. Color white, or colorless; to grayish, yellowish, reddish to red. Transparent to translncent. Optically+. Ai. pi.] 5. Bi J, c. 3E, = 93" 28'. /?,= 1-4797; y-a = 0012.

V«T.— Ordiaary. Commonly citber (a) in groups of Blender colorless prlBmatic crystals, varyini; but little In angle froni square prisma, often actcular, or (A) In Qbrousdiregent or radiated massea. vitreous In luster, or butslfgbtly pearly (these radiated forms often Tesemble those of tbomsotilte and pectolile); often also (is) solid amygdules, usually radiated flbrouB, Ktid aomewbiit silky In luster within; (d) rarely compact massive. OaiaetiU is ordinarily natrollte. In colorless needles from southern Scotland.

BergmanniU. tprtutUin, brmUite, are names wblcb have been given to the natrollte from the augite-syenlte of southern Norwar, on the Langesund Bord. in the "Brevik" region, nhcre it occurs Bhrous. massive, and in long prismatic crystallizations, and from white to red In color. Derived Iti part from elxollle, In part frum aodalle. Iron-natrolile la a dark green opaque variety, either crystalline or smoi-pbous, from the Brevik region; the iron ia due to iriclusiins.

Comp.—Nft,Al,Si.O,-f 2H,o or Na,O.Al,o,.3SiO, + 2H,o = Silica 474, alnmina 26-8. Na,o 16-3, water 9-5 = 100.

Pyr., etc.—In the closed lube whliens and becomes opaque. B.B, fuses qule'ly at 2 to a colurleM glass. Fusible In the flame of an ordinary steartneur wax candle. GelaliniKes with acids.

Diff — Distingulsbed from aragonile and peclolite by Its easy fuslblllly and gelatinIzation with acid.

Obs.—Occurs in cavities in amygdaloidal basalt, and other related igneous rocks:

•Iran

9 cases the crystals seem to be monoclinic.

^Aiolc

BomellmeB fn Eeums In greoite, gDClea, and avenile. Found Al Ansaig and Teplltz la Bohe-miit; ill fine L-rystiila In Auvergiie; Fussatliial, Tyrol; KapLiik, Huugury. In red amygdules ferocalite) iu miyygtiHliiId of Ireland, Bcutlaud uiid Tyrol; tbe amygdaloid o' BLshopton {salactit«)\ at Glen Farg l/argiie) iti Fifesliire. Common in (he Bugite-sjenite of &.e Laugc-euBd Dixd, neiirBrevIk, guilhein Norwiyy.

Ill Norb Americii, in llh trupof Novu Bcolia; at Bergen Hill, N. J.; at Copper Fulb, Lake Superior.

Named Metoiypt by HaOy, from HtVoS, middle, and rvteoi, type, because the form of the cryBial—iiiltis viev n square piiiai—was intermediate between the forma of atllblte aud auHlcite. NatroliU, of KlaprolU, la from natron, toda; It alludes to tlie presence of soda, - wheuce also the name waa-me»oiYpt, iu contrast with scolecite, oi Unw-mMolj'jM.

SOOLEOITB.

Monoclinic. Axes $h : a = 0.9764 : 1 : 0.3434$; $\beta = 89^\circ 18'$

Crystals slender prismatic ($mm = 88' 3''$), twins Bhowing a featherlike stmtion on &, diverging upward; also as penetration-twins. Crystals iu diverfrent groups. Also massiTe, fibrous and radiated, and in nodules.

Cleavage: m nearly perfect. H. = 5-5'5. G. = 216-'J4. Lustervitreous, or silky wHen fibrous. Transparent to sub translucent. Optically —. Ax. pi. and Bx., J, ft. Bi, a ($\wedge = 15M6^\circ$. $2H_{\parallel} = 33'' 26'$. $\rho = 1.4953$.

Comp.— $CuAl_2Si_2O_{10} \cdot 3H_2O$ or $CaO \cdot Al_2O_3 \cdot 3SiO_2 \cdot 3H_2O$ - Silica 45.9, alumina 36.0, lime 14.3, water 13.8 = 100.

Pyr.. Bto. —B,B. someltmeg curls up like a worm (whence the name from okh^Xii^, a vorm. nblcb gives tuMdU. and not icoleiile or toUtiU); otlier TiirlencK JntumeEce but sligblly, and ail fuse at 3-3 3 to a white btebby euamei. Gelatinizes with adds like nBtroitte.

Obi.—Occurs in the Beruflord. Iceland; in amygdaloid at Slaffa; in Skye. lit TuliakN: near Fisenni'h, Saxony; in Auvergne; common in fine crystalIfzatlong iu the Decenn trap area, in British India. In tbe United Stales, in Colorado at Table Mountain near Oolden In

cavities in basalt. In Canada, at Black Lake, Hégoulic Co., Quebec

Mesolite. Intermediate between iatrolite and scolecite (see p. 453). In acicular and capillary crystals: delicate divergent tufts, etc. G. = 8.39. White or colorless. In amygdaloidal basalt at numerous points.

THOUBONITE.

Orthorhombic. Axes $S : i : i = 0.9932 : 1 : 1.0066$.

Distinct crystals rare; in prisms, $mm = 89^\circ 37'$. Commonly columnar, structure radiated; in radiated spherical concretions; also closely compact.

Cleavage: i perfect; n less so; c in traces. Fracture uneven to subconchoidal. Brittle. H. — 5-5-5. Q. = 2-3-2-4. Luster vitreous, more or less pearly. Snow-white; reddish, green; impure varieties brown. Streak uncolored. Transparent to translucent. Pyroelectric. Optically -f. Ax. $pi \perp c$. BxJ, 6. Dispersion $p > v$ strong. $2E = 83$. $\mu^i = 1.503$.

Var.—1. Ordinary. (a) In regular crystals, usually more or less rectangular in outline, prismatic in habit. (b) Prisms sometimes often vesicular or radiated, (c) Radiated fibrous, (d) Spherical concretions, consisting of radiated fibers or slender crystals. Usually massive, granular to impalpable, sometimes whitish to reddish brown, sometimes (often) greenish in thin sections. The spherical massive forms also radiated with several centers and of varying colors, hence of much beauty when polished. Opatka is a white massive thomsonite from Arkansas.

Comp.— $\{Na, Cs\}Al_8Si_6O_{46} + nH_2O$ or $(Na, Ca)O \cdot Al_2O_3 \cdot 2SiO_2 \cdot 5H_2O$. The ratio of Na : Ca varies from 3 : 1 to 1 : 1. If Ca : Na = 3 : 1 the percentage composition requires: SiO₂ 37.0, Al₂O₃ 31.4, CaO 12.9, Na₂O 4.8, H₂O 13.9 = 100.

Pyro. etc.—B.B. fuses with intumescence at 2 to a white enamel. Gelatinizes with hydrochloric acid.

Diff.—Resembles some natrolite, but fuses to an opaque not to a clear glass.

^~.0066

Obs.—Found in cavities in lava in amygdaloidal igneous rocks, sometimes white as a result of alteration. Occurs near Eilpairick, Scotland; in the lavas of Somma (comptonite); in basalt at the Pfafferskaute in Saale Weimar; in Bohemia, in the Cyclopean Islands, Sicily; near Brevik, Norway; the Pfafferskaute; Iceland (carpholite, straw-yellow): at Mt. Monzon, Fölsch.

Occurs at Peter's Point, Nova Scotia. In the XI. 8. at Magnet Cove (Oparka) in the Ozark Hills, Arkansas; in the amygdaloid of Grand Uraire, L. Superior, which yields the water-worn pebbles resembling agate. In part green (tinonite); in the basalt of Table Mt. near

Oolden, Colo,

Hydion^{hcUU}. HNa,Al.SitO₃+SH₂O. Hasdve, rodUted. H. = 4-6-0. a = 2-268. Color white; also dark gray. From Utcbfield, Maine, BaitiU from the Lsngmund flord, Norway, is limllar.

il. Hlca Division.

The species embraced onder this DiTiBioiii fall into three gronps: 1, the Mica Group, including the Micas proper; 2, the Clintonitb Gbodp, or the Brittle Micas; 3, the Chloeitb Group. Supplementary to these are the VermiculiteB, hydrated compounds chieSy resulte of the alteration of some one of the micas.

All of the above species have the characteristic micaceous structure, that is, they have highly perfect basal cleavage and ^ield easily thin lamiee. They belong to the monoclinic system, but the position of the bisectrix in general deviates but little from the normal to the plane of cleavage; all of them show on the basal section plane angles of 90° or 130°, marking the rela-tive position of the chief zones of forms present, and giving them the appearance of hex-agonal or rhombohedrai symmetry; further, they are more or less closely related among themselves in the angles of prominent forms.

The species of this Division all yield water upon ignition, the micas mostly from 4 to 5 p. c, the chlorites from 10 to 13 p. c; this is probably to be regarded in all cases as water of constitution, and hence they are not properly hydrous silicates.

More or less closely related to these species are those of the Serpentine and Talc Division and the Kaolin Division following, many of which show distinctly a mica-like structure and cleavage and also pseudo-hexagonal symmetry.

1. Mica Qronp. Monoclinic, MuMOTlte Potassium Mica H,KAl,(SiO₂),

ii:i:i = 0-57735 ; 1 : 3-3138 fi = 89° 54' Paragonite Sodium Mica H,KaAl,(SiO₂),

LepidoUte Lithium Mica KLi[Al(OH,F),]Al(SiO₂), pt.

Zinnwaldite Litbiuro-iron Mica

Biotite Magnesium-iron Mica (H,K),{Mg,Fe),(Al,Fe),(SiO₂), pt.

&:\$:i = 0 57735 ; 1 : 3-2743 >3 = 90' 0' Phlogopite (H,K,{MgP}),Mg,Al{SiO₂),

Magnesium Mica; usnally containing fluorine, nearly free from iron, lepldomelane Annite

Iron Mioaa. Contain ferric iron in large amount. The species of the Mica Qrottp crystallize in the monoclinic system, but with a close approximation to either rhombohedrai or

orthorhombic symmetry; the plane angles of the base are in all cases 60° or 130° . They are all charac-

DESCRIPTIVE HTNBALOOT.

terised by highly perfect basal cleavage, yielding very thin, tough, and more or less elastic laminae. The negative bisectrix, a , is very nearly normal to the basal plane, varying at most but a few degrees from this; hence a cleavage plate shows the axial interference-figure, which for the pseudo-rhombohedral kinds is often uniaxial or nearly uniaxial. Of the species named above, biotite has usually a very small axial angle, and is often sensibly uniaxial; the axial angle of phlogopite is also small, usually 10° to 12° ; for muscovite, paragonite, lepidolite the angle is large, in air commonly from 50° to 70° .

The Micas may be referred to the same fundamental axial ratio with an angle of obliquity differing but little from 90° ; they show to a considerable extent the same forms, and their isomorphism is further indicated by their not infrequent intercrystallization in parallel position, as biotite with muscovite, lepidolite with muscovite, etc.

A blow with a somewhat dull-pointed instrument on a cleavage plate of mica develops in all the species a six-rayed percussion-figure (Fig. 901. also Fig. 477, p. 149), two lines of which are nearly parallel to the prismatic g_1 edges; the third, which is the most strongly character-

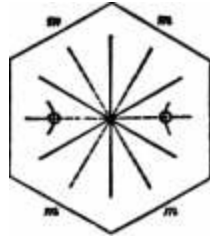
ized (Leitstrahl term.), is parallel to the clinopinacoid or plane of symmetry. The micas are often divided into two classes, according to the position of the plane of the optic axes. In the first class belong those kinds for which the optic axial plane is normal to b (010), the plane of symmetry (Fig. 901); in the second class the axial plane is parallel to the plane of symmetry. The percussion-figure serves to fix the crystallographic orientation when crystalline faces are wanting. A second series of lines at right angles to those mentioned may be more or less distinctly developed by pressure of a dull point on an elastic surface, forming the so-called pressure-figure; this is sometimes six-rayed, more often shows three branches only, and sometimes only two are developed. In Fig. 901 the position of the pressure-figure is indicated by the dotted lines. These lines are connected with gliding-planes inclined some 67° to the plane of cleavage (see beyond).

The micas of the first class include: Muscovite, paragonite, lepidolite, also some rare varieties of biotite called anomite.

The second class embraces: Zinnwaldite and most biotite, including lepidomelaene and phlogopite.

Chemically considered, the micas are silicates, and in most cases orthosilicates, of aluminium with potassium and hydrogen, also often magnesium, ferrous iron, and in certain cases ferric iron, sodium, lithium (rarely rubidium and cesium); further, rarely,

barium, manganese, chromium. Fluorine is prominent in some species, and titanium is also sometimes present. Other elements (boron, etc.) may be present in traces. All micas yield water upon ignition in consequence of the hydrogen (or hydroxyl) which they contain.

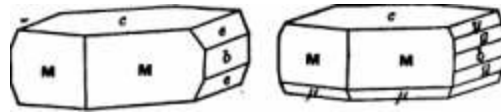


B9U800VITB. Commoo Hlca. Potash Mien. EaKglimmer ffortn.

Monoclinic. Axes $d:i:i = 057735 ; 1 : 3-3128; j8 = 89^\circ 54'$.

Twins common according to the mica-lato: tw. plane a plane in the zone cV normal to c , the crystals often united by c . Crystals rhombic or hexagonal in outline with plane angles of 60° or 120° . Habit tabular, passing into tapering forms with planes more or less rough and strongly striated horizontally;

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Ticinal forms common. Folia often very small and integrated in stellate, plumose, or globular forma; or in scales, and scalj maasive; also crypto-crjstatline aud coihpact massive.

Cleavage: basal, eminent. Also planes of secondary cleavage as shown in the percnsioQ-figure (see pp. 464 and 149); natural plates hence often yield

$0M. 001 A 231 = M^\circ 8(r. 6^\wedge. 001 A ill = 81' 80''$. $MM'. 221 A aSl = W 48^*$. $ntt \setminus ill A ill = 69^\circ 1 \ll J'$.

narrow strips or thin fibers | axis h , and less distinct in directions inclined 60° to this. Thin laminie fiezible and elastic when bent, very tough, harsh to the touch, passing into kinds which are less elastic and nave a more or less unctuous or talc-like feel. Etching-figures on c monodinic in symmetry (Fig. 481, p. 150).

II. = 'i-i-b. G. = S'76-3. Luster vitreous to more or less pearly or silky. Colorless, gray, brown, hair-brown, pale green, and violet, yellow, dar£ olive-green, rarely tose-red. Streak uncolored. Transparent to translucent.

Fleochroism usually feeble; distinct in some deep-colored varieties (see beyond).

Absorption in the direction normal to the cleavage plane (vibrations 1 b, c) at right angles, much more so than transversely (vibrations | a); hence a crystal unless thin is nearly or quite opaque in the first direction though translucent through the prism. Optically —. α , β , γ , δ and nearly \perp to c . β , ($= \gamma$) inclined about -1° (behind) to a normal to c . Dispersion $p > v$. Axial angle variable, usually about 70° , but diminishing to 50° in kinds (phengite) relatively high in silica. Birefringence rather high, $\gamma - \alpha = 0.033$; $\gamma = 1.5941$.

V*1.—1. Ordinary Xvioletite. In crystals of above described, often tabular, thin, tapering with vertical faces rough and striated; the basal plane often rough unless as developed by cleavage. More commonly in plates without distinct outline, except as developed by pressure (see above); the plates sometimes very large, but passing into fine scales. Arranged in plumose or other forms. In normal muscovite the thin laminae spring back with force when bent, the scales are more or less harsh to the touch, unless very small, and a pearly luster is seldom prominent.

2, Damourite. Including margarodite, graphite, hydrothermal, and most hydrothermal mica in general. Foliation less elastic; luster somewhat pearly or silky and feel unctuous like talc. The scales are usually small and it passes into forms which are flake scaly or fibrous, spongy, and finally into the compact crypto-crystalline kinds called oneonite, including much phengite. Axial angle for damourite varies from 60° to 70° . Often derived by alteration of cyanite, topaz, corundum, etc. Although often spoken of as hydrothermal micas, it does not appear that damourite and the allied varieties necessarily contain more water than ordinary muscovite; they may, however, give it off more readily.

Margarodite, as originally named, was the talc-like mica of Mt. Qrener in the Zillerthal; granular to acaly in structure, luster pearly, color grayish white. It occurs in whitish, silky forms from the tin mine of Bt. Austell, Cornwall. Berthelinite is a fine scaly muscovite united in fibrous aggregates and characterized by its silky luster (hence the name from *berthel*, its discoverer).

Comp., Tar.—For the most part an orthosilicate of aluminum and potassium $(H,K)AlSi_3O_{10}$. If, as in the common kinds, $H : K = 1 : 1$, this becomes $H_2KAl_2(Si_3O_{10})_2 \cdot 3H_2O$, = Silica 45.3, alumina 38.5, potash 11.8, water 45 = 100.

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Some kinds contain a larger amount of silica (47 to 48 p. c.) and correspond to a normal orthosilicate, and they have been called phengite. As shown by Clarke, these acid muscovites can be more simply regarded as molecular mixtures of $H_2KAl_2(Si_3O_{10})_2$ and $B_2KAl_2(Si_3O_{10})_2$.

Iron is usually present in small amount only. Bismuth is rarely present, as in the variety. Chromium is also present in the variety from Bismuthstein, Zillertal, and elsewhere.

Fyr., etc.—In the closed tube gives water, which with Brazil-wood or tea reacts for fluorine. B.B. whitens and fuses on the thin edges (F. = 67, ▼. Kobell) to a gray or yellow glass. With fluze* gives reaction for Iron and sometimes manganese, rarely chromium. Not decomposed by acids. Decomposed on fusion with alkaline carbonates.

Diff.—Distinguished in normal kinds from all but the species of this division by its perfect basal cleavage and micaceous structure; the pale color separates it from kaolinite; the laminae are more flexible and denser than those of phlogopite and still more than those of the brittle micas and the chlorites.

In thin sections recognized by want of color and by the perfect cleavage shown by fine lines (as in Fig. 907, p. 44) In sections 1 e. In a direction parallel to e. By reflected light under the microscope the same sections show a peculiar mottled surface with a net-like luster; birefringence rather high, hence interference-colors bright.

Occ.—Muscovite is the most common of the micas. It is an essential constituent of mica schist and related rocks, and is a prominent component of certain common varieties of granite and gneiss; also found at times in fragmental rocks and limestones; also volcanic tuffs. It is rare and appears only as a secondary product. The largest and best developed crystals occur in the pegmatite dikes associated with granitic intrusions, either directly cutting the granite or in its vicinity. Often in such occurrences in enormous plates from which the mica or "sheet mica" of commerce is obtained. It is then often associated with crystallized orthoclase, quartz, albite; also apatite, tourmaline, garnet, beryl, columbite, etc., and other mineral species characteristic of granitic veins. Further, muscovite often encloses flattened crystals of garnet, tourmaline, also quartz in thin plates between the sheets; further not infrequently magnetite in dendrite-like forms following in part the directions of the cleavage-figure.

Some of the best known localities, are: Abul in the Sulzbachthal, with adularia; Rohenkopf in the Zillertal, Tyrol; Soboth, Styria; Bt. Oolhard, Binnenthal, and elsewhere in Switzerland; Mourne Mts., Ireland; Cornwall; Uppsala, Falun, Sweden; Skutterud, Norway. Obtained in large plates from Greenland and the East Indies.

In Italy, at Mount Mica in the town of Paris; at Buckfield, in fine crystals. In If. Bamp., at Acworth. Oraftou. In Md., at Chesterfield; South Royalton; at Goshen, rose* red. In Conn., at Honrore; at Litchfield, with cyanite; at the Middletown feldspar quarry; at Haddam; not abundant with albite. etc.: New Milford. In N. York, near Warwick; Edenville; Edward. In Penn., at Pottsville, Chester Co.; at Unionville, Delaware Co., in the Middletown. In Maryland, at Jones's Falls. Baltimore. In VA, at Amelia Courthouse. In No. Carolina, extensively mined at many places in the western part of the state; the chief mines are in Hittell, Yancey, Jackson and Macon Cos. The mica mines have also yielded many rare species, as columbite, samarskite, hatchettolite, uraninite, etc.; in good quantities in Alexander Co. In S. Carolina, there are also muscovite deposits; also in Georgia and Alabama.

Muscovite mines have also been worked to some extent in the Black Hills, South Dakota: In Washington, at Lockford, Spokane Co.: in Colorado.

Muscovite is named from *Muscovium* or *Muscovy-glatt*, formerly a popular name of the mineral.

Finite. A general term used to include a large number of alteration-products especially of biotite, also spodumene, nephelite, acapule. It is similar to other minerals. In composition essentially a hydrous silicate of aluminum and potassium corresponding more nearly to muscovite, of which it is probably to be regarded as a massive, compact variety, usually very impure from the admixture of clay and other substances. Characters as follows: Amorphous; granular to cryptocrystalline. Rarely a fibrous cleavage. H = 2-5. G = 3-6. Luster feeble, waxy. Color grayish white, grayish green, pea-green, dull green, brownish. Translucent to opaque. The following are some of the minerals classed as mica: *giant mica*, *giant mica* (see p. 85), *giant mica*, *giant mica*, *giant mica*. Total. *polyargyrite*, *giant mica*.

Agalmatolite (p) like ordinary mass. It is an amorphous compact variety, luster, and other physical characters, but contains more silica, which may be from free quartz or feldspar as impurity. The Chinese has H = 2-2.5; G = 2-7.5-3-8.5. Colors usually grayish, greenish green, brownish, yellowish. Named from *ajaxia*, *animag*;

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Pogonite is from pagoda, the Chinese corundum the stone into miniature pagodas. It is a part of the so-called *sfumatolite* of China. It is in composition, a part of compact *pyrobylute*, and still another *steatite* (see these species).

Pangolite, A sodium mica, corresponding to muscovite. Its composition; formula, $H.KsAl_2(SiO_3)_2$. It is a fine pearly mica; also compact. Q = 78-200. Color yellowish, grayish green; translucent. It is found at *Moule Cumplone* near *Faldo* in *Oanton Tesain*, [Bwllzer], containing *cyauite* and *staurolite*; called *paragonite-schist*. Occurs associated with *tourmaline* and *corundum* at *Unioarille*, *Delaware Co., Pa.*

LiBPIDOUm. Lithia. *Lithionite* *Otrm*.

In aggregates of short prisms, often with rounded terminal faces. Crystals sometimes twins or triplets according to the mica law. Also in cleavable plates, but commonly massive scaly-granular, coarse or fine.

Cleavage: basal, highly eminent. H = 2.5-4. G = 3.8-2.9. Luster pearly. Color rose-red, violet-gray or lilac, yellowish, grayish white, white. Translucent. Optically —, Ax. parallel to c. J. 6; rarely b. Bx, (o) inclined $V = 47^\circ$ red, and 133° yellow (Na) to normal to c. Axial angle large, from $50^\circ-72^\circ$, $\mu = 1.5975$.

Coup.—In part a metasilicate, $R, Al(SiO)$, or $KLi[Al(OH,F), Al(8iO),]$. The ratio of fluorine and hydroxyl is variable. The following are analyses (HiggB):

SIO. $Al_1,0, Fe,0. FeO HoO K.O Ll,0 Na,0 H,0 F Paris, SO-fti S499 0 80 0-2S (r. ll-»8 420 2-11 196 6-28=103-88 Hebron, 48-80 38-30 0'3» 009 008 13-81^* 4'4S 074 1-78 4-96$

[CaO.MgO 017 = 101-86

■with 0-17 (Rb,C»),0.

Pyr., etc.—In the closed tube gives water and reaction for fluorine. B.B. fuses with Intumescence at 2-26 to a white or grayish glass, sometimes magnetic, coloring the flame purplish red at the mouth of the tube. With the fluxes some varieties give reactions for iron and manganese. Attacked but not completely decomposed by acids. After fusion, gelatinizes with hydrochloric acid.

Occ.—Occurs in granite and gneiss, especially in granitic veins: often associated with tourmaline; also with amblygonite, spodumene, casiterite, etc.; sometimes associated with muscovite in parallel position.

Found near Ulo in Sweden; Peabody, Saxony; Rozena (or Rozna). Moravia, etc. In the United States, common in the western part of Maine, in Hebron, Auburn. Purl, etc; at Cheshirefield, Mass.; Middleboro, Conn.: with niobite near San Diego, California.

Named lepidolite from Aeffel, after the earlier German name Sdiupptstein, alluding to the scaly structure of the massive variety of Rozena.

LEPIDOLITE is a micaceous mineral occurring in rounded aggregates on tourmaline, also with lepidolite, tourmaline, etc., at Hebron, Me. Composition $Li[Al(OH), (8iO),]$.

Zinnwaldite. An iron-lithium mica in form near hiotite. Color pale violet, yellow to brown and dark gray. Occurs in Zinnwald and Altenberg; similarly in Cornwall.

Criophyllite is a lithium mica from Rockport, Mass, Polykitkimitite is a lithium mica from Kaogerdluarsuk, Greenland.

BIOTITE,

Monoclinic; pseudo-rhombohedral. Axes $a:b:c = 1.057735 : 1 : 3.3743$; $\beta = 90^\circ$.

Habit tabular or short prismatic; the pyramidal faces often repeated in oscillatory combination. Crystals often apparently rhombohedral in symmetry since $r(110)$ and $z(110)$, which are inclined to c at sensibly the same angle, often occur together; further, the zones to which these faces belong are inclined 130° to each other, hence the hexagonal outline of basal sections. Twins, according to the mica law, tw. pi. a plane in the

prismatic

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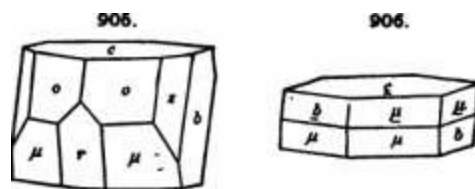
xoneJ_c. Often in disMminated scales, sometimeB in massive aggregations ot cleavable scales.

ec, 001 A US = 78" I'. w. 001 A 101 = 80' 0".

ejf. 001 A 321 = 86* 88". et. 001 A 183 = 80* O'.

e/i. 001 A ill = 81* IB'. JfJf', 331 A 321 = 6r-4Si'. 004.

@35



Cleavage: basal, highly perfect; planes of separation shown in the percussion-figure; also gliding-planes p(S05), C {i'^^} shown in the preesure-fi^ nre inclined about 66 to c and yielding pseudo-crystalline forms (Fig. 475,p. 148). H. = 2-5-3. G, = 2-7-3'l. Luster splendent, uiii] more or less penrly on a cleavage surface, and sometimes aubmetallic when black; liiter»] gurfiices vitreoiib vben sinootb and shining. Colors usually green to black, often deep black in thick crystals, and sometioies even in thin iamirse, unless the lamiiw are very thin; such thin taminie green, blood-red, or brown by transmitted light; also pale yellow to dark brown; rarely white. Streak nncolored. Transparent to opaque.

Pleochroism strong; absorption b = c nearly, for a much stronger. Hence sections y c (UOI) dark green or brown to opaque; those Xc lighter and deep brown or green for vibrations | c, pale yellow, green or red for vibrations J_c. Pleochroic halos often noted, particularly about microscopic inclusions. Optically —. Ax, pi. usually { b, rarely _L />. Bi» (— a) nearly coincident with the normal to c, but inclined about half a degree, sometimes to the front, sometimes the reverse. Axial angle usually very small, and often sensibly uniaxial; also up to 50°. Birefringence high, y — a = 0"04 to 006,

Comp., Tar.—In most cases an orthosilicate, chiefly ranging between (H,K),(Mg,Fe). {Al,Fe),(SiO.). and (H,K),(Mg,Fe)^l.(SiO.). Of these the Becond formula may be said to represent typical biotite. Toe amount of iron varies widely.

Biotlle U divided into two clanea by Tschermak:

I. Merops. A. iul plane I b, II. Anorth. Az. pi. i b. Of these two kinds, merops is nearly all orthoclase, while the other is. so that its yet observed, of real occurrence, the typical localities being Greenwood Furnace, Orange Co., N. Y., and L. Baikal in E. Siberia. Merops is a name early given to the Muscovite. Ancillite is from Avafioi, contrary to law.

Himantite and Sitouropaj/litK are kinds of biotite containing much Iron.

Manganophyllite is a manganese-bearing biotite. Occurs in aggregations of thin scales. Color bronze- to copper-red. Streak pale red. From Pajsberg and Långban, Sweden; Piedmont, Italy.

The following are typical analyses of biotite: 1, by Berwerth; 3, by Bammesberg; S, by Smith and Brush.

BIO. Al₂O₃, Fe₂O₃, FeO, MgO, GaO, E₂O, Na₂O, B₂O₃, F

1. Vesuvius 89-80 1605 0.4e B-4S 21'8» 0-83 7'T» 0-40 402 089

[=10108

2. Mlask. & foet 83-49 12'84 650 3018 539 - 969 088 2 43 1-fl]

[TiO₂, 4 03 = 100-34 8. .dniMnft. Qieewood F. 89*88 14-99 7-68 - S8'«9 - 911 112 1-30 0'S5

[O10 44 = 9816

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Borne varieties give the reaction for 907.

SILICATES.

Pyr., etc.— If the doHcd tube gives a little water. fluorine is the open tube; some kinds give little or no reaction for Iron with the fluxes, while others give strong reactions for iron. B.B. white and fuses on the thin edges. Completely decomposed by sulphuric acid, leaving the silica in thin scales.

DIB.—Distinguished by its dark green to brown and black color and micaceous structure, usually nearly uniaxial.

Recognized in thin sections by its brown (or green) color; strong pleochroism and strong absorption parallel to the elongation (unlike tourmaline). Sections I e are nonpleochroic, commonly exhibit more or less distinct hexagonal outlines and yield a negative sensibly uniaxial figure. Sections X e are strongly pleochroic and are marked by

floe parallel cleavage lines (Fig. 907); they also have nearly parallel extinction, and show high polarization colors; by reflected light they exhibit a peculiar mottled or watered sheen which is very characteristic and aids in distinguishing them from brown hornblende.

Oba.—Biotite is an important constituent of many different kinds of igneous rocks, especially those formed from magmas containing considerable potash and magnesia. Common in certain varieties of gabbros, syenite, diorite, etc., of the massive granular type; also in rhyolite, trachyte, and andesite among the lavas; in micrites, kerstites, etc. It occurs also as the product of metamorphic action in a variety of rocks. It is not infrequently associated in parallel position with muscovite, the latter, for example, forming the outer portions of plates having a nucleus of biotite.

Some of the prominent localities of crystallized biotite are as follows; Yesuvius, common particularly in ejected limestone masses on Monte Somma, with uugite, clirysiite, uepielite, humite, etc. The crystals are sometimes nearly colorless or yellow and then usually complex in form; also dark green to black; Mt. Monzon in the Fassatbal; Bchwnrz-eusteln, ZUIerthal; HezbSuya and Horawitza in Hungary; Schelingen and other points in the Kaisersluhl; the Laacher See; on the west side of X. Ilmen near Minsk, etc.

In the United States ordinary biotite is common in granite, gneiss, etc.; but notable localities of distinct crystals are not numerous. It occurs with muscovite (wh. see) as a prominent constituent of the pegmatite veins in the New England States; also

'." 'Vnrh Pjirnlina .Qsbn>n£>i/;>rt< is frnm Ihs Plrn'o PbhIe rsirliin

Section i e.

PHLoooprm.

Monoclinic. In form and angles near biotite. Crystals prismatic, tapering; often large and coarse; in scales and plates.

Cleavage: basal, highly eminent. Thin laminae tough and elastic $H. = 2-5-3$. $G. = 3T8-286$. Luster pearly, often submetallic on cleavage surface—Color yellowish brown to brownish red, with often something of a copper-like reflection; also pale brownish yellow, green, white, colorless. Often exhibits asterism in transmitted light, due to regularly arranged inclusions, pleochroism distinct in colored varieties; c brownish red, b brownish green, a yellow. Absorption $C > b > a$, Burgess. Optically —, Ax. pi. j b. Bi, nearly J_c , ' Axial angle ($3E$) small but variable even in the same specimen, from 0° to $17^\circ 25'$ for red. Dispersion $p < v$. The axial angle appears to increase with the amount of iron.

A magnesian mica, near biotite, but containing little iron; potassium is prominent as in all the micas, and in most cases fluorine. Typical phlogopite, according to Clarke, is $R,Mg,Al(SiO_3)_2$, where $E = H,K,Mg,F$. Analyses: 1 by Penfield; 2 by Clarke and Schneider.

BIO, Al.O. Te,o, FaO KgO VaO K,Q V^Q H.O F

Edwnrda 2792 | 44S1 10(J7 — OBI 38-90 — 8W O'itC ola — Ign (100°) o-M

(=10018 Burgeu SSOO 17-00 027 0-30 36-49 063 997 0'60 399 334 Ti,o 0-60

[= 100 « • wiLh 0-0e u.o.

Obs.—Pblogoplte la eipednly cbaracteristic oC crystullioe lineatone or dolomite. It la oFttiu assoctaled wiib pyroxee, Hinhphole, aerpenllne, etc TLua aa at Farm, Fiu-laiiit; Id Si. Lawrence Co. aad JeSersoo Co., N. Y.; also Burgess, Ontario, ood elMwkere iu CauEnla.

Named from tfiinj'mnoi, fire-like, in allusion to the color.

The nalerUm of pUiogopite, seen when a caadle-flame la Hewed through a tbin sheet, is a'commoQ cbanicter, parllciiiarly prominent in the kinds from Dortbeni New York anil Cannda. It bns been aiiown to be due lo miauie aclculnr IncluBlons, rutlle or tourmaline, arranged chiefly in tile dlrectloD of the rays of Uie pressure-flnire, producing a distinct six-rayed star; nlao parallel to the lines of tho percuatdoa-flgure, giving a aecondary atar, uaitally less promluent than the other.

Lepidomelan*. Npar blotlte, but characterised by the large smouul of ferric Iron preaent, Annite from Cape Ann, Haaa., belonga here. Iu small aii-slded lables. or an ugeregate of minute scales. H. = S. O. — S'o-8'3. Color black, with occosionnly a leek-green reflection.

Alnrgite. A manganese mica from St. Marcel. Piedmont.

Rosco«llto. A vanadium mica; formula doubtful. In roioule tcalesi structure micaceous. O. = 2-93-2-94. Color cloTe-brown to greenish brown. Occurs at the gold mine tX Qraolte Creek, El Dorado Co., California.

2. Clntonite Group. Monoclinio.

The mineralB here included are sometimes called the Brittle Micah. They are near the niicae in cleavage, crfstalline form and optical properties, hut are marked physically hy the hrittleness of the lamiufe, aad chemically by their basic character.

In several reepecte they form a transition from the micah proper to the cblorites. Margarita, or calcium mica, is a basic silicate of aluminium aiiid calcium, while Chloritoid is a basic silicate of alnminium and ferrous irou (with magnesium), like the chlorites.

MARaAIUTB. Ealkgllnimer (7«r)n.

Monoclinic. Rarely in distinct crystals. tJaually in intersecting or aggregated laminf;e;

Sometimes massive, with a scaly structure.

Cleavage; basal, perfect. Laminae rather brittle. H. = 3-4-5, G. = 2.99-3.08. Luster of base pearly, of lateral faces vitreous. Color grayish, reddish white, pink, yellowish. Translucent, but subtranslucent.

Optically —. An. pi. X *• Bx» approximately J. c, but varying more widely than the ordinary micas, a c = $\pm 6^\circ$. Dispersion $p < v$. Axial angle large, from 100° to 120° in air. Refractive index $\mu_3 = 1.64-1.65$.

Comp. — H, Ca, Al, Si, O, = Silica 30-2, alumina 51-3, lime 14, water 1-5 = TOO.

Pyrr., etc.—Yields water in the closed tube. B. 6. whitens and fuses on the edges. Blowingly imperfectly decomposed by boiling hydrochloric acid.

Obs.—Associated commonly with corundum, and in many cases obviously formed directly from it; thus at the emery deposits of Quinch-ding in Asin Minor, the island of Naxos, Sicily, etc. Occurs in chlorite of Mt. Greiner, Sterzing, Tyrol. In the U. S., at the emery mine at Cheate, Mna.; at Unionville, Cheater Co., Va.; with corundum in

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Sand Co. from the Mwhera Id North Carolina; at Galenille, Hall Co., Georgia; Unionville, Alabama,

Idamed Marganite from /lapyapirtii, ptarL

SXTTBUUTITU, Clintonville. BnuKJUte.

Monoclinic, near biotite in form. Also foliated massive; sometimes lamellar radiate.

Cleavage: basal, perfect. Structure foliated, micaceous. Laminae brittle. Percussion- and pressure-figures, as with mica. H. = 4-6. O. = S-Sl. Luster pearly submetallic. Color reddish brown, yellowish, copper-red. Streak uncolored, or slightly yellowish or grayish. Pleochroism rather feeble. Optically —. An. pi. J. o seyherlike; \ b braitdisite. Bi» nearly X c. Axial angles variable, but not large. Birefringence high.

Var.—1. The Clintonite (clintonite) is in reddish-brown to copper-red brittle foliated masses: the surfaces of the folia often marked with equilateral triangular like wme mica and chlorite. Axial angle $8^\circ-18^\circ$.

2. Braadite* (diatriite), from the Fassathal, Tyrol, in hexagonal prisms of a yellowish.

Sea or leetgreen color to reddish gray; H. = 6 of base; of side, 4-6.5. An. plane | b. axial angle $10^\circ-12^\circ$. Some of it pseudomorphous after fassathite. Comp.—In part H,

(Mg,Ca),Al,Si,O,₃ = 3H₂O.10(Mg,Ca)O.5Al₂O₃.4SiO₂. Fyr., etc.—Tields water. B.B. iofusible alone, but whitens. In powder acted on t⁺.concentrated adds.

o\M.—8eii>trilt occurs at Amity. N. T., in limestone with serpentine, nssociated with «mphibole. spinel, pyroxene, graphite, etc.; also a chlorite near leuchleuberj⁺ite. Brandmtt -occurs OD Mt. Honzoni in the Fasuthal, I'yrol, In white limestone, with fassaie and black

XanthophylUU. Perhnps H.(Mg,Ca),<A1₂,.SI.O.>. The original xanthophyUiU is In trusts or in implanted globular forms. Optically negntiye. Ax. ADgle usually very small, ■or sensibly uniaxial; sometimes 80°. From near Zlatoust iu the Unil.

WaluwittSa the same species occurringin distinct pseudo-rbombobedral crystals. Folia brittle. H. = 4'8. G. = &'093. Luster vitreous; on cleavage plaue penrly. Color leek-to botrie-green. Transparent to transluceat. Pleocbroism rutber feeble: | h fine green; i i reddish brown. Optically -. Ax. pi. I b. Bx sensibly i c. 2B = 50°-40". Axial angle 17° to tS', Fonnd with perovsklte and other appecies in chloritic schists near Achma-lovsk, is the southern UraL

CHLORTTOID. Chloritspath. Ottrelile. Phyllite.

Probably triclinic. Barely in distinct tabular crystals, usually hexagonal in outline, often twinned with the individuale turned in azimuth 130° to each ■other. CryBtals grouped in roaettee. TJeually coarsely foliated maseive; folia often curved or bent; and brittle; also iu thin scales or small plates disseminated through the containing rock.

Cleavage: basal, but less perfect than with the micas; also imperfect parallel to planes inclined to the base nearly 90° and to each other about 60°; b diffi-■cuft. Lamina brittle. H. = 65. G. = 3-52-3-57. Color dark gray, greenish gray, greenish black, grayish black, often graas-greeii in very thin plates. Btreak nncolored, or grayish, or very slightly greenish. Luster of surface of cleavage somewhat pearly.

Pleocbroism strong; c yellow-green, 6 indigo-blae, a olive-green. Optically 4-. Ai. pi. nearly | ft. Bi» inclined about 12° or more to the normal to c (001). ^Dispersion p > v, large, also horizontal. Axial angles large, in air 100° to 118 . Birefringence low, y — a = 0-015.

Comp.—For chloritoid H,(Fe,Mg)Al,SiO₃. If iron alone is present, thia le⁺uires: Silica 23-8, alumina 40'5, iron protoxide SS'S, water 7'2 = 100.

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Diff.—Recosnlzed by the crjBtiil oiitlfnea and general micaceous sppefiniDce; high relfet; green colors: dlatluct cleavage; frequent tnioniDg; strong plileochrolsm and low iuletfer-euce-coloiB. Bjr the last characlar rtsdily diitiugiiiabed from ihe micas; also by the high relief aud eitlnclion obllqno lo tlie cleavage from the chlorili⁺s.

Obs.—Chloritoid (otrelile, etc.) are characteristic of stibular rocks which have suffered dynamic metamorphism, especially in the earlier stages; it is found in argillites, mica-schists, etc., which have assumed the schistose condition, with a more advanced degree of metamorphism it disappears. Often grouped in fan-shaped, leaf-like forms, also in irregular or rounded grains.

The original chloritoid from Eosibiria, near Ekaterinburg in the Ural, is in the form of curving laminae or plates, grayish to blackish green in color, often spotted with yellow from mixtures with limonite. Other localities are in the Groix (Moribau); embedded in large crystals at Vanlup, Shetland; in schists with orthopyroxene from Upper Uchigau; Leeds, Canada, etc.

Sillimanite (H, Fe, Al, Si, O) is found at Tom St. Marcel; it occurs also with glaucophane at Zermatt in the Valais, Switzerland, and elsewhere.

Salmite is a manganese variety occurring in irregular masses, having a coarse saccaroidal structure and grayish color. O. = S₂S₈. From Vielbalm, Belgium.

Matonite, from Kalic, R. I., is in very broad plates of a dark grayish-green color, but bluish green in very thin laminae parallel to c, and grayish green at right angles to this; occurs in argillaceous schist.

Orthopyroxene is generally classed with chloritoid, though it is not certain that they are identical; it seems to have the composition H, (Fe, Mn)Al₂Si₂O₇. It occurs in small, oblong, shining scales or plates, more or less rhombic in form and gray to black in color; in argillaceous schist near Ottrez, on the borders of Luxembourg, and from the Aude: also near Serravezza, Tuscany; Tintagel in Cornwall. Venatguite is from Venasque in the Pyrenees, and from Teulade, Finistère. Phyllosilicate is from the schists of New Buglia.

3. Chlorite Group. Monoclinic,

The Chlorite Group takes its name from the fact that a large part of the minerals included in it are characterized by the green color common with silicates in which ferrous iron is prominent. The species are in many respects closely related to the micas. They crystallize in the monoclinic system, but in part with distinct monoclinic symmetry, in part with rhombohedral symmetry, with corresponding uniaxial optical character. The plane angles of the base are also 60° or 120°, marking the mutual inclinations of the chief zones of forms. The mica-like basal cleavage is prominent in distinctly crystallized forms, but the laminae are tough and comparatively inelastic. Percussion- and pressure-figures may be obtained as with the micas and have the same orientation. The etching-figures are in general monoclinic in symmetry, in part also asymmetric, suggesting a reference to the triclinic system.

Chemically considered the chlorites are silicates of aluminium with ferrous iron and

magnesium and chemically combined water. Ferric iron may be present replacing the aluminium in small amount; chromium enters similarly in some forms, which are then usually of a pink instead of the more common green color. Manganese replaces the ferrous iron in a few cases. Calcium and alkalis—characteristic of all the true micas—are conspicuously absent, or present only in small amount.

The only distinctly crystallized species of the Chlorite Group are Clinocllore and Penninite. These seem to have the same composition, but while the former is monoclinic in form and habit, the latter is pseudo-rhombobedral and usually uniaxial. Prochlorite (including some ripidolite) and Corundophilite also occur in distinct cleavage masses.

Besides the species named there are other kinds less distinct in form, occurring in scales, also fibrous to massive or earthy; they are often of more or less undetermined composition, but in many cases, because of their extensive occur-

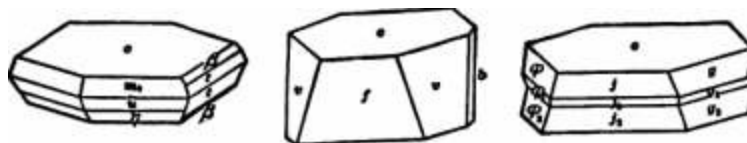
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rence, of considerable geological importance. These latter forms occur as secondary minerals resulting from the alteration especially of ferro-magnesian silicates, such as biotite, pyroxene, amphibole; also garnet, vesuvianite, etc. They are often accompanied by other secondary minerals, as serpentine, limonite, calcite, etc., especially in the altered forms of basic rocks.

The rock-iiiakin|: chlorites are recognized in thin sections by their characteristic appearance in thin leaves, scales or fibers, sometimes aggregated into spherulites; by their greenish color; pleochroism; extinction parallel to the cleavage (unlike chloritoid and ottrelite); low relief and extremely low interference-colors, which frequently exhibit the "ultra-blue." By this latter character they are readily distinguished from the micas, which they strongly resemble and with which they are frequently associated.

OIJNOOHZ.OBB. Ripidolite pt.

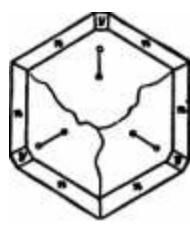
Monoclinic. Axes: $a:b:c = 0.57736 : 1 : 2.2772$; $\beta = 89^\circ 40' . 908$. MM. 910.



Pfltkh.

Schwabenst. u.

Zillertal.



Crystals usually hexagonal in form, often tabular. Plane angles of the basal section — 60° or 120° , and since closely similar angles are found in the zones which are separated by 60° , the symmetry approximates to that of the rhombohedral system.

Twins: (1) Mica law, tw. pi, X in the zone cm.; sometimes contact-twins with c as comp.-face, the one part revolved 60° or a multiple of 60° in azimuth with reference to the other; also in threefold twins, (2) Penninite law, tw. pi. c, contact-twins also united by c (Fig. 910); here corresponding faces differ 180° in position. Massive, coarse scaly granular to fine granular and earthy.

Cleavage: c highly perfect. Laminae flexible tough, and but slightly elastic. Pencil-like figure and prismatic figures orientated as with the micas (p. 464). H. = 2-2½ Luster of cleavage-face somewhat pearly. Color deep grass-green to olive-green; pale green to yellowish and white; also rose-red. Streak greenish white to uncolored. Transparent to translucent. Pleochroism not strong, for green varieties usually a green, c yellow. Optically usually +. Ax. pi. in most cases | b. Bx, inclined somewhat to the normal to c, forward; for Achmatovsk $2^\circ 30'$. Dispersion $p < v$. Axial angles variable, even in the same crystal, $2E = 20^\circ - 90^\circ$: sometimes sensibly biaxial. Birefringence low. $n_g = 1.588$; $y \sim a - 0.011$.

$V_r = -1$, Ordinary; green clinohedral, pi. L. in the zone $c^{TM}pn$: ") '^ crystal, ns described. usually with triclinic symmetry; (h) foliated: (e) massive.

Contains usually little or no iron. Color white, pale green, yellowish; often resembles talc. From near Zlatoust in the Ural.

Contains several per cent of chromium oxide. Crystal rhombohedral habit. Color rose-red. From the Ural.

Achmatovsk.

G. = 2-65-2-78.

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DE3CEIPTITB KIKEEALOOT.

Comp.—Normally $H, Mg, Al, 8i, o, = 4H, 0.5MgO, Al, o, 3SiO, =$ Silica 32.5, alumina 18.4, magnesia 3.1, water 13.0 = 100. Ferrous iron usually replaces a small part of the

magnesia, and the same is true of manganese rarely; sometimes chromium replaces the aluminium,

Pyr., etc.—Yields water, B.B. in the platinum forceps white and fuses with dil. HCl (III) the edges into grayish-black glass. With borax a clear glass colored by iron and sometimes chromium. In sulphuric acid wholly decomposed.

Obs.—Occurs in connection with siliceous mica rocks or schists and serpentine: sometimes in parallel position with biotite or phlogopite. Prominent localities are: Achnacross in the Ural; Ala in Piedmont; the Zillerthal: Zertatt, Switzerland; Miesbach, Saxony; Zschiz, Urvila. A manganese variety occurs at Pajsberg, Sweden.

In the U. States. at Westchester, Penn., in large crystals and plates; also in Devon and Texas, Penn.; at the magnetic iron mine at Brewster, N. Y., in part changed to serpentine.

PEIKNINIT. Pseudomorph.

Apparently rhombohedral in form but strictly pseudo-rhombohedral and monoclinic.

Habit rhombohedral: sometimes thick tabular with c prominent, again steep rhombohedral; also in tapering six-sided pyramids. Rhombohedral

faces often horizontally striated. Crystals often in curved groups. Also massive, consisting of an aggregation of scales; also compact cryptocrystalline.

Cleavage: c highly perfect. Laminae thin, flexible. Percussion on figure and

figure as with clinoclone but less easy to obtain; not elastic. H. — 3-2-5. G. = 2-6-285. Luster of cleavage-surface pearly; of lateral plates vitreous, and sometimes brilliant. Color emerald- to olive-green: also violet, pink, rose-red, grayish red; occasionally yellowish and silver-white. Transparent to subtranslucent. Pleochroism distinct: usually J t-p-reen; J.c yellow. Optically -f .also —, and sometimes both in adjacent laminae of the same crystal. Usually sensibly uniaxial, but sometimes distinctly biaxial (occasionally $2E = 61^\circ$) and both in the same section. Sometimes a uniaxial nucleus while the border is biaxial with $2E = 36^\circ$, the latter probably to be referred to clinoclone. Indices 1.576 and 1.579 Levy-Lox.

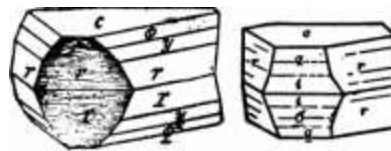
Var.—1. Pseudomorph, as first named, included a green crystallized chlorite from the Piedmontine Alps.

Kammereck. In hexagonal forms bounded by steep six-sided pyramids. Color kermes-red; peach-blossom-red. Pleochroism distinct. Optically - from L. Itkiil, Bieersk. Perm. Ruesselt: + Texas, Pa. Uniaxial or biaxial with axial angle up to 20° . Rhodochrosite from Texas, Pa., and rhodochrosite from L. Itkiil belong here.

Pseudomorph is compact massive, without cleavage, and resembles serpentine. Comp.—Essentially the same as clinoclase, $H_2(Mg,Fe)_2Al_2Si_2O_{10}$, H_2O . Pyr., «tc.—In the closed tube yields water. B.B. exfoliates somewhat and is difficultly fusible. With the fluxes all varieties give reactions for iron, and many varieties react for chromium. Partially decomposed by hydrochloric and completely by sulphuric acid.

Obs.—Occurs with serpentine in the region of Zermatt, Valais, near Mt. Rosa, especially in the moraines of the Fudalen glacier; crystals from Zermatt are sometimes 2 in. long and 1 in. thick; also at the foot of the Simplon; at Ala, Piedmont, with clinoclase: at Spangenberg in Tyrol; at Thiergen in Wernland; at Snarum, greenish and foliated.

Kammererite is found at the localities already mentioned; also at Miasma in the Ural;



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U Haroldswick la Unst, SbetUDI) Isles. In large crystals enclosed in the talc of the chlorite from Kraubach, Styria. Abundant at Texas, Lancaster Co., Pa., along with clinoclase, some crystals being embedded in clinoclase, or the reverse. Also in N. Carolina, with chromite at Cillsage, Haddon Co, in Webster, Jackson Co.; and other points.

PROOF. ORITB. Rippled pt.

Monoclinic In six-sided tables or prisms, the cleavage planes strongly furrowed and dull. Crystals often implanted by their sides, and in divergent groups, fan-shaped, vermicular, or spheroidal. Also in large folia. Massive, foliated, or granular.

H. = 1-2, G. = 2.78-3.96. Translucent to opaque; transparent only in very thin folia. Luster of cleavage surface feebly pearly. Color green, grass-green, olive-green, blackish green; across the axis by transmitted light sometimes red. Streak uncolored or greenish. Laminae flexible, not elastic. Pleochroism distinct. Optically -}- in most cases. Bz inclined to the normal to c some 2°. Axial angle small, often nearly uniaxial; again $2\epsilon = 23^\circ - 30^\circ$. Dispersion $p < i$.

Corap.—Lower in silicon than clinoclase, and with ferrous iron usually, but not always, in large amount. Analysis by Egger:

BIO. Al_2O_3 , Fe_2O_3 , FeO , MgO , CrO_3 , H_2O

Zellerthal 26 03 2016 1-07 2808 16'SO 044 9 65 = 10093 '

Obs.—Like other chlorites in modes of occurrence. Sometimes in implanted crystals, HR

at SI..Oi>lliBrd. cnvelopiDg often adulsrin, etc.; Ht. Qreiner In the Zillertbal, Tyrol; Ruurla iu Salzburg; Traversella in Piedmont; at Mtn. Sept Lacs and St. Cristophe In DiiupbinI: in Siyna. Bohemia. Also nmssfve in Corowall, lu tin velug; at Arendiil in Norway;'Salberg and Diinuemora, Sweden; Dogiincaka, HimsHry. Occatiluually formed from at)ipbi1)ote. In Scotland at varinua piitits. Iu tbe U. States, near '^ashiugton; on Castle Mt., BatesTille, Va., a mnssive form resembling soapstone, color grayish green, feel sreasy; Steele's mine, Montgomery Co., N, C.; also with corundum at tbu Culsagee mine, m broad plntes of a dark green color and flue scaly; it differs from ordinary prochlorile in the small amount of ferrous iron.

AmsTTX. H.(Hg.Fe)tAliBiO>. Silica 21-4 p. c. In beiagonal plaiea, foliated, resembling the green talc from thie Tyrol. H.= 2 5-8. G. = 2-71. Color apple-green. Luster pearly on cleavage face. Optically 4-, sensibly uniaxial. Occurs with diaspore at Chester,

Otbbr Chloriteb. Besides the cblorites already described which occur usuallr In distinct crystals or plates, there are. ns noted on p. 4Xi, forms varying from fine scaly to fibrous aud earthy, which as already noted are prominent in rocks, In some cases lley

J belong to the species before clescribed, but frequently the want of sufficient pure malerlal has left their composition in doubt. These cblorites are commonly characterized by their green color, disliuct pleoclirolsm and low birefringence (p. 478).

The foliowlug lire names which have been given particularly to the cblorites filling cavltiesorseamslu basic igneous rucks: a'phroti^nUt^iaba'atUa, ii>lettita,tpichbril«,miralite, i ehlorep/imU, hullite. \

Th>e following are other related minerals.

Oronstadite. 4Feo.2Fe,Oi.SSiO,,4H,o. Occurs taperingln hexagonal pyramids; also tn diverging groups; amorphous. Cli'avage: basal, highly perfect. ThIo laminte elastic. O. = 8-84-<-3S. Color coal-black to brownish black; by transmitted light in thin scales emerald-green. Bbreak dark olive-green. From Pllbram in Bohemia; also In Cornwall.

Thnringite. 8Feo.4rAl.(Fe),Oi.eSIO,.9B,o. Massive; an aggregation of minute

Karly scales. Color olive-green to plslachio-green. From nenr Siialfeld, lb Thnringla: >t Springs. Arkansas, etc.; from the metamorphic rocks on the Potomac, near Harper's Ferry (inMnA*).

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Chamobitb. Conlains iroD (FcO) wllh but little HgO. Occun compact or oOliilc wiih H. nbuiitS; G. = 3-3'4; color ereeQisb gray to bluck. From CbnmoMD, neiir Si'. Muurict;, in the Valuis.

Stilpnomelane. Au irou alHcate, In folbted platfa: also librous. oc as & velvety coaU Ing. G. — a'77-a-B6. Color black, giwoisli bWk. Occurs al Obergruud atid tlwwliete Id Sllletiiii; ulso lu HorsTlu: Lear Weilbuig. Nusaiiu. C/ialecditt, from the Bterling Iron uiiue. iu Auiwerp, Jefferson Co., N. Y., ctUlug Ijetnatite and calcite, ts Uie saine miuerol in velvety coating of niico-like Bcalea wiib a broaze color.

Btrigovltv. H.Fe,(Al,Ff)iSj,o. In nggregutious of miuul« crysUtls. Color dark

freeii. Occurs as a fine coaling over Ibe iiiiuerala in cavities in the groalta of Striegau la lleiia. Rumpfite, H,iHgTAl«BlitOii. HaKive; granular. consiatlDs of vsry floe Kales. Color greenlab wliie. Occurs with talc near St. Hicliuel tu Upper Stjria.

APPENDIX TO THE MICA DIVISION.-VERMICUUTES.

TLe Vbrhicolite Group Includes a number of micaceous minerals, all hydrated silicates, in pari closely related to the chloriles, but varying scimewlirit widely in compod-tiou. They are tilteraiion-products cbleti; of tlie micas, bioiite, pblogoplte, etc.,aiid r^aln more or less perfectly tlje micaceous cleavage, anil often show tlie negative optical character and small iixiil angle of the original species. Kany of tlien aiie of n more or less ludefluile chemical nature, and the composition varle> with tbat uf tiie original mineral and with the degree of iilteratiou.

The lamtnn iu general are soft, pliable, and Inelastic; the luster |>early or bronze-like, and tlie color varies from while to yellow and brown. Heated to 10u°-110° or dried over sulphuric acid most of tlie vermiculiles lose considerable water, up tii 10 p. c, which is probably hygroscopic: at itOO* nnothei' portion la often ^ven oS; and ai a ri'd heat a some' what large!' amount is expelled. Connected with the loss of water upon ignition is the common physical cliaracter of exfoliation; some of the kinds especially show this to a marked degiee, slowly opeolng out, when heated gradually, Into long worm-like tbreada. This character has given the name to thegrono, from the Iiatin Dermieulari, to bread uormt. The minerala Included can hardly rank as distinct species and only their names can be given here: JafferiiOa, wrmieaiiU, euitageviU, k«mt», Itnjtilile, AoUtU, phOaMpAUt, taalitt, maamila, dvdityilt, piftvteler&t.

III. Serpentine and Talo Division.

The leading species belonging here. Serpentine and Talc, are closely related to the Chlorite Group of the Mica Division preceding, as noted beyond. Some other magnesium Bilicates, in port amorphons, are included with them.

SEBPBNTINII.

Monoclinic, In distinct crystals, but only as psendomorphs. SometimeB foliated, folia rarely separable; also delicately fibrous, the fibers often easily separable, and either flexible or brittle. Usually massive, but microscopically finely fibrous and felted, also fine

granular to impalpable or cryptocrystalline; slaty. Crystalline in structure but often by compensation nearly isotropic; amorphous.

Cleavage 6 (010), sometimes distinct; also prismatic (50°) in chrysothile. Fracture usually conchoidal or splintery. Feel smooth, sometimes greasy. H. — 2.5-4, rarely 5.5. G. = 2.50-2.65; some fibrous varieties 2.3-2.3; retinite, 2.36-9.55. Luster siliceous to greasy, pearly, earthy; resin-like, or wax-like; usually feeble. Color leek-green, blackish green; oil- and siskin-green; brownish red, brownish yellow; none bright; sometimes nearly white. On exposure, often becoming yellowish gray. Streak white, slightly shining. Translucent to opaque.

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Pleochroism feeble. Optically—, perhaps also—]io chrysothile. Double refraction weak. Ax. pi. | a (100). Bx (a) J. b (010) the cleavage surface; c || elongation of fibers, biaxial, angle variable, often large; $2V = 20^\circ$ to 90° . Indices:

Anilgorit « = $\frac{J}{O} \frac{f}{l} \frac{O}{S} = 1.570 \text{ } 7 - = 1.571 \text{ } r - 'r = 0.011 \text{ } L \text{ } f \text{ } i \text{ } v \text{ } y - L \text{ } c \text{ } x$.

V*T.—MsLy UDSuslBiaed Bpecies linve been m^a out of urpeDlIDS, illfferlug Id sUuc-
lure (muMivti, elaty, foliated, Hbroua), or, as supposed, in chemical composition; and these

DOW, ill puTt, stHud as rarieties, alutig witli some others based on varlationalia iu textuie, etc.

A. In Crystalh —PsKtiDOMcsPHB. The most common have the form of chrysolite. Other kinds are pseudoniopha after pjuzeue, aoiptibole, spinel, chondrodite, gartiet, pblogupite, etc. BoMtiU ot tkhiiUr Spur la euetatite (hypentbeoe) altered more or less* completely/ to serpentine. See p. SSfl.

B. Mabbite. 1. Ordinary maitint.- (a) Predtnu or JiebU Serpentine is of a rich oil-green color, of pale or dark shades, and translucent even when in thick pieces. (A) Ckmtmon Serpentine is of dark shades of color, and Bubtranslucent, The former has a hardness of 2.5~li: the latter of 4 or beyond, owing to impurities.

Seinoui. JielinaliU. Massive, hue^'-yellow to iTgbt oil-green, waxy or resin-Uke

Bowenite (Nephrite Botnen). Massive. of very fine granular texture, and much resembles nephrite, and has long so called. It is apple-green or greenish white in color; Q. = 5.4-2.7»7, Bowenite; and it has the unusual hardness 5.5-6. From Smithfield, R. I.; also a kind from New Zealand,

C. Lamellar. Antimite, thin lamellar in structure, separating into translucent folia: B. =

2"); O = 2'622; color bronuish green by reflected light; feel smoolh, but not greasy. From Auligorio valley. Piedmont.

I). Thin Poliatbd, Murmoiiile, tbiu fitliuted; the lamlns brittle but i^parable. Q. = 2'41: colors greenish white, bluish while to pale aspiiragus-green. From Hiibuken, N. J.

E. FiBBOua. ChrytolUe. Delicately hbrous, ibe hhers usually Qexible and eHsily separating: luster silky, or silky nielnlllic; color greenish white, green, olive-green, yellow and brownish: G = 2'21B. Often conslilutts seuins in serpeniue. It includes moat ol the silky iimianaiut ot serpentine Tocks ntid mucli of wliaC is popularly called atbettut (asbestos). CI. p. 401.

Picroiile. coUiinnar, but fibers or columns not easily flexible, and often not easily separntble. or nffording only 11 apliutery fracture; color diiik green to mo ud lain-green, gray, brown. The ori^iial was from Tuberg, Sweden. BaltimoriU Is plerolile from Bare Hills, Md.

F. SBHPBNTinB RoCH8. Serpentine often constitutes rock-masses. It frequently occurs mixed with more or less of dolomite, magnesite. or culcite, making a rock of clouded green, somelimea veined with wbtle or pale green, called mtJ antique, ophiolitft, or opMeaMte. SerjientiDe rock is sometimes mottled with red. or has something of the aspect of a red porphyry; the reddish porliooos coDtainliig an unusnal amount of oxide of iron. Any serpentine rock cut Into slabs and polished Is called t«rpe)itine marble.

Microscopic examination has established the fact that serpentine In rock-masies has been largely produced by the alteration of cbrvsollte, and many apparently homogeneous serpentines show more or less of tliis original mineral. In other casee it has resulted from the alierstion of pyroxene or amphibole. Sections of the gi^

aerpentine derived from chrysolite often show a peculiar structure, like the met<bes of a net (Fig. 614): the line* marked by grains of magnelile also follow the original cracks ana cleavage directions of the chrysolite (Fig. 916, a). The serpentine from amplilbole and pyroxene commonly shows an analogous structure; the Iron particles following the former cleavage lines Hence the nature of tlie original mineral can often be Inferred. Cf. Fig. 915, a, b. e (Plrasoa).

Comp.—A magnesium silicate, H,Mg,Si,o, or 3Mgo.3SiO,.aH,o = Silica 44-1, magnesia 43-0,

water 12*9 = 100. Iron protoxide often replaces a small part of tbe mu^ nesinm; nickel in small amount is sometimes present. The water is chie^ expelled at a red heat.

Pyr., etc.— Id the closed tube yields water. B.B. fuses on the edges with diffloully.

F. = 6. GWea i:

acldB. From cbrysotife tbe silica la left in floe ll>erB.

Diff. —Characielrzed by Boftness, absence of cleavage anii feeble nsxy or oily luater; low specific gravity; by yteldiug much water B.B.

Readily recoguized iu tliin sectloDs by its greecUli or yellowlsta-gretQ color: lo* relief and aggregate polarization due U> Ita fibrous Bstructure, Wbea the flbera are parallel, (be

916. a h e

a, Berpeutine derived from chryBollte; b, from amphibole; e, from pyroxene.

Interfereace-colors are not very low, but tbe coufused aggregates may show the " ultra blue" or even be Isotropic. The conatant associatlua nftli otber magnesia benring minerals like chrysolite, pyroxene, hornblende, etc.. U also cbaracteristic. The presence of lines of Iron parllclea as noted above iFii;. 916) is cbai-acledstlc.

Obs.—Serpentine la always a secondary mineral resulting, us nored above, from the alteration of silicates coutaluiig magnesia, pankiilarly cb)ysolite, amphibole or pyroxene. It frequently forms large rock-maoses, then being derived frani the aheralion of peridoiito. duultes and otber basic rocks of Igneous orlcm: also of amphibolites. or pyroxene and chrysolite rocks of metnmorpblc origin. In the first case it is usually Bccom]mnied by spinet, garnet, cbrnmlce and sonieiiuiea nickel ores; lu the second case by various cni'boDiites such as doloinile. miignesite, breunnerile. etc.

Crystals of serpeulie. pseudomorphous after monticelllte. occur In the Fnssalbal. Tyrol; near Miask at Liike Ausbkul. Barwvka, Ekaterinburg, itud elsewhere; In Norway. at Siiiruni; etc. Fine precious serpnlines come from Falun and Gulsjo In Sweden. Ilie Ihle of Mnu, the neighborbood of Foitsoy in Al^erdeensblre, the Lizflrd in Coruwall, Uor* sIcB. Slberi'i, Saxony, etc.

In N. America, in Maint. at Deer Isle, precious serpentine. In Vrrmont. at New Fnne, Rozbiiry, etc. In Mau . fine at Newburyport. In H. Uland, at Newport; boweiille nt SmiChlied. In Gonu., uear New Haven and Milford. at the verd-antique quurrles. In N. York, at Port Henry, Essex Co.; nt Antwerp, JelTeraou Co., in crystals; in Oouver-neur, St. Lawrence Co , in crystals; In Cornwall. Monroe, and Warwick. Oranire Co.. somelimes in large crystals nt Warwick; and from Richmond to New Brighton, Richmond Co. In N. Jerity, at Hobokeu. with bnicite. magnesite, etc.; nt Mootville, Morris Co., cbvysotlle and retiunlite. wiib common serpentine, produced by the nllemtinn of pymxcne. Iu Penn., ma.'jsive, fibrous, and foliated, at Texas. Lancaster Co.; nt West Chester. Chester Co., icilliaiiiale; iil Mineral Hill, Newtonn. Marple. nnd Middlelown, Delaware Co. Id

Maryland, at Bait; Hills; at Cooplow, Harford Co., with diallage. In California, at various points in the Cotist Range.

In Canada, abundant among the metamorphic rocks of the Eastern Townships and Gaspe Peninsula, Quebec; at Thetford, Coleraine, Brookfield, Orford, St. Ham, Bolton, Sliplou, Melbourne, etc. The fibrous variety chrysotile (asbestos, bostonite) often forms several inches in thickness in the massive mineral, and is now extensively mined for technical purposes. Massive Laurentine serpentine also occurs in Grenville, Argenteuil Co..

Quebec. and North Burgess, Lanark Co., Ontario. In N. Brunswick, at Crow's Nest in New Brunswick. The names serpentinite, Ophiolite, Zapolite, eolubrinia, allude to the green serpent-like cloudings of the serpentine marble. Uralite is from Ural; Pektolite, from Pektol, bitter, in allusion to the magnesia (or Bitlererde) present; Thermophyllite, from Thermopylae, heal, and tpevnt; leaf, on account of the exfoliation when healed; Chrysothite, from jiro'u?, golden, and lithon, stone; Meibomite, from Meibom, silk; Marmolite, from Marmora, to thine, in allusion to its peculiar luster.

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Uralite. A magnesian silicate near serpentinite but with more water. Formula perhaps $4\text{H}_2\text{O} \cdot 8\text{SiO}_2 \cdot \text{Mg}_3\text{O} \cdot \text{MgO}$. Amorphous, resembling gum arabic, or a resin. H. = 9-3-5. G. = 2.0-2.2. Color white, yellow, reddish, brown. Occurs with serpentine in the Pleimahlul. Tyrol; also at Texas, Peou., and Uie Bure Hill, Ud, Qymnite of Tbcmsou, named from ytivoi, naktd. In allusion to the locality at Ban; Hills, Md., is the same species.

Owttblite. Nickel-Oxide. A gemstone with part of the magnesium replaced by nickel. $2\text{NiO} \cdot 2\text{MgO} \cdot 8\text{SiO}_2 \cdot 6\text{H}_2\text{O}$. Amorphous, with a delicate stalactitic surface, incrustations. H. = 3-4; sometimes very soft. G. = 3.4-3.8. Luster resinous, Color pale apple-green, or yellowish. From Texas, Lancaster Co., Pa., in thin crusts on chromite.

Oamierite. Nouraeite. An important ore of nickel, consisting essentially of hydrated silicate of magnesium and nickel, perhaps $\text{Ni}(\text{Mg}, \text{Fe})\text{SiO}_4 \cdot n\text{H}_2\text{O}$, but very variable in composition, particularly as regards the nickel and magnesium; not always homogeneous. Amorphous. Soft and friable. G. = 2.8-3.8. Luster dull. Color bright apple-green, pale green to nearly white. In part unctuous; sometimes adheres to the tongue, (as in serpentinite rock near Noumea, capital of New Caledonia, associated with chromic iron and steatite, where it is extensively mined. A similar ore occurs at Biddle in Douglas County, southern Oregon; also at Webster, Jackson Co., N. C.

TAXITE.

Orthorhombic or monoclinic. Rarely in tabular crystals, hexagonal or rhombic with prismatic angle of 60° . Usually foliated massive; sometimes also globular and stellated

groups; also granular massive, coarse or fine; fibrous (pseudomorphous); also compact or cryptocrystalline.

Cleavage: basal, perfect. Sectile. Flexible in thin laminae, but not elastic. Percussion-figure a six-rayed star, oriented as with the micas. Feel greasy. H. = 1-1.5. G. = 2.7-2.8. Luster pearly on cleavage surface. Color apple-green to white, or silvery-white; also greenish gray and dark green; sometimes bright green perpendicular to cleavage surface, and brown and less translucent at right angles to this direction; brownish to blackish green and reddish when impure. Streak usually white; of dark green varieties lighter than the color. Subtransparent to translucent. Optically negative. Ax. π . γ a. Bx J, c. Axial angle small. \wedge — \ll = 0035-0-050.

Tar.—Folin Ud. Talc consists of folia, usually easily separated, having a greasy feel, and presenting irregularly fluted faces, greenish white, and white colors. O = 2.75-2.78.

Minerals, Steatite or Soapstone (Serpentine Oerm.). a. Coarse granular, grayish green, and brownish gray in color; U. = 1-SC. It is an ordinary soapstone, more or less impure, b. Fine granular or cryptocrystalline, and soft enough to be used as chalk; as the French chalk, which is milk-white with a pearly luster. e. Indurated talc. An impure aluminous talc, harder than ordinary talc.

Pseudomorphous. a. Fibrous, fine to coarse, altered from enstatite and tremolite. b. Jlimite, having the form of pyroxene from northern New York and Canada.

Corap.—An acid silicate of magnesium, $H_2Mg_3(SiO_4)_2$, or $H_2O \cdot 3MgO \cdot 4SiO_2 = SiO_2, 63.5$, magnesia 31.7, water 4.8 = 100. The water goes off only at a red heat. Nickel is sometimes present in small amount.

Pyr., etc.—In the closed tube B.B. when intensely ignited, most varieties yield water. In the platinum forceps it exfoliates, and fuses with difficulty on the thin edges to a white enamel. Moistened with cobalt solution, assumes on ignition a pale red color. Not decomposed by acids. Rensselaerite is decomposed by concentrated sulphuric acid.

Diff.—Characterized by extreme softness, soapy feel; common foliated structure; pearly luster; it is flexible but inelastic. Yields water only on intense ignition.

Obs.—Talc or steatite is a very common mineral, and in the latter form constitutes extensive beds in some regions. It is often associated with serpentine, talcose or chloritic kyanite, and dolomite, and frequently contains crystals of dolomite, beryl, also asbestos, actinolite, tourmaline, magnetite.

Steatite is the material of many pseudomorphs, among which the most common are those after pyroxene, hornblende, mica, scapolite, and spinel. The most common mineral are

tbweffhlich commouly ufFord steatite by alteration; while IhowHIie BcapolfleBndnepbelite, wlfch contain soda and no magnesia, most frequently yield pIultG-like pseudoaiorplis. Thieie nm uIhu ste&tilic peeudomorpbs after quartz, dolomite, topaz, cliuistolte, Hiaurokite, cyanlte, garnet, veauviaulle. cbiysotile. gehlealte. Talc in the fibrous form Is pseudomorp b after enstatite atid Iremolite.

Apple-L^reeu talc occurs at UL Orelner in the Zlllerthal, Tyrol; Id the Valals and St. Ootliard III Sntzlerlaud; Id Cornwall, near Lizard Pcdot, with serpeuUue; the Slietland Ulauds.

In N. America, foliated talc occurs in Maine, at Dexter. In Vermont, at Bridgewaler, handsome green talc, with dolomite; Newfane. In Mau., at Middlefield, Windsor, Binnford, Auiluver, atid Cheater. In B. hland, at Bmlthfield, delicate green and white in a CTjstiilline limestone. In N. York, at Edwards, St. Lawrence Co., a due fibrous talc {agalite) associated with pink tremollte; on Staien Island. In N, Jarvey, Sparta. In Ftnn., at Texas, Nottngliam, TJtiluiville; in South Mountain, ten miles south of Carlisle; at Chestnut Hill, on tbe Schuykill, tolc and also soapstone, the latter quarried eitenslTely. Id Maryland, at Cooplow, of green, blue, aud roae colors. In N. Cur., at Webster. Jack. son Co. In Canada, in the towusljips Bolton. Sulton, and Potton, Quebec, with steutite in beds of Cambrian age; In tbe towiuhip of Elzevir, Hastings Co., Ontario, an impure grayial) var. in Archtean rocks.

SBPIOLTZt. Meerscham Q&rm. L'&sume de mer Fr.

Compact, with a smooth feel, and fine earthy texture, or clay-like; tXm rarely fibrous. H. = 2-2-5. G. = 2. Impressible by the nail. In dry masses fioate on water. Color grayish white, white, or with a faint yellowish or

reddish tinge, bluish green. Opaque.

Comp.—H.Mg,Si o., or $3H_2O \cdot 2MgO \cdot 3SiO_2$, = Silica 60-8, magnesia 2T-1, water 12-1 = 100. Some analyses show more water ($2U_2O$), which is probably to bo regarded as hygroscopic. Copper and nickel may replace part of the magnesiinm.

Pyr., etc.—In the closed tube yields first hygroscopic moisture, and at a higher temperature gives much water and a burnt smell. B.B. Booie vjirietlea blacken, then bum white, and fuse with difficully on the thin edges. Wiib cobalt solution a pink color on Ignitiou. Decomposed by hydrochloric acid with gelaliuizuiion.

Obi,—Occura in Asia Hmor, in masses in Etratified earthy or alluuiial deposits at the glains of Esklbi sber: at Unibscliilz in Horavia; in Morocco, called ic French Pierre de awn de Maroe; nt Vallecas In Spain, in extensive beds.

A flbious miuernl, baving llie composition of sepiolile, occurs in Utab.

The word meericlavja is Oermau fur tea-froth, iind alludes ii> lis liglitness and color. Sepiolile Glocker Is from cryjzia, euUle-fith, the bone of,whicb Is llgbl aud porous; and being also a production of the sea, " dclnde jpumotn marinam sigtilQcubat," sa_i s Olocker.

GoDDaiite, A hydrous nickel silicate, perhaps $H_2Ni_2Si_2O_8 \cdot 8H_2O$. In small fi-oglle gisins. O. = $3'4S9-3'S19$. Color yellowish, green. From RSttis, in Saxon Votgtlnnd.

Spadaite. Perhaps $5MgO \cdot 6SiO_2 \cdot 4H_2O$. Massive, amorphous. Color reddish. From Capo di Bove. near Rome.

SAPOHITEI. Plotine. .

Massive. In nodules, or filling cavities. Soft, like bntter or cheese, but brittle on drying. G. = $2-24-3'30$. Luster greasy. Color white, yellowish, grayish green, bluish, reddish. Does not adhere to the tongue.

CoBp.—A hydrous silicate of magnesium and alnminium; but the material is amorphous and probably always impure, and hence analyses give no uniform results. Contains SiO_2 , 40-45 p. c, Al_2O_3 , 5-10 p. c, MgO 19-26 p. c, H_2O 19-21 p. c; also Fe_2O_3 , FeO , etc.

Fyr., etc.—B.B. gives out water very readily and blackens; thin splinters fuse with difficulty on the edge. Decomposed by sulphuric acid.

Obs.'—Occurs in cavities In basalt, dlabfise, etc.; also with serpentine. Thus at Lizard Point, Cornwall, in veins in serpentine; nt various localities In Scotland, etc.

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Saponile la fn»D Mipo, toap; and piotlne from iciorifi, fat.

Caladonlta. A silicate of Iron, magDeelum and potasaium. Earthy or In mioute Males. Very soft. Color green. From cavities in amygdaloid at MU. Biildo tiat

OUuconit«. Essentially a bydrouB atllcate of iron and potassium. Amorphous, nud Tesembliiig aartby cliloite; either in cavities !□ rocks, or loosely Krauular massive. Color dull green. Abundant Id the "green sand," of the Clialk formatloD, sometimea constituting 75 to M p. G. of the whole,

PhoUdoUtc, Corresponds approiimntely to $5H_2O \cdot E_2O_3 \cdot 13(Fe,Hg)O \cdot Al_2O_3 \cdot 18BiO_2$. In minute crystalline scalea, G, = $S'406$. Color grayish yellow. From Taberg in Werm-land, Sweden, with garnet, dlopside, etc.

IV. Kaolin Division.

KAOUHITEI. EaollQ.

Moaoclinic; in thin rhombic, rhomboidal or hexagonal scales or plates with angle of 60° and 120° . Usually constituting a clay-like mass, either compact, friable or mealy.

Cleavage: basal, perfect. Flexible, inelastic. H. = 2-3.5. O. = 2-6-3-63. Luster of plates, pearly; of mass, pearly to dull earthy. Color white, grayish white, yellowish, sometimes brownish, bluish or reddish. Scales transparent to translucent; usually unctuous and plastic.

Optically biaxial, negative. $Bz \perp b$. Bx. and ax. π . inclined behind some 20° to normal to c (001) Dick. Axial angle large, approx. 90° .

Tar.—1. Kaolinite. In crystalline scales, pure white and with a satin luster like the mass. 2. Ordinary. Common kaolin, in part in crystalline scales but more or less impure including the comparative lithomarge.

$C^* \ll p$.—H, Al, Si, O, or $2H_2O \cdot Al_2O_3 \cdot 2SiO_2$, = Silica 46.5, alumina 39.5, water 14.0 = 100. The water goes off at a high temperature, above 330° .

Pyrophoric.—Takes water. B.B. Infusible. Gives a blue color with cobalt solution. Insoluble in acids.

DilF,—Characterized by its ductile, soapy feel and the alumina reaction B.B. Resembles infusorial earth, but readily distinguished under the microscope.

Obs.—Ordinary kaolin is a result of the decomposition of aluminous minerals, especially the feldspar of granitic and gneissoid rocks and porphyries. In some regions where these rocks have decomposed on a large scale, the resulting slurry remains in vast beds of kaolin, usually more or less mixed with free quartz, and sometimes with oxide of iron from some of the other minerals present. Pure kaolin in scales often occurs in connection with iron ore of the Coal formation. It sometimes forms extensive beds in the tertiary formation, as near Richmond, Va. Also met with accompanying diaspor and emery or corundum.

Occurs in the coal formation in Belgium; Schlan in Bohemia: in argillaceous schist at Loddve, Dept. of Hérault, France: as kaolin at Diendorf (Bodenmaia) in Bavaria: at Schemnitz; with fluor at Zinnwald, Yrieix, near Limoges, is the best locality of kaolin in Europe (a discovery of 1760): it affords material for the famous Sevres porcelain manufactory. Large quantities of clay (kaolin) are found in Cornwall and West Devon, England.

In the U. States, kaolin occurs at Newcastle and Wilmington, Del.; at various localities in the limestone region of Vermont (at Brandon, etc.), Massachusetts, Pennsylvania; Jacksonville, Ala.; Edgefield, S. C: near Augusta, Ga.

The name Kaolin is a corruption of the Chinese Katiling, meaning /[^]iBh[■]ridgt. the name of a hill near Jaucbau Fu, where tbe material la obtained.

Pholetlta Near kaolinite, but some analyses give 10 p. c. water. The original wu from tbe coal mines of Fins. Dept. of AUier, France.

BAIXOTSTTB.

Massive, Clay-like or earthy.

Fracture conchoidal. Hardly plastic. H.= 1-2. G. = 2'0-2-20. Luster

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BOmewh&t pearly, or waxj, to dull. Color white, grayisb, greenish, yellowish ^ bluish, reddish. Translucent to opaque, sometimea becoming translucent or even transparent in water, with an increase of one-fifth in weight.

Var.— Ordinary. E&rthj or waxy [a lusrer. aEi<l opaque niMsive. OalapeeU'it li liallojs-ite of Anf;lar. PteudotttaUU of TbomaoD & Binney la ad Impure vurleiy. dark j^ roen in color, with H. = 3-26. O. = 2'46i>. Itidianaiit \a a white poi-celulD clay from Luwreace Co., ludiana, where It occura with allophane in beds four to ten feet thick.

Simetile Is greeulah. aud iu certaiu atalea of humidity appears tranepareut aud almost gelatinous; it is from Condi, near Houdan, Frsuce.

Bole, iu pnrt, may beluug here; tUat is, those colored. UDCTuoits clays coutalolug more or less iroD oiide. which also have about 34 p. c. of water; the Irou gives them a browniah, yellowish or reddish L-olor; but they may be mixtures. Here beiongaBergteift (niouutaln-soap).

Cornp.—A silicate of aluminium (Al₂O₃.3SiO₂) like kaolinite, but amorphous and containing more water; the amount is somewhat uncertain, but the formula is probably to be taken as H₂O.Al₂O₃.2SiO₂.+ aq = Silica 43-5, alumina 36-9, water 19-6 = 100.

Pyr., eto.—Yields water. B.B. iafutible. A toe blue with cobalt solution. Decom* posKd by acids.

Obi.—Occurs often In veins or beds of ore, iia a aecoiiidarv product; also In grtuite and other rocks, buiog durived from the dcompoaitiou of some nluiiiious minerals.

Newtoalt*. HuAliSiiOii-l-nq. In soft white compact masses resembling kaolin. Fouud on Sued's Creek in the oortUeru part of Newton Co., Arkansas.

Oinolite, A hydrous silicate of aluminium. $2\text{Al}_2\text{O}_3 \cdot 9\text{SiO}_2 \cdot \text{SH}_2\text{O}$. Amorphous clay-like, or chalky. Very soft. $\text{O} = 2-18-280$. Color white, grayish white, reddish. From the island of Argentiern (Elmolos of the Greeks),

Montmorillonite. Probably $\text{H}_2\text{Al}_2\text{Si}_4\text{O}_{20} \cdot \text{nH}_2\text{O}$. Massive, clay-like. Very soft and tender. Luster feeble. Color white or grayish to rose-red, and bluish; also pistachio-green. Unctuous. Montmorillonite, from Montmorillon, France, is rose-red. Gonfolensite is paler rose-red; from Coufolens, Dept. of Charente. See Bl. Jean-de-Cale, near Thiviers.

Stolpniite is a clay from the basalt of Stolpen. Saponite of Nickel is a white, plastic, soap-like clay from the fountain from which issues one of the hot springs of Plombières, France, called Soap Spring; It was named saponite by MaunDaun.

PYROPHYLLITE.

Monoclinic? Foliated, radiated lamellar or somewhat fibrous; also granular to compact or cryptocrystalline; the latter sometimes slaty.

Cleavage: basal, eminent. Laminae flexible, not elastic. Feel greasy. $\text{H} = 1-2$. $\text{G} = 2-8-2-9$, Luster of folia pearly; of massive kinds dull and glistening. Color white, apple-green, grayish and brownish green, yellowish to ochre-yellow, grayish white. Sub transparent to opaque. Optically — $\text{B}_1 \times$ cleavage. Ax . angle large, to 108° .

Varieties.—(1) Foliated, and often radiated, closely resembling talc in color, feel, luster and structure. (2) Compact massive. white, grayish and greenish, somewhat resembling compact steatite, or French chalk. This compact variety includes part of what has gone under the name of agalmatolite, from China; it is used for slate-pencils, and is sometimes called pectolite.

Comp.— $\text{H}_2\text{Al}_2(\text{SiO}_4)_3$, or $\text{H}_2\text{O} \cdot \text{Al}_2\text{O}_3 \cdot 4\text{SiO}_2$, = Silica 66-7, alumina 28-3, water 5-0 = 100.

Pyro. etc.—Yields water, but only at a high temperature. B.B. whitens, and fuses with difficulty on the edges. The radiated varieties exfoliate in fan-like forms, swelling up to many times the original volume of the assay. Heated and moistened with cobalt solution gives a deep blue color (chromium). Partially decomposed by sulphuric acid, and completely on fusion with alkaline carbonates.

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note—BaMblite is talc, but distinguished by the method for testing with cobalt solution.

Obaite.—Compact pyrophyllite is the material on the base of some schistose rocks. The foliated variety is often the chief constituent. Occurs in the Ural; at Weatou, Sweden; near Orléans, Lorraine; Ouro Preto, Brazil.

Also !□ white stellate H^gregatious in Cottonstoae Hin., Heckleoburg Co.. N. C ; in Chaalerfeld Dist., B. C, with lazulile and cyanlte; in Lincoln Co., Oa., on Qiaves Utn. The compact kind, at Deep Rtver, N. C, is exlensively used for making alaie-peocils and reaemlea the so-called agalmatolile or pagodite of Ohina, often uted for omBmenlil earrings..

AIXOPBANB.

AmorphonB. In iooruBtations, asaally tbio, with a nammillaTy surface, and hyalite-like; aonietimea Btalaotitic. Occasionally almost pnTernlent.

Fractnre imperfectly conchoidal and shining, to earthv. Very brittle. H. = 3. ■ G. = 1-85-1'89. Luster yitreous to aubresinoos; bright and waxy internally. Color pale sky-blue, sometimes Kreenish to deep green, brown, yellow or colorless. Streak nncolored. Tran^ucent.

Comp.—Hydrous aluminium silicate, Al,SiO,-|- SH,o = Silica 238, alumina 40-5, water 35*7 = iOO. Some analyses give 6 equivalents of water = Silic» 22-2, alumina 37'8, water 400 — 100.

Impurities are oftvn present. The coloring matter of the blue varietj Is due to trmc«s of chrysncola. nnd substances Intermediate between allopbane and chrjsocola (mixtures) are not uncommon. The green variety Is colored by malachite, and the yellowish stHl brown by iron.

Pyr., ato.—Yields much water In the closed tube. B.B. cnimbiei. but Is infusible. Qlves a blue color with cobalt solution. Qelatlnizes with bydr ochloric scld.

Obs.—Allopbane ia regarded as a result of the decomposition of some aluminous silicate (feldspar, etc.); and it often occurs incnisting fissures or cavities in mines, especla!]; those of copper and limonite, and even in beds of coal.

Named from aXXoS. other, and i/Mtiyf<j€ai, to apptar. Id allusion to Its change of appearance under the blowpipe.

OoUyiits. 3AliOi.8iO,.9HiO, A clay-like mineral, white, with a glimmering luster, greasy feel, and adberiuq to the longiie G. = 3-3-15. From Ezquerra In the Pyrenees.

BohrBttarits. 8AliOf8aiOi.8DH,o. Resembles allopbane; sometimes like ^m in speaiance. H. = 3^'G, Q. = 1-05-3 OG. Colorpale green or yellowish. FromDulUu-

gvc mountain, near Freiensteln, In Styrla; at the Falls of Little River, on the Sand Htn., herokee Co., Alabama. The following are claylike minerals or mineral substances: Siopile, smectite, catlinite.

OnuMita. H.Cat(T,Er),C8i.on. Color yellowish brown. From HhterO, Norway.

Thanmaslte. CBaiO>.CBCOi.CaSO₄.16H₂O. Haesive, compact, cirstalline. Cleavage in traces. H. = 8 6. Q. = 1*877. Color white. Occurs fling cavftits and crevices at Uie BJeIke mine, near Areskuta, Jemtlaud, Sweden: nt first soft, but hardens on exposure to the nlr. Also in fibrous crystnlliiie masses at Paterson. N. J.

Uranophane Uranotil. Ca_{0.3}U_{0.2}SiO₄ + 6H₂O. In radiated aggrentions; massive, fibrous. Q. — 3-81-8-90 Color yellow. From the granite of Eupfetberg, Slle^ Uranotil occurs at Wslsendorf, Bavaria; Mitchell Co., N. C.

OHBTSOOOIXA.

Cryptocrystalline; often opal-like or enamel-like in texture; earthy. In-crasting or filling seams. Sometimes botryoidal.

Fracture conchoidal. Rather sectile; translucent varieties brittle. H, = 2-4. G. = 2-2 238. Luster vitreous, shining, earthy. Color moiintaiQ-green, bluish green, passing into sky-blue and turquoise-blue; brown to black when impure. Streak, when pnre, white. Translucent to opaque.

Comp.—True chrysocoUa appears to correspond to CnSiO₄ - 2H₂O = Silica

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34'3, copper oxide 45-3, water 205 =: 100, the water being double that of dioptaBe.

CompositloD variei much tliTOUgh impultles; free bHIcr, sIbo iilumlna, black oxide of copper, oxide of iron (or Ilmonlte) luid oxide of munguneee m&y bu preaeuti The color con-Beqieotl; varies rnDia bluish green lo blown and black.

Fyr , etc.— Id the closed tube blackeoB uod yields water, B.B. deci«|)itBt«*. colon the fiame ernernld-green, but la Infualble. With the Uuxes gives the reactioDB for copper. Wiih soda and charcoal a globule of metallic copper. Decompoaed by acids without gelalinizatlon.

Oba —Accompanies other copper ore«, occurring especially in the upper part of Teios. Found ill copper mines la Cornwall^ Hungary; Siberia; Saxony; South Australia; Chllf,

In the n. B., similarly at the Schuyler's mloes, New Jeney; atHorgantowa, Pa.; at the Clifton mines, Qrabam Co., Arizona; Kmna mine, Utah.

C^rt/toeoBa is from ;[f>u(TDS, foM.and K6\Xa,glvt,aii<l was the name of a material used Id Bolderlug gold. The name is often applied now to borax, which Is so eniplttyed.

OHLOROPAL.

Compact massite, with an opal-like appearance; earthy. H, = 2-5-4-5. Or. = 1-727-1'870, earthy varietal, the second a conchoidal specimen; 3105, CeYlon, Thomson. Color greenish yellow and pistachio-green. Opaque to blue translucent. Fragile. Fracture conchoidal and splintery to earthy. Feebly adhering to the tongue, and meager to the touch.

Var.—Ohtoropal has the above-mentioned characters, and was named from the Hungarian mineral occurring at Unghwar.

Jfanirouit is pale straw-yellow or canary-yellow, and translucent, with an unctuous feel; flattens and grows lumpy under the pebble. and is polished by friction; from Noiron. Dept. of Dordogne. France. Pinguite is siskin- imd oil-green, entirely siliceous like new-made soap, with a slightly resinous luster, not adhering to the tongue: from Wolkenstein in Saxony. Gramine has a grass-green color (whence the name), and occurs at Meixenberg, in the Siebengebirge. In thin fibrous seams, or as delicate lamella.

Corap.—A hydrated silicate of ferric iron, perhaps with the general formula $H.Fe.fSiO_2$, + $211,0$ or $Fe_2O_3 \cdot 38iO_2 \cdot 5H_2O =$ Silica 41-9, iron sesquioxide 37-2, water 20'9 = 100. Alumina is present in some varieties.

The water and silica both vary much. The Hungarian chloropal occurs mixed with opal, and transitions into it, and this accounts for the high silica of some of its varieties.

Obs.—Localities mentioned above chloropal occurs also at Meuser Selenberg near Gottin; pinguite at Sternberg, Moravia. On Lehigh Mt., Pa., south of Allentown, occurs in connection with iron deposits.

UmPERITB. An iron silicate near chloropal. Color green. From Kitz, Bohemia.

Hlaingarita. A hydrated ferric silicate, of uncertain composition. Amorphous. compact. Fracture conchoidal. H.=8. G. = 2-5-3 0. Luster (resinous). Color black to brownish black. Streak yellowish brown. From Riddarbyttan, Tunaberg. Sweden; Liogban, etc., Norway.

The following are hydrous manganese silicates. Bementite

gillie. G.

ye. N. J.

OaryopllU. Approximately $4MnO \cdot 38iO_2 \cdot 3H_2O$. In stalactitic and reniform shapes. G. = 2-83-2-91. Color brown. From the Haratig mine near Pajsberg, Sweden.

Heotocte. A hydrated silicate of manganese and iron, of doubtful composition, usually

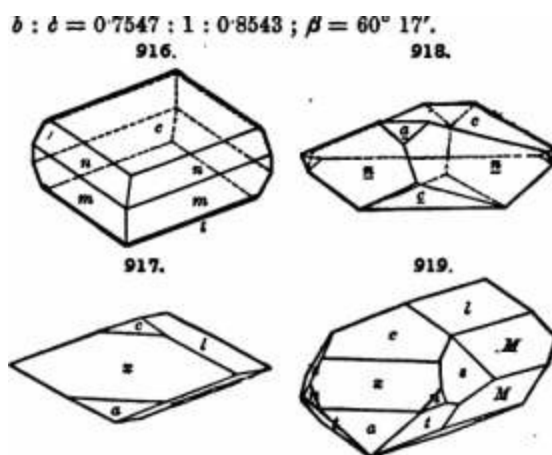
derived from the altheiatlou of rhodonite. Amorphous Color black to dark brown and Uver-brown.

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TITANO-SILICATES, TITANITES.

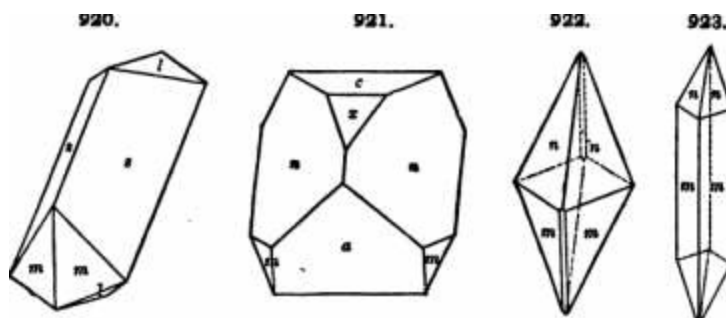
This section includes the common calcium titanate, Titanite; also a number of silicates which contain titanium, but whose relations are not altogether clear; further the titanate, Perovskite, and niobio-titanate, Dysanallyte, which is intermediate between Perovskite and the species Picrochlore, Micro-lite, Koppite of the following chapter.

In general the part played by titanium in the many silicates in which it enters is more or less uncertain. It is probably in most cases, as shown in the preceding pages, to be taken as replacing the silicon; in others, however, it seems to play the part of a basic element in schorlomite (p. 419) it may enter in both relations.



Twins: tw. pi, a rather common, both contact-twins and cruciform penetration-twins. Crystals very varied in habit; often wedge-shaped and flattened; also prismatic. Sometimes massive, compact; rarely lamellar.

Cleavage: m rather distinct; a, l (Il2) imperfect; in greenovite, $\{111\}$ easy, t (111) less so (Dx.). Parting often easy $\{221\}$ due to twinning



lamellar. H. = 5-5.5. G. = 3.4-3.66; 3.54. Chester, Pirsson. Luster adamantine to resinous. Color brown, gray, yellow, green, rose-red and black. Streak white, slightly reddish in greenovite. Transparent to opaque.

Pleochroism in general rather feeble, but distinct in deep-colored kinds: c, red with tinge of yellow; b, yellow, often greenish; a, nearly colorless. Optically +. A_i . ρ_i . i b. Bx nearly X (102), i.e., $B_x A_i = + 51^\circ$. Dispersion $p > v$ very large, and hence the peculiarity of the axial interference figure in white light. Axial angles $2E = 50^\circ$ to 90° . $\gamma = 1.594$. Birefringence high, $\gamma - \alpha = 0.121$.

Vai.—Ordinary. (a) Titanite; brown to black, the original being thus colored, also opaque or translucent. (b) Spinel (named from *Tipfer*, a *Ti*); of light greenish, as yellow, greenish, etc., and often translucent; the original was yellow. Ligustrum is greenish yellow. Spinel (from *tiemline*) a greenish black. Titanite is brown, opaque, translucent, of the form in Fig. 916.

Titanite is a white mostly granular alteration-product of rutile and ilmenite, not uncommon in local crystalline rocks; it also belongs to leucosomes (see p. 47).

Titanite; Qreemite. Red or rose-colored, owing to the presence of a little manganese; from St. Marcel.

Containing yttrium or cerium. Here belong grossular, androite, andradite.

Comp.— $CaTiSiO_6$, or $CaO \cdot TiO_2 \cdot SiO_2$, = Silica 30.6, titanium dioxide 40.8, Lime 28.6 = 100. Iron is present in varying amounts, sometimes manganese

and also yttrium in some kinds.

Pyr., etc.—B.B. some varieties change color, becoming yellow, and fuse at 8 with lathumescence, to a yellow, brown or black glass. With borax: they form a clear yellowish-green glass. It is perfectly soluble in heated hydrochloric acid; and if the solution be concentrated along with tin. It becomes of a violet color. With salt of phosphorus in R.F., gives a violet bead; varieties containing much iron require to be treated with the flux on charcoal with metallic tin. Completely decomposed by sulphuric and hydrofluoric acids.

Diff.—Characterized by its oblique crystal habit, a wedge-shaped form common; by resinous (or adamantine) luster; hardness less than that of staurolite and greater than that of apatite. The reaction for titanium is distinctive, but less so in varieties containing much iron.

Distinguished in thin sections by its acute-angled form, often lozenge-shaped; is generally pale brown to a very high relief and remarkable birefringence, causing the section to show double of the higher order; by its biaxial character (showing many lemniscate curves); and by its great dispersion, which produces colored hyperbolus.

Obs.—Titanite, as an accessory component, is widespread as a rock-forming mineral,

though confined mostly to the acidic feldspathic Igneous rocks; it is much more common than the plutonic granular types than in the volcanic forms. Thus it is found in the more basic bornblende granites, syenites, and diorites, and is especially common and characteristic in the nepheline-syenites. It occurs also in the metamorphic rocks and especially in the schists, gneisses, etc., rich in magnesia and iron and in certain granular limestones. It is also found in beds of iron ore; commonly associated minerals are pyroxene, amphibole, chlorite, scapolite, zircon, apatite, etc. In gneiss and granite, it often accompanies adularia, smoky quartz, apatite, chlorite, etc.

Occurs at various points in the Orisons, Switzerland, associated with feldspar and chlorite; Tavetsch; in the Bt. Golhurd region: Zermatt in the Valais; Maderanertal in Uri; also elsewhere in the Alps; in Dauphin (France): at Ala, Piedmont (Italy); at St. Marcel, in Piedmont: at Bchwdrzeusien and Rothenkopf in the Zillerthal, Pflsch, Tyrol; Zopau, Moravia; near Tavistock: near Tremailoc, in North Wales; in titaniferous iron at Arendal, in Norway; with magnetite at Nordmark, Sweden: Achmatovsk, Ural. Occasionally found among volcanic rocks, as at Lake Laach (Germany) and at Andemach on the Binn.

In Maine, in fine crystals at Sandford. In Mass., in gneiss, in the east part of Lee: in Bolton with pyroxene and scapolite in limestone. In N. York, at Roger's Rock on Idike George, abundant in small brown crystals; at Gouverneur, in black crystals in granular limestone; in Diana near Natural Bridge. Lewis Co., in large dark brown crystals, among which is the variety Udrill; at Rossie, Fine, Pitcairn, St. Lawrence Co.; in Orange Co., in limestone; near Edenville, in light brown crystals in limestone; at Brewster, at the Tilly Foster iron mine. In N. Jersey, at Franklin Furnace, honey-yellow. In Penn., Bucks Co., three miles west of Attleboro', associated with wollastonite and graphite. In N. Carolina, at Statesville, in Edell Co., yellowish white with staurolite; also Buncombe Co., Alexander Co., and other points.

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Occurs in Giniubi. stOrenrille. Argenteuil Co.; also Bucklefield. Wakefield. Hull, Ottawa Co.: titaniferous yellow: near Eganville, Redfrew Co., Ontario, in very large dark brown crystal* with apatite, amphibole, zircon.

Keilhanite. A titanate-silicate of calcium, aluminium, ferric iron, and the yttrium metals. Crystals near titanite in hornblende and angles. H. = 0.5. O. = 8-09-8-77. Color bluish-black. From near Arendal, Norway.

Quarantite. CaTiSiO_6 . u for titanite. In minute thin plates, flattened | b (010), nearly tetragonal in form. H. = 6. G. = 8-487. Color lilac-yellow, honey-yellow. Found in a grayish trachyte on Monte Boio.

Tschermakite. A titanate-silicate of the cerium metals, iron, etc.. It is an alteration-product,

more or less heterogeneous, and (lie compo«ltlon of the original mineral is very tincerlain. Haaslve, amorphous. H. = 5-S-O. o. = 4-So8-4'Mil Color velvet-black. From the Ilroen mouDlains In the Ural. Also from 8. Indiit. probably KHnjonutlai Hill, Salem distr. Au Uolated moss weighing 30 lbs. has been found on Hat Cre^ . near Hna-ale'B Hills. Nelson Co , Vlrglua; also found, south of this point. in Bedford Co.

ArtrophTlllU. Probably K,R,Tl(BiU,). >vlib H = H, Nu, E, and R = Fe. Hn cbiefly, including also Fcid. In elongated crystals; also in Ihln strips or bUdes; sometimes la stellate groups, Cleavase: 6 perfect like mica, but lamlDee brilile. H. = 8. Q. s 8 3-8'4. Luster submetallic, pearly. Color bronze-yellow to gold-yellow.

Occurs oo tlie small islsuds in the Langeaund fiord, near Brevik. Norway, In elsoljte* syeuile, embedded In feldspar, with catnpleiite, eegirtte, black mica, etc. Similarly at Kaueerdluarsuk, Oreenland. Also with arfvedsouite and zinwu at SL Peter's Dome, Pike's Peak. Ul Paso Co., Colo,

Johnstraplta. A silicate of the cerium metals, caldum and sodium chiefly, with titanium and fluorine. In prismatic crystals. O. = S-29. From near BarkevIk, Norway.

Mosandrita. Near jobaslrupile In form and composition and from the same region.

Rinkite, also near johnstruplle. Is from Oreenland.

Neptiinite. A titRno-silicate of iron (niantninese) and the alkali melals. In prismatic monoclinic crystals. H. = 5-6. O. = 8-38. Color black. Southern Greenland.

PBROTSKITB. Perofskite.

Isometric or paeu do-isometric. Crystals in general (Ural, Zermatt) cubic in habit and often highly modified, but the faces often in-egularly distribnted. Cubic faces striated parallel to the edges and apparently penetration-twins, as if of pjritohedral individuals. Also in reniform masses showing small cubes.

Cleavage : cubic, rather perfect. Fracture Uneven to aobconchoidal. Brittle. H. = 5-5. G. = 4 017-4-039 Zermatt. Luster adamantine to metallic-adamantine. Color pale yellow, honey-yellow, orange-yellow-. reddish brown, grayish black. Streak colorless, grayish. Transparent to opaque. Usually exhibits anomalous double refraction.

Geometrically considered, perovsklte conforms to the isometric system; optically, however, it is uniformly blailal and usually positive. Tlie molecular structure (also as developed by eichiug. Baumbsuer) seems to correspond to onhorbonible symmetry. Cf. Art. 411.

Comp—Calcium titanate, CaTiO, — Titanium dioxide 58-9, lime 41-1 = 100. Iron is present in small amount replacing the calcium.

Pyr.. etc.—In the forceps and on cobalt infusible. Willi salt of phosphorus In O.F. dissolves easily, giving a greenish bead while hot, which becomes colorless on cooling; in H.P. the bead changes to grayish green, and on cooling assumes a violet-blue color, j Unlikely decomposed by boiling Sulphuric acid, I

Ob(,—Occurs in small crystals, associated with chlorite, and magnetic Iron In chlorite i slate, at Acmatovsk. near Znloust, in the Urnl: at Schellngen in the Kaiserstuhl.in graniilnr timeslone: in the valley of Zermatt, near Ibe Findelen glacier; at Wildkreuzjub, between Ptilsi'b and Pfuuders In Tyrol, SpmelimeB noted in microscopic octahedral crystals iis a rock coiistiliicnt; iilus in nephellite- and melilite-basalts; also in serpentine (altered perido-tite) lit Syracuse, N, T.

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Eiioipit«. Near perovskite but contains cerium. In black isometric crystals. From AluQ, Sweden.

Dywuialyto. A titanate of calcium and Iron. In cubic crystals. From the granular limestone of Vo ' ' ' ■ ■ ■ ■ " ■ ■ "

, ferroTskite, but is in fact pyrochlore and koppelite.

A related mineral, which has also long passed as a* perovskite, occurs with m brookite, niobite, etc., at Magnet Cove, Arkansas. It is octahedral or cubo-octa black or brownish black in color and highly metallic in lustre.

See also the allied titanate, niobite, mentioned on p. 548.

OkiaUU. Huxnesium titanate, MgTiO.. Massive, as rolled pebbles. G. = 4, Color bluish or brownish black. From Ceylon.

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KIOBATBS, TANTALATES,

Oxygen Salts. 8. NIOBATBS, TANTA^^TES.

The Niobates (Golumbates) and Tantalates are chiefly salts of metaniobio and meta tautalic acid, BNb₂O₇, and RTa₂O₇; also in part iron niobate B, R, Nb, O₇, etc Titanium is present in a number of the species, which are hence intermediate between the niobates and titanates. Niobium and tantalum also enter into the composition of a few rare silicates, as wohlerite, laTenite, etc.

The following groups may be mentioned:

The isometric Ptrochlobe Group, including pyrochlore, microlite, etc The tetragonal Fkhgusosite Group, including fergusonite and aipylJto The orthorhombic Columbitk Group, including colnmbite and tantalita Also the orthorhombic Sauabskitb Gbouf, iQclading yttrotantalite, samarsfc ite, and ^snerodite.

The species belonging in this class are for the most part rvio, and are hence but briefly described.

PTROORLORB.

Isometric. Commonly in octahedrons; also in grains. Cleavage: octahedral, sometimes distinct. Fracture conchoidal. Brittle. H. = 5-5'5. O. = 4'2-4'36. Luster vitreous or resinouB, the latter on fracture surfaces. Color brown, dark reddish or blaactiab brown. Streak light brown, yellowish brown. Subtranslucent to opaque.

Comp.—Chiefly a niobate of the cerium metals, calcium and other bases, with also titanium, thorium, fluorine. Probably essentially a metaniobate with a titanate, RNb,o.,R(Ti,Th)o.; fluorine is also present.

The followlDg ate aiuilyseB bj Ranimelsberg :

G. Nb,o, TIO, ThO, Ce,o, CaO FeO MgO NaO, P Mlask 4-859 f 58-19 1047 TM 7-00 14-21 1-84 o 22 SOI ~

FredriksvBrn 4-228 47 18 18 62 — 780 15-84 10-03 019 3 13 2-SO

[fgn. 1-39 = 101-52

Oba.—Occura In elEeotlte-ayeuite at FredriksvBrn and I>iurvIk. Norwnv; od the iiland L5vo. oppoaileBrevIk, aod at several points in the Lujpesuud fiord; uearHiaak In tbsUral. NuDied from nCp, fin. aod jfAiupdS. green, because B.B. it becomes yellowisli green.

Kopplla. Bssentialij a pjroDiobste of cerium, calcium, etc., near pyrochlore. In miuute brown dodecahedrons. O. = 4-45-4-50. From Scbellngen, Eaiserstuhl, embedded In llmeEtoue,

BatclMttollta. A tantaloniobale of uranium, near pyrochlOre. In octahedrons wllb a (100) and n> (811). G. = 4-T7-4'90. Color yellowlBli brown. Occura wllh samai-sklte, at the mica mioea of Hitchell Co., North CaTofiaa.

BUorolite. Essentially a calcium pyrotantalate, CaiTaiOi. but contalnin); hIm niobium, fluorine and a variety of bases In small nmount. Isometric. Habit octahedral; Cfyalals often very small and highly modified (Fig. 109, p. 40). H. = 5 5. G. =5-486-5-562; 8 18 Virginia. Color pale yellow 10 browo, rarely byacynth-red. From Cbesterfeld, Hbbs.. in albite: Brancbville, Coun.; UtO, Sweden. Also in fine crystals up to 1 in. in diameter at the

mica mines at Amelia Court-Houae, Amelia Co., Ya.

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DSaCBIPTIVK UINEBA.LOGT.

Ptrkhitb. Probsblf a mlcrolita. Occore In minute orange-yellow In ihe Ural.

FEIRanBOHITB. Tyrite. Braglte.

Tetragonal-pyramidal. Axis t = 1'4643. OrTStals pyramidal or prismatic in biibit.

GleATage: j|[III) in traces. Fracture snbeenboidal. Brittle. H. = 5'5-6. G. = 5'8, aiminiahinK to 4*3 when largely hydrated. Luster externally dull, on the fractare brilfiantly ritreouB and Bubmetallic. Color brownish black; in thin sc^es pale liyer-brown. Streak pale brown, Subtranelucent to opaque.

Comp,—E^entially a metaniobate (and tantalate) of jrtrrium with erbinm, cerium, uranium, etc., in varying amounts; ^bo Iron, cscliam, etc. Qeneral

formula $K\{Nb,Ta)O$, with $R = T^Er.Ce$.

Water U usually preBent and sometimea In conaidenble amount, but probably not an orlgiDai conatituent; the epeclflo grarity falls as the amouut Increaaes. Aualyses by Bam-mebberg;

G. NbiO. Ta,o. UO, WO, SnO, Y.O, Er.O, Ce,o, FeO CbO H,o Greenland, F»g. S-sn 41-46 68o 3-58 o-lB o-47 24-87 981 7-6S* o-T4 O'Sl 1'49

[= WIO Tttert, yw. 4-774 38-14 27-04 213 — — 24-46 S-M — 072 417 613

[=100-03

• iDcl. G'U T>t,o,,U,Oi.

Oba.—Prom Cape Farewell In Qreeulaad, In quarU; also at Tlterbv, Sweden, and E&rarfvet. Tyrit» ia associated with euxeulte nt Hampemyr on Ibe lalana of TromO, aod Helle on the mainland; bragite la from Eelle, Narestn, elc, Norway.

Found Id the U. S., at Kockport, Mass., In granite; in the Brindlstown gold district,

Splyllta. A niobate of erbium chiefly, also the cerium melala, etc., near fergusonite In form. Rarely In octahedral cryatala. IJaiially In Irregular niasaea. G. =: 4*89. Color bTOwnish black to brownish orange. Occurs sparingly with allanlte In Amberat Co..

Virginia.

COIiUMLUTXI-TAHTAUrXI.

Orthorhombic Axes d:h:i = 0-8286 :1 : 0-8898.

W", 210 A 210 = 45" 0'. et. 001 a 031 = 60° 40*.

mm ", 110 A liO = 78' IT. ao. 100 A 111 = 61' 18'.

ffg', 130 A I80 = 43" W. en, 001 a 138 = 48° 48'.

ek, 001 A lOS = 19- 43'. ««', 138 a 188 = 29' ST.

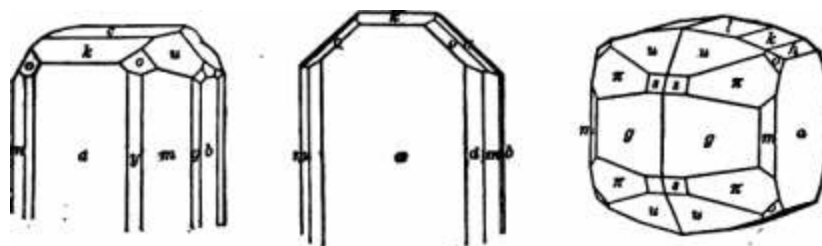
eg, 001 A 038 = 80° 41'. w." \ 188 a iSS = 79* 64'.

Twina: tw, pi. e (021) common, usually contact-twin a. heart-shaped (Figs. 347, p. 118), also penetration-twina; further tw. pi. g (053) rare (Fijr. 404, p. 281). Crystals abort prismatic, often rectangular prisma with the piracoida, a h c, prominent; also thin tabular || a; the pyramids often but slightly developed, sometimes, however, acutely terminate by « (133) alone. Also in large groups of parallel crystals, and maaaiye.

CleaTage: a rather distinct; 6 leas ao. Fracture anbnconchoidal to nneven. Brittle. H. — 6. Q. = 5-3-r-3, varying with the compoition (see below). Luster submetallic, often very brilliant, sub-reainous. Color iron-black, grayidi

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and brownish black, opaque; rai-elj reddiah brown and tranalncent l frequently iridescent. Streak dark red to black.



Comp^iTar.—Niobate and tantalate of iron and manganese, (Fe,Mii)(Nb,Ta),o^ paseiug by insensible gradations from normal Ooluhbite, the uearl; pure niobate, to normal Tantalite, the nearly pure tantnlate. The iron and manganese also vary widely. Tin and wolfram are present in small amount. The percentage composition for FeNb o, = Niobium pentoiide 82'7, iron protoxide 17-3 = 100; for FeTa,o, = Tandtum pentoiide 861, iron protoiide 139 ^ 100

In some varieties, mattffanoeolumbiU or mangantantaiiie, the Iron is largely replaced b;

The coonectloD between the specific gravity and the percentage of metalllcacids 1b ahowD

in the following table :

a. Ta, o. Q.

6-86 3-3 Bodeumals G-93

G'66 ia-8 flnddam 6-09

S'TO 18-8 BodcDinalB 9-M

6-74 13'4 Haddam 618

5-86 10-0

Tantalitt 7-08

Gre«o1aad Acworlh, N. H. Limoges

Bodenmals (JXaniU) Bad dam

Tb, o. 271 30-4 8S4 815

«S'S

Diff. — DlstinguUbed (from black tourmaline, etc.) by orthorbomblo crstalllzation, rectangular forms common; high specific gravity; submetallc luater, often with Iridescent surface: cleavage much less distinct than tor volframlte.

PyT., eto, — For (aii((ii((«.B.B. alone unaltered. With salt of phosphorus dissolves slowly, glvJDg an iron glsss, which In R.F, is pale yellow on cooing; treated with tin on charcoal It becomes green. Decomposed on fusion with putasalum bisulpbate la the platinum spoor, and gives on treatment with dilute hydrochlorlc add a yellow solution and a heavy wbl e powder, which, on addition of metallic zinc, assumes a smalt-blue color; on dilution with water the blue color soon dlsappean. ColTimbite. wlien decomposed by fusion with caustic potash, and treated with hvdruciloric and sulphuric acids, gives, on the addition of zinc, a blue color more lasting than with tnttalte. Partially decomposed when the powdered mineral Is evaporated to dryness with concenlrnted sulpliurlc add. Its color is changed to while, light gray, or yellow, and vfer boiled with hydrocblortc acid and metallic Bine it gives a beautiful blue.

Obs — ColumbilB occurs at Ribenateln and Bodenmala. Bnvnrla. in granite; Tammela. in Finland: Chnnteloube. near LImogea. in (le^matite wlili tntnlile; near Miask, In the Ilmen Mis., with samarskite; in the eold-washings oT the Sanarka region in the Ural; In Oreeiland. In cryolite, at Ivlgtili (or EvigtokV In hrllHant crystals.

in the United aiates, Id Maine, at Standfsh. in splendent cryalals in gmnlle; also at

Stoneham with cassiterite, etc. in N. Bismuth. at Acworth, at the mica mine. In Mass. at Chesterfield; Northfield, Mass. In Connecticut, at Haddam, in a granite vein; near Haddam: at Branchville, Fairfield Co., in fine vein of albite granite, in large crystals and aggregates of crystals, also in minute translucent crystals (manganocolumbite). upon monomene. In H. Tork, at Greenfield, with chrysoberyl. In Fenn., Hioeml Hill, Delaware Co. Id Virginia, Amelia Co., in fine Bplendeit cnsUtli with microcline, etc. Id JV. Carolina, with Mmsrabit at the mica mine of Milcllell Co. Id Colorado, in microcline at the Pike's Peak region; Turkey Creek, Jefferson Co. In S. Dakota, in Ube Black Hills region, common in the granite veins. In California, in Eio'g' B Cnek dLBtr., Fresno Co.

Manganotantalite (Nordenskiöld) from Utö, Sweden, occurs with pettite, lepidolite, microcline, etc. Hanganotantalite (Arzruni) is from gold-washings in the Buivka region in the Ural.

Manganotantalite occurs in Finland, in Tammela, at HsrlcBsaari in Mr Torro; in Ktmlto, at SkogbÖle; in Somero at Eaidiuuo, and in Euortane at Katlala, with lepidolite, tourmaline, and beryl; in Sweden, near f^luD, at Broddbo and Fliibo; in France, at CLanteioubes near Limoges, in perniatite. In U. S. In Tazcey Co., N. C.; CoomCo., Ala.; also in the Black Hills. S. Dakota.

SzoaOLiTK is essentially FeTaO₄, like normal tantalite, but occurs in prismatic crystals of disphenoidal form the prism is near that of samarskite. From HfitkBatari in Tammela, Finland also from SkogbÖle in Elmto. Ixiolite, from Skogifile, is a niobio-tantalate of iron and manganese: also contains tin; reliability doubtful.

Tapiolite. Fe(Ta,Nb)₂O₇, Like tantalite, but occurring in square octahedra. G. = 7.496. Color pure black. From the Eulmala farm, Tammela, Finland.

Orthorhombic Axes a:b:c = 0.5413 : 1 : 1.330. Crystals prismatic, mm = 56' 50'.

TTTOTAHTAU.

Cleavage: very indistinct. Fracture email conchoidal. H. = 5-5*5. G. = 5'5-5'9. Luster submetallic to vitreous and greasy. Color black, brown, brownish yellow, straw-yellow. Streak gray to colorless. Opaque to sub-translucent.

Coup.—Essentially R₂(Ta,Kb)₂O₇ + 4H₂O, with R = Fe, Ca, H = Y, Er, Ce, etc. The water may be secondary. Analysis by Rasmussen:

Ta₂O₅, Nb₂O₅, WO₃, SnO₂, TiO₂, Er₂O₃, Ce₂O₃, UO₂, FeO, CaO, H₂O. O. = 54.36, I 46.25, 12.83, 3.56, 11.2, 10.02, 6.71, 3.23, 1.61, 8.00, S-78, 6.31

[= 98 95

The so-called yellow yttrantinite of Ytterby and Elrarfvet belongs to fergusonite. Obs.— Occurs in Sweden at Ytterby, near Vaxholm, in red feldspar; at Finbo and Broddbo, near Falun.

SAMARSKITE,

32, Orthorhombic. Axes $a:b:c = 0.5456 : 1 : 0.5178$. Crys-

tals rectangular prisms {a, b}, with c (101) prominent. Angles, $mm' = 57^\circ 14'$; $cc' = 87^\circ$. Faces rough. Commonly massive, and in flattened embedded grains.

Cleavage: b imperfect. Fracture conchoidal. Brittle. H. = 5-6. G. = 5.6-5.8. Luster vitreous to resinous, splendent. Color velvet-black. Streak dark reddish brown. Nearly opaque.

Comp.—R, R, {Nb, Ta}, O, according to Rg., with R = Fe,

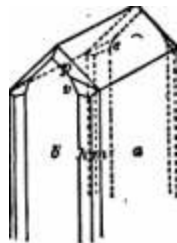
Ca, UO, etc.; R= cerium and yttrium metals chiefly. Analyses by Rammelsberg:

3. Ta. O. N. J. O. 8dO, WO, UO, Cf, o, * T, o, Er, o, FeO TiO,

N. Carolina 5-839 u'si 4107 016 lOW 337 610 10-80 14-61 0-5ff> = 100-98

Miask 5-673 f — 56-84 033 11-94 488 8-80 8-83 14-80 108 = 99-88

• Inol. EUtOfl-atO, ■■ Idcl. SIO,.



.dbyGoogle

NIOBATES, TANTALATES.

.. _ _ 11 emerald-

graeii bead. With todii yields a nmngsueBe reoclon. DecuDipoBed oo fusiim with poiasaium blsulpbate, yieldiag ik yelluw uuaa which oo treatoieit witli dilute hydmcbloric ftcid eeparalea while laotalic acid, and un boiling with metuUic zluo gives a Uae blut color. Id powder sti&lcieilly decoiipoEeil od boUiug with coDcentrated Bulphuric acid to give the blue rerluctlno test when ibe acid fluid is treated with metallic zinc or liu.

Obi.—Occurs !□ reddish brown feldspar, wilh eescliynile aud columbite !□ the Ilmeu Diounimtis. near Miask. In the United Stulea, rntber sbundunt and sometimes in large DiMses, up to 30 lbs., ut the mica mluee in Stillchell Co., N. Carolina, iDtImalely sasocluded ■ -with columbite; sparingly slsewhere.

AnnarlSdlU. Esaenially a pyro-siobate of unuluinand yttrium. lu prismatic crystals. ofK'n resembling columblte. E. = 6. o= o'7. Color black. From the pegmatite vein at Aunerddi uearHoss, Norway.

Hlolmito. A Stan no-tan ts late (and nfobate) of yttrium. Iron, mauganese, calcium. Crystals (orthorhombic) usually rough; massive. O, = 6'88, Color pure black. From the EfirwfTet mine. Fulun, Btveden.

JEiohyntlts, A niobate and litauaie (thoraI«) of the cerium metals chleUy, also In smal'i amount iroa, calcium, etc. Crystals prismulic. oitliorhombic Fracture small conchotdat-

From HiLtrO, Norway. In gn.i>iiewlih gadolinite: at SWitSkra. Smftlaod. Sweden. In the U. Slalea, in N. Carolina, in the gold-washinm on Davis Innd. Henderson Co. with zircon, monazite, xenolime. magnetite; also in 8. Carrollua. four miles from Marietta in Greenville Co. Named from xokvi, many, and Kfidtrti, mialurt.

.dbyGoogle

DESCBIPTIVB UIKEBALOQY.

Oxygen Salts.

4. PHOSPHATES, ARSKBTATES, VANADATES, AIVTIMONATES.

A. Anhydrous Phosphates, Arsenates, Vanadates, Anitmonates. Kormal phosphoric acid ia H_3PO_4 , and consequently normal phosphates

have the formulas E_3PO_4 , $B_3(PO_4)_3$, and RPO_4 , and similarlj for the arsenates, etc. Only a comparatively small number of species conform to this simple formula. Most speciee contain more than one metallic element, and in the prominent Apatite Oroup the -radical (GaF), (CaCl) or (PbCl) enters;

in the Wagnerite Group we have similarly (RF) or (ROH).

Tetragonal Axis i =: 0-6187, a 33'.

930.

$z' (111 A 111) = 55^\circ 30'$, $zz'' (111 A III)$ In crystals resembling zircon in habit; \wedge , $\wedge\wedge$ sometimes compounded with zircon in parallel position

$/Wty^\wedge /fXX$. (Fig- 314- p. 131). In rolled grains. $\langle \wedge \rangle - /1 fV^\wedge\wedge\wedge$ Cleavage: m perfect. Fracture uneven and splintery. Brittle. H. = 4-5. G. = 4 i5-4 56. Lnster reainons to vitreons. Color yellowish brown, reddish brown, hair-brown, flesh-red, grayish white, wine-yellow, pale yellow; streiik pale brown, yellowish or reddish. Opaque. Optically 4-

Coinp.—Essentially yttrium phosphate, YPO_4 , or

*<:s \wedge >y VS- \wedge Y,O..P,o. = Phosphorus pentoiide 38 6. yttria 61-4

$\wedge \setminus X^\wedge \wedge V^{**\wedge} \sim 'oo$, The yttrium metals may include erbium in

large amount; cerium issometimes present; also silicon

and thorium as in monazite.

Pyr.p «to,—B.B. Infusible. When mnlsreneil wllb sulphuric ncl d colors tbe flftme bluish

^reeu. Dllflciilly mliib'e In snlt of phtHplionis. Ingoltihle in ncids,

Diff—Resembles zircon in its (etnigonal form, but disUiiguished by inteHor Imnioen and perfect prismntic cleavage.

Obi.—Occurs as an accessory mineral in grinite veins: somelimes in minute embedded crystnla genenillj distributed in RTnaitic nnd gneiasnld rocks. Fnund at HilterQ; nt Hnss, KragerS. anil from pegmatite veins (it ntber points in Norway; nt Ttterby, Sweden; the FIBitt Berg. S.W. from St. Golhard and tbe Binnenthal. Switzerlaud. An accessory cod-Etituent in the miiBcoTite-graniles of Brazil.

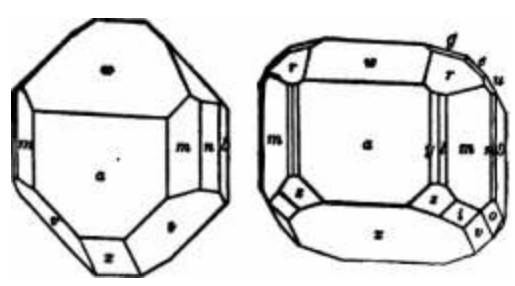
Id the United States, in the gold washings of Ctorksville, Georgia; In N. Carolina, Burke Co.. Henderson Co . Mitchell Co.; In brilliant cry-tals in Alessnder Co. wilb nitiie, etc.; with tysoDite near Pike's Peak, Colorado; rare on New York laland.

.dbyGoogle

PHOSPHATES, ABSKKATES, ETC,

495

933.



Norwich, Ct.

BwltUTlMld.

BSOHAZITXI. MoDocitnic. Axes a-A: 6 = 0-9693 :1: 0-9256j fi = 76° 20'.

mm", 110 A 110 = 88° 84'.

OM. 100 A 101 = 8»° I2i'.

a'x. 100 A 101 = 58* 81',

fttr", on A Oil =88*56'.

rr". Ill A lil=i»o"40'.

vtf, ill A iil =78*19'. Crystals commonly small, often flattened || a or elongated | axis b; eome-times prismatic br extension of V (111) (Fig. 334, p. 103)i uIbo large and coarse. In masses yielding angular fragments; in rolled

Cleavage: c sometimes perfect (parting?); also, a distinct; h difficnl; sometimes showing parting { c, m. Fracture conchoidal to uneven. Grittle. H. =■ 5-5'5. G. = 4-9-5'3; mostly 5'0 to 5"2. Lnster inclining to resinous. Color hyacinth-red, clove-brown, reddish or yellowish brown. Snbtransparent to Bubl ran sin cent. Optically+, Ax. pi J. 6 and nearly [a. Bi, A<' = 4-1° to 4°. Dispersion p < v weak; horizontal weak. SE, = 29° to 31°.

Coup.—Phosphate of the cerium metals, essentially (Ce,La,Di)FO,.

Host uiitlfBes show ibe pre«eDce of TbOi luyi] SiOi, usually, but not alwayt, Id thft

«roper nmouDt lo fonn lUoiium silicate; thiit this U nriehaulcsllj present is not c«rlBln ut possible, lypcal Baalyeea: 1, Blomatnodi 2, Peofield.

G P.O. Ce,o, La,o, TrfErj.O. 810, TbO,

1. Arandal SIS 37-65 39-30 26-20 882 I'86 057 X3'84 = 100 60

2. Burke Co., N. C. 610 2B-2S 8188 8088 — 1-40 S'49)gDo-30= 9968

X = Fe.O,118. CaOo 89. H,Oo 53.

_ rt.. rto.—B.B. Inl the ^

coolfuR: s satunlecl be«d becomes eQamel-white od flaming. Difficultly soluble luibydra chloric SClD.

Ob«.—Ratber ubundnntly distributed as an accesaoiy coDStltueul of goelsoid rocks ia reruln re^oos, thus Id Norlb Cnrolinii and BrHZil. Occurs uear Zlatiuisl In the Ilroen Hts.,ii) ^niDlte. In Norway, near Arendnl. and at AnnerOd. In brisII yellow or brown crystsla (tuTTierite) in Dsuphine sod Switzerland., Foued also <q the gold wasliingsof Aoioqnia: in the diamond gravels of Brazil.

In the United Stales, formerly found with the sl]IImanite of Norwich, Conn.: also at Yorktown, N. Y. In laree coarse cn'stnls nnd mauet in albitic ^ninlio with nilcn)liie. i-lc, at Amelia Court-Ho use, Virginia. In Alexander Co., N. CarollnA, in splendent cryaislsi in Hitchell, Hadison, Burke, and McDowell counties, obtained in large quautiies in rolled groins by washing thie eravels.

Monazlle Is named from /loyaf^fir, to be toltitarf, In allusion lo its rare occurrence.

Cryptolita occurs in wine-yellow prisms and grains in the grsen and red apulile of Arendal,

Norway, and is discovered on putting the apatite in dilute nitric acid. It is probably monazite.

Beizellite. R,As,O , ($R = Ca, Mn$). Isometric, usually massive. $G = 4.08$. Color .dbyGoogle

Monimite. An antimonate of lead, iron, and sometimes calcium; in part, R,Sb,O , usually in octahedra; massive, lustrous. $G = 4.88$. Color yellowish or brownish gray. From the Haratig mine, Pujaberg.

Orminite. Perhaps Pb,As,O . In clusters of fine needles; also in spheruloidal forms. $G = 4.106$. Color carmine to lilac-red. From the LuUe mine at Horhausen.

Pucheranite. Bischofite. A small orthorhombic crystal. $H = 4$. $O = 24$. Color reddish brown. From the Puelz Mine, Schneeberg, Saxony.

Triphylite Group. Orthorhombic

Triphylite $Li(Fe, Mn)PO_4$ 4348 : 1 : 5865

Lithiophilite $Li(Mn, Fe)PO_4$.

Satohite $NaMnPO_4$.

Orthophosphate of (an alkali metal, lithium or sodium, with iron and manganese).

TRIPHYLITE-LITHIOPHILITE.

Orthorhombic. Axes $a:b:c = 4348 : 1 : 5265$. Crystals rare, usually coarse and faces uneven. Commonly massive, cleavable to compact.

Cleavage: c perfect; b nearly perfect; a interrupted. Fracture uneven to subconchoidal. $H = 4.5-5$. $G = 3.43-3.56$. Luster vitreous to resinous. Color greenish gray to bluish in triphylite; also pale pink to yellow and clove-brown in lithiophilite. Streak uncolored to grayish white. Transparent to translucent.

Comp., Tar.—A phosphate of iron, manganese and lithium, $Li(Fe, Mn)PO_4$, varying from the bluish-gray triphylite with little manganese to the salmon-pink or clove-brown lithiophilite with but little iron.

Typical triphylite is $LiFePO_4$. = Phosphorus pentoxide 45.0, iron protoxide 45.0, lithium 9.5 := 100. Typical lithiophilite is $LiMnPO_4$, = Phosphorus pentoxide 45.8, manganese protoxide 45.1, lithium 9.8 = 100. Both Fe and Mn are always present.

Pyro., etc.—When heated in a closed tube sometimes decrepitate, turns to a dark color, and gives

oatriioes of water. B.B. fuses at $\backslash H$, coloring the flame beauiiifil lilia-red In streaks. with a pale blutob green on tlie exterior of ibe cone of flame. With the fluxes reacU for iron and manganeae; the iron renclion is feeble in pure lithiophilite. Soluble in hydrochloric acid.

Q\ a—Trip\ yIH« Is often assocnipd wl'h mwdunipnp; occurs a' RnbenstHn. near Zwiessel, in Bsvftrfa; KellyO. Fliilanrt; No»wich Mas*. ; Ppru. Me : Gmfto", W. H Nnmed from rpii, thretfoid. and i/iuItJ. ffmilj/. in nllusioo lo Us contalninff tbr'e phoanhaio^.

LMiiiipMtiU occurs at BmnchTllle, Fairfield Co . Conn., In a vein of nLMHc granite, with spodiimene, mans^ nese phosphates, etc.; also at Norway, Me. Named from lithium and ipi\oi, friend.

Natrophillt*. $NaMnPO_4$. Near triphyllite in form. Clileflv massive. elpavaWe. H. = 45-5. G. = 8-41. Color deep wine-yellow. Occura sparingly at Bmnchville. Pair-field Co., Conn,

Beryllioilte. A phosphate of ■odium and hprylllnm, $NaRePO_4$. Cryntals abort prismatic to tabular, oiltaorliombic. H, = 5'5-A. G = SM-I. Luster vitreous: on e pearly. Colorless to white or pale yellowish. From Sloneham, Maioe.

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PHOSPHATES,

497

Apatite Group.

General fomnla $R.(F,Cl)[\{P,Aa,V\}O_4]_2 = (R\{F,C1\})R_2[\{P,A8,V\}O_4]_2$;

Apatite $(CaF)Ca_3(PO_4)_2$, Flnor-apatfte " <! = 0-7346

or $(CaCl)Ca_3(PO_4)_2$, Chlor-apatite

Pyromorphite $(PbCl)Pb_3(PO_4)_2$, 07363

Himetite $(PbCl)Pb_3(AsO_4)_2$, 0-7234

Tanadinite $(PbCl)Pb_3(VO_4)_2$, 0-7183

lex, tlicre are alsu cerlain iutermidiate compounds conlaln-ill) pbospLoriis ami arsenic, or anuQlcaod Tanadiuin, ua re culcium nvaieiate, Svnbittt, also seeras to belong In this

III iidditiOH to llic ubove api lug lead ftiid calcium 1 others noted beyond. Ftirlher the i group.

The species of the Apatite Group cryBtallize in the hexagonal system, but nil show, either

by the Biibordinnte facca, or in etching-figures, that they belong to the pyramidiil group (p. 71). They are chemically phosphates, arsenates, vanadates of calcium or lead (also manganese), with chlorine or fluorine. The latter element is probably present as a univalent radical CaF (or Cl), etc., in general RF (or RCl), replacing one hydrogen atom in the acid $R_3(PO_4)_2$, so that the general formula is $(RF)R_2(PO_4)_2$, and similarly for the arsenates. This is a more correct way of viewing the composition than the other method sometimes adopted, viz., $3R_2(PO_4)_2 \cdot RF_2$, etc.

APA'ITfU. Hexagons-pyramidal. Axis $i = 0.7346$.

933. 934 93G 936.

^^

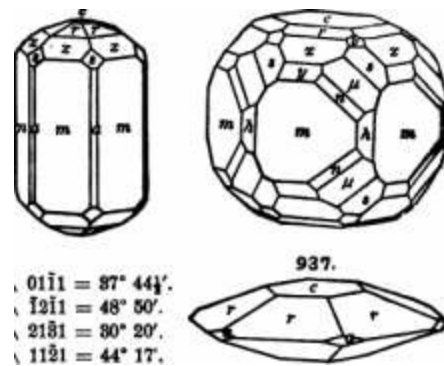
f^S>

fr, $0001 \wedge 1013 = 22' SV$.

fx. $0001 \wedge 10\bar{1}1 = 40^\circ 18'$.

/T/. $0001 \wedge 2021 = 56^\circ 36'$.

rr'. $1013 \wedge 011\bar{3} = 23^\circ SI''$.



. $10\bar{1}1 \wedge 011\bar{3} = 23' Mi'$. , $1131 \wedge 131\bar{1} = 48^\circ BC$. (, $1010 \wedge 3181 = 80^\circ 20''$. , $1010 \wedge nS1 = 44^\circ ir$.

Crystals varying from long prismatic to short prismatic and tabular. Also globular and reniform, with a fibrous or imperfectly columnar structure; massive, structure granular to compact.

Cleavages: c imperfect; w more so. Fracture conchoidal and uneven. Brittle. H. — 5, sometimes 4.5 when massive. G. = 3.17-3.33 cryst. Luster vitreous, inclining to resinous. Streak white. Color usually sea-green, bluish green; often violet-bine; sometimes white; occasionally yellow, gray, red, flesh-red and brown. Transparent to opaque. Optically—. Birefringence low, $Oj \wedge 1.6461$, $6j \wedge 1.6417$.

Tip.— 1. Ordinary. Crystallized, or cleavable and gnuislr maBBve. Colorleu to green, blue, yellow, flesh-red. (o) The a»paraguitone, origiQallj from Hurcts. Spulu, is yellowBh green. MoroxiU, from AreDiliiJ, is in greuUh blue and bluish crjlsals. (b) LaruravaiUe in a sky-blue varieiy wiiU iMpig-liiziili in Siberia, (e) FrancoUu, from Wheal Fraucu, near Tavistock, DevonBhire, occura iu BUiall cryalullne stalactitic a^usBes and io minute curviug crjlsliis.

Ordinary apatite w fluor-apatUe. containing Uuorlne often with only a trace of chlorine, up to O'S p. c; rarely chlorine preponderates, and someilmes flunrine Is enilrtly absent

3. MangaTutpatUe contains munguuese replacing calcium to IU'G p. c. HuO; culur durk bluish erecu.

8. Fibrou, eonerttionm-y, ttaUetUie, PhotphoriU iucludes the fibrous cuocrelionar;

i partly scaly mineral from Esremadura, Spain, and elsewhere. EvpvrehroiU, fio "~'nt, N.T
'- ' ■ — " • ---■- ' - -■-

Crown Point, N. T., iwlongs here; it is concentric in airucture. .. _

Ing the phosphorite of Sl^el in bolryoidol, renifonn, or sialactitic'masses, fibroua and radiating. Seep. 469.

i. Earth]/ apatite; OiUolile. Mostly altered apatite; coprolilea are Impure calcium phosphate.

Comp—For Fluor'apatite $(CaF)Ca_3(PO_4)_2$; and for Cklor-apatite $(CaCl) Ca_3(PO_4)_2$; also written $3Ca_3P_2O_8 + CaF_2$ and $3Ca_3P_2O_8 + CaCl_2$. There are also intermediate compounds containing both fluorine and chlorine. The percentage composition for these normal varieties is as follows:

Fluor-apatilB P₂O₅ 42.8 CaO 55.5 F 8.8 = 1016 or Ca₃P₂O₈, 82.25 CaF₂, 7.75 = 100 GhUnr-
apaliU P₂O₅ 41.0 CaO 58.8 Cl 68 = 1018 or Ca₃P₂O₈. 8»-4 OaCl, 108 = 100

Fluor-apatite !s much n the Alps, Spain, St. Lawre iuent are rare; Ibis is true of some Norwtgian kinds,

Pyr., etc.—B.B. in the foreeps fiiues vilh difficulty on llie edges (F. = 45-5), coloring the flame rediiiBb 3-ellow; moistened with sulphuric acid and heated colors lbe flame pale bluish green (phosphoric acid); some varieties reatt for chlorine wiih salt of phosphorus, when.the bead has been pveriouBly saturated with copper oxide, while others give fluorine when fused with this salt in an open glass tube. Otves a phosphide with the sodium test. Dia'wlvcB in hydrochloric and nitric acids, yielding with sulphuric acid a copious precipitate of calcium sulphate; the dilute nitric acid solullon gives with lead acetate a white precipitate, which B.B. tin chnrconi fuses, giving a globule wlih crjstnlloes

facets on cooling. Some varieties of apatite phosphoresce on heating.

131ff~ChHraclerfzed by the common hixngonul form, but softer than beryl, being scratched by a knife: does not eflin^rvcsce in arid (like cnlcite); difficultly fusible: yields a green flame B.B. after hein^ moistened with sulphuric acid.

Recognized in thin sections by lis mndirately high relief; extremely low birefringence (hence not often showiuR a dislnct axial flgure in basal sections}, llie Inlerference colors in ordinary sections scaroely rising above gray of the first order; parallel extlncUon and negative extension; columnar form: lack of color and cleavage; and by the rude cross parting seen as occasional cracka crossing the prism,

Oba.—Apatite occurs In rucks of vai-ious kinds and ages, but Is most common in metamorphlc crystalline rocks, especially in granular limestone and in many metalliferous veins, particularly those of tin. in gneiss, syenite, hornblendlc gneiss, tulca schist, beds of Iron ore; occaslonRlly In RPrpentine. In tlie form of minuie microscopic crystals it has bd atmnst universal distribution as a rock-forming mineral. It is found in ail kinds of igneoiaa rocks and Is one of the eariist producW of crystallization. In Inrger cryslalB it isespeciully characteristic of the pegmatite fiicies of igneous rocks, psrticularly the graniles, and occurs there associated with qiiariz. feldspar, tourmaline, muscovite. beryl, etc. * '■ ■"•"—■—— prweot In ordinary sttatified limestone, beds of sandstone - -■- '- ~' •'- iferoua, Jurassic. Cretaceous, or Tertiary. "' "" ' ' of wood. . L TT

of wood. . L TT

Among Its localities sre Ehrenfrledersdorf in SBxony: BchwarzenBlem, the Knappeo- ■wand In Unteranzbachlhnl in the Tyrol ; Rt- Qothard, Tavelscb. etc.. in Switzerland; MuBsa-Alp in Hedmont, while or colorles': Zinnwald and Schlackenwald n a.hemla: id England. In Cornwall, with tin ores; in Cumberland, al Carrock Pells; in Devonshire, cream-colored a' Bovey Trncpv, «"d at Wlpr.l Franco [fmneoliU). The atparag^uilojit or wpairgtitltin of Juroilla, In Murcla, Spain. Is pale yellowish green in color. Large quantl-

;CoogIc

ties of ap&tftte are mined Id Noiycay at EnigerO; also at Odegurd, near Bamle, and elsewhere.

In Maine, on Long IslaQd. Blue-hill Bny. In JT. Hamp., WeslmorelHiid. !□ Mau., at Norwich; at Bolton ubundau. Id Conn., at Bniuchvllle (manganapatiu), aitia greeolBh witite and colorleas. In Nme York, commoa iu Bt. Luwreuce Cu., in grauukr Ittiiestoue: nl.to Jefferson Co.; Stuidford mine. Eiibl Horialj, Essex Co.. in magtielltt; Dear Edeovtlli.; Orange Co. In Penn., at Lelperville. Delaware Co.; iu CLeater Co. In N, Carotina, at Slouy Point, Alaxander Co., etc.

In extensive beds in the Laurentian gneiss of Canada. usually associated with Juraestoué, and associated by pyroxene, amphibole, lithium, zircon, garnet, vesuvianite and many other species. Prominent mines are in Ottawa County, Quebec, in the townships of Buck-ingham, Templeton, Portland, Hull, and Wakefield. Also in Itenrew county, Ontario, and in Lanark. Leeds, and Frontenac counties.

Apatite was named by Werner from the district of Ivry, in the province of Auvergne, having referred it to the amphibole, arctite, Urt, scoria, etc.

Besides the above mineral phosphate. Including normal apatite, phosphorite, etc., there are also extensive deposits of amorphous phosphates, consisting largely of "bone phosphate" (Calcium phosphate), of great economic importance. though not having a definite chemical composition and hence not strictly belonging to pure mineralogy. Here belong the phosphate nodules, coprolites, bone beds, guano, etc. Extensive phosphate deposits also occur in North Carolina, Alabama and Florida. One is a bone phosphate of lime, mixed with siliceous hydrous phosphate. and generally with some calcium carbonate, and often a little magnesia, alumina, iron, silica, gypsum and other impurities.

Staphelite. a carbonated calcium phosphate. Occurs encrusting the phosphorite of Staveland. In bituminous shaly masses, nodules and druses; It is the result of the action of carbonated waters. U. = 4. Q. = 8 123 Color black to dark green, greenish yellow. Duhren, from Barale. Norway, is similar.

PTROHORPHITE. Green Lead Ore. Orthorhombic form. Hexagonal-pyramidal. Axis $a = 0.7362$.

Crystal prismatic, often in rounded barrel-shaped forms; also in branching groups of prismatic crystals in nearly parallel position, jagged

tapering down to a slender point Often globular, reniform, — ■

and botryoidal or verruciform, with usually a subcolumnar structure;

also fibrous, and granular.

Cleavage: $\{100\}$ in traces. Fracture uneven, conchoidal, 'uneven.

Brittle. H. = 3-5-4. G. = 6-5-7-1 mostly, when |

pure; 5-9-5, when containing lime. Luster resinous. Color green, yellow and brown, of different shades; sometimes wax-yellow and fine orange-yellow; also grayish white to milk-white. Streak white, sometimes yellowish. Subtransparent to translucent.

Optically —. t

Var. —1. Ordinary, {a} / « erythrite as described : sometimes yellow and in rounded forms

resembling camnlyite {p» «udo-campylHe). (b) In nekular and WHM»-(»t< aggregations. Id) (»nflj'e(i«n"-y groups or masses of crya-tnis, linving ti>« surfuce angular, (rf) FiSroxu. (e) Granular mauit». (/) Earths ,* inctusiig.

3. Foiytpharite. Containing lime: color brown of different shades, yellowish gray. pale jellow to nearly while; streak white; G. = 5-89-e-44. Rarely in aeprate cryalala; usually ill groups, globuliir. miimmillnry, verruciform. Mietite. from Mies in Bohemia, is a brown variety. NumieriU is similar and impure, from NiisslSre, near Beaujeu, France; color yellow, greenish or grayish; G. =6-043. 8. CliTamiferous; color hrillinnt red and orange. 4. Arien\f»rou»; color green to white; G. = 55-88. 5. J^eudotnorphmit; (a) after galena; lA) cenissite.

Cemp.—(PbCl}Pb.(PO,). or also written 3Pb,P,o..PbCl = Phosphoms pentoxide 13-7, lead protoxide 82'3, chlorine 2-6 = 1005, or Lead phosphate 897, lead chloride 103 = 100.

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TUe plinspbonis is often replncert by arsenic, and as the amooDt Increases the species puiMS iuiu iiiiiiuiilt!. Culclutn also tepliLces ibe lead to a cuDsiderable extent.

Fyr.. etc.—lu Uiu closed tube ^ives a wjlte sublimate of lead chloride. B.B. la the forci-ps fiis.s easily (F. = 1."!), colmiig the flame bluish greeu; on charcoal fuses without reduction lu » globule, wblcbon cooling assumes a crystalliue polyhedral form, nhile the coal is coined wbiie from cbloride and, uearer the assay, yellow from lead oxide. With stiila iin charcoal yields inetidlic lend; some varielies contain arsenic, and give the odor of garlic iti K.F. ou charcoal. With salt of phosphorus, previously saturated wltli copper .oxhie, (jivcs iiD nzure-blue color to the flame when treataid in O.F. (chlorine). Soluble in uitic mid.

Diff,—Distinguished by its hezagoDul form; high specific gr&vityi resinona luster; b!ow[Hp4! cliariiciers.

Obs.—Pyromorjiblte occurs principally In.veins, and accompaues other ores of lead. Al Puulluoueu nud Huelgoel in BriUany; at Zschopnu and other places in Saxony; Ht Pfihrnni, Bldsudt. in Bohemia; In fine cryslab at Ems, Braubach, in Nassau; also «t Dernbiich in Nassau; Berezov in Siberia; in Che Nerchinsk mining district; Cornwall, grt-en aud bronii; Dtvon. gray; Derbyshire, green and yellow; Cumberland, golden yellow, in EiiEluud; Leudhills, red and orange, lu Scntlitud.

lu the v. S.. has been found very fine at PheDixTille, Pa.; also In Maine, at Lubec and Lenox; In New York, a mile south of Sing Sing; in Davidson Co.. N. C, also iu Cabarrus and Caldu-ell Cos.

Named from itvp, fire, nofiipij, form, alluding to the cryslallfne form the globule assumes on cooling. This species passes Into mImelile.

HIMBTm.

Hexagonal-pyramidal. Axis $a = 0.7324$.

Habit of crystals like pyromorphite; sometimes rounded to globular forms. Also in micaceous crusts.

Cleavage: $a(1011)$ imperfect. Fracture uneven. Brittle. $H = 3.5$. $O = 70-75$. Luster resinous. Color pale yellow, passing into brown; orange-yellow; white or colorless. Streak white or nearly so. Subtransparent to translucent.

Var.—I. Ordinary. (a) In rounded aggregates. (6) Capillary or fibrous, especially marked in a variety from St. Prix-sous-Beuvray, France; somewhat like asbestos, and straw-yellow in color, (e) Concretionary.

Campylite, from Drygill in Cumberland, has $Q = 7218$. and is in barrel-shaped crystals (whence the name, from *Kittling*, *euTwd*), yellowish to brown and brownish red; contains 3 p. c. PbO .

Comp.— $(PbCl)Pb_3(A_8O_5)_2$, also written $3Pb_3A_8O_{15}PbCl_2$ = Arsenic pent-oxide 23.2, lead protoxide 74.9, chlorine 3.4 = 100.5, or Lead arsenate 90.7, lead chloride 9.3 = 100.

Phosphorus replaces the arsenic in part, and calcium the lead, Enderskite (p. 501) is intermediate between mimetite and vanadinite.

Fyr., etc.—In the closed tube like pyromorphite. B.B. fuses at 1, and on charcoal gives. In R.F. an arsenical odor, and is easily reduced to metallic lead, coating the coal at first with lead chloride, and then with arsenic trioxide and lead oxide. Gives the chlorine reaction as under pyromorphite. Soluble in nitric acid. ■

Obs.—Occurs near Riddulph and elsewhere in Cornwall; Beer Alston, Devonshire; in Cimolite; near Pontgibaud, Puy-de-Dôme; at Johannegeorgensliedl. in line yellow crystals; at Nerchinsk, Siberia; at Zinnwald; Långbau, Sweden; at the Brookdale mine, Pheasantville, Pa.

Named from *imitator*. It closely resembles pyromorphite,

VANADINITE.

Hexagonal-pyramidal. Axis $a = 0.7122$.

Crystals prismatic, with smooth faces and sharp edges; sometimes cavernous, the crystals hollow prisms; also in rounded forms and in parallel groupings like pyromorphite. In implanted globules or inclusions.

Fracture uneven, or flat conchoidal. Brittle. H. = 2-75-3. G. = 6-66-7*10. Luster of surface of fracture resinous. Color deep ruby-red, light

brownish yellow, straw-yellow, reddish brown. Streak white or yellowish. Subtranslucent to opaque.

Comp.—(PbCl)Pb.(VO₅), also written 3Pb,V,O..PbCl, = Vanadium pent-oxide 19-4, lead protoxide 78-7, chlorine 2'5 ~ 1006, or Lead vanadate 90-2, lead chloride 9-8 = 100.

Phosphorus is sparingly present, also sometimes arsenic, both replacing vanadium. In endliche the ratio of V : As = 1 : 1 nearly.

Pyr., atc^ID the closed tube decrepitate and yields a fluffy white sublimate. B.I.J. fuses easily, and on cooling to a black lustrous mass, which in R.F. yields metallic lead and a coating of lead chloride; after completely oxidizing the lead in O.F. the black residue even with heat of phosphorus an emerald-green bead in R.F., which becomes light yellow in O.K. When the chlorine reaction with the copper test. Decomposed by hydrochloric acid.

Obs.—First discovered at Zimapan in Mexico. Later obtained at Wanlockhead in Dumfriesshire; also at Berezovo in the Ural, with pyromorphite; and near Eppel in Carinthia, in crystals; at Udensjö, Blekinge, Sweden. In the Sierra de Córdoba, Argentine Republic

In the U. States, sparingly near Sing Sing, N. Y. Abundant in the mining regions of Arizona and New Mexico, often associated with wulfenite and desclozite; in Arizona, in the mines of Yuma Co., in brilliant deep red crystals: Vulture, Pinal Co., etc. in Maricopa Co.; the Mammoth gold mine, near Oracle. Pinal Co. In New Mexico, at Dike Valley, Sierra Co. and the Mimbres mines near Georgetown.

Habit. From Freiberg, Sweden; has ordinarily been included as a calcium variety of wulfenite, but is now made monoclinic. Massive, cleavable. Color yellowish white.

Structure. A calcium arsenate, related to the species of the Apatite Group. Crystals hexagonal prisms; colorless; $d = 0.7143$. H. = 6. G. = 8-53. From the Harstig mine, Pajsberg, Sweden.

Crystallographic Group. Monoclinic.

Formula

Wagnerite (MgP)MgPO₄, 1-9145 : 1 : 1-5059; 71" 53'

Tripbte (Rr)RPO,, R = Fe : Mn = 2 :1,1: 1, etc.

■ftipbidite (ROH)RPO., R = Mn:Fe = 3:l 1-8572 :1 : 14925; 71M6'

AdeHte {MgOHjCaAsO. 2-1978 : 1: 1-5642; 73" 15'

Tilasite (Mgr)CaA8o.

Barkinite (MnOH}MnABOj 2-0017 : 1: 1-5154; 62° 13^'

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Phoepbates (and srsenates) of maEneaium (calcium), iron and mangaceBA containing flnonne (also hydroiyl). PormnU B,FPO, or (B1")RPO,, etc.

WAGHEBmi.

Mouoclinic. Axes, see p. 501. Crystala Bometimes large and coarse. Also massive.

Cleavage: a, m imperfect; c in traces. Fracture uneven and splinter;. Brittle. H. = 5-5-5. G. = 307-3-14. Luster vitreous. Streak white. Color yellow, of different shades; often grayish, also flesh-red, greenish. Translucent.

Comp.—A fluo-phosphate of magnesium, (MgFjMgPO, or Mg,I\o,.MgP, = Phosphorns pentoiide 43'8, magnesia 49'3, fluorine ll'S = 1049, deduct (o = 2F) 4-9 = 100. A little calcium replaces part of the magnesiain.

PjfT., oto,—B.B. ill the forceps fuses al 4 to ngreeiish-gray slasa; iitoisteDed with ■iilpliiric acid colurs the flame bluisli green. With borax leacle for itdd. Ou fusion with soda efFervesci^, but is uot completely dissolved; g\jea a faint maugauese reactioo. Fused with Bait ut pbnspliorus in an open gliua lube reucts for fluorine. Soluble in nitric aud hydrochloric acids. Witli sulphuric acid evolves fumes of liydrofluoric acid.

Ob*.— WagneriU (tu small bigbly moditled crystals) occurs in the vaJley of HOLLen-graben. neiir Werfen, in Salzburg, Auslrin. E^^~<Ufina (musaivi!, cleavable; auo in coarse crystals) is from KJORreetod, near Bamle, Norway.

Bpodloalts. A calcium fiuo-pbosph'ite, perhaps (CaF)CaPOi. In flattened prismatic crystals. Q. = 294. Color ash-gray. Prom Ihe Kraugrufva, Wermlaud, Sweden.

TRIFUTB.

Monoclinic. Massive, imperfectly crystalline. Cleavage: uneqnal in two directions perpendicular to each other, one much the more distinct. Fracture small conchoidal. H. = 4-55. G. = 3'44-38. Luster resinous, inclining to adamantine. Color brown or blackish brown. Streak yellowish gray or brown. Subtranshicient to opaque.

Comp., Tat.—(RF)EPO or B,PX)..EF, with R = Fe and Mn, also Ca and Mg. The ratio varies widely from Fe : Mn = 1 : 1 to 2 : 1 (zwieaelite); 1:3; 1:7.

TnlktrippiiU is a variety from HorrsJOIwrg; coulaius magnesium and calcium In larga

Pyx., eto.—B.B. fusee easily nt 15 lo a black magnetic globule: moistened with Bulpburic acid cilors Ihc flame lihiisli green. Wltli borax in O.F. gives an smelhytinc-colored gloss (magnuesL'): !□ R.F. a strong reaction for iron. With soda reacts for miugimcse. Wllh sulphuric acid evolves hydrofluoric acid. Soluble in liydroooliloric acid.

Oba. —Found by AUuaud at Limoges iu France; Helslagfors, Finlnud: Stnneham, Miiliie: nrrnchville, Codu. ZurittfUU. a ciuve-browu variety, is from AalwDstein. near Zwlesel in Bavaria.

Griphitb. a problemallcal phosphate related to tHpItle occurring In embedded reni-form uinsses. From (lie Riverton lode near Harney City, Pennington Co.. 8. Dakota.

TriploidltB. Like Irlplile. but with Ihe F replared by (OH). Commniily In crys'alline aggregnten. Fibrous lo columnar. H. = 4-5-5. O. = 3687. Color yel'owish to reddish brown. From Braochville, Fairfield Co.. Conn.

Adellta. (MgOHK^aAsO.. MonocHnic. Axes, see p Ml: also massive. H. = 5. G. = 8'74. Color gray or grayish yellow. Prom Nordmark and LAnghan, Sweden.

Tilasits. Like ndelite, but contains fluorine. From L&ngban. t

SarklnltB. (MnOHLMnAsO,, In monoclinic crystals; al=0 In spherical forms. G, -^ 4 17. Ci'lor rose-red. flesh-red, reddish yellow. From the iron-manganese mlnen of Pajs-berg. SwedpD. Polyantnite from the 'B]6 mine. Grytb^t« parish, Orebro. Sweden, is essentially the same.

,ab,GoOgIc

PHOSPHATES, ARSB1IATE3, &TC. ftOS

Berdvrit*. A fiuo-pboepli&te of beryllium uud calcium, (CaF)BePO,nltb(CBOG)BePOt. In priamiitic crystals, monoclinic wltli complex twimilug. H. = 5. Q. = a-9B-8'01. Luitur viireouB, Color yellowisli uitd KreeuisL wltte. From the tiu mlaeH of Eliren-fiit.'dtirsilurf, Sulouy; also at Stoueliaui, Auburn and Hebrou, HalDS.

Hunllnite. A basic phosphate of alumiuium and atronlium. In colorless riombobedral crysiutii. fl. — 40. Q. — U-16-tt'!8. Occurs wltb berderile, bertiaadite, etc.. at Slone-ham. Maine.

Dar*Dgit«. A fluo-ai crj»iala. Q. = "" ' ""

Triclinic. Crystals large and coarse (Fig. 20, p. 10); (orine rarelj distinct. Usually clearable to columnar and compact masaWe. Polysyntbetic twinning laiuellB common.

Cleavage: c perfect, with pearly luster; a somewhat less so, vitreous; € (o3l) sometimes equally distinct; M (lIO) difficult; ca ^ 75° 30', ce = (4° 40', cM=9-i° iO'. Fracture uneven to snbconchoidal. Brittle. H = 6. G. = 3'0l-3'09. Luster vitreous to greasy, on c pearly. Color white to pale greenish, bluish, jollowish, grayish or brownish white. Streak white. Sob-transparent to traufilucent.

Comp. — A fluo-phosphate of aluminium and litbinm, Li(AlF)PO, or AlPO.-LiF = PhosplioniB peutoxide 47!), alumina 3'14, lithia 10 1, flnorine 12 9 = 105-3, deduct (o = 2F) 5-3 = 100 Sodium often replaces part of the lithium, and hydrosyl piirt of the fluorine.

Pyr., sto —In ttiic HowkI tulie yields water, whicli nt a high brat Is acid and corrodes thu glass. B.B. riises eaaily (iit i) with liitumesceuce, and becomes opuque while on c<>oliug. Colors Ibe fliitae yellowlBli red with trncea of green; the Hebron variety gives nti iuteusa litbia-red: moistened with siilpbiiirio ncl d gives a bluish green lo the tiame. With boniz and suit of phospliorua forms n iraospareiit colorlass glass. In Hue powder ilisBClves easily 111 sulphuric acid, more slowly hi hydrochlorlc acid.

nff —Diatingufahed by Its easy fusibility and by yielding n red flame B.B., from feld-spsr. barite. calcite, etc.; alao by the acid water In the lube ffom apodumene.

Obi.—Occurs near Penig in SuKony; Arendal, Nomay; Moutebras. Creuze, Frauce. In the U. States, iu Maine, at Ilebrnu; also at Paris, Peru, etc.; Branchville, Cono.

The name ambipgonite is from d/i(S^ti, Hunt, and yoyv, anglt.

B. Basic Phosphates. This section includes a series of well-characterized basic phosphates, a number of which fall into the Olivinite Group. Acid phosphates are represented by one species only, the little known monetite, probably HCaPO., see p. 507.

Ollrenlte Group. Orthorbombic.

OUvenite Ca,(OHLAsO, 09396 : 1 : o-6726

Libethenite Cn,{OH)Po. 0-9601 i 1: 0 7019

Adamite Zn,(OH)AsO. 09733 : 1 : o-7158

DeKloizite (Pb,Zn),(OH}VO,

S\l:ii = o'6368 : 1 : o-8045 oT^:hi{: = 0-9552 : 1 : o-8045 Cnprodescloizite (Pb,Zn,Cu),

(OH)VO,

The Olivenite Group includes several basic phosphates, arsenites, etc., of copper, zinc, and lead, with the general formula $(\text{BoH})\text{EPO}_4 \cdot \{(\text{BoH})\text{ItAsO}_4\}$,

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DEaCBIPTITE MINBBALOOT.

etc. They crystallize in the orthorhombic system with similar form. It is to be noted that this group corresponds to a measure to the monoclinic Wagnerite Group, p. 501, which also includes basic members.

OLIVENITE. Orthorhombic Axes $a : b : c = 0.9396 : 1 : 0.6731$

$m = 101$. $n = 67$. $d_{100} = 10.1$ A. $d_{110} = 7.3$ A.

$d_{111} = 6.7$ A. $d_{112} = 5.1$ A. $d_{113} = 4.7$ A. $d_{114} = 3.4$ A.

Crystal habit: prismatic, often acicular. Also globular and reniform, indistinctly fibrous, fibers straight and divergent, rarely irregular; also curved lamellar and granular.

Cleavage: m, b, c (OH) in traces. Fracture conchoidal to

uneven. Brittle. H. = 3, G. = 4.1-4.4, Luster adamantine to

vitreous; of some fibrous varieties pearly. Color various shades of

olive-green, passing into leek-, siskin-, pistachio-, and blackish

green; also liver- and wood-brown; sometimes straw-yellow and

grayish white. Streak olive-green to brown. Sub transparent to opaque.

W. — (n) Oyl («U» ii. (6) F. Siroum; finely and divergently fibrous, of gray, yellow, brown and gray, to white colors, with the surface sometimes teluretic or acicular found (investing the common variety of passing into U; called itood copper or toud a ttena U. (c) Earthy; nodular or massive; sometimes soft enough to roll the Uagers.

Comp. — $\text{Cu}_2\text{As}_2\text{O}_7 \cdot \text{Cu}(\text{OH})_2$, or $4\text{CuO} \cdot \text{As}_2\text{O}_5 \cdot \text{H}_2\text{O}$ = Arsenic pentoxide 40-T, cupric oxide 56.1, water 3.2 = 100.

F. — In the closed tube gives water. B.B. fuses at 2. color blue; the flame bluish

green, and on cooling the fused mass spatters crystalline. B.B. on charcoal fuses with borax, giving off arsenical fumes, and yields a metallic residue which with soda

yields a globule of copper. With the fluxes reacts for copper. Soluble in nitric acid.

Obs.—The crystalline varieties occur in Cornwall, at various mines; Tavistock, in Devonshire; in Tyrol; in Banat; Nizhni Tagilsk in the Ural; Chili. In the U. S., in Utah, at the American Esfie and Mammoth in the B. T. district, both in crystals and in druses. The name alludes to the olive-green color.

Orthorhombic. Axes $a : b : c = 0.9601 : 1 : 0.7019$. $\alpha = 90^\circ$, $\beta = 111^\circ 11'$, $\gamma = 59^\circ 44'$.

Fracture, on a or b = 'o' y. «', m a in =er47j'.

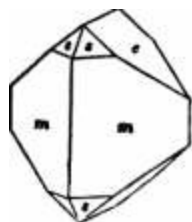
In crystals usually small, short prismatic in habit; often united in druses. Also globular or reniform and compact.

Cleavage: a , 6 very indistinct. Fracture subconchoidal to uneven. Brittle. $H. = 4$. $G. = 3.6-3.8$. Luster resinous. Color olive-green, generally dark. Streak olive-green. Translucent to subtranslucent.

Comp. — $Cu, P, O, Cu(OH)$, or $4CuO.P, O, H, O$, = Phosphorus pentoxide 29.8, cupric oxide 66.4, water 3.8 = 100.

Pyro. etc.—In the closed tube yields water and turns black. B. B. fuses at 8 and colors the flame emerald-green. On charcoal with soda gives metallic copper, sometimes also an arsenical odor. Reduced with metallic lead on charcoal is reduced to metallic copper, with the formation of lead phosphate, which treated in R.P. gives a crystalline polyhedral bead on cooling. With the fluxes reacts for copper. Soluble in nitric acid.

Obs.—Occurs with chalcopyrite at Libelben, near Neusohi, Hungary; at Kleinbilchenbach and Ehl on the Rhine; at Nizhni Tagilsk in the Ural; in small quantities in Cornwall.



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Adamite. $Zn_3As_2(OH)_4$. In small crystals, often grouped in crusts and druses. Aggregate. $H. = 3.5$. $U. = 4.34-4.35$. Color honey-yellow. Tint, rose-red, green. From Cofre de Perote, Chili; Cupquero, FraDcei at the ancient zinc mines of Idrija, Q. recce.

Deidizite. $R, V, O, B(OH)$, or $4RO.V, O, H, O$; $R = Pb, Zn$ chiefly, and usually in the ratio 1:1. In small crystals or in druse; also massive, fibrous and dimorphous with mamillary

surface. U. = S'5. ti. = 5'ft-B'2. Color cherry-red and brownish red, to ligil or dark brown, black. Streak orange lo browuish red or ^ellowbh gray.

From the Sierra de Cordoba, Argentina; Kappel in CannlhJa. Abuudant at Lake Viilley. Sierra Co., New Mexico, also near Oeorgetown; in Arizona near Tombsloue; iu VaTHpai Co.; at tbe Maminolh Gold mine, near Oracle, Final Co,

A IUflsiiiTe variety, contiiiuiug copper (0-5 to.9 p. c), in ciusIb, and reuiform masses with radiateil structure, occurs iu Siii Luis Poiosi, also iu a vein of argentiferous ^lenn in Zncalccas, Mexico; il lias been vnriously named Irilochorilt, evprodetelointe, ramirttt. A diuilnr variety (11 p. c. CuO] uccurs as an Incrustation on quartz at tbe Lucky Cuss mine. Tombstone, Cochise C'>., Arizona.

EcsTNCHiTB may be identical with deacloizite, HoaalTe: hi nodular, staUctltlc forms. G. = 6 596. Color yellowish red, reddish brown, greenish. From Hofsgnind near Freiburg in Baden. The same may be true of armoxtn« from Dahn near Nieder-Bchletteubacli. Rhenish Bavaria,

Dbchknitb. Composillou usually accepted as PbV.Ot. MaasiTe, borryoldal. nodular. G. = 5 6--j*81. Color deep red lo yeliuwish red and brownish rod. From Nieder-Schlet-teubach iu the Lauterthiil. Khenlsh Bavaria.

Oalclovolborthite, Probably (Cu,Ca).ViOi.(Cu,CaKOH),. In thin greeu tables; also gray, fine cryslalline granular. From Friedriclisrode, Thuringla.

Brackebiuchita. Near descloizite (monoclinlcj. From the State of Cordoba. Argentina,

Paittacinite. A vanadale of lead and copper, from the Silver Star District, Hoiitana. In liiiti coatings: also pulverulent. Color siskin- to oKve-green.

MoTTRAMiTii:. A vanadate of lead and copper; possibly Identical with psittaciuite; in velvety black incrustallouB. From Mollram St. Andrew's, Cheshire, EugUnd.

OUNOai.ASrrai. Eilnokloa. Aphanfse.

Monoclinic. Axes A : J : ^ = 1-9069 : 1: 3-8507; j8 = 80° 30'. Crystals prismatic (m); also elonfjated Q i; often grouped in nearly apherical forms. Also massive, hemispherical or reniform; Btructnre radiated fibrous.

Cleavage: c highly perfect. Brittle. H. = 2-5-3. O. = 4-19-4-37; 4-37 Utah. Luster: c pearly; elsewhere vitreous to resiououB. Color internally dark verdigris-green; eiterially blackish bine-green. Streak bluish green. Sub transparent to translucent.

Comp.—Cn,As,o..3ou(OH), or 6CnO.As,o..3H,o = Arsenic pentoiide 30-3, cupric oxide 62-6, water 7-1 = 100. Pyr., etc.—Same as for oliveille.

Obi. - Occurs in Cornwall, with other ores of copper. In Umh, Tintic district, at the Mammoth mine. Niimed in al!ii«Ton to the basal cleavage beina; oblique to the sides of the prism.

Brinlt*. Ca,As,o..2Cu(OH) , In mammillnted rryatallne groups. Color fine emerald-green. From Corn wall; also the Tintic district. Utah.

Dihydrlte. $\text{Ci}^i\text{P,Oi.3Cu(OH)i}$ In dark emerald-green crvslals (monoclinic). $H. = 4-5-^{\wedge}$. $G. = 4-4-4$. Prom Elil near Linz on the Rhine; the Ural eic.

Piaudomalachlte. In part $\text{Cn.P.O. SO}^n\text{OHl- Mnaslve. refembliiiiB}$ malachite In color and structure. From Rhelnbrei ten bach: Ntzhtit Tngilak, etc. BMile Is closely allied.

Ohoadraisente. Perlinpa $\text{MniAS}^|o.,8\text{Mn(OH),}$. In small embedded grains. Color yellow to ri'ddish yellow. Prom the PajsWrg mines. Sweden.

X*WTH*ReBKiTB Nenr fhondrnrsenile. but contains more water. Iu sulphur-yellow grains; majsive. From the 8j& mine, parish of Orylhyle, Sweden.

;CoogIC

So6 DBSOBIPTtTE MINEBALOOT.

DUFBNNTXI. Eraurile.

Orthorhombic. Crystals rare, small and indistinct. Usaally i nodules; radiated fibrous with drnsj surface-Cleavage: a, probably also b, but iudistinct. $H. = 3'5-4$. $G. = 3''2-3''4$. Luster silky, weak. Color dull leek-green, oliTe-green, or blackish green; alters on exposure to yellow and brown. Streak siskin-green. Subtranslucent to nearly opaque. Strongly pleochroic.

Comp.—Doubtful; in part $\text{FePO.,Fe(OH),} = 2\text{Fe,o.,P,o.,3H,o} = \text{Pho8-phoruB}$ pentoxide $2\text{T}'5$, iron sesquioxide $63-0$, water $10'5 = 100$.

Pyr., ate.—Same as for Tivianite, but lesa w&Ier Is giveo out in the closed tube. B.B. fusen easily to a alag.

Oba.—Occureiiiifar Aselar, Dept. of Haute Vieuue, and at Hlrschbei^ lu Westphalia; from the KotblSufclieD mfLe near Waldelrmes; St. BeDlGDa, Bobemia; EasI Curnwall. lu the United States, at Alleotown, N. J.; in Rockbridge Co., Vs., in radiated coenelf

IiAZULmt.

Monoclinic: Axes $A : J : t^{\wedge} = 09750 : 1 : 1-6483$; $/\text{?} = 89^{\circ} 14'$.

M4. at. $100 A 101 = 80^{\circ} 24'$. a^{\wedge} , ill a ill = $80^{\circ} W$.

fpp". Ill A lit = TB" 40'. pa. Ill A 111 = 82* Sff.

Crystals usually acute pyramidal in habit. Also massive, granular to compact, Cleavage: prismatic, indistinct. Fracture uneven. Brittle. H. = 5-6. G. = 3.057-3.122. Luster vitreous. Color azure-blue; commonly a fine deep blue viewed along one axis, and a pale greenish blue along another. Streak white. Subtranslucent to opaque. Comp.—EA],(OH),P,O, or $2AlPO_4 \cdot \{Fe,Mg\}(OH)_2$ with Fe : Mg(Ca) = 1 : 13, 1 : 6, 1 : 2, 2 : 3 (Eg). For 1 ; 2 the formula requires: Phosphorus pentoxide 45.4, alumina 32.6, iron protoxide 7.7, magnesia 8.5, water 5.8 — 100. Pyr., Bto.—In HCl tube whitens and yields water. B B. will cobalt solution the blue color of Ibe mioenil U restored. In HCl forte pH » whitens, turns open, swells up. and without fusion fuses to a glass, coloring the flame bluish green. The green color is made more intense by reduction; the assay with sulphuric acid. With the fluxes gives an iron glass; with soda on charcoal in infusible mass. Reduced upon yields, retains perfectly its blue color.

Obs.—Occurs near Wevfeu in Snuburg; Krieglach, in Styria; also Horrafiberg, Swt-deu. Abundant with corundum at Crowder's Mt., Gaston Cu., N. C.; nod on Oraves Mt., Lincoln Co. Qn., with calcite, etc.

The name lazulite is derived from an Arabic word, ami, meaning Amethyst, and alludes to the color of Igemist

Tavistockite. $Ca_2P_2O_7 \cdot 2Al(OH)_3$. In microscopic acicular crystals, sometimes stellate groups. Color white. From Tavistock. Devonshire.

Olivite, Perhaps $Ca_2(PO_4)_2 \cdot Al(OH)_3$. Comp. G. = 3.08. Color pale yellow. Occurs at the iron mine at Westau, in Scania, Sweden.

Anatase, $Ca_2Fe(AsO_4)_2 \cdot 8Fe(OH)_2$. In yellowish brown fibrous concretions. O. = 8-530. From Romaneche, near Micon, France; also at Schneeberg, Saxony,

Allanite, $Mn_2As_2O_7 \cdot 4H_2O$. In small brownish-red prismatic crystals. From the Hoss mine, Nordmark. and at Långban, Sweden.

Bynadelphite. $2(Al,Fe)AsO_4 \cdot 5Mn(OH)_2$. In prismatic crystals; also in grains. G. = 3.45-3.48. Color brownish black to lilac. From the Moss mine, Nordmark. Sweden.

Fluorite. $MnAsO_4 \cdot 2Mg(OH)_2$. In minute orthorhombic crystals, tabular or grouped

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in feather-like aggregates. O. = 3-87. Color greenish brown. From the Hsnttg mine, Pajsberg, Sweden.

Bemakite. $(Al,Fe)AsO_4 \cdot 4Mg(OH)_2$. In rhombohedral crystals. O. = 3.80-3.84.

Color brownish red, black on the surface. From the Hoia mine, Nordma^ 8we<leD.

Retslui. A bade arseoate of Ihe ytlriuni eartbB, maDganese and caldum. In ortho-rhombic ciystala. H. =4. Q. = 4'IS. Color cliocolule- lo cheslnul-brown. From lb« Hum iiiie, Nonliiiarh, Sweden.

Ananiopleita. Perliaps U,R,lOH),(AsO.)i H = Hu, Ca, also Pb. Ug; K = Mu.

also Fe. HiisbItc, cleavabU. Color brownUli red. Uccurs at ihe SjO mlue, Qrythytta pariah. Sweden, wltb rhod'inlte iu crystalline limeslone.

Hangaastblit*. Hemntoalibilte. Highly banic manganese antimouales. Iu embedded gntiuB. Color black. Maj^ anotliSult occurs at Nordmark, Sweden; htmatotlibiite la from Ihe 8J6 mine, Grytlytte parish.

Atalestita. Basic bUmiitb arsenate. HiBI.AsO.. In minute tabular crystals. O. =: 6'4. Color sulphur-yellow. From Schoeberg, Saxony.

C. Normal Hydrous Phosphates, etc.

■taut group t monoclinic Vivianite Gaoop.

StmTits. Hydrous ammoDium-miigneslum ptiosphate. In orlhorbombic.hemlmorphlc crjBtlns (Fig. 307. p. 9S'; white or yellowUh; slightly soluble. From gunno deiwsils.

OoUophaiiite. CaiPiOi + H,o. In layers resemhllng gymnite or opal. Colorless or snow-whlli,-. From the isliind of Somlirero. Monite U aimllar. fmin the Islnnds Moia and Monetaiu (heWest ludies, where It Is associated with monftUe, HCaPOi. occurring in yellow-Isb-while tricllulc crystals

PVBOFHOBPHORITB!. Mg,P,o, + 4(Ca,P,o, + Ca,P,o,). Massive, earthy. Color SDOW-wliite, dull. From the West ludies.

. . s 7.inc rliMphitle. Zn,P,o,-f H,o. In minute prismatic

crystals. Also in reniform masses. G. = 3"6-285. C<t!or grayish wliiie. Found in caTiies is cilainiDB at the zinc miues of Allenberg, uear Alxla-Clnpelle. The composition given is that of tLu urtlflrliil mlueriil.

DioWnsonil*. 31l,P,o, + H,o wiih R = Mii, Pe, Na, chiefly, also On, K,, LI, Id

tiiuliir, psemlorliomljoheilral cryatsls; commonly follsleri to micaceous. G = 3538-

3 843. Color olive- to oil-greeii, grass-green. From Branchville, FnirfieUi Co., Coun.

Fillowit«. Formula as tor dickinsonite and nlaofrom BrauchiTille, but differing In

„io T., 1... „ iu,,=,~ , Q _ 8.^ (.jIjjj. wai-yel!ow, yellowish lo red-

The three following triclinic species are related in composition and may be in crystalline form.

(Ca,Co,Mg),As₂O₇ · 2H₂O. In small crystals; often in druses and spherical aggregates. Q. = 8-5-3 6. Color light to dark rose-red. From Achenbach, Saxony.

Brandvite. Ca₂MnAs₂O₇ · 3H₂O. In prismatic crystals; crystals often united in radiated groups. O. = 3671-3-673. Colorless to white. From the Hntstig mine, near Pajsberg, Sweden.

Fairfieldite. A hydrous phosphate of calcium and manganese, Ca₂MnP₂O₇ · 3H₂O. In prismatic crystals; usually in foliated or fibrous crystalline aggregates. G. = 8-07-8'16. Color white or greenish white to pale straw-yellow. From Branchville, Fairfield Co., Conn.; Kabenstetn, Va. (leilaomanganiu).

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Reddickite. H₂PiOg · 8H₂O. In orthorhombic crystals near Brodick. En angle; also granular. Q. = 810:1. Color pinkish white to yellowish white. From Brodick, Cotland.

Picrophormite. R₂A₂O₇ · 6H₂O, with R = Ca: Mg = 1:1. In small prismatic form. Color white. From Biechelsdorf, Freiberg; Joplin, Mo.

Triclinite. Cu₂AsiOg · 5H₂O. In radiated groups, columnar; dendritic. Color terdligriagreen. From the Turgusk copper mine.

Vivianite Group. MoDOclinic.

Vivianite, + 8H₂O a:t:i = 0-7498 :1 :0-7015 yS = 75° 34'

Symplectite, + 8H₂O 07806:1:0-6812 72= 43'

Bobierite, + 8H₂O

Hffinibite, + 8H₂O

Brythrite, + 8H₂O 0-75 :1:0*70 75"

Aimabergite, -f-8H₂O

Cabrerite, Ab, + 8H₂O

Kbttigite, + 8H₂O

The Vivianite group includes hydrous phosphates of iron, magnesium, cobalt, nickel and zinc, all with eight molecules of water. The crystallization is monoclinic, and the angles so far as known correspond closely.

Vivianite.

Monoclinic. Crystals prismatic ($\alpha = 71^\circ 58'$); often in stellate groups. Also reniform and globular; structure divergent, fibrous, or earthy;

also encrusting.

Cleavage: d highly perfect; a in traces; also fracture fibrous nearly X -flexible in thin laminae; sectile. H. = 1-5-2. G. = 2.58-2.68, Luster, 6 pearly or metallic pearly; other faces vitreous. Colorless when unaltered, blue to green, deepening on exposure. Streak colorless to bluish white, changing to indigo-blue and to liver-brown. Transparent to translucent; opaque after exposure. Pleochroism strong.

Comp.—Hydrous ferrous phosphate, $\text{Fe}_3\text{P}_2\text{O}_8 \cdot 8\text{H}_2\text{O}$ = Phosphorus pentoxide 28.3, iron protoxide 43.0, water 28.7 = 100.

Chemical analyses show the presence of iron sesquioxide due to alteration.

Pyro., etc.—In a test tube yields neutral water, whereas, sod extricate. B.B. fuses at 1'S, coloring the fluo blue to green, to a grayish black in a globule. With the fluxes reads for iron. Soluble in hydrochloric acid.

Occurrence—Occurs associated with pyrrhotite and pyrite in copper and lin vitreous; sometimes in narrow veins with gold, traversing graywacke; both friable and crystallized in beds of clay and sometimes associated with iron ore; occurs in cavities of fossils or buried bones.

Occurs at St. Agnes and elsewhere in Cornwall; at Bodenmais; in the gold mines of Yerespatiek in Transylvania. The Siirtby variety, sometimes called blue iron-mine or nad'w Prussian blue (Fer azure), occurs in Greenland, Carinthia, Cornwall, etc.

In N. America, in Seie Je-wy, at Alleitown. Monmouth Co., both crystallized, in nodules, and earthy; at Mullica Hill. Gloucester Co. (Massachusetts), in cylindrical masses. In Virginia, in Seward Co. In Kentucky, near Eddyville. In Canada, with limonite at Vaudreuil.

Symptoms. Probably $\text{Fe}_3\text{P}_2\text{O}_8 \cdot 8\text{H}_2\text{O}$. In small prismatic crystals and radiated spherulic aggregates. G. = 2.58-2.68. Color pale indigo, inclined to celadine-green. From Lobenstein in Volga; Hutteberg, Carinthia.

Bobierite. $\text{H}_2\text{Fe}_2\text{P}_2\text{O}_8 \cdot 4\text{H}_2\text{O}$. In aggregates of minute crystals; also massive. Colorless to white. From the guano of Mexillones on the Chilean coast. Hauflerite is like

bobierrite, but contains calcium. From Bamle, Norway.

Hyacinthine. $\text{Hg}_2\text{As}_2\text{O}_7 \cdot 2\text{H}_2\text{O}$. In crystals resembling gypsum; also columnar; sometimes foliated. Color snow-white. From Ilie Banat, Hungary.

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BBTmTB. Cobalt Bloom. $\text{Co}_3\text{As}_2\text{O}_{11} \cdot 8\text{H}_2\text{O}$.

Monoclinic. Crystals prismatic and vertically striated. Also in globular and reniform shapes, having a drusy surface and columnar structure; sometimes stellate. Also pulverulent and earthy, micaceous.

Cleavage; cleavage highly perfect. Sectile. $\text{H} = 1-1.5$; least on 6. $G = 3-4$. Luster of cleavage faces pearly; other faces adamantine to vitreous; also dull, earthy. Color crimson- and peach-red, sometimes gray. Streak a little paler than the color. Transparent to subtranslucent.

Comp.—Hydrous cobalt arsenate, $\text{Co}_3\text{As}_2\text{O}_{11} \cdot 8\text{H}_2\text{O}$ = Arsenic pentoxide 38.4, cobalt pentoxide 37.5, water 24.1 = 100. The cobalt is sometimes replaced by nickel, iron, and calcium.

Pyro., etc.—In the closed tube with water heat a green leaf and turns blue; on cooling gives off white fumes. In the cool glass, and in the dark gray or black color. B.B. in the forceps fuses at 2 to a gray bead, and colors the flame light blue (arsenic). B.B. on charcoal gives an arsenical odor, and fuses to a dark gray mass, which with borax gives the deep blue color characteristic of cobalt. Soluble: in hydrochloric acid, giving a rose-red solution,

Obs.—Occurs at Schönbühl, Saxony, in white scales; Wolfach in Baden; Modum in the U. S., in Peon., sparingly near Philadelphia; in Nevada, at Lovelock's station. In California. Named from its red color.

Annabergite. $\text{Ni}_3\text{As}_2\text{O}_{11} \cdot 8\text{H}_2\text{O}$. In micaceous crystals; also massive and disseminated. Color apple-green. From Annaberg near Döhlen; Annaberg near Schneeberg; Riechelsdorf; in Colorado; Nevada.

Brookite. $\{\text{Ni}, \text{Mg}\}_2\text{As}_2\text{O}_7 \cdot 2\text{H}_2\text{O}$. Like erythrite in habit. Also fibrous, radiated; reniform, granular. Color apple-green. From the Blerra Cabreni. Sp. no.: at Lauron, Greece.

Köttigite. Hydrous zinc arsenate, $\text{Zn}_3\text{As}_2\text{O}_{11} \cdot 8\text{H}_2\text{O}$ Massive, or in crusts. Color light carmine- and peach-blossom-red. Occurs with unaltered at the cobalt mine Döhlen, near Schneeberg.

Hydroxanthite. Scorodite. A hydrous phosphate of the cerium and yttrium minerals.

Massive, small mamillary: as an iocruslaion. O. = 8't>4-4'01. Color brown, pinkish or j-ellnwiab while, tthabdophanite is from Coruwall; SeovUiiU is ritim the Scoville (llDionite) ore bed ig Salisbury, Coon.

Ohnrohitfl. A bjdroua phosphate of cerium and calcium. As a tliiu coating of mlutte crystals. o. = 8 14. Color pale smokc-giay tinged with flesh-red. From Cornwall.

SOOHODlTti. »*» Orthorhombic. Aiea ll:l-A = 0-8658 : 1 : 0-9541.

*f, 130 A l20 = 60' 1'. pji'. Ill A ill = 111° 6'.

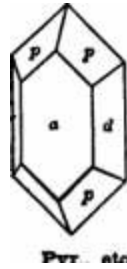
pj/ 111 A 111 = ""!°8' pff.ni fym= 65*20'.

Habit octahedral, also prismatic. Also earthy, amorphous. Cleavage: d imperfect; n, b in tnices. Fracture uneven. Brittle. H. = 3-5-4. G. = 31-IV3, Luster vitreous to sub-adamantitie and subresinous. Color pale leek-green or liver-brown. Streak white. Subtransparent to translucent.

Comp.—IIVdrouB ierric arsenate, $\text{FeAsO}_4 \cdot 2\text{H}_2\text{O}$ = Arsenio pentoxide 49'f<, iron sesquiojide 34 6, water 15'C = 100.

—, — In the closeii liil* yields neutral water and turns yellow. B.B. fuses

«aaiW, coloring Ibe-Unmv blue. B B. on charcoal gives arsenical fumes, and with soda a black magnetic scoria. With the fluxes reacts for iron. Soluble in hydri)chloric acid. OIm.. — Often associated with arseuopyrllle- From Schwarzeoberg, Saxooy; Demhacb.



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Nassau: LOLLiag, CariothU: Nerchinsk, Siberia, in flne cryiials; leek-green, in the Corniih mluea.

Occiii-B near Edeuville, N. Y., with araenopvdle; in Umb, Tintic dlitric, at th« Haiti tuoih iilue ou enarglte. As sn tncrusialiou on BiliceuiiBsiatev of the Yellousluue geysers.

Named from criediJoSoy, garlic, Hlludlug to tlje odor before tlii; blowpipe.

Strengita. $\text{FePO}_4 \cdot 3\text{H}_2\text{O}$. Crj'slala rare: lu habll and augli- nmr scorodiie: gi-neralljr III spberival and botryoidnl forms. G. = 2'6T. Color jwle red. From iron miues dcht Qlessen;

also in RiickOridge Co.. Va.. wtlh dufreuite.

PhoBphoaldarlte. $SFePO_4 + 8H_2O$. Au iron pliospbutc oeirr Btreut;ie, but wjtb $8H_2O$. From the Blegea luinDg district, Qermiiny.

Bairandlta (Al.Pe)FO₃ + SH_2O . Id ipberoidal coucretiun«, color pitle shades of gniy. From Bohemia.

Tarlictte. $AlPO_4 + 5H_2O$. Commonly in crystalline aggrt^tes and IncniBiatloDS witli renifnnn surface. Color greeo. From Mesebach in Siixou voigtlaDil; Homgomer^ Co., Arkansas, ou quart/.; iu uodulnr masses from Utah.

Oallalnlta. $AlPO_4 + 34H_2O$. Massive; wax-like. Color apple- to emerald-green. From a Celtic grave Id Lockmariaquer.

Zephuroviohlte. $AlPO_4 + 5H_2O$. Cryslalliue to compact. Color yellowish or grayish Tblie. From Treolc Id Bohemia.

Add Hydrous Phosphates, etc. .ooum.

Monoclinic. Crystals rare. Commoiily in delicate silky fibers; also botryoidal, stalactitic.

Cleavage I 6 perfect. Fracture uneven. Flexible in thin lamioie. H. = 2-25. G. = 3'64-2-73. Luster viti-eous; on b inclining to pearly. Color white or grayish; frequently tinged red. Streak white. TranBiucient to opuque.

Comp.—Probably $H_2CaAs_2O_7 + 2H_2O$ = Arsenic pentoiide 53-3, lime 25-9, water $2U^*8 = 100$.

Obs.—Fouod with arsenical ores of cobalt and silver, also with arsenopyrite: at AndrciLiliert; fn tbe Hurz; Riecheistlorf lu Hesse; JoachlmstLal iu Bohemia, etc. Named from oaiiU'iKor, poison.

Haidingortto. $UCaAs_2O_7 + 4H_2O$. Ill niliniio crystal aggregates, botryoidal and drusy. G. = 3 848. Color white. From JoiicLimslial with pbarmacnjlle.

Wapplarita. $H_2CaAs_2O_7 + 3H_2O$. In minute crystals; also in incrustatioDS. Colorless lo wliite. Fimiid wiib pliarmucollie at Joacliimellial.

Bruihlte. $H_2CaAs_2O_7 + 3H_2O$. In smnll slender prisms; coticretiotiHry massive. Color-le.ts to pile vellonish. Occurs in guano. Motabnuhlte, similmily B&suciated, is aHCuPO₄·8H₂O.

Martinite. $H_2CaAs_2O_7 + 3H_2O$. From pbospriorite deposits (from guano) in ibe islaiul iif CLir99oii, W. Indies.

Newberylte. $H_2MgPO_4 + 8H_2O$. In white onln'rlioiniile crystals. From guao of 8kit>toii Ciives, Victurta. Haanaylt*, from same li (iilUy, is a hydrous phosphate of amiiKiiiuiin

and magnesium.

Stercorite. Microcosmic salt. $\text{HNb}(\text{NH}_4)\text{PO}_4 \cdot 4\text{H}_2\text{O}$. Phosphorazul Germ. In white crystalline masses and nodules in gneiss.

Hursanite. $\text{H}_2\text{Mg}(\text{PO}_4)_2 \cdot 4\text{H}_2\text{O}$. In short prismatic crystals (monoclinic). Also may be compact, or imperfectly fibrous. Color yellowish, orange-red, rose, frequently. From Limoges, commune of Bureaux, France, in the U. States at Branchville. Codd.

Forbesite. $\text{H}_2(\text{Ni}, \text{Co})_2\text{As}_2\text{O}_7 \cdot 8\text{H}_2\text{O}$. Structure fibrous-crystalline. Color grayish white. From Atacama. •

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Ba²⁺ Hydrous Phosphates, etc.

Laolite. $\text{Ca}_2\text{P}_2\text{O}_7 \cdot 4\text{H}_2\text{O}$. In minute white crystals; also columnar. From Joachimsthal.

Hemifibrite. $\text{Mg}_3\text{As}_2(\text{OH})_6 \cdot 8\text{H}_2\text{O}$, + As_2O_5 . Commonly in a fibrous radiated group. Color brownish to garnet-red, becoming black. From the Hoes mine, Norrmark, Sweden.

SITOHROITB.

Orthorhombic. Habit prismatic $\text{mm}^3 = 63^\circ 40'$. Cleavage: m, n (011) IQ traces. Fracture small conchoidal to uneven. Rather brittle. $H. = 3.5-4$. $G. = 3.389$. Luster vitreous. Color bright emerald- or leek-green. Transparent to translucent.

Comp.— $\text{Cu}_2\text{As}_2\text{O}_7 \cdot 6\text{H}_2\text{O} =$ Arsenic pentoxide 34.3, cupric oxide 47.1, water 18.7 = 100.

Obs.—Occurs in quartz and mica schists at Lelbeten in Ungary, in crystals of considerable size, having much resemblance to diopside. Named from its color.

Orthochalcolite. Perhaps $(\text{Cu}, \text{Ca})_2\text{As}_2(\text{OH})_6 \cdot 4\text{H}_2\text{O}$. Resembling and massive, resembling diopside. Color pistachio-green to emerald-green. From Audulusia, Spangville district, Utah.

Baydonite. $(\text{Pb}, \text{Cu})_2\text{As}_2(\text{OH})_6 \cdot 4\text{H}_2\text{O}$. In massive concretions, drusy. Color green. From Cornwall,

Tagilite. $\text{Cu}_2\text{P}_2\text{O}_7 \cdot 3\text{H}_2\text{O}$. In reniform or spheroidal concretions; earthy. Color verdigris- to emerald-green. From Irbil, Ural.

Chalcocite. Probably $\text{Cu}_2\text{O} \cdot \text{Cu}(\text{OH})_2 \cdot 3\text{H}_2\text{O}$. In white, silky acicular crystals. From the Wilhelmie mine in the Spessart, Germany.

Volborthite. A hydrous vanadate of copper, bismuth, and calcium, in small six-sided tables; in globular forms. Color olive-green, siliceous-yellow. From the Ural.

Oronovite. $\text{Cu}_4\text{S}_2\text{O}_{12} \cdot 2\text{Cu}(\text{OH})_2 \cdot \text{H}_2\text{O}$. Massive, resembling malachite. Color emerald green. From Cornwall.

Tyrolite. Tyrolite. Kupferschaum Germ. Perhaps $\text{Cu}_8\text{O}_{13} \cdot 3\text{Cu}(\text{OH})_2 \cdot 7\text{H}_2\text{O}$. Usually in fan-shaped crystalline groups; also massive. Cleavage

circled. Very soft thin flexible brittle. Color pale green inclining to sky-blue. From the Ural. Hungary; Nerchinsk, Siberia; Palkestein, Tyrol; etc. In the U. States, in the Tintic district, Utah. Some analyses yield CaCO_3 , usually regarded as an impurity, but it may be essential.

OHAIPOPHITIM.

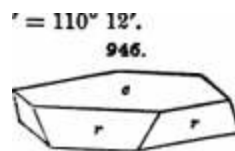
Rhombohedral. Axis $l = 2.5538$. $\alpha = 71^\circ 16'$, n

In tabular crystals; also foliated massive; in druses.

Cleavage: c highly perfect; r in traces. H. = 3. G. = 2.4-3-C6. Luster of c pearly; of other faces vitreous or subadamantine. Color emerald- or grass-green to verdigris-green. Streak somewhat paler than the color. Transparent to translucent. Optically —.

Comp.—A highly basic arsenate of copper; formula uncertain, perhaps $7\text{CuO} \cdot \text{As}_2\text{O}_5 \cdot 14\text{H}_2\text{O}$.

Pyro., etc.—In the closed tube decrepitates, yields much water, and gives a residue of olive-green scales. In other respects like olivine. Soluble in nitric acid, and in



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12 DESCRIPTIVE MINERALOGY.

WAVELLITE.

Orthorhombic. Axes $d:h:d = 0.5049 : 1 : 0.3751$. Crystals rare. Usually in aggregates, hemispherical or globular with crystalline surface, and radiated structure.

Cleavage: p (10L) and b rather perfect. Fracture uneven to subconchoidal. Brittle. H. = 3-3.5-4. G. = 3.16-2-3.37, lustrous, inclining to pearly and resinous. Color white, passing into yellow, green, gray, brown and black. Streak white. Translucent.

Comp. $4\text{AlPO}_4 \cdot 2\text{Al}(\text{OH})_3 + 9\text{H}_2\text{O} = \text{Phosphoria pentoxide } 35.2, \text{ alumina } 38.0, \text{ water } 26.8 = 100.$ Fluorine is sometimes present, up to 2 p. c.

E^{yr.} etc.—In the closed tube gives off much water, the last portions of which meet acid and color Benzil-yellow (Uorioe), and also etch the tube. B.H. in the forceps swells up and splits into Hue Infusible particles, coloring the flame pale green. Gives off blue vitriol cobalt solution. Heated with sulphuric acid gives off fumes of hydro-fluoric acid, which etch glass. Soluble in hydrochloric acid, and also in caustic potash.

Obt.—From Birsfield in Devonshire; at Zblrow in Bohemia; at Finkenbergr, Saxony; Minas Geraes, Brazil, etc.

In the United States at the slate quarries of York Co., Pa.; White Horse Station, Chester Valley R. R., Pa.; Magnet Cove, Arkansas. color white. $\text{AlPO}_4 \cdot \text{Al}(\text{OH})_3 + 2\text{H}_2\text{O}$. In crystals. Color green. From Nizhni Tagilsk in the Ural.

Pegmatite*. $\text{Al}_2(\text{PO}_4)_3 \cdot \text{Al}(\text{OH})_3 + 11\text{H}_2\text{O}$. Occurs in green crystals, of an all prismatic crystals at Striegis, near Freilberg, Saxony.

TITRQUOITE. Turquoise.

Massive; amorphous or cryptocrystalline. Reniform, stalactitic, or incrusting. In thin seams and disseminated grains. Also in rolled masses.

Cleavage none. Fracture small conchoidal. Rather brittle. H. = 6. G. — 2.6–2.83. Luster somewhat waxy, feeble. Color sky-blue, bluish green to apple-green, and greenish gray. Streak white or greenish. Feebly sub-translucent to opaque.

Comp. —A hydrous phosphate of aluminium colored by a copper compound, $\text{Al}_2(\text{PO}_4)_3 \cdot \text{Al}(\text{OH})_3 + 11\text{H}_2\text{O} = \text{Phosphorus pentoxide } 32.6, \text{ alumina } 46.8, \text{ water } 20.6 = 100.$ The copper salt probably has the composition $2\text{CuO} \cdot \text{P}_2\text{O}_5 \cdot 4\text{H}_2\text{O}$.

Pyro. etc.—In the closed tube decrepitates, yields water, and turns brown or black. B.H. in the forceps becomes brown and assumes a glassy appearance, but does not fuse; colors the flame green; instead with hydrochloric acid the color is at first blue (copper chloride). With the fluxes reacts for copper. Soluble in hydrochloric acid.

Obt.—The highly prized oriental turquoise occurs in narrow seams (8 to 4 or even 6 mm. in thickness) or in irregular patches in the brecciated portions of a porphyritic trachyte and the surrounding sky slate in Persia, not far from Nishapur, Elisorassan; in the Megara Valley, Sinai; in the Kaira-Tube Mis. in Turkestan. .50 crystals from Samarkand.

In the U. States, occurs in the Los Cerillos Mis., 20 m. S. E. of Santa Fe, New Mexico, in a trachytic rock, a locality long mined by the Mexicans and in recent years reopened and extensively worked; in the Burro Mts., Grant Co., N. M.; pale green variety near

Colitmbin. Nevadn.

Natural turquoise of inferior color is often artificially treated to give it the tint desired. Moreover, many stones which are of a fine blue when first found retain the color only so long as they are kept moist, and when dry they fade, become a dirty green, and are of little value. Much of the turquoise (now artificially used in jewelry in former centuries, as well as the present, and thus described in the early works on minerals, was boiu-turquois* called also othntolite, from osioii, Uiofi). which is fossil bone, or tooth, colored by a phosphate of iron. Its organic origin becomes manifest under a microscope. Moreover, true turquoise, when decomposed by hydromehloric acid, gives a fine blue color with ammonia, which is not true of the odontolite.

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PHOSPHATES, ARSENATES, ETC.

518

Siderite. $FeCO_3$. In globular druse concretions. Color light gray, bluish. From near Si. Beulgiu, Bohemia.

Liakmirdito. $(Al, Fe)_2AsO_4 \cdot 2H_2O$. In intersecting layers, white or bluish. From Lukeard, Cornwall.

Evansite. $2AlPO_4 \cdot 4Al(OH)_3 \cdot 12H_2O$. Massive; reniform or botryoidal. Colorless, or milk-white. From Zseizuk, Ungary.

Citronelle. $Al_2P_2O_7 \cdot 10H_2O$. Crypto-crystalline; milk-white to light copper-blue. Found near Katzenellenbogen, Nassau; also East Whitford Township, Chester Co., Penn.

Augelite. $3Al_2O_3 \cdot P_2O_5 \cdot 8H_2O$. In tabular rhombic crystals and masses. O. = 27. Colorless to white. From the mine of Westai, Sweden. The same locality has also yielded the following aluminophosphate.

Berberite. $2Al_2O_3 \cdot P_2O_5 \cdot 8H_2O$. Compact, massive. G. = 2-64. Colorless to grayish

Trollite. $4Al_2O_3 \cdot 3P_2O_5 \cdot 8H_2O$. Compact, distinctly cleavable. G. = 810. Color pinkish.

Attacolit. $P_2O_5, Al_2O_3, MnO, CaO, H_2O$, etc.; formula doubtful. Massive. G. = 500. Color salmon-red.

PHARMACOSIDERITE.

Isometric-tetrahedral. Commonly in tubes; also tetrahedral. Rarely granular.

Cleavage: a imperfect. Fracture uneven. Bather 947.

sectile. H. = 2'5. G. = 2'9-3. Luster adamantine to greasy, not very distinct. Color olive-, graas- or emerald-green, yellowish brown, honey-yellow. Streak green tn brown, yellow, pale. Subtransparent to sub-translncent. Pyroelectric.

Comp.—Perhaps $6\text{FeAsO}_4 \cdot 2\text{Fe}(\text{OH}) + 12\text{H}_2\text{O}$ -Arsenic pentoxide 431, iron sesquioxide 40-0, water 16-9 = 100.

Pjrr., etc.—Same hs for srorodite.

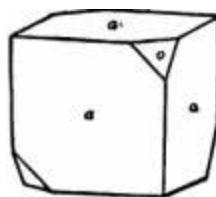
Obi—Obt! lined Hi the mines In Cornwall, with ores of cop-pc-r; at Sdinceberg and Brhwjirzplibere, Saxony; at Kftuigsbert, gary. In Uiab, at Ihe Mimimoth mine, Tintlc district. Named fi «Dd <TiSTj/}as, iron.

Iiudlamite. $\text{LiFe}_2\text{P}_2\text{O}_7 \cdot \text{Fe}(\text{OH})_2 + 8\text{H}_2\text{O}$. Occurs In small green tabular crystals {monoclinic), iie:ir Truro, Cornwall.

Oaooxenita, Kakoxen Oerm. $\text{FePO}_4 \cdot \text{Fe}(\text{OH})_2 \cdot 4\text{H}_2\text{O}$. In radiated lufta of ft yellow or brownish color. From near St. Beuigun in Bobcmla; Laucaster Co., Pi'nn.

Berannlte. $\text{Pm-hnps SFePC.FKOH}$. $\cdot 3\text{H}_2\text{O}$. Commonly in druses and in radiated glolmlesand cniuls. Color reddieli brown to dark liyacfnlh-red. From Si. Benlguii, near Beraun, in Bohemia BUonoriU. in tabular cryslAla, is the same mliiemi. From the Elentiore iniie nenr Oiessen.

Globobitk, PiCiTE, Dkltauxite are other hydraled ferric phosphates.



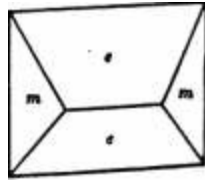
, near Schemnitz, Hun-
gari φάρμακον, ροίον,

OHILDRBNITB. Orthorhombic. Axes till: mm'' , no A lio = $7.V4n'$.

80 : 1 : 0-52570.

rr'' , $131 A |2| =]fl6^\circ V. m''$. lai A 131 = $46^\circ 56i'$. Only known in crystals. Cleavage; aimperfect. Fractureuneven. H. = 4'5-5. O. = 3.'18-3-24. Lnster vitreous t.o resinous. Color yellowish white, mle yellowish l>rown, brownish black. Streak white to yellowish. Translucent.

CoBp.~In general $2\text{AlPO}_4 \cdot 3\text{Fe}(\text{OH})_2 + 2\text{H}_2\text{O}$. Phosphorus pentfliida



30-9, alumina 22-2, iron protoxide 313, water 15-6 = 100. Manganese replaces part of the iron and it hence graduates into eoephorite.

Pyr.. etc.— In the closed tube gives off neutral water. B.B. swells up into ramification, and fuses on the edges to a Wack mass, color of the flame pale green. Heated on charcoal turns black and becomes magnetic. With soda gives a reaction for manganese. With borax and salt of phosphorus reacts for iron and manganese. Soluble in hydro-chloric acid.

Ob*.—From Tavistock, Devonshire. In U. States, at Hebron, H^{ts}.

Boopitite. Form and composition as (or chlorenite, but contains chiefly manganese instead of iron. In prismatic crystals; also massive. Color rose-pink, yellowish, etc. From B. ranch ville, Conn.

Hauasite. $\text{CaFe}_2(\text{AsO}_4)_2 \cdot 8\text{H}_2\text{O}$. In slender prismatic crystals. G. = 3.588. Color black. From Mazapil, Mexico.

Olofenite. $\text{CaFe}_2(\text{PO}_4)_2 \cdot 8\text{H}_2\text{O}$. Occurs in yellow to green nodules in clay at Battenberg, Rhenish Bavaria,

Borckite. Perhaps $\text{Ca}_2\text{Fe}_2(\text{PO}_4)_3 \cdot 13\text{H}_2\text{O}$. (Reniform massive) compact. Color reddish brown. From Leoben in Syria; Bohemia.

Rickite. Perhaps $\text{Ca}_2\text{Fe}_2(\text{PO}_4)_3 \cdot 8\text{H}_2\text{O}$. Massive, compact or foliated. Color yellow. From Richele, Belgium.

Uboonite. Monoclinic. Axes 1949.

Crystal resembling rhombic octahedrons. Rarely granular. Cleavage: m, e indistinct. Fracture subconchoidal to uneven. Imperfectly sectile. H. = 2-2.5. G. = 2.852-2.985. Luster vitreous, inclining to resinous. Color and streak sky-blue to verdigris-green.

Comp.—A hydrous arsenate of aluminium and copper, formula uncertain; analyses correspond nearly to $\text{Cu}_2\text{Al}_2(\text{AsO}_4)_3 \cdot 20\text{H}_2\text{O}$ = Arsenic pentoxide 289, alumina 103, cupric oxide 35-9, water 24.9 = 100. Phosphorus replaces part of the arsenic.

Pyr., etc.—In the closed tube gives much water and turns olive-green. B.B. cracks open,

but does not decreplntnt; fuses leaa readily lliau olivenlte ton dark gray slag; on charcoal cracks opeu. dedagralea, and gives reactions like oUventte. Soluble in nitric acid. Obi,—
Prom Curuwall; Herregrund in Hunirary,

Ohansvixita. Perhaps $\text{Cu}, \langle \text{FeO} \rangle, \text{As}, \text{O}, (-3\text{H}|\text{o})$. Massive to compact. Color dark green to greenish yellow. From Cornwall; UImIi,

Hbkwooditb. a hydrated phosphate of aluminhim and copper. In botryoidal globular masses. Color tunjuois-blue. From Cornwall.

Ohalooalderlte. $\text{CuO } 8\text{Fe}, \text{o}, .2\text{P}, \text{o}, .\text{gH}, \text{o}$. In sbeaf-llke crystalline groups, as in-cruslatioUB. Color llglit slab In-green. From Cornwall. AnDBEWHITS, also from Cornwall. Is near chalcosld^rite.

Eshoelts. A hydiKted phosphate of aluminium, zinc, etc. Massive. $Q. = 2 \ 34$. From Galenii, So. Dakota.

Qoyaslt. Perhaps $\text{Ca}, \text{Al}, \text{iF} > \text{oii}, .9\text{H}, \text{o}$. In small rounded grains Color yellowish white. From Hinaa Geraes, Brazil.

Plambognmmlt*. A hydrated phosphate of lend and alumlohim. Resembles drops or coalings of gum; as iDcrustatlona. Color yellowish, brownish. With galena at Huelgoet, Brittany, and elsewhere.

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PHoBPEATB8, ARSENATES, ETC.

TJrauite Group.

TOBBBRNITB. Copper Unmite. EapfBruraoit Qerm.

Tetragonal. -Axis $6 = :^{\wedge} -9361$. Crystals usually square tables, sometimes very thin, again thick; less ofteu pyramidal. Also foliated, micaceous.

Cleavage: c perfect, micaceouB. Lamiose brittle. $H. = 2-2'5$. $G. = 3'4-^{\wedge}6$. Luster of c pearly, other faces subadamantine. Color emerald- and grass-green, and sometimes leek-, apple-, and aiskin.green. Streak paler than the color. Transparent to eubtranslucent. Optically uniaxial; negative.

Comp.—A hjdroua phosphate of uranium and copper, $\text{Cu}(\text{UO},), \text{P}, \text{o}, (-8\text{H}, \text{o}) = \text{Phosphorus pentoxide } 15'1, \text{uranium trioiide } 61'2, \text{copper } d'4, \text{ivatet } 15''3 = 100$. Arsenic may replace part of the phosphorus.

Pyz., etc.—Ta Ibe cliwed lube yields wnter. In ibe Torceps fuses ut 30 to a blackish masa,

and colors the flame green. With salt of phosphorus gives a green bead, which with tina on charcoal becomes on cooling opaque red (copper). With soda on charcoal gives a globule of copper. Soluble in nitric acid.

Obs.—Froidefontaine elsewhere in Cornwall; Schneeberg, etc., Saxony; Joachimsthal, Bohemia.

Zennerite. $\text{Cu}_2(\text{UO}_2)_2\text{As}_2\text{O}_7 \cdot 8\text{H}_2\text{O}$. In thin tabular crystals resembling torbernite. In form and color. G. = 8.2. From Schneeberg, Saxony; Joachimsthal; Cornwall.

Lime Uranite. Ealluranit Qerm. Orthorhombic. In thin tabular crystals, nearly tetragonal in form and deviating but slightly from torbernite in angle; also foliated, micaceous.

Cleavage: basal, eminent. Laminar brittle. H. = 2-2-5. G. = 3.05-3.19. Luster of c pearly, elsewhere subadamantine. Color lemon- to sulphur-yellow. Streak yellowish. Transparent to translucent. Optically —. Ax. pi. β . Bx J. c. $\mu = 1.572$.

CoBp. —A hydrous phosphate of uranium and calcium, probably analogous to torbernite, $\text{Ca}(\text{trO})_2\text{P}_2\text{O}_7 \cdot 8\text{H}_2\text{O}$ or $\text{CaO} \cdot 3\text{UO}_2 \cdot \text{P}_2\text{O}_7 \cdot 8\text{H}_2\text{O}$ = Phosphorus pentoxide 15.5, uranium trioxide 63.7, lime 5.1, water 15.7 = 100.

Borne analyses give 10 and others 12 molecules of water, but it is not certain that the additional amount is essential.

Pyr., etc.— Same as for torbernite, but no reaction for copper.

Obs.—With uranite, as at Johanngeorgenstadt, Falkenstein. In the U. States, at Middletown and Branchville, Conn. In N. Carolina, at mica mines in Sutchell Co.; in Alexander Co.; Black Hills, S. Dakota.

Unosptnlte. Probably $\text{Ca}(\text{UO}_2)_2\text{As}_2\text{O}_7 \cdot 8\text{H}_2\text{O}$. In thin tabular crystals rectangular in outline. Color skin-green. From near Schneeberg, Saxony.

Uranocircs. $\text{Ba}(\text{UO}_2)_2\text{P}_2\text{O}_7 \cdot 8\text{H}_2\text{O}$. In crystals similar to uranite. Color yellow-green. From Falkenstein, S. Saxony.

PhosphuranyUt*. $(\text{UO}_2)_2\text{P}_2\text{O}_7 \cdot 8\text{H}_2\text{O}$. As 8 pulverulent in crusts. Color deep lemon-yellow. From Mitchell Co., N. C.

Tiigerit«. $(\text{UO}_2)_2\text{As}_2\text{O}_7 + 13\text{H}_2\text{O}$, in thin druses of tabular crystals. Color lemon-yellow. From near Schneeberg, Saxony.

WalpurgitB. Probably $\text{Ba}(\text{UO}_2)_2(\text{TiO}_2)_2(\text{OH})_2(\text{AsO}_4)_2$. In thin yellow crystals resembling gypsum. Q. = 5.78. Color yellow. From near Schneeberg, Saxony.

Rhaglt«. Perhaps $3\text{BiAsO}_4 \cdot \text{SbI}(\text{OH})_2$. In crystalline aggregates. Color yellowish green, wai-yellow. From near Scbne^rg, Baiony.

Mlzlte. A hydrated basic arsenate of copper and bls'^'Uth, formula $\text{Cu}_2(\text{OH})_2(\text{AsO}_4)_2 \cdot 4\text{H}_2\text{O}$. In acicular crystals: as an Incrustation, Color green to whitish. From Joachimsthal: WillicUen, Baden; Tinlic distr., Utah.

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Antimonates; also Antimonites, Arsenites.

A number of antimonates have been included in the preceding pages among the phosphates, arsenates, etc

Atopite Perhaps $\text{Cu}_2(\text{OH})_2(\text{AsO}_4)_2 \cdot 4\text{H}_2\text{O}$. Isometric orthorhombic. H. = 5.5. $\rho = 5.03$. Color yellow to resin-brown. From LAogbaD. Sweden.

Bindhaimlte. A hydrated arsenate of iron. Amorphous. reniform: also earthy or fibrous. Color green, brownish, yellowish, A result of the decomposition of other arsenates; in Cornwall; Sevier county, Arkansas.

Romeite. An arsenate of calcium, perhaps $\text{Ca}_2(\text{AsO}_4)_2$. In groups of minute square crystals, H. above 5.5. $G = 4.7$. Color bright yellow or bony-yellow. From St Martel, Piedmont.

Nadorito PbClSbO_4 . In orthorhombic crystals. H. = 8.6. $G = 7.02$. Color brownish yellow. From Djebel-Nador. Constantine, Algeria

Elcdemlt«. Helopbyllite. Perhaps $\text{Pb}_2\text{As}_2\text{O}_7 \cdot 2\text{PbCl}_2$. In crystals. mass, and as an Incrustation. O. = 6.8. $G = 7.14$. Color bright yellow to green. From LAogbao, Sweden; also Pajsberg (Finland).

OohroUte. Probably $\text{Pb}_2\text{Sb}_2\text{O}_7 \cdot 3\text{PbCl}_2$. In small crystals, united in diverging groups. Color Sulphur-yellow, From Pajsberg, Sweden.

TilpplMite. Essentially an arsenate of copper. In small bluish-green, tetragonal crystals. From CopUpo. Chile.

Tripnhylte. An iron antimonite. $2\text{FeO} \cdot \text{Sb}_2\text{O}_3$. In microcrystalline aggregates of a dull greenish-yellow color. From Tripuhy. Brazil.

Darbylita. An arsenic-litanate of iron. In prismatic, orthorhombic crystals. H. = 6. $G = 4.53$. Color black. Tripuliy, Brazil.

Lewisite. $\text{Sb}_2\text{O}_3 \cdot 2\text{TiO}_2 \cdot \text{Sb}_2\text{O}_3$. In massive yellow to brown isometric crystals. Tripuliy.

Brazil.

Mauielllle. A titanite of lead and calcium, related to lewisite. In dark brown isometric octahedrons. Jakobsberg. Sweden.

Amimutk. A doubtful antimonite of mercury; forming a scarlet earthy mass. From Chill.

Phosphates or Arsenates -with Carbonates, Sulphates, Borate*.

Sfelfel and Dahill (p. 459) may belong here.

Diadochite. A hydrated phosphate and sulphate of ferric Iron: Thuringia. Dettingen; from Belgium.

Pitticite. A hydrated arsenate and sulphate of ferric Iron. Reniform and fibrous. Yellowish and reddish brown. From Saxony, Cornwall, etc.

Svanbergite*. A hydrated phosphate and sulphate of aluminium and calcium. In rhombohedral crystals. Color yellow to yellowish brown, rose-red. From Hornsjöberg, Sweden.

Bendantite*. A phosphite or arsenate with sulphate of ferric Iron and lead: formula doubtful. In rhombohedral crystals. Color green to brown and black. From Cork; Dernbach and Horhausen, Nassau.

Undackite. Perhaps $3\text{NiO} \cdot \text{CuO} \cdot \text{SO}_2 \cdot \text{As}_2\text{O}_3 \cdot \text{Bi}_2\text{O}_3$. In rosettes, and in reniform masses. Color venetian- to apple-green. From Joachimsthal.

Lilaebergite $5\text{MgO} \cdot \text{RiO}_2 \cdot \text{P}_2\text{O}_5 \cdot 8\text{H}_2\text{O}$. In flattened masses, fibrous to earthy structure. From Lüneburg, Hannover.

A hydrous iron arsenate and lead sulphate from Laurion, Greece.

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The Nitrate being largely soluble in water plays an important role in Mineralogy.

BODANITE.

Rhombohedral. $a = 0.8276$; $\alpha = 73^\circ 30'$. Homobimorphous with calcite. Usually in massive form, as an incrustation or in beds.

Cleavage: r perfect. Fracture conchoidal, seldom observable. Rather sectile. H. = 1.6-2. Q. = Z-M-2.9. Luster vitreous. Color white; also reddish brown, gray and lemon-yellow. Transparent. Taste cooling. Optically - $c^{\wedge} = 1.5874$, $e_j = 1.3361$.

Comp.—Sodium nitrate, NaNO_3 , = Nitrogen pentoxide 63.5, soda 36.5 = 100. Pyr., etc.—
Deflagrates on charcoal with violence, throwing out nitric acid, causing a yellow light, and alu-
minium oxide, coloring the flame yellow. Dissolves in three parts of water at 60° P.

Occ.—From Tarapaca, northern Chile, and also the neighboring parts of Bolivia; also in
Humboldt Co., Nevada; near Concho, Big Horn Co., Wyo.

Tests.—Sodium nitrate, NaNO_3 . In thin white plates and silky tufts.

Nitroalloy*. Hydrogen calcium nitrate, $\text{Ca(NO}_3)_2 \cdot \text{H}_2\text{O}$. In efflorescent silky tufts and
masses. In the limestone caverns, as those of Kentucky.

Nitramagite. $\text{Mg(NO}_3)_2 \cdot \text{H}_2\text{O}$. In efflorescences on limestone caves.

Nitrobarite. Barium nitrate. $\text{Ba(NO}_3)_2$. In hexagonal prisms. From Chile.

Nitrochalcite. Basic cupric nitrate, $\text{Cu}_2(\text{NO}_3)_2 \cdot \text{Cu(OH)}_2$, in pyramidal orthorhombic crystals.
No. = 8426. Color emerald-green. From the copper mine at Jerome, Arizona.

Nitrocalcite. $\text{Ca(NO}_3)_2 \cdot \text{H}_2\text{O}$. In square plates. From Chile.

Nitroglauconite. $6\text{NaNO}_3 \cdot 3\text{Na}_2\text{SO}_4 \cdot 8\text{H}_2\text{O}$. From

Nitroiodate*. Calcium iodate. $\text{Ca(IO}_3)_2$.—In prismatic crystals, colorless to yellowish. From the
sodium nitrate deposits of Atacama.

Nitroiodate. A calcium iodo-chromate. Monoclinic; commonly fibrous or columnar, H. = 2-3.
No. = 8-70. Color dark gold-yellow. From the same region as the Carite.

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DBSCBIPTITE UINERALOOT.

Oxygen Salts. 6.

The aluminum borates, ferrates, etc., allied chemically to the borates, but have already been introduced
among the oxides. They include the species of the Spinel Group, pp. 337-341, also
Cobaltite, p. 842, etc.

In fibrous seams or veins. H. = 3. G. = 3.43. Luster silky to pearl. Color white with a tinge
of pink or yellow. Translucent.

Comp.— HRBO_3 , where R = Mn, Zn and Mg = Boron trioxide 34.1, manganese protoxide,
41.5, magnesia 15.6, water 8.8 = 100. Here Mn (or Zn) : Mg = 3 : 1.

Pyr., etc.—In the closed tube darkens in color and yields neutral water. If turmeric paper

is moistened with this water, and then with dilute hydrochloric acid. It assumes a red color (boric acid). In the forceps fuses in the flame of a candle ($F = 2$), and B.B. in O.F. yields a black crystalline mass, coloring the flame intensely yellowish green. With the fluxes reacts for manganese. Soluble in hydrochloric acid.

Obt.—Found on Mine Kill, Franklin Furnace, Sussex Co., N. J., with franklinite, Eucrite, willemite, etc. An intimate mixture of zircon and calcite, not uncommon at Mine Hill, is often mistaken for sussexite, but the ready fusibility of the genuine mineral is distinctive.

Andwigite. $3MgO \cdot B_2O_3 + FeO \cdot Fe_2O_3$. In finely fibrous masses. $G. = 3.91-4.02$. Color blackish green to nearly black. From Horowitz, Hungary.

Plankovite. $KMgO \cdot B_2O_3 + MnO \cdot H_2O$. Small rhombic crystals. $H. = 6$. $G. = 3.881$. Luster metallic. Color black. From Laugban, Sweden.

Nordenfalkine. A calcium-titanate borate, $CaTi(BO_3)_2$. In tabular rhombohedral crystals. $H. = 5.5$. $G. = 4.20$. Color sulphur-yellow. From the Langensund fiord, Norway.

Jeremejevite. Eichwaldite. Aluminium borate, $AlBO_2$. In prismatic hexagonal crystals. $H. = 6.5$. $G. = 3.28$. Colorless to pale yellow. From Mt. Soktuj, Aldan-Central range in Eastern Siberia.

Hambergerite. $Be_3(OH)BO_3$. In greyish-white prismatic crystals. $H. = 7.5$. $G. = 2.34$. From Langesund-fiord, southern Norway.

Szaibelyite. $3MgO \cdot B_2O_3 \cdot 8H_2O$. In small nodules; white outside, yellow within. From Rezbany, Hungary.

BOBACITE.

Isometric and tetrahedral in external form under ordinary conditions, but in molecular structure orthorhombic and pseudo-isometric: the structure becomes isotropic, as required by the form, only when heated to 265° . (See Art. 411.)

Habit cubic and tetrahedral or octahedral; also dodecahedral. Crystals usually isolated, embedded; less often in groups. Faces smooth and bright, dull or uneven.

Cleavage: $0, 0$, in traces. Fracture conchoidal, uneven. Brittle. $H. = 7$

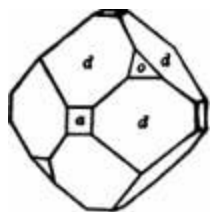
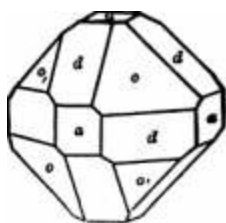
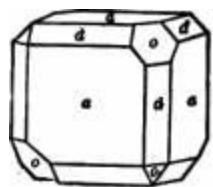
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in crystals. $G. = 2.9-3$. Luster vitreous, inclining to adamantine. Color white, inclining to gray, yellow and green. Streak white. Subtransparent to translucent. Commonly shows double refraction, which, however, disappears upon heating to 265° , when a section

becomes isotropic. Refractive index,

- a = 0.0107.

949.



Strongly pyroelectric, the opposite polarity corresponding to the position of the + and - tetrahedral faces {see pp. 234, 235}. The faces of the dull tetrahedron a, (111) form the analogous pole, those of the polished form o (111) the antipodal pole, Kose.

Comp.—Mg, Cl, B, O, or 6MgO.MgCl₂.8B₂O₃= Boron trioxide 63.5, magnesia 31.4, chlorine 7.9 = 100, deduct {O = Cl} 1.9 = 100.

V>r.—1. Ordinary. In crystals of varied habit. 2. Mosaic with sometimes a subcolumnar structure; it is of the rhombohedral form. It is a uniaxial crystal with the optical axis parallel to the c-axis. Faratite of Yolger in the plumotte lenticles of some crystals of boracite. S. £^rM(a^urtiJ« coDtninsBome Fe.

Pyr., etc.—The massive variety gives water in the closed tube. B.B. both varieties fuse at 2 with lumefedCt: 10 a white crystalline pearl, coloring the flame green: heated after moistened with cobalt solution presents a deep pink color. Mixed with nitride of copper and heated on charcoal colors the same deep azure-blue (copper chloride). Soluble in hydrochloric acid.

Alters very slowly on exposure, owing to the magnesium chloride present, which takes up water. It is the frequent presence of the deliquescent chloride in the massive mineral, originating, that led to the view that there was a hydrous boracite (at the first), Paratite of Yolger is a result of the fine kind of alteration in the interior of crystals of boracite; this alteration giving it a somewhat plumose character, and introducing water.

Obi.—Observed in l^{ds} of anhydrite, gypsum or salt. In crystals at Ealkberg and Scbildslcin in LOneburg, Hannover; at Segeberg, near Kiel, in Holstein; at LunevilTe, La Meurthe. France: massive, or as part of the rock, also in crystals, at Blassfurt, Prussia.

Ascharite. A hydrous magnesium borate. In white lumps with boracile. From Aschersleben, Germany.

Rhodsite. A borate of aluminium and potassium, with cesium and rubidium. Tetrahedral: in white, translucent dodecahedrons. H. = 8. Q. = 8'41. Found on red iron ore from the vicinity of Ekaterinburg in the Ural.

Warwickite. Perhaps $\text{CaMg}_2\text{Si}_2\text{O}_{10} \cdot 3\text{H}_2\text{O}$. In elongated prismatic crystals, G. = 8'38. Color dark brown to dull black. From Edenburg, N. T.

Howlite, A silico-borate of calcium, $\text{Ca}_2\text{B}_6\text{O}_{17}$. In small white rounded nodules; also earthy. From Nova Scotia.

Idgonite. $\text{Fe}_2\text{Si}_2\text{O}_7$. In crusts at the Tuscan lagoons.

Larnite. $(\text{NH}_4)_2\text{B}_4\text{O}_7 \cdot 4\text{H}_2\text{O}$, From the Tutcao lagoons:

OOLBMANITE.

Monoclinic. Axes $a : b : c = 0.7748 : 1 : 0.5410$; $\beta = 69^\circ 5'$. Crystals usually short prismatic ($a = 0.72$). Massive to granular and compact.

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cr>N

620 DESCRIPTIVE MINERALOGY.

Cleavable: b highly perfect; c distinct. Fracture uneven to subconchoidal. H. = 4-4'5. G. = 2.12. Luster vitreous to adamantine, brilliant. Colorless to milky white, yellowish white, gray. Transparent to translucent,

Comp $\text{Ca}_2\text{B}_6\text{O}_{17} \cdot 5\text{H}_2\text{O}$, perhaps $\text{H}_2\text{Ca}(\text{BO}_3)_2 \cdot 2\text{H}_2\text{O}$ = Boron trioxide 50-9,

lime 27-2, water 21-9 = 100.

Pyr.—B.B. does not heat, exfoliates, sinters, and fuses imperfectly, coloring the flame yellowish green. Soluble in hot hydrochloric acid with separation of boric acid on cooling.

Oba.—First discovered in the Death Valley, Inyo Co., California; later in the Calico district. See Bernardt's country.

pRiCmTB. Hear colemaiite. HassJTe, fdahle aad chalky. Color suow-wbile. From Curry Co., Oregon. PandermiU is similar; in compact nodules from Asia Minor.

Plnnohte. HkB,o..8HiO. Tetrsgonul-pyramidal. Usually ia nodules, radiated fibnius. Color sulphur- or straw-yellow. From Stassfurt

BBintEite. Hinzelte. EBlil>orlte. A bydrous borate of magnesium and potassium. In small crystals, someilmes aggregated togeUiep H. = 4-S. O. = 3'18. Colorless to white. From Leopolds ball, Stassfurt.

BORAX.

Monoclinic. Axes A : t : i = 10995 : 1 : 0-5632; /? = 73° 25'.

ea. 001 A 100 = 7Z° Sff. «. 001 A 531 = 64° ff.

mm", 110 A lio = 98° 0". . a/, lU A ill = 67° 27. eo, 001 A ill = 40° 81'. «", §21 A ^1 = 88° 28'.

Crystals prismatic, sometimes large; reEembling pyroi-eue in habit and angles.

Cleavage: n perfect; t» less so; i in traces. Fracture coDchoidaf. Rather brittle. H. = 2-2 5. Q. = 1-69-1-72. Ltister fitreous to resinous; sometimes earthy. Color white; Bometimes grayish, bluish or greenish. Streak white. Translucent to opaque. Taste sweetish-alkaline, feeble. Optically-. Aipl. J.J. Bx.X*- Bi., A** = - 56° 50'. /3 = 1 -470. 2E, = 59° 30'.

Comp Na,B,O,₁₀H₂O or Na₂O 2B₂O₃·10H₂O = Boron trioxide 366, soda

16 2, water 47-2 = 100.

Pyr., eto.—B.B. puffs up and afterward fuses to a tranpparent globule, called the glass of borax. Fused wilb fliioilte and poiaasium blsulphate, It ciilors tbe Baine arouuJthe assay s clear green. Soluble iu water, yielding a falutly alksUne solullou. Boiling water dissolves double Its weight of this salt,

Obs,—Obtained from the lalt lakes of Tibet; Ihe crude mineral is called lineal. In California, abundant in Lnke Co., iit Boihi Lake and Hnchlbnbiimn, two small alkaline lakes In ilie Immediate THuity of Clear Lnke; present in Kolutlon in the lake waters, and obtained also in targe quantities In flne crystals embedded in the lake mud and the surrounding marsby soil; also found in fire largfi clear crvstnls st Borax Lake, Ban Bemardiini Co.: at Death Valley, Inyo Co. Also Rbodes Hnrsh. etc.. Esmeralda Co.. Nevada.

Named borax from tbe Arabic iuro?. whicb Included also the niltr (sodium carbonate) of ancient writers, tbe natron of the Egyptians. Borax was called chrysocolia by Agricola because used in soldiring gold.

Ullmannite. Boronocalcite. Natron borocalcite.

Usually in rounded masses, loose in texture, consisting of fine fibers, which are acicular or capillary crystals. H. = 1. G. = 1.65. Luster silky within. Color white. Tasteless.

Comp.—A hydrous borate of sodium and calcium, probably $\text{NaCaB}_3\text{O}_7 \cdot 8\text{H}_2\text{O}$ = Boron trioxide 43.0, lime 13.8, soda 77, water 35.5 = 100.

^

ab,GoOgIc

Ullmannite. B.B. fuses M 1 with Intumescence to a clear blebby glue, coloring the Ullmannite deep yellow. Molten with sulphuric acid the color of the flame Ullmannite usually bluish deep blue. Not soluble in cold water, and but Ullmannite so in hot water the solution alkali is its reaction.

Occ.—From the dry pit in or Iquique, Chile. In Nevada, in large quantities in the salt marshes of the Columbus Mining District, Esmeralda Co.

Named after the German chemist, G. L. Ullmann

$\text{CaB}_2\text{O}_7 \cdot 4\text{H}_2\text{O}$. In crusts, as a deposit from springs in Tuscany.

Hydroboracite. $\text{CaB}_2\text{O}_7 \cdot 4\text{H}_2\text{O}$. It resembles fibrous and foliated gypsum; color white. From the Caucasus.

Uranate.

Uraninite. Cleveite. Bröggerite. Nivenite. Uranocerite. Isometric. In octahedrons, also with dodecahedral faces {d}; less often in cubes with o and d. Crystals rare. Usually massive and botryoidal; also in grains; sometimes columnar, or curved lamellar.

Fracture conchoidal to uneven. Brittle, H. = 5.5. G. = 5.0 to 5.4 of crystals; of massive altered form from 4 upwards. Luster submetallic, to greasy or pitch-like, and dull. Color grayish, greenish, brownish, velvet-black. Streak brownish black, grayish, olive-green, a little shining. Opaque. Comp.—A uranate of uranyl, lead, usually thorium (or zirconium), often the metals of the lanthanum and yttrium groups; also containing the gases nitrogen, helium and argon, in varying amounts up to 26 p. c. Calcium and water (essential?) are present in small quantities; iron also, but only as an impurity. The relation between the bases varies widely and no definite formula can be given.

Tar.—The minerals provisionally included under the name uranates are as follows:

1. Crytlaltiteri. DVonrtio*iW'mm Norway. la crystals, iisuaDy octahedral, with G, varying for liie most part from 9 0 to 0'7 ; occurs as an originiil constituent of coarse gratiitcs. The variety from Branchville. wliich is as free fToro alicratiou as any yet exiimiiiieii. cuutuuiis chiefly L'O, with a relatively sniHll amount of UOi. Thorln Is prom-ineiii. whitv till: earihs of tlic inntliuDum nnd yltrlum groups uru oiily sparingly reprcseuted.

Bro-igtrite, as nuulyzed by Hillebrund, gives the oxy^'cn ratio dF UO, to other bases of ubonr 1:1; it occurs in octahedral crystals, also with d ond a. G. = 8oS.

OleMUe add iineniU coutaiu UOi in larger amounL than tlie other Tarieties mentioned. aiul are charuclerized by coDtaiuiog about 10 p. c. of tlie yttrium eavtlig. Clevelte is ft variety from tLe Arendai region occurring in cubic crystals modified by the dodecahedron and octubedrou. Q. = 7'49. It is particularly rich iu Ihe gas helium. Nivenile occurs iias-sive, with Indistinct cryBtallizalinn. Color velvet-black. H. = SS. Q. = S'Ol. It is more solitbie than other kinds of uranltilte, beioq completely decomposed by the action for one liour of very dilute sulphuric acid nl 100°.

3. ifouAis, probably nmorphous. PlIchblende. Uranpecherz o«mi. Coutains oo thoria; the rare earths also aljsent. Water Is prominent and the specific gravity is much lower, in some cn<<s not above 65; these last diSei'ences are doubtless largely due to altei'nllonii. Ili-re belong the kinds of pitchblende which occur In metnlliferous veins, with sulnhideaof silver. Ipad. cobalt, nickel, iron. zinc, copper, a-a that from Jolinnggeorgetaitt, Pribram, etc.: probably al«o that from Black Ha»k. Colomilo (Hillebraud).

Pyr.. etc.—B.B. Iiifuaible, or only sliglMly rounded on the edges, sometimes coloring the outer flame green (copper). With borax and salt of pliosphonis gives a yelliiw bead In O.F . becoming creen In R.F. (uranium). With soiln mi charcoal gl^ea a coaline of lead oxide, »ud frequently Ibe odor of HTsenic. Many specimens give reactions for Sulphur and arsenic In the open tube Soluble In nitric and sulphuric acids; the solubility differs widely In different varieties, being greater in those kinds containing the rare earths. Not Attractable by the magnet.

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Ob*.— Ab DOfed kbove, unninlte occurs either m a primsir conatitueDt of granittc rocks

or as II secondsrj minenl witL ores of sllTsr. lend, copper, etc. Under the latler condition ll ia found at JohanDgeoigenatudl. Uarieuber)^, and ticnneeberg id Saxooy, at Joochimithal and Pfibrem Id Botiemla, and Rezbaoya in HuuKurv. Occuib In Norway Id pegmsiile TeiuB at Bcverai poluls uear Mosa, viz.; AuDerttd (briggarite), Elvestad, etc; also oeax Areudal at the Oarta feldspar quarry (titteite), aaauciated wtlh orlliUe, fergiuoolte, thorite.

pi'giimiite veil], as Bmall uctalitdral crjatuls, euibediled in alblLe. In N. Carolina, at th« Flat Rock mlue and other mica miuea In Mitchell Co., rather abundant, but usually altered, in part or entirely, tn guiuDiie and uranopbsoe; tlie crystala are Bometimes an inch or more across and cubic in hnbtt. In S. Carolina, al Marietta. In Texas, at the gadolinile Incalily In Llano Co. (ninsnde). In targe quuutities at Black Hank, near Central City, Colorado. Rather abundant in the Bald Mountain district. Black HIIIb, 8. Dakota. Also vi'Uh inonazlt', etc., at the Villeueuve mica veins. Ottawa Co., Quebec, Canada.

auinmlta. An alteration-pro duct of urauinlte of doubtful compoaiiJoD. In rounded or Uittted pieces, looking much like gum. G. = 8-e-420. Luster greasy. Color reddish yellow to orange-red, reddish brown. 3^om Johnngeorgenstadt, also Mitchell Co., N. C.

Yttroocmmittb. Occurs with cleveite aa a decorapoaltion-product

Thosoolmhitb. Occurs with fergusonite, cyrtollte, and other species at Ibe gadoli-nite locality lu Llano Co., Texas.

tTruioBpliaTlt«. (BiO)',U.OT.8H(o. In half-globular aggregated fom. Color orange-fellow, brick-red. fVom near StJineeberg, Saxony.

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BULPHATSS, CHKOUAIES, BIO.

Oxygen Salts.

6. SUIfHATES, OHROMATES, TEUASRATEB.

A. Anhydrous Sulphates, etc.

The important Babite Group is the only one among the anhydrons Boi-phates and chromatea.

BCuoagnlt*. AmmoBium aulphaU, (NH,)]SO,, Usually In cruBta and italaclltlo forms. Uccunt aUiut volcunoes, ns ot Etua, VnuvUts, etc.

TaylotiU. 3E|SO,,(NH.)i8o,. Id Hmull compact lumps or coDcrettons. From ths g'jntui of ibe Clilticba Islaads.

Thsnardlt*. Aiiiydrous swllum sulphale. NiiiSO,, In orlhorhomble crjatals, pyramidal, short prismuli ■ or labiiiar: also ns twins {Fig. 846. p. IIS). Wliile lo brownlab. Soluble Id wnter. Ofti-ii observed !□ connectlou wiib salt lakea, us od the sbor«e of Lake B.'ilkuhash, C«ulral Asia; similarly cUi'Wlitrc; also in B. America in Tarapaca. Id tbe U. S. forms uxtCDsive depo:jita on Lhe Rio Verde, ArizoDii. In Caliruroia, at Borax Lake, Saa BeTuardimi Co.

AphtUtallt*. ArcaDlte, (E.Na),8oi. RboinboheclTal; also massive, (d crusts. Color white. Protn Veauviuu, upon lava; atDouglaballnearWesturegelD in blOdUe; Rocalmuto, Sicily.

aLAUBBRTTB.

Monoclinic. Axes $A:t:i = 1-2200 : 1 : 1-0275$; $/3 = 67^\circ 49'$. ea. $001 A 100 = 87'$ W. et. $001 a Hl = 48^\circ \wedge . \dots$,

mm", $111 A 111 = \ll 6^\circ 58'$. em, $001 A 110 = K' 30^*$ '.

In crystals tabular | c; also prismatic.

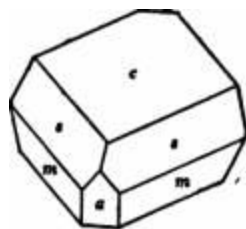
Cleavage: cperfect. Fractnreconchoidal. Brittle. H. = 20-3. G. = 2 7-2-85, Luster vitreous. Color pale yellow or gray; sometimes brick-red. Streak white. Taate slightly saline. On the optical properties (see p. 225).

Comp.—Na,SO..CaSO. = Snlphur trioxide 57-6, lime 20-1, aoda 223 = 100; or. Sodium sulphate 61-1, calciam sulphate 489 = 100.

Fyr., ato.—B.B. decreptatea, liima wlfte, and fusea at 1-5 to a white enamel, coloring the flame latenseiy yellow. On nliarcoa) fuses In O.F. to a dear bead: In R.F. a portion ia nbaorbed by the cbarconl. leaTlng iin infiaible hepatic residue. Soluble la hydrochloric iicld. lu water it loaea its tranapareDcy. Is pnrlfally dlasolved, leaving a reddue of calcium Bulphata, nnd in n large excecs this Is completely diiwnlveil.

Oba.—In crystals Id rock (alt at Tilla Rnbla, In New Caalle, Spain; alao at Aussee, Upper Austria; Berchteagaden, Bavaria; Westeregelu: Btassfiirt. In crystals la the Kio Verde Valley, Atizona, with Ibenardlte, mirabilite, etc.; Botax lake, San Bemardluo Oo., California.

I.iiifb«bilta. E|Hri<SO.)i. laometrhc-tetartobedral. In highly modified colorleai crystals. O. = 9'8.^ From Wetteregeln. Oermauy.



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DESCBlPnrE HINEBALOQY.

Barlte Group. RSO,. Orthorhombic.

The Bakite Group includes the sulphates of barium, slroittium, and lead, three species

which are closely isomorphous, agreeing not only in axial ratio but also in crystalline habit and cleavage. With these is also included calcium sulphate, anhydrite, which has a related but not closely similar form; it differs from the others conspicuously in cleavage. It is to be noted that the carbonates of the same metals form the isomorphous Aragonite Gkoup, p. 353.

BAHTTB. Heavy Spar. Baryite. Schwerspath Oerm. Orthorhombic. Axes $d:h:i = 0.8152 ; 1:1.3136$.

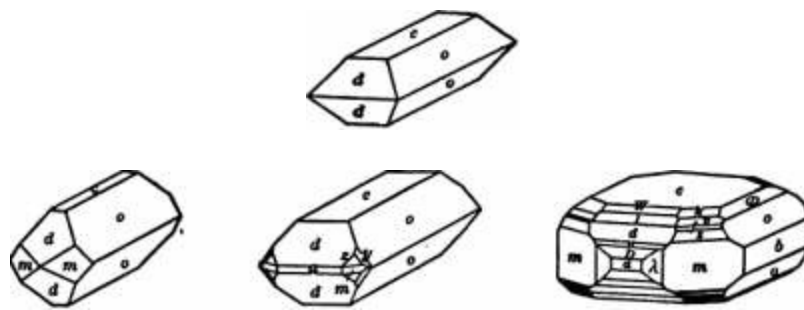
mm'' . $110 \text{ A } 103 = 83^\circ 51'$. $CO. 001 \text{ A } 011 = 62^\circ$ iz:

$M' \setminus 103 \text{ A } 105 = 103'$ IT. eo' , on $\text{A } 011 = 74'84'$. $ct, 001 \text{ A } 111 = 64^\circ$ W.

966. 967.

^

/f3



Crystals commonly tabular $\setminus c$, and united in diverging groups having; the axis h in common; also prismatic, most frequently Q axis $J, d (102)$ predominating; also I axis i, m prominent; again $|| d$, with o (Oil) prominent. Also in globular forms, fibrous or lamellar, crested; coarsely laminated, lamina

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BOLPHATES, CUBOHATES, ETC. »29

Convergent nod often curved; granular, resembling white marble, and earthy; colors sometimes banded as in stalagmite.

Cleavage: c perfect; m also perfect. Fig. 954 the form yielded by cleavage; also 6 imperfect. Fracture uneven. Brittle. $H, = 3.5-3.5$. $G, = 4.3-4.5, 4.489$ cryst. Luster vitreous, inclining to resinous; sometimes pearly on c , less often on m . Streak white. Color white; also inclining to yellow, gray, blue, red, or brown, dark brown. Transparent to translucent to opaque. Sometimes fetid, when rubbed. Optically $+ \bullet$ Ax $pi. \setminus b. BiX^{**} - SE, = 63^\circ 6'$, $p, = 1-6371$, (see further p. 224).

Y>x.—Ordinary, {a) Crystals UBually broad or atout; aiimeituea very large: again In ■leader ueedles. (A) Crated; massive a^gregutious of labular ciytlals, tbe crytala project-iai; nt surface ID 10 cresl-like forms, (c) Coiumnar; tie columns often coaise and tnosely aggreguled, and eilber radiated or parallel: larry fine fibrous, (d) In globular or nodular coucrtiloDB. Bubtlbrous or columaar witbiu. Bologna lslon4 (from ueur Boloena) is here iincludeil; it was early a source of woudtr btcaise of ibe pbuHphorescence ii eiubiied after heating mltb cbarcoal. "Bologna pluspborus" «'iia made from it. (s) Laoiellar, eiiher atraight or curved; the latter sometiies as »g(jregiilious of curved scelC'like plates. (/) Orauular. (g) Compact or cry piocrys lull Inc. (/<) Earthy. (o Stilaclic itod slaliig-mltic; similar In structure and origin to calcnreous etuiactiiea and stalagmilea aud of niucb iMiiuty wbeu polished. (A) Fetid; so called fioiii llie odor given off vheu struck or vbin two pieces are rubbed together, which odor may be due to carbonaceous matlers present.

The bBL'ite of Muzsaj, Hungary, and of Betier. uenr Rosenau, was early <'nllled Wolnyn. Cawk [a the ordlDary barite of the Derbyshire leaii mines. Drtetitt, supposed to be ihoni-boJiedral, ta simply barite. Mielhel-levyte from Ptrktii's Mill, Templeton, Quebec (deecrlbed aa moDOclitic), is peculiir lu its pearly luster on m, twinning striations, eic

Comp Barium sulphate, BaSO. = Sulphur trioxide 34-3, baryta 6B-7 = 100.

StroDliim sulphate Is often present, also calcium sulphate; further, aa Impurllles, silica, clay, bituminous ur carbonaceous subBlances.

Pyr., etc.—B.B. decrepitates and fuses at 8, coloring the flame yellowfeh fCTtta: the fused mass reacts alkaline with lest paper. On charcoal reduced to a sulphide. Wllb soda gives at tlrst a clear penrl. but on continued blowing yields a hepatic niam. which tprends out and soaks into the coal. This react>< for sulphur (p. 20S). loEoluble in acids.

Dlff.^Characterized by high specific gravity (higher thnn cealite. aiagonite. niblle, csclite, gypsum, etc.) : cleavage; insolubility : griei^ coloration of the blowpipe flame. Ablte is harder and calcite eServesces vnith neid.

Obfl.—Occurs commonly in connection with beds or veins of metallic cres, eFpecially of lead, also copper, silver, cobalt, manganese, as part of the gnnguc of the ore; also often ai^companies stibnite. Sometimes present la massive forms with hemallie deposits. It is ini:l wiib In secondary limestones and sandstones, sometimes forming distinct veins, and la the former often In crystals along with enlccte and celcslie; lu ihe latter often with copper ores. Sometimea occupies the cavities of nmyednloldat basalt, porphyry, etc.; forms eaitly masses in beds of marl. Occurs aa the petrifying material of fossils and occupying caviiea in them.

Pine crystals are obtained at the Duflon lead mines, Westmoreland, England; also In Cumberland and Lancashire; in Derbyshire, Staffordshire, etc.; Cleator Moor: Alston Mont. In Scotland, In Argyleahlre, at Strontian. Some nf the most Important of the many

European localities are PelsObfinya, Nngybinya, Schemnitz. and Eremnilz, in Hungary, often with stibnite; HCitenbern, Carinthin; Freiberg. Marienberg. in Saxony; Cliiusllial In IheHarz; Ffibrum, Bohemia; wllh the manganese ores of Ilefitd, (Ehrenstock, etc: Auvergne.

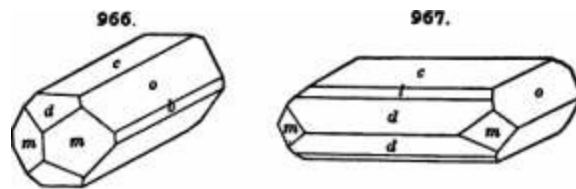
In Hie U. States, formerly In Conn., at Cheshire. Intersecting the red sandstone In veins with cbalcocite and malacliite. In N. Turk, nt I'llar Point, opnoaite Sackttl's Harbor, ma'^lve: at ScohaHe. fibrous; in St. Lawrence Co., crystals at De Kalb; the cre^ti'd variely nl Hammond. In Thnn., in cryatala st Perkiomen lead mine. In Virginvj, at Eldrigrige a jrold mine in Buckingham Co. In N. Carolina, white mnsaive at Cmwilers Mt.. Qaslon Co., PIC. In Tenn , on Brown's Creek: at Haysboro', near Nashville; In large veii)B in sandstone on Ihe W, end nf I. Royale, h. Superior, and on Spar Id., N. shore. In 3lit»tmri, not uncommon with the lend ores: In concretionary forms ni Salina. Ballne Co., Kansas. In Colorado, at Sterlini;, Weld Co ; ApTslinpa Creek; also In El Paso and Fremont Cos. In fine crystals, near Fori Wallace, New Mexico.

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In Oulario. Id Bathurat, and N. Burgew, Lanark Co.: Gslnaf. Peterboroueh Co.; u large veina on Jarvli, HcEellaT*, and Pie Ulande, In L. Superior, and near Fort Wflllam. Tbunder Bay. Ill Nova Scotia, Id velna In tbe alales of But Riv«r of the Five Islaoda, ColcLester Co.

Named tiaai fiapvi, AMny.

OELEffrrm. Ccsleatine. Orthorhombic Axes S:h:i = 0-7790 :1:1-2800,



mm''', 110 A lIO = n° no: ed, 001 A lOa = 89° 241'.

et. 001 A 104 = 2i° aC. 60. 001 A OH = 62' o''.

GrystalB resembling those of barite in habit; commonly tabnlar | c or prismatic Q axie d or S; also more i-arely pyramidal by the prominence of the forms ip (133) or x (1^)- ^^ fibrous and radiated; sometimes globular; occaeioDally granular.

Cleavage: c perfect; m nearly perfect; b less distinct. Fracture uneven. H. = 3-3'5. Q. = 3'95-3'9i. Luster vitreous, sometimes inclining to pearly. Streak white. Color white, often faint bluish, and sometimes reddish. Transparent to subtraaalucet. Optically +, Ax. pi, lb. BxJ. a. 2E, = B9° 36\ ^j = 1-6237.

V«r.—1, Ordinary, (a) In crjBtaU of varied habit as noted above: a tinge of a dellcnie blue
Is very cominna and eometimea belongs to only a part of a crystal. The variety fmm
MnDttntartre. r-allsd apoioTnt. \% prismatic liy extension of o 'Oil) aod doubly terminated
l>f tbe pyraiiilil it (IDS)- {*) Fibriuis, either parallel or radiated, (e) Lamellar; of rare
ocpur-iBOce. idi Orauulur. {») Coucreiiooary. (f) Saitiiy; impure usually with carbonate of

Comp.—Strontium sulphate = SrSO_4 , = Snlphur trioxide 43*6, etrontia Ei6'4 = 100.
Calcium and barium are sometimes present.

Pyr., etc.—B.B. frequently liecrepltalea, fuses at 8 to a wblte pearl, coloring the Same
stroutiii-red; llie fused muss resets alkaline. On cbar&Tal fuses, and in E.P. U converted
Into :i dilHciLltly fusible hepulic mass; this treated with hydrochloric acid and alcoli"!
gives iiQ inteisi"ly red flume. With soda on clinrcoal reacts like barite. Tnsoluble in acids.

Diff.—Characterized by form, cleavnge, high speciSc gravity, red coloration of the
blowi>i|>e ll^ime. Does not effervesce with acids like the carbonates (e.g., stroDtianite);
specilip gravity lower than that of barite.

Obs.^Usu lily associated with llmeatone, or sandstone of various ages; occasionally with
metaltiferotia ores, as with galena and splinlerite at Condorcet, France; nt lte7.iiiiiv»,
Hiiagiiry; nlsio in beds of gypsum, rock salt, as at Bei, Swizerlsnd, laobl. Austria. Liioe-
Iwrg. Hannover; sometimes fills cavities in foasile. e.g., ammonites; with sulphur in some
volcanic regions as at Qirgentl, Sicily.

Specimens, finely crystalized, of a bluish tint, are found In limestone about Lake Huron,
particularly on Drummond Isiaod, also on Btrontian Is., Put-in-Bay, L. Erir-, and at
Kingston In Ontiirlo, Canada; Chaumont Bnj. L. Ontario, Schoharie, and Lnchport, N. Y. A
blue fibrous oelestite oeciim at Bell"- Mills. BInIr Co . Peon. In MinemI Co.. TT. Virginia,
a few miles south ot Cumberland, Md., in pyramidal blue ciyslalf. In Texas,

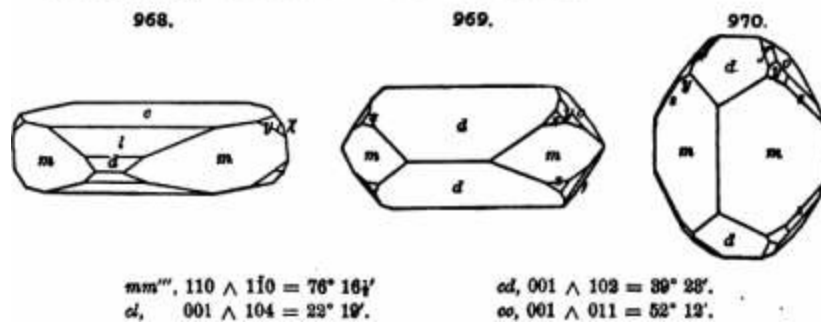
;,CoogIC

BL Idmpoau, large crystals. Witb coletnaaltA at Death Valley, Bac BeniArdliiO Cu.,
Califoruia. lu Cauoda, m cryatuiltne masBea at Kingston, Froatenac Co.; Lansdowu^
Leeds Co ; In radiatlnj^ fibrous musses in llie Laureutlnn of Renfrew Co.

Named from ealtttu, aU*Ual, in uHusIod to the faint shades of blue often piesent.

Airai^SITE. BlelTltrlol, VHriolblelerz Germ. Orthorhombio. Axes &xh\i = 0-7852 :1 : 13

Orthorhombic. Axes $d : b : c = 0.7852 : 1 : 1.2894$.



mm'' , $110 \wedge 1\bar{1}0 = 76^\circ 16'$
 cd , $001 \wedge 104 = 22^\circ 19'$

cd , $001 \wedge 102 = 39^\circ 28'$
 oo , $001 \wedge 011 = 52^\circ 12'$

Crystals sometimes tabular | c; more often prismatic in habit, and in all the three axial directions, m, d, o, predominating in the different cases; pyramidal of varied types. Also massive, granular to compact; stEilactitic; nodular.

Cleavage: c, m distinct, but interrupted. Fracture conchoidal. Very brittle. H. = a'70-3. Q. — 63-6'39. Luster highly adamantine in Home specimens, in others inclining to resinous and vitreous. Color white, tinged yellow, gray, green, and sometimes blue. Streak uncolored. Transparent to opaque. Optically +. Ax. pi. [J. Bx J. a. $2H^\wedge = 89^\circ 53'$, Di. Indices measured by Arzrnni:

2V

1-86061 1-87709 1'6»H9 1-87838

1-87539 187260

SV (calc.)

TM_i^*

100.

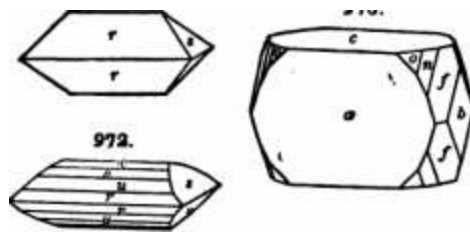
Comp.—Lead sulphate, $PbSO_4$, = Sulphur trioxide 26.4, lead oxide 73.6 = 100.

Pyr., etc.—B.B. decrepitates, fuses in the flame of a candle (P. = 1B). On charcoal in O.F. fuses to a clear pearl, which on cooling becomes milk-white; in R.F. is reduced with effervescence to metallic lead. With soda on charcoal in R.F. gives metallic lead, and the soda is absorbed by the charcoal. Difficultly soluble in nitric acid.

XMff. —Characterized by high specific gravity adamantine luster; cleavage; and by yielding lead B.B. Cerussite effervesces in nitric acid.

mine in Anglesea; at Leadhills; in Derbyshire and in Cumberland in crystals; in the Harz; near Siegen in Prussia; Barchin and Badenweiler in Baden; Felsobánya and elsewhere in Hungary; Nerchinsk, Ilberia and Mt. Mente Poni, Sardinia; Oranada; massive in Siberia. Andalusia: in Almadén, whence it is exported to England. In the Sierra Gorda, Mexico, in immense quantities mostly massive.

Id the United States In cmtals at Wheatlcly's mine, Phenizvllle, Pa.; in Missouri lead mlues; in ciystala of varied habit at tbe Mouniain View rolue, Carroll Co., Maryland. In



Colorado at TffriouB poiU. but less connuon than ceruaalle. At the Ceno Oordo miaea of CeKforuU large□ liferous galeoa). wUli other lead mlueraU. In AiiXoot,, In the mines o1 the Castle Dome district, Tuma Co., and elsewhere.

Named from tbc locailly, Aiiglesea, nhere it was flrit found.

ANHTDRITB.

Orthorhombio. Aiea d:h:i = 0-8933 :1:1-0008.

t^ . Oil A Oil = 90* 8* io. o10 A HI = W* 18*

Twins: 1, tw. pi. d (012); 2, r (101) occaaiOD&lly as tw. lamellfe. Crystals oil not common, thick tabular,

^^3 also prismatic | ititia h. TJsn-

ally massive, clearabie, fibrous, lamellar, granular, and sometimes impalpable.

Cleavage: in the three pin-acoidal directions yielding rectangular fragments bnt with varying ease, thus, c very perfect; b also perfect; a somewhat less so. Fracture uneven, Wl, 97a, Stastfurt, Hbg. V3S, Ausae, Id. sometimes splinterv. Brittle.

H. = 3-3-5. G. = 2-899-2-985. Luster: c pearly, especially after beating in a closed tube; a somewhat greasy, b vitreous; in massive varieties, vitreous inclining to pearly. Color white, sometimes a grayish, bluish, or reddish tinge; also brick-red. Streak grayish white. Optically +. Ax. pi. o J. Bx ± a. SE^ = TO. /3 = 1-576.

Vat. —1. Ordinary, (n) Crysiallized; crysiils rare, more commoul; massive and cleav-able in ils Ibree reetauguiar directiona. (6) Fibrous; either pm-allel. i-adiated or plumose.

ie) Fine graular. (d) Scaly granular. Vulpiniit is a scnly granular kind from Vulpino n Lomliard); it is cut and jiolisbed fnr ol-nanitntnl purposes. A kind in contorted coucre- tioDary forms is the iripestone [OeirdtsUin or SfJilnngjeuilabatier). a. Pitudomorphout; in cubes after ruck-salt.

Comp. —Anhydrous calcium sulphate, CaSO_4 , = Sulphur trioxide 58.8, lime 41.3 = 100.

ID enainpl-vriih schIa lloes not fuse to a clear globule, and is uot aisorbed by the coal like be'ritr; is, lioweir, decomp'ised. niid .vielids a muss which binckpns silver. i>oliible in hydrochloric atid.

riff.—Chitrncteriieil by its cleavage in Ihi-ee reclsnguliir dlTeclions {pseudo-cubic tii napeon: harder than gyraum; doesnrt effervesce with acids like the carbonates.

Oba.— Occurs in tockb of various nges. especinlly in limestone stmla.aDd often the sniie that conlnii ordinary gypsum, and also very commonly In beds of rock-salt: at llie salt mine near Hall In Tvrol; of Bex, Switzerland: iil Aussce, cryBlulllzed and massive; Lnneburjr. Hannover; Eapnik in Hungary; Wieliczka In Polsnd; Ischl in Upper Aiiisria: Berchlexgnden in Bavaria; Slassrurt. In line crystals, embedded fn kieserile; In cavities in lava at Sant'Hn.

In ilieU. State*. MLockporl. N, T.,flneblue. Id geodes of black limestone, wilL calcile and gypsum; in limestone at Nashville, Tenn., etc In Nova Scotia it forms extensive

Anhydrite by absorption of moisture changes to gypsum. Extensive beds are sometimes thus altered In part or throughout, as at Bex, In Switzerland, where, bydigginj: down 60 to 100 ft., the unaltered nnliydrite may be found. Sometimes specimens of anhydrite are altered between the folia or over the exterior.

Zinkoalte. ZnSO_4 . Reported as occurrtng at a mine In the Sierra Almngrera, Spain.

Hrdocranite. CuSO_4 . Fouod at Vesuvius as a pale green to blue Incrustation after the eruption of 1868.

3ULPHA.TS8, CHBOHA.TES, STO.

OROOOITB. Rothblelerz, Oerm. Monoclinio. A^es & :i li: =

: 1 : 0-9159; 0-11"

974.

Crystals nauall; prismatio, habit varied. Also imperfectly columnar aod granular.

Cleavage: m rather distinct; c, a less so. Fracture small con-choidal to uneven. Sectile. H. = 2-5-3. 6. = 6*9-6'l. Luster ndamantioe to vitreoue. Color various shades of bright hyacinth-red. Streak orange-yellow. Translucent.

Gomp.—Lead chromate, PbCrO_4 . = Chromium trioxide 31.1, lead protoxide 68.9 ^ 100.

Pyr , •to.—lu the closed tube decrepitates, blackens, but recovers Its ■irij^iuikl color on

cooling. B.B. Fuses At 1'5. and on charcoal U reduced to metallic lead with deflagration, leaving a residue of chromium oxide, and giving a lead coating. With salt of potassium borate gives an emerald-green bead in both flames.

Ob*.—First found at Berezov, in crystal in quartz veins; also at Hirsfeld and near Nizhni Tugilak in the Urals; in Brazil, at Couffonlândia do Campo; at Rezbánya in Hungary, Holdam in Hungary; on Luzon, one of the Philippines; from Tasmania; in the Vulture district, Haricopa Co., Arizona.

The name Oroonite is from Kofion, saffron.

Phonochroite. PbCrO₄. A basic lead chromate, 9PbO.2CrO₃. In crystals and massive. Color between cochineal- and hyacinth-red. From Berezov in the Ural.

If lead, perhaps 3(Pb,Cu)CO₃.(Pb,Cu)₂SO₄. Color green to brown, from Berezov in the Sulphates with Chlorides, Carbonates, etc.—In part hydrous.

Crystallography, p. 17

Monoclinic. Axes $a : b : c = 1.000 : 0.7476 : 2.2154$; $\beta = 90^\circ$

$\alpha = 90^\circ$

$\alpha, 001 \text{ } \angle = 68^\circ 31'$. $w, 001 \text{ } \angle = 80^\circ 64'$.

Twins: tw. pi. m, analogous to aragonite. Crystal commonly tabular on c.

Cleavage: c very perfect; a in traces. Fracture conchoidal, scarcely observable. Rather sectile. H. = 2-5. G. = 6.26-6.44. Luster of c pearly, other parts resinous, somewhat adamantine. Color white, passing into yellow, green, or gray. Streak uncolored. Transparent to translucent.

Comp.—Sulphato-carbonate of lead, perhaps 4PbO.S₂O₃.2CO₂.H₂O=Sulphur trioxide 7-4, carbon dioxide 8-2, lead oxide 82.7, water 17 ~ 100.

Pyro., etc.—B.B. Intumesces, fuses at 1'6, and turns yellow; but becomes white on cooling. Easily reduced on charcoal. With soda affords the reaction for sulphuric acid. Effervesces briskly in nitric acid, and leaves white lead sulphate undissolved. Yields water in the closed tube.

Obs.—Found at Leadhills, with other ores of lead: Red Gill, Cumberland; at Matlock Derbyshire. From the Haiti Calxelt lead mine near Iglesias. Sardinia (malme). Observed from Arizona, at the Scobuk gold mine with wulfenite, vanadinite, cerussite; partly altered to cerussite.

SoBAKiHTB. Regarded at one time as rhombohedral and dimorphous with leadhillite, but probably only a modification of that species. From the Busantia mine, Leadhills, Scotland.

by Google

DESCRIPTION.

Color. Pale greenish yellow dodecahedral. From Bonanza Lake, Benavente Co., Cal.

Occurrence. Perhaps $Pb(OH)Cl$. As crystals. Colorless. From Atlatl.

Crystallography. $UgSO_4 \cdot KCl + 8H_2O$. UbuqUj granular massive and crusts. Color white to dark blue. From St. Hefurt; Kulusz, Gulick.

Occurrence, Arizona, San Bernardino Co., Cal. In lixiviation pits, about prismatic to tabular in quartzoids (Figs. 202, 205, p. 60). Color white to yellow. From Bonanza Lake, San Bernardino Co., California; also from Death Valley, Inyo Co.

B. Acid and Basic Sulphates.

In fibrous form and white color.

DESCRIPTION.

Orthorhombic. Axes $a:b:c = 0.7739:1:0.4871$.

In groups of prismatic acicular crystals ($mm = 75^\circ 28'$) and drusy crusts; massive with reniform structure.

Cleavage: b very perfect; m in traces. Fracture uneven. H. = 3.5-4.

G. = 3.907. Luster vitreous; a little pearly on the cleavage-face b. Color

emerald-green, blackish green. Streak paler green. Transparent to translucent

Comp.—A basic sulphate of copper, $CuSO_4 \cdot 3Cu(OH)_2$, or $4CuO \cdot SO_3 \cdot 3H_2O$ =

Sulphur trioxide 17.7, cupric oxide 70.3, water 12.0 = 100.

Pyrolysis, etc.—Yield water, and at higher temperature sulphuric acid, in the closed tube, and becomes black. B.B. to fusa, and on charcoal affords metallic copper. With soda gives

the reaction for sulphuric acid.

Oba.—Occurs in the Urni; the kOnigloe (or ki/nigiU) was from Qtlmesfaevsk; near Ronghlen Gill, in Cumberland; in Cornwall (in nnn teaHngUmiU); tit Kezbinjra: In small hedH at EriBiivIg In Iceland (kriruiiffitey. in Mexico (bronffnartinti; AtBcama. In Ibe U. IJlBleB. At Monarch mlDC, Chaffee Co., Colorado; In Uinb, Tlnllc district, at the Mam> moth mine.

Iianarklte. Basic lead silicate. Pb_3SO_4 . In monoclinic crystals. Color greenish white, pale yellow or gray. From Leadhills, Scotland; Siberia; the Harz.

Dol*rophanita. A basic cupric sulphate. Cu_2SO_4 (?). In small brown monoclinic crystals. From Tesuvlus (eruption of 1869).

OalAdonits A basic silicate of lead and copper, perhaps $2(Pb,Cu)_2O \cdot SO_3 \cdot H_2O$. In small prismatic crystals. Color deep verdigris-green or blue-green. From Leadhills; Bed Qili, Cumberland. etc: Inyo Co., Cal.

Unarite. A basic silicate of lead and copper. $(Pb,Cu)_2O \cdot (Pb,Cu)_2(OH)_2$. In blue monoclinic crystals. From Leadhills; Cumberland; the Ural, etc. Also Inyo Co., California.

AKTLnRiTE. Perhaps $3CuSO_4 \cdot TCu(OH)_2$. In light green soft lumps. From the Antler mine. Hohave Co., Arizona.

Alnmlan. Perhaps $Al_2O_3 \cdot 3BO_2$. White crystalline or massive. Bleira Almagrem, Spain.

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C. Nonnal Hydrous Sulphates.

Three well-characterized groups are included here. Two of these, the $EP_8O_{11}T_8$ GuotJp and the Mklantbiute Ghodp, have the same general formula, $K_2SO_4 \cdot nH_2O$, in the first the crystallization is orthorhombic, in the second monoclinic. The species are best known from the artificial crystallization of the laboratory; the native minerals are rarely crystallized. There is also the isometric Alum Qboep, to which the same remark is applicable.

Glauber Salt.

Monoclinic. Crystals like pyroxene in habit and angle. Usually in efflorescent masses.

Crystal: a perfect; c, A, in traces. H. = 1-5-2. G. = 1'481. Luster vitreous. Color white. Transparent to opaque. Taste cool, then feebly saline and bitter.

Comp.—Hydrous sodium sulphate, $Na_2SO_4 \cdot 10H_2O$ = Sulphur trioxide 24-8, soda 19'3,

water $55-9 = 100$.

Pyr., rto;— Id the closed tube much water; gives an Intense yellow lo ttae Bame. Very loluble <n water. Loses Its water on exposure to dry nlr and falls to powder.

Obs.—Occurs at Isclil. Hallstndt, huJ Auasee fn Upper Austria: also fn Huncary, SwilEerlaod. Imlv; nt the bot sprlugs at Csrl-^bad, etc. Laige quantities of this sodium Biilpbate are oblalii'd from the waters of Qreat Salt Lake, Utnb.

Klaserlte. MgSO. + H₂O, Usually massive, (jgranular to compact. Color white, gmyish, yellowish. From StassFurt; Hallstadli ludla.

SuniUt*. HuSO₄ + H₂O. StalacUllc. WhItUb, reddisb. From FelsOb&nya, Hungary.

OTPSUM.

Monoclinic Axes ii:\$:6 : =:(>•

mm^{'''}, no Alio = 88° SO*.

ed. 001 A 101 = 28" ir,

et, 001 A 101 = 88* 8i',

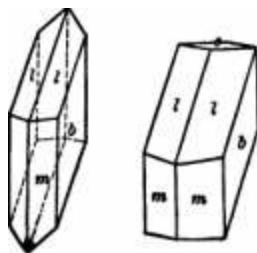
ft), 001 A 103= 11° aff.

w*. on A Oil = 44* 17V,

»7S. B7«,

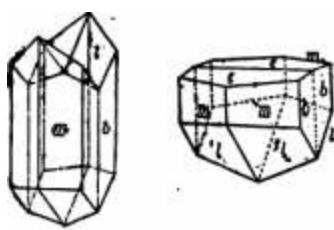
:1:0-41S4; /? = 80'«'. le. Ill A 111 = 86° ly. n< 111 A 111 = 41° 20'. ml. lloAlll = 48' V. mn. 110 a ill = 69° ly.

»77.



^^

Hs^



Crystals small, simple in habit, common form flattened | b or prismatic to acicular | i; again prismatic by extension of I (111'). Also lenticular by

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rounding of I (111) and e (103). The form e (103), whose faces are usually rough and convex, is nearly at right angles to the vertical axis (edge m/m'), hence the apparent hemimorphic character of the twin (Fig. 979). Simple crystals often with warped as well as curved surfaces. Also associated massive; lamellar brittle; often granular massive; and sometimes nearly impalpable. Twins: tw, pi. a, very common, often the familiar awallow-tail twins.

Fracture: b eminent, yielding easily thin polished folia; a (100), giving a surface with conchoidal fracture; n (111), with a fibrous fracture [t (101); a cleavage fragment has the rhombic form of Fig. 549, p. 322, with plane angles of 66° and 114°. H. = 1-5-3. G. = 3-314-3-328, when in pure crystals. Luster of J pearly and shining, other faces subvitreous. Massive varieties often glistening, sometimes earthy. Color usually white; sometimes gray, flesh-red, honey-yellow, ochre-yellow, blue; impure varieties often black, brown, red, or reddish brown. Streak white. Transparent to opaque.

Optically +. Ax. pi. | J, and Bi A t! = -i- 52° (at 9-4° C), (cf. Figs. 549, 550, p. 222}. Dispersion p > v; also inclined strong. Bi, a bit, ^ 0° 30'. 2Er ^ 95° 14' at 20". /S, = 1-5226, y - a := 0-009. On the effect of heat on the optical properties, see p. 225,

Var.,—1. Oryzallite or Selenite/ colorless, transparent: in distinct crystals. or broad folia, often large. Usually flexible and yields a fibrous fracture | t (101), but the variety from Hottentotta rather brittle.

2. Fibrous; common or Selenite. Called Batin spar, when fibrous. with pebbly opal-

8. Mastichite; J. Offer. a fibrous variety, white or delicately shaded earthy quartz rock, a dull-colored rock, often impure with clay, calcium carbonate or silica. Also, in caves, curious curved forms, often grouped like rosettes and other shapes.

Composition.—Hydrous calcium sulphate, CaSO₄ · 2H₂O = Sulphur trioxide 46.6, lime 32.5, water 20.9 = 100.

' Pyr., etc.—In a closed tube gives off water and becomes opaque. Fuses at 2-5-3, coloring the flame reddish yellow. For other reactions see Abhandl., p 528. Infused at a

temperature not exceeding 290°C , it again combines with water when moistened, and becomes firmly solid. Soluble in hydrochloric acid, and only in 400 to 600 parts of water.

Distinguished by its softness in all varieties, and by its crystallized kinds; does not effervesce with acids like calcite, nor gelatinize like the zeolites; harder than talc and yields much water in the tube.

Gypsum often forms extensive beds in connection with various stratified rocks, especially limestones, and shales or clay beds. It occurs occasionally in crystalline rocks. It is also a product of volcanoes, occurring about fumaroles, or where sulphur gases are escaping, being formed from the sulphuric acid generated, and the lime afforded by the decomposing lavas. It is also produced by the decomposition of pyrite when lime is present. Gypsum is also deposited on the evaporation of sea-water and brines, in which it exists in solution.

Fine specimens are found in the salt mines of Bex in Switzerland; Hall in Tyrol; the sulphur mine of Sicily; in the Hay of Sibley Hill, near Oxford; and large lenticular crystals at Montmartre, near Paris. A noted locality of alabaster occurs at Canselino, 8.5 m. from Leghorn, whence it is taken to Florence for the manufacture of vases, figures, etc.

Occurs in extensive beds in several of the United States, and more particularly in New York, Ohio, Illinois, Virginia, Tennessee, and Arkansas, and is usually associated with salt springs, also with rock salt. Also on a large scale in Nova Scotia, etc.

Handsome solenoidal and snowy gypsum occurs in New York, near Lockport in limestone. In Maryland, large grouped crystals on the Mt. Mary's in clay. In Ohio, large transparent crystals have been found at Ellsworth and Canfield, Trumbull Co. In Tenn., see and alabaster in Davidson Co. In Kentucky, in Mammoth Cave. It has the forms of rosettes, or flowers, vines, and shrubbery. Also common in isolated crystals and masses. In the Cretaceous clays in the western U. S. in enormous crystals, several feet in length, in Wayne Co., Va. In N. Scotia, in Sussex, Kings Co., large single and grouped crystals, which mostly contain much symmetrically disseminated sand.

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Named from gypsum, the Greek for the mineral, but more especially (or the ancient mineral). The derivation is from $\gamma\psi\mu\sigma$, earth, and $\sigma\mu\sigma$, to cook, corresponding with the use of the word among the Greeks.

Burnt gypsum is called Plaster-Paris, because the Montmartre gypsum quarries, near Paris, are, and have long been, famous for affording it.

$(\text{Mg}, \text{Zn}, \text{Fe})\text{SO}_4 \cdot 4\text{H}_2\text{O}$. In loosely adherent aggregates. Color clear green. From Colorado.

Epsomite Group. $\text{MgSO}_4 \cdot 7\text{H}_2\text{O}$. Orthorhombic.

Epsomite $\text{MgSO}_4 \cdot 7\text{H}_2\text{O}$ d:h:i = 0-9902 : 1 : 05709

(Fe,Mg)SO₄ · 7H₂O Oodardite ZnSO₄ · 7H₂O 0-9807 :1: 0-5631

Ferro-goslarite (Zn,Fe)SO₄ · 7H₂O Morenoiite NiSO₄ · 7H₂O 0-9816:1: 05655

EPSOMITE. Epsom Salt.

Orthorhombic. Usually in botryoidal masses and delicately fibrous crusts. Cleavage: b very perfect. Fracture conchoidal. H. = 2.0-2.5. G. = 1.751. Luster vitreous to earthy. Streak and color white. Transparent to translucent. Taste bitter and saline.

Comp.—Hydrogen sulphate, $\text{MgSO}_4 \cdot 7\text{H}_2\text{O}$ = Sulphur trioxide 32.5, magnesia 163, water 51.2 = 100.

Obs.—Common in mineral waters, and as a delicate fibrous or capillary efflorescence on rocks, in the galleries of mines, and elsewhere. In the former slate it exists at Epsom, England, and at Bedlitz and Snidschitz (or Saischitz) in Bohemia. At Iclis in Cornwall it occurs in silky fibers, and is hence called hair by the workmen. Also obtained at the gypsum quarries of Hontmartre, near Paris. Also found at Vesuvius, at the eruptions of 1850.

The floors of the limestone caves of Kentucky, Tennessee, and Indiana, are in many instances covered with epsomite, in minute crystals, mingled with the earth. In the Mammoth Cave, Ky., it adheres to the roof in loose masses like snowballs.

Other salts. $\text{ZnSO}_4 \cdot 7\text{H}_2\text{O}$. Commonly massive. Color white, reddish, yellowish. Formed by the decomposition of sphalerite. First found in the passages of mines, as at the Itammelsberg mine near Ooslar, in the Harz, etc. In Montana at the Gagnon mine, Bunte. Ferrous sulphate (48 p. c. FeSO_4) occurs with sphalerite at Webb City, Jasper Co., Missouri.

Morenoite. $\text{NiSO}_4 \cdot 7\text{H}_2\text{O}$. In acicular crystals; also fibrous, and an efflorescence. Color apple-green to greenish white. A result of the alteration of uterine ores, as at Cape Hordale, in Galicia; Röchelsdorf, in Hesse, etc.

Cupromagnesianite $(\text{Mg,Fe})\text{SO}_4 \cdot 7\text{H}_2\text{O}$

D.,.,.,cb,Google

Chalocite $\text{Cu}_2\text{SO}_4 \cdot 5\text{H}_2\text{O}$ Triclinic

d:h :i = 0-5656 : 1 : 05507; a = 82° 21', /b = 73" 11', c = 77° 37'

The species here included are the ordinary vitriole. They are identical in general formula

with the species of the Epsomite group, and are regarded as the Slime compound essentially under oblique crystallization. The copper Sulphate diverges from the others in crystallization, and contains but 5 of water.

HELANDTEIRITB. Coppeiu.

Monoclinic. Usually capillary, fibrous, atalactic, and concretionary; also massive, pulverulent. Cleavage: c perfect; m less so. Fracture conchoidal. Brittle. H. = 2. G. = 1.89-1.90. Luster vitreous. Color, various shades of green, passing into white; becoming yellowish on exposure. Streak uncolored. Subtransparent to translucent. Taste sweetish, astringent, and metallic.

Comp.—Hydrous ferrous sulphate, $\text{FeSO}_4 \cdot 7\text{H}_2\text{O}$ — Sulphur trioxide 28.8, iron protoxide 25.9, water 45.3 = 100. Manganese and magnesium sometimes replace part of the iron.

Ofa.—Proceeds from the decomposition of pyrite or marcasite: thus near Ooslar in the Harz; Bodefeld in Bavaria; Felsa, Bnei-lei, and Elmberg. Usually accompanied by pyrite in the U. States. ■■ by biorescence. Lueke (IB p. c. CuO) is from the "Lucky Boy" mine, Butterfield County, Utah.

Massive. $\text{H}_2\text{SO}_4 \cdot 7\text{H}_2\text{O}$. Fibrous, massive; colorless. From the mine "Lucky Boy," south of Salt Lake, Utah.

Pisumite. $(\text{Fe}, \text{Cu})\text{SO}_4 \cdot 7\text{H}_2\text{O}$. CuO 10 to 16 p. c. In concretionary and atalactic forms. Color blue. From Turkey.

Basic copper-iron vitriol near Pisumite. From the mine "Lucky Boy," Utah.

Berberite. $\text{CoSO}_4 \cdot \text{TH}_2\text{O}$. Usually in stalactites and crusts. Color flesh- and rose-red. From Berber, in Hesse, etc. Capromaffone is from Vesuvius.

OHAIANTHTE. Blue Vitriol. Kupferblau Oerm.

Trigonal. Crystals commonly flattened on (111) (Figs. 339-341, p. 109). Occurs also massive, stalactitic, reniform, sometimes with fibrous structure.

Cleavage: M, m, p imperfect. Fracture conchoidal. Brittle. H. = 3.5. G. = 2.12-2.30. Luster vitreous. Color Berlin-blue to sky-blue, of different shades; sometimes a little greenish. Streak uncolored. Subtransparent to translucent. Taste metallic and nauseous.

Comp.—Hydrous cupric sulphate, $\text{CuSO}_4 \cdot 6\text{H}_2\text{O}$ = Sulphur trioxide 32.1, cupric oxide 31.8, water 36.1 = 100.

Pyr., etc. — In the closed tube yields water, and at a higher temperature sulphur trioxide. B.B. with charcoal yields metallic copper. With the fluxes reacts for copper.

Soluble In water; a drop of Ibe solulton plscond on a surface of ii-on cnat» it with

; Parys mine, Anglesen: at various mines In Co. of Wlcklow; Rio Tiiito mine. Spain. Trom
tbe Hlwnssee copper mine, alan In larice quan-ttlea Ht other mines, in Polk Co.,
Teimeaaee. In Arizona, near CUftou, Graham Co.. and Jerome, Tavapal Co.

SyngMilta. KaUiaztte, CaSO₃,R₂gO. + HiO. In prlsmalte (monocllDfc) eryitala. Colorless or
milky-white. From Ealusz, Qalicia.

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Xittwalt*. HgSO₄.Nai8ot + 3iH₂o. Massive, clWTable. Color paUvellow. From Ischl,
Aiturin.

BlSdita. Hg60<.Niii6oi + 4HiO. Ciystals abort priaoi&ttc, monocllDic; also massive
gmnular or coinpscl. Colorless lo greenUb, jellowisb, red. From the salt mines of lacbl; «t
Uallsladl (rimorailey, at Slassfurt; tie salt lakes of Aslrakbao iaiWalianiU), Asia; Iodia,
Cbili. tic.

Itsonlte. Hg8o₂,Et8O. + 4HiO. In monocllnic crystals from Westeregela and Leo-
polilsbiill. Germtiny.

Boosaliiganltte. (KH₂),SO₄.Mg8o4 + 6H₂o. From the boric acid logoonl, Tuscany.

Piciotneiita. MgBO₂,E,BO₂, + eH₂o. As a whils cTTStalline Incnislatlou. From TesiiTiiis
with eyanoehroitt, iin isomorpboiis species la nblch copper replaces the mag* Desium.
Also at liilassfurt {lekotuitt); at Ascberslebea; Galusz la JCast Galicia.

Polyhalite. gCaSO₄, MgSO₄.E.SOt + 3H₂o. Usually In compact fibrous or lamellar inasses.
Color flesb- or brick-red. Occurs at tbe mines of Ischl, HalUtadt, etc. Id Austria;
BurchlesgadeQ, Bavaria; Slassfitrt.

Alum Oroup. Isometric.

RA1(So₄), + 12H₂o or R.SO₄.Al₂(SO₄)₃. + 24H₂o. Sali&ite Potash Alum EAl(SO₄)₂, + 12H₂o

TMli«niiigito Ammonia Alum (NH₄)Al(SO₄)₂, + 12H₂o

Xendozite Soda Alum liaAl(SO₄)₂, + 13H₂o

The Alums proper are isometric in cryatallizatioD and, cheroioally, are faydrona sulphates
of sItimiDium with an alkali metal and 13 (i.e., if the formula is doubled, 24) molecules of
water. The speciea here included occur very spiiringly in nature, and are best known in
arttfioiol form in tbe laboratory.

The Halotrichites are oblique in crystallization, very commonly fibrous in structure, and

are hydrous sulphates of aluminium with magnesium, manganese, etc.; the amount of water in some cases is given as 23 molecules, and in others 34, but it is not always easy to decide between the two. Here belong:

Pickeringite, $MgSO_4 \cdot Al_2(SO_4)_3 \cdot 22H_2O$. In long fibrous masses; also in efflorescences.

Halotrichite. $Al_2(SO_4)_3 \cdot 8H_2O$. In yellowish silky fibrous

crystals. $MnSO_4 \cdot Al_2(SO_4)_3 \cdot 94H_2O$. In fibrous masses; also as efflorescences. Diatloite.

$(Zn, Fe, Mn)_2(SO_4)_3 \cdot 32H_2O$. A fibrous alum from Egypt supposed to contain a new element, maarlum.

Opuntia. $Al_2(SO_4)_3 \cdot 9H_2O$, rhombohedral. Crystals. Color white, yellowish, or brownish. From the Tierra Amarilla near Coquimbo, Chile (not from Coquimbo.)

Quartzite. $Al_2(SO_4)_3 \cdot 10H_2O$. In reddish tabular crystals. From Coquimbo, Chile.

Ilmenite. $FeTiO_3$. An orange yellow efflorescence on graphite. From Himmelsberg, Bohemia.

Almandine. $Al_2(SO_4)_3 \cdot 18H_2O$. In dense fibrous masses or crystals; massive. Color white, or tinged with yellow or red. From near Binn, Tyrol; Bodenbach, Prussia; Pnau, Tyrol.

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DE&CaiPTITE HINBaALOST.

Philofix. Perhaps $CuSO_4 \cdot Fe_2(SO_4)_3 \cdot nH_2O$. In blue fibrous masses. Chile. Found at the copper mine in the Cordillera of Coquimbo, province of Santiago, Chile.

Forronatrium. $8Na_2SO_4 \cdot Fe_2(SO_4)_3 \cdot 6H_2O$. Rarely in acicular crystals; usually in prismatic forms. Color greenish or gray to white. From Sierra Gorda near Coahuila, Mexico.

Chile.

Basic Hydrous Sulphates.

Langite. $CuSO_4 \cdot 5Cu(OH)_2 \cdot H_2O$. Usually in fibrous-lamellar, crystalline masses. Color blue to greenish blue. From Cornwall.

Herrenbergite. $3(CuOH)_2SO_4 \cdot Cu(OH)_2 \cdot 5H_2O$ with one-fifth of the copper replaced by iron. In tabular crystals; usually in prismatic groups. Color emerald-green, bluish green. From Uerregründ, Hungary.

KaoxvUIite. A, hydrous basic sulphate of chromium, ferric iron, and aluminium. In rhombic plates. Color greenish yellow. From the Redbiou mercury mine, Enovill[^] California.

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ftBDnroTOinTE. A hydrous chromium sulphate, in thin fibrous masses of a pale[^] purple color. Same locality as Knox^{*}ite.

Oypnulto. Perhaps $7\text{Fe}_2\text{O}_3 \cdot \text{Al}_2\text{O}_3 \cdot 10\text{B}_2\text{O}_3 \cdot 14\text{H}_2\text{O}$. An aggregation of microscopic crystals. Color yellowish. From the Isobod of Cyprus.

Alniliite (Websterite). $\text{Al}_2\text{O}_3 \cdot \text{SO}_3 \cdot \text{H}_2\text{O}$. Usually in white earthy reniform masses, compact. From near Halle, in clay: also at Nehen[^]n, Suasei. and elsewhere.

Paralimnites. Near aluminite. but supposed to be $2\text{Al}_2\text{O}_3 \cdot \text{B}_2\text{O}_3 \cdot 15\text{H}_2\text{O}$.

Felsparite. $2\text{Al}_2\text{O}_3 \cdot \text{SO}_3 \cdot 10\text{H}_2\text{O}$. Ussairei-ite; concretions. Color snow-white. From near Felspar, Hungary.

$\text{Fe}_2\text{O}_3 \cdot 4\text{SO}_3 \cdot 18\text{H}_2\text{O}$. Usually in reniform and botryoidal, Sweden: [^]sofroi Pen

Sideroatite. $2\text{Na}_2\text{O} \cdot \text{Fe}_2\text{O}_3 \cdot 4\text{SO}_3 \cdot 7\text{H}_2\text{O}$. Fibrous, massive. Color yellow. From the province of Tampico, Chiapas. Also on the Urus plateau, near Sitakiyu, on the island, Cht-lekei, in the Oispiun Set, (vntilei,

Voltaite. Perhaps $5(\text{K}, \text{Fe})\text{O} \cdot 2(\text{Al}, \text{Fe})_2\text{O}_3 \cdot 10\text{SO}_3 \cdot 18\text{H}_2\text{O}$. In octahedrons, etc. Color dull orange-green to brown or black. From the solfataras near Naples; also Persia.

Matschertite. Perhaps $6(\text{K}, \text{Na}, \text{Fe})\text{O} \cdot \text{SO}_3 \cdot 18\text{H}_2\text{O}$. In aggregates of minute yellow crystals. Occurs with Tolu in Persia.

Alumstone. Alumstone. Alumstone.

Rhombohedral. Axis $a = 1.3520$. In rhombohedrons, resembling cubes ($\alpha = 90^\circ 50'$). Also massive, having a fibrous, granular, or impalpable texture.

Cleavage: c distinct; r in traces. Fracture flat conchoidal, uneven; of massive varieties splintery; and sometimes earthy. Brittle. H. = 3-5-4. G. = 3.08-3.25. Luster of r vitreous, basal plane somewhat pearly. Color white, sometimes grayish or reddish. Streak white. Transparent to subtranslucent. Optically positive.

Comp. -Hydrous sulphate of aluminium and potassium, $\text{K}_2\text{O} \cdot 3\text{Al}_2\text{O}_3 \cdot 4\text{SO}_3 \cdot 6\text{H}_2\text{O}$ -Sulphur trioxide 38.6, alumina 37.0, potash 11.4, water 13.0 = 100. The formula may be written $\text{K}(\text{AlO})_3(\text{SO}_3)_2 \cdot 3\text{H}_2\text{O}$.

Pyr., etc.—B.B. decrepitates, and is Infusible, In the closed tube yields water, sometimes also aminoum sulniiate. and at a higher temperature sulphurous nud sulphuric oxides. Heated with cobalt solution sffotils a fine bine color. With soda and charcoal infusible, but yields a hepucic mass. Soluble in stitphuiic acid.

Ohi.—Forms seama in IracLytlc and allied rocks, where it has been formed as a result of the alteralloD of the rock by mnnns of siilphiirous vnnors; as at Tolfn, nenr Civira Vucchia; in HiiQcary; on Milo, Grecian Arcliipelago; at Ml. Dore. Prance. In tbc U. 8,, associ:iled with Ulaipore, in rhombohedral ciyslals, tabular through-the presence of e (0001) lit the Rosila Hills, Custer Co., Colorado,

JAROSTTB. OelbelseDers.

Rhombohedral. Axis 6 = 1-2492; rr' = 90" 45', cr = 55° 16'. Often in druses of minute crystals; also fibrous, granular massive; in nodules, or as an incrustation.

Cleavage: (J distinct. Fracture nneven. Brittle. H. = 25-3 5. G. = 3'15-326 cryst. Luster vitreous to subadamuntine; brilliant, also dull. Color ocher-yellow, yellowish brown, clove-brown. Streak yellow, shining.

Comp.—K,o.3Fe,O,.4SO,.6H,O = Sulphur trioxide 31-9, iron aesquioxide 47-9, potash 94, water 10'S = 100. The formula may be written K(FeO> (8o > + 3H.o. '■ *"

Ob«.—The original Oelbeisendrzs was from Lusnitz. between Eolosomk and Bllin, Bohemia, lu brown coal; and Inter from Modum. Norway, in alum slate. The iarosite waa from BarraDCoJaroso, Id the Sierra Almagrera. Spain: also from other points. In the U.S.. on quartz In tlie Vulture mine, Arizona; In ChaOee County, Col.; Tiutlc district, Utah.

..Coogic

S38 DESCBIPTIVE MINBEALOOT.

Lttwiflto. F«TbapBKtO.8&.1,O..4So..9H,O. la rouoded

«luDii«. Found in a coal bed at Tabrze Id ll[]per Silesia; also witli aluoite at Tolfa.

Bttringlt*. Perliaps eCaO.AUOi.SSOi.SSHjO. !□ mlnate colorless acfcular crysUlt. From limeBioiteiDClusfonB is lava, Ufui Hayen, Rlienlsh Prussia.

Quetonita. HgO.Fe,o.,3SOi.l8H,o. HassiTe, Id Indistinct cirslals. Color reddiib browu. Fi^im Quetena, Cliill.

Zinoalumnita. 3ZiiSo..4 lolor wliile. bluish. From L

Johannlta. A brdrous sulpLate of uranium and copper. In druses or raDiform maisM of a

green color. From JoHChimatlia].

DnnopiUta. Perhaps CaUtStOii.25H]o . Id velTety IncnurUtlODS; yellow. From Johanngeoi^en itad t.

ZiipgaUt, mogUar^, ttraeonitt Are uncertain uranium sulphateo, also from Joachlmrthal.

Tellorates; also TellnilteB, Selenltes.

Bi,o,.TeO,.2B,o In earthy incrustations; yellowish to whlU. From Bightand, Moalonn, with tetradymite.

Bmmoaalta. Probably a hydrated ferric tellurite. In thin yellow-green scales. Prom near Tombstone. Arizona.

Dnrdsnite. Hydrous ferric tellurite. Fe,(TeOi)i + 4H,o . In small mammtllary forme; greenbh yellow. Honduras.

Obaloomente. Hydrous cupric selenite, CuSeOi + SH,o . In small blue monoclnic 'Crystals. From the Cerro de Cacbeula. Argentina, with silver, copper selenides.

MoLTBDOMENiTB is lead selenite and CoBuyrouBKim probably cobalt aelenlte, from the same locality as cbalcomenfte.

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TUKGaTA,TBS, 1[oLTBDA.TZS.

Oxygen Salts.

7. TUiraSTATES, XCOLTBDATEB.

The monoclinio Wolframite Group end the tetragonal Scheelite Group are included here.

Wolframite Hiibnerite

Wolft^mlte Group.

$(\text{Fe,MD})\text{WO}$, a:'b:i = 08300

I : 0-8678 ft = $89^\circ 23'$ 0-8362 : 1: 0-8668 89" 7f

WOU'RAMTTEI. Wolfram.

Monoclinio. Axes: a-Azi = 0-8300 : 1 : 0-8678j /?= 6 mm". 110 A lIO = 78° aff. aj/, 100 a 105 = fla" M'. at. 100 A 102 = $81^\circ 54'$. /', Oil A Oil = 81' M'.

/G\

Twins: (1) tw. axis i with a as comp.-face; (2) tw.

f\ k (033), Fig. 419, p. 130. Crystals commonly tabular a; also prismatic. Faces in prismatic zone vertically striated. Often bladed, lamellar, coarse divergent columnar, granular.

Cleavage : b very perfect; also parting Q a, and It (102). Fracture uneven. Brittle. H. = 5-5.5. G. = 7.2-7.5. Luster submetallic, Color dark grayish or brownish black. Streak nearly black. Opaque. Sometimes weakly magnetic.

Comp., Tar.—Tungstate of iron and manganese (Fe,Mn)WO₄. Fe ; Mn = chiefly 4 : 1 {FeO 18.9, MnO 4.7 p, c.) and 2 : 3 (FeO 9.5, MnO 14.0).

Pyr., ate.—Fusion B.B. easily (F. = 26-8) to a globule, which on cooling is brittle and metallic. With inflow of plimptonite five times a clean red yellow glut while boiling, which is paler on cooling; if the bead becomes black, if not too antimonial, the bead assumes on cooling a green color, which continued treatment in R.F. «becomes» to reddish yellow. With soda and silver or platinum foil fused to a bluish gray. Decomposed by aqua regia with separation of tungstic acid as a yellow powder.

8. Decomposed by concentrated sulphuric acid, or even hydrochloric acid, to give a colorless solution, which, treated with metallic zinc, becomes intensely blue, but soon bleaches on dilution.

Obs.—Wolframite is often associated with tin ores; also in quartz, with telluride, blende, pyrite, arsenic, sphalerite, etc. In fine crystals at Bismackwald, Zimmern; Schneeberg, Freiberg, Altenberg; at Nerchinsk, Siberia; Ghantaloup, near Limoges, France; near Redruth and elsewhere in Cornwall with tin ores. In S. America, at Oruro in Bolivia. Wilfite tin stone at various points in New South Wales.

In the States at Lane's mine, Monroe, Conn.; Flowe mine, Hecuba Co., N. C. with scheelite; in Missouri, near Mine la Motte.

H. Near wolframite, but contains 20 to 25 p. c. MnO. Usually in bisect form*, rarely in distinct terminated crystals. Color brown to red to black-brown to nearly

..Ca>o'

QIC

block. Streak yellow to brown, greenish gray. Often translucent. Uamotb dist, Nevada; Olanj County, Col.; near Silverlode, Sui Juan Co.; Black Hills, S. Dakota, etc. Also in Peru, and in uranovite at Adirondack in the Pyrenees.

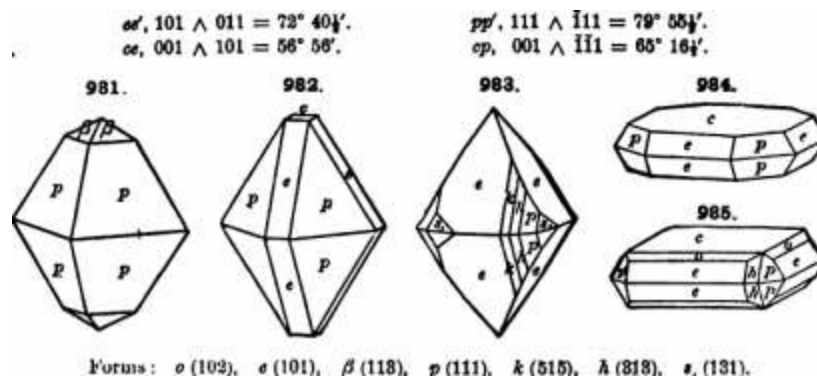
Bcbeelite Ororp. Tetrsgouol-pyramidal.

pp' (111 A ill) = 79° 55' i = 1-5360

80" 1' 6 = 1-5445 80° 15' 6 = 1-5667 80° 23' i = 1-5T71

The ScaEELiTE Group includes the tangBtates and molybdatea of calcium and lead; also copper. In crytallization they belong to the Pyramidal Group of the Tetragonal System. Wulfenite is probably hemimorphic.

Tetragonal-pyramidal. Axis i = 1'5356.



(10aj, it(101). (5(118), p(111), *(51S), A (813). 1,(131).

Twins: (1) tw. pi, a, both contact- and penetration-twins (Fig. 'ITS, p. 125). Habit octahedral, also tabular. Symmetry shown by faces k, h, k (E\g- 983). Also reniform with columnar structure; massive granular.

Cleavage; p (H^ most distinct; e (101) interrupted. Fracture nneyn. Brittle. H. = 4-5-5. G. = 5'9-6'l. Luster vitreous, inclining to adamantine. Color white, yellowish white, pale yellow, brownish, greenish, reddish. Streak white. Transparent to translucent. Optically)-

Indices: $oa^{\wedge} =: V9\backslash9$. e, = 1 934.

Comp.—Calcium tungstate, CaWO, = Tungsten trioxide 806, lime 19-4 = 100.

Molybdenum is usually present (to 8 p. c.). Copper may replace calcium, see cupro-gcheelite.

Pyr., ato,—B.B, Id the forceps fuses at S to a semi-transpaTent glass. Soluble with borax l» a traaspareut glass, whicb afterward b^ comei npnque and crystnlllDe. With itilt of pliciaploruB forms a eIbss. colorless in outer flume, lu inner green when bot. aud fine blue wLeu cold ; Tarleties cootainiug iron require to Iw treated on charcoal with tin before

,i,:cc ..Google

inNOSTATES, MOLTBDATBS.

the blue color appeals. In hydrochloric or nitric acid decomposed, leaving a yellow powder soluble in ammonia.

Ob*.—Bismuth is associated with siliceous rocks, and is commonly found in contact with cassiterite, topaz; fluorite, apatite, molybdenite, or wolframite, but quartz is also associated with gold. Thus at Schacketwald and Zlittwald, Bohemia: Alsenberg, Saxony; Kieselgebirge: the Kupferberg in the Unter-Bohemia (Tyrol); Cinnabar in the Tyrol; Tmversella in Piedmont; Meynig, Corrivaux. France (L'ouisaingTHiOi) Surtout; Pikkimäki in Finland. In New South Wales, at Adolphus, from a gold mine; New Zealand, near Mt. Kanisav, Tasmanian, with cassiterite.

It is found in the U. States, at Leadville, Colorado, and in Trumbull, New York; Mecklenburg Co., N. Carolina; the Mammoth mining district, Nevada; with gold at the Chryse mine, Wilkes Co., Colorado. In quartz veins in Ribblesdale and Harlow, Leicestershire, Quebec.

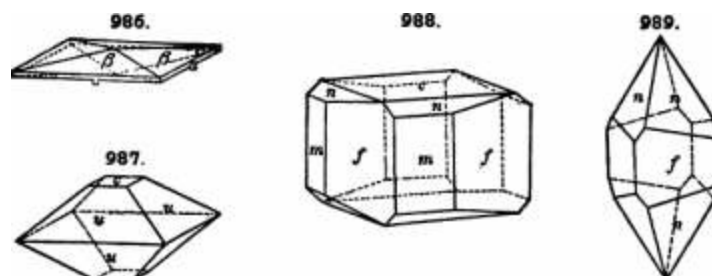
Opuntite. Cupric tungstate, CuWO_4 . From the copper mines of Llanuoch, near Santiago, Chile. Cuprobismutite, from the vicinity of La Paz, Lower California. $\text{Cu}_2\text{W}_2\text{O}_7$, with 6 a. p. c. Cu_2O ; color green.

Powellite. Calcium molybdate with calcium tungstate (10 p. c. WO_3), $\text{Ca}(\text{Mo},\text{W})_2\text{O}_8$. It is a yellow rhombic pyramidal. G. = 4.84. From western Idaho; Houghton Co., Michigan.

Stolzite. Lead tungstate, PbWO_4 . In pyramidal tetragonal crystals. H. = 3-7-8. G. = 7.8-8.13. Color green to gray or brown. Zinnwald.

Raspite. Has the same composition as stolzite, but is referred to the monoclinic system. In rhombic tabular crystals. Color brownish yellow. From the Broken Hill mine, New South Wales.

Wulfenite. Yellow, Molybdenite. G. = 5.77. Tetragonal-pyramidal; hemimorphic. Axis $c = 1.577a$.



ca. $001 \text{ A } 102 = 38^\circ 18'$. $001 \text{ A } 012 = 51^\circ 58'$.

et. 001 A $101-67^{\circ} 87'$. teT, 101 a Oil = $78^{\circ} SC$.

en, 001 All! = $65^{\circ} 61'$. nn', 111 a ill = $80^{\circ} 28''$,

Crystal commonly square tabular, sometimes extremely thin; less frequently octahedral; also prismatic. Hemimorphism sometimes distinct. Also granularly massive, coarse or fine, firmly cohesive.

Cleavage: n (111) very smooth; c, s (113) less distinct. Fracture subconchoidal. Brittle. H. = 2.75-3. G, = CT-T'O. Luster resinous or adamantine. Color wax- to orange-yellow, siskin- and olive-green, yellowish gray, grayish white to nearly colorless, brown; also orange to bright red. Streak white. Subtransparent to subtranslucent. Optically negative. Indices: (W_r = 2.402, e, = 3.304.

Comp. —Lead molybdate, $PbMoO_4$ = Molybdenum trioxide 39.3 , lead oxide 60.7 = too. Calcium sometimes replaces the lead.

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PTt. , Bto,—B.B. decrepitates at $300^\circ C$ below 2 . With nitric acid of phosphoric acid in O.F. forms a yellowish-green gelatinous precipitate in R.F. becomes a yellowish-green. With hydrochloric acid, with the formation of lead chloride and molybdic oxide; on moistening with water and addition of a few drops of hydrochloric acid, it gives a yellowish blue color, which does not fade on dilution of the liquid.

Obs.—Occurs in various ores of lead. At Bleiberg, Carinthia; Hezli, Hungary; Pöchlarn; Uoldarn in the Baa; Aaaaberg, Sclineeberg. etc.

In the U. States, especially at the Southampton tail mine, Massachusetts, and near Hingham, N. Y.; near Phenixville, Pa.; at the Cobden lode in Nevada; in large orange-yellow tables at the Tecoma mine. Utah. In New Mexico, yellowish white in the Orea Hills. In Arizona, large deep red crystals at the Hamburg and other mines, Yimii, often with red Tanadulite; also at the Castle Dome district; at the Mnm-motb gold mine near Oracle. Pinal Co., with vanadinite and desclozite.

Named after the Austrian mineralogist W. Seliwsky (1818-1880).

Reddish-brown. FeWO₄. In blackish-brown tetragonal pyramids, pseudomorphous. H. = 1. G. = 6-6.4. Eimbosan, Japan.

Perhaps HgWO₄— In minute acicular crystals at Yesu.

YH. SA.1.TS OF ORQAJSnO AOXIDES.

Oxalates, Mellates.

Whewellite. Calcium oxalate, CaC_2O_4 , + H₂O. In small colorless monoclinic crystals. From Boxony, North Carolina.

Oxammite. Ammonium oxalate, $(NH_4)_2C_2O_4$ + 2H₂O. From the guano of the Guafiapu Islands, Peru.

Hydrated ferrous oxalate, $3FeC_2O_4 \cdot 8H_2O$. Color yellow. From Dearfield, Bohemia.

Hellite. Hydrated aluminium oxalate, $Al_2(C_2O_4)_3 \cdot 9H_2O$. In square prisms; also massive, granular, q. — ■ — " ■ , , ■ TM,

Prague, Bohemia. etc

massive, granular. Q. = 1'S5-1'6C. Color honey-yellow. Ocean Id brown coal in Tbu--"--■-Boheir'- -'

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HTDBOCABBOIT COUPOniTItB.

ym. ETDROOARBON OOMFOUNDB.

1 The Hydrocarbon compounds In general, with few exceptions, are not homogeneous, but mixtures, which by the action of solvents or by fractional distillation may be separated into two or more component parts. They are hence not definite mixtures and do not strictly belong to pure Mucilage, rather, with the recent gums and resins, to Cinnabar or, so far as they are of practical use, to Economic Geology. In the following pages they are treated for the most part with great brevity. >

1. Simple Hydrocarbons. Chiefly members of the Paraffin Series $\text{C}_n\text{H}_{2n+2}$.

Styrene. In white monoclinic crystals. Perhaps a polymer of marsh-gas (CH_4). Found in brown coal at Usnab, Switzerland.

Hatchettite. Mountain Tallow. In thin plates, or masses. Like soft wax. Color yellowish. Ratio of C to H = nearly 1 : 1. From the Coal-measures near Merthyr-Tydfil in Glamorgan, England.

Paraffin. A simple crystallized paraffin has been described as occurring in basaltic lava near Palermo, Sicily.

Ozocerite. Hard wax pt. Like waxes or spermaceti in appearance and consistency. Colorless to white when pure: often leek-green, yellowish, or brownish yellow, brown. It is a paraffin, and consists chiefly of one of the higher members of this series. Occurs in beds of coal, or associated bituminous deposits, as at Slanik, Moldavia; Boryslaw in the Carpathians. Also occurs in southern Utah on a large scale.

Zinckite. Crystalline. Found in ore near Ozocerite.

Terpene. In white crystalline labile crystals. Perhaps $\text{C}_{10}\text{H}_{16}$. Occurs in thin layers of pine wood from peat-beds near Uedwiltz. In the Fichtelgebirge, Bavaria. It has a similar occurrence.

2. Oxygenated Hydrocarbons.

Amber. Resin, Q. In irregular masses, with conchoidal fracture. H. = 2-S-S. G. = 1096. Luster resinous. Color yellow, sometimes reddish, brownish, and white. It is solid, sometimes fluorescent. Transparent to translucent. Heated to 150° begins to

soften, and finally mells at 380° - 300° , Ratio for C, H, O = 40 : M ; 4.

Part of the so-called amber is separated mineralogically as terebinthine (yielding succinic acid). Other related fossil resins from many other regions (e. g. the Atlantic coast of the United States) have been noted. Some of them have been called retinite, gerdane, glutite*, ruminite, tinu, terebinthine, chemawinit, etc.

Amber occurs abundantly on the Prussian coast of the Baltic from Driizig to Memel :* also on the coasts of Denmark, Sweden, and the Russian Baltic provinces. It is mined extensively, and is also found on the shores coast up by the waves after a heavy storm. Amber and the similar fossil resins are of vegetable origin, altered by fossilization : It is inferred both from its native situation with coal or fossil wood, and from the occurrence of insects incased in it. Amber was early known to the ancients, and called *elektron*, whence, on account of its electrical susceptibility, has been derived the word electricity.

Copal or Highgate resin, is from the London blue clay. It is like the resin copal in hardness, color, luster, transparency, and difficult solubility in alcohol. Color clear pale yellow to dirty gray and dirty brown. Emits » resinous aromatic odor when broken.

The following are oxygenated hydrocarbons occurring with coal and peat deposits, etc.:

Bathillite. Occurs in dull, brown, porous lumps in the torbanite or Bngbead coal (of the Carboniferous formation) adjoining the lands of Torbane Hill, Bathville, Scotland.

.dbyGoogle

It may be » altered rosolite, or else material which has filtrated into the cavity from the surrounding torbanite.

Tasmanite. In minute reddish brown scales disseminated through a laminated shale; average diameter of scales about 0.3 in. Not dissolved at all by alcohol, ether, benzene, turpentine, or carbon disulphide, even when heated. It is remarkable as yielding 5 B p. c sulphur. From the river Mersey, north side of Tasmania; the rock is called combuttite

Dye analysis gave 3.5 p. c. sulphur and 1.7 p. c. bicily, but elsewhere.

Oeocrite. a white, wax-like substance, separated from the brown coal of Qesterwiz, near Wuiasefels. Queomyrid and geoenil are other products from the same source.

Leccofetrite. Also from the Gesttrwiti brown coal. Between a resin and was in physical contact.

PvKUHETiHiTB. From brown coal near Aussig, Bohemia.

DoppLERiTB. In elsslic or partly jelly-like masaea; brownish black. An acid sub-Biaaire. or mixture of different acids, related to hiimie acid. Italio for C, H, o. nearly 10 : lli: ,'). From pent beds near Aussee in Styria, etc.

iDRtAUTE. Occurs wiith the ciunabarof Idria. In the pure state white andcryslnllie In stt-uciurp. In nature fiiund only impure, being mixed with cinnabar, r.lny. aaid some pyriie and gypsum In a brownish-black earthy maleHal, called, from lis combustibility aud llie presuuce uf uierciivy, infiamfnabl cinnabar {QuMkti'Uwbrander*}.

FoBuPMYTK. Dccun in Lard, brittle plates or nodules, light greeu in color. From Ifae Great Weitiern mercury mine, LakeCo., California. See also napollie, p. M8.

The followio^sresllll more complex native hydrocarbon compounds of great importance from an economic standpoint.

PelxoleDm. Napbtha; Pbtrolbox. Mineral oil. Eeroaene, Erdfil, BergSl, StelnOl, Qerm. *

PiTTASPBALT: Maltha. Mineral Tar. Bergtheer Otrm.

Liquids or oils, in the crude state of disagreeable odorj raryjug widely in color, from GiilorlesB to dark yellow or brown aud nearly blauk, the greenish-brown color the most comiiti: also in consisency from thin flowing kinds to those that are thick and viscous; aud in apeciflc gravity from o'6 lo 09. Petroleum, proper, passes by insensible gradations htio ptIfatpAaH OT mallfia (vlaciA bitumen); aud the latter as luaensibly Into (up/iaf(or solid bitumen.

Chemically, petroleum consists for the most part of mem bets of the pai-offln series, Cntlin-1-1, Varying tram minih gns, CUt, to the solid forms. The olefines, CbHih, are Tihn preaeDl in smaller amount. This is especially true of the American oils. Those of the Ciuii':i!>ii.4 have n higher deiisliy. the volatile coasiituentis are less prominent, they distill at abi>ut 150° and couiaiu the benzenes, C>,Hin'«, in consideinble amount. There are present also membei-s of the series CnH,»-.. Tlie Gerinao pelnileiim is iniermediiile between the Amnrleim iind the Caucasinn. ThuCnnmllan petroleum isespechill^-richln the solid parafflna.

Petroleum occurs In locks nr deposits rif nenrly all geological ages, from the Lower Silurian lo the present'epiich. It is Hssnclaled most nbundiuitly Willi argillaceous shales, sands, and sandstones, but is found also permeaiing limestones, ^Iviig them a bituminous ciilor, and rendering ibem sometlines n confiderahle source of oil. From these oleiferous shales, saods and limestones the oil ofteo exudes, and spears floating on the streams or lakes of the region, or rises iu oil springs. It also exists collected in subterranean cavities in certain rocks, whence it issues In Jets or fountains whenever an outlet is made by boring. The oi! which fills the cavities has ordinarily bten derived from the subjacent rocks; for the strata in which the cavities exist are frequently barren

sandstones. The condition a requireil for the production of such subterranean accumulations would be therefore a bltn-mlnou' oil-bearing or else oil-producing stratum at a greater or less depth below; cavittea in receive the oil; an overlying stratum of close-grained shale or limestone, not allowing of the ea«y escape of the naphtha vapors.

The two regions which now furnish the chief part of the petroleam Bt«, flist Id linport-ance, western Pennsylvania, with parts of New Tork and Ohio, and, second, the Baku region on the Caspian Bea, at the eastern end of the Caucasus. The oil Las been known to

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Stales. lu Ciiiaaia, oil ii produced in Lttuibton couDty, Ootario. Other oll-priHluciDg regiouH are found lu Burma, Jupau.tNew ZealEuid, etc.

Aiphaltmn. Mioeral Pitch. Asphalt, BerKpecli, Erdpech, Germ.

Aspbulium, or mlueral pitch, is a mixture oidtflerent liydrocarboni, part of which are oxygtualed. lis ordiuarj characters are as foUuns ; Amorphous. O. =<l-l'8i Home-tioies higlier fmin impurities. Luster liho that of lilacii pitch. Color brownish black and blat^s. Odor bitumlaous. Melts ordluarly at 80° to 100°. aud burns with a bright flame. Buluble mostl/ or wholly iu oil of turpeutiue, and partly or wholly in other; commonly partly in alcohol. The more solid kiods graduate into the plttosphaltB or mlueral tar, and through these there is a gradation to petroleum. The Suid kinds change luto Lhe solid by the loss of a vaporizable portloa on exposure, aud also by s process of oxidation, which L-unsJsts flnt In a loss of hydrogen, and fiually Id the oxygenation of a poriiioD of the mass. The action of heat, alcohol, eiber, naphtha and oil of turpemtne, as well as direct aolyscs, show that the ao-called asphaltum from different localities is vejij various in composition.

Aspbaltuiii belongs to rocks of no particular age. The most abnndaDt deposits are superficial. But these are generally, if not always, connected with rock deposits containing some kind of bituminous material or vegetable remains. Borne of the noted localities of asphaltum are the region of the Dead Sea, or Lake Aspbaltltes, whence the most of the aspbiitlum of iificent writers; a Inke on Trinidad, 1} m. in circuit, which Is hot at the center, but is solid and cold townrd the shores, and has its borders over a breadth of } m. covered w llh the hardened pilch with trees flotrishlag over it; at various places In South America; Iti California, near the coast of Si. Barbara; also in smaller quatituies, elsewhere.

Elatbrits. Elastic Bitumen. Mineral Caoutchouc. Soft, elastic, sonistimes much llfee India-rubber; occasionally bard and brittle. Color usually dark browu. Found at Castletou lu Derbyshire, and elsewhere.

Albbhtite. Differs from ordinary asphaltum in being only partially soluble in oil of turpentine, and in its very imperfect fusion when heated. H. = 1-2. G. = 1-0.97. Lustre brilliant. Pitch-like; color jet-black. Occurs filling an irregular fissure in rocks of the Lower Carboniferous in Nova Scotia.

Graphite. Resembles asphaltum in its pitch-black, lustrous appearance. H. = 2. G. = 1-0.115. Soluble mostly in oil of turpentine; partly in ether, naphtha or benzene; not at all in alcohol; wholly in chloroform and carbon disulphide. Melts only imperfectly, and with a decomposition of the surface. Occurs in W. Virginia, about 30 m. S. of Parkersburg, filling a fissure in a Carboniferous sandstone.

ULTRONITE, also called "oil shale" (oil shale). A variety of asphalt from near Ft. Duchesne, Utah, which has found many applications in the U.S. Occurs in veins several inches in diameter, with conchoidal fracture; very brittle. H. = 2-2.5; G. = 1.085-1.070. Color black, brilliant and lustrous; in powder a rich brown. Fuses easily in the flame of a candle and burns with a brilliant flame, much like sealing-wax. Named after Mr. S. H. Henson of Salt Lake City.

Bituminous coal. Compact massive, without crystalline structure or cleavage; breaks with a degree of regularity, but from a jointed rather than a cleavage structure. Structure lamellar: often finely and delicately banded, successive layers differing slightly in color. Fracture conchoidal to uneven. Brittle: rarely somewhat sectile. H. = 0.5-3-5. G. = 1-1.80. Lustre dull to brilliant, and either earthy, resinous or sub-metallic. Color black, grayish black, brownish black, and occasionally iridescent; also sometimes dark brown. Opaque. Infusible to subfusible; but often becoming a soft, pliant or paste-like mass when heated. On distillation most kinds afford more or less of oily and tarry substances, which are mixtures of hydrocarbons and paraffin.

The varieties recognized depend partly (1) on the amount of the volatile ingredients afforded on destructive distillation; or (2) on the nature of these volatile compounds, for ingredients of similar composition may differ widely in volatility, etc.; (3) on structure, lustre and other physical characters.

Coal is in general the result of the gradual change which has taken place in geological history in organic deposits, chiefly vegetable, and its form and composition depend upon the extent to which this change has gone on. Thus it passes from forms which still retain the original structure of the wood in lignite and through those with traces of volatile or bituminous matter to anthracite and further to kinds which approach graphite.

^AlOglc

1. Anthracite. muckkohle. H. = 2-2.5. G. = 1.82-1.7. Lustre bright. Opaque. Sub-metallic, iron-black, and frequently iridescent. Fracture conchoidal. Volatile matter after drying 8-6 p. c. Burns with a blue flame of a pale color. The anthracites of Pennsylvania contain

ordinarily 8a-B3 per cent. of carbon; those of South Wales, SS-95i of France, 80-88; of Saiony, 81; of uoulherii Russia, sumetlmea 9i per cent. Aulbmcite gradiiiLles througli aemi-auUiracile iulo bitumiaua coal, becoming less bard and coniaiuing more volatile mutter; and an lulermediate variety la ci\et\ frt»-l>uming anthracite.

3. BITUMINOUS Coal. Burns in the fire with a yellow, smoky flame, and gives out on distillation hydrocarbon oils or tars; hence the name Bituminous. The most bituminous coals contain from 5-15 p. c. (rarely 18 or 17) of oxygen (ash excluded); while the so-called brown coal or lignite contains from 30-86 p. c. after the expulsion, at 100°, of 15-86 p. c. of water. The amount of hydrogen in each is from 4-7 p. c. Both have usually a bright, pitchy, greasy luster, a firm compact texture, are rather fragile compared with anthracite, and have U. = 1'14-1'40. The broken coals have often a brownish-black color, hence the name, and more oxygen, but in these respects and others they shade into ordinary bituminous coals. The ordinary bituminous coal of Pennsylvania has O. = 1'25-1'87; of Newcastle, England, 137; of Scotland, 147-1-8a; of France, 1'2-1'83; of Belgium, 107-118. The most prominent kinds are the following :

(a) Caking or Coking Coal. A bituminous coal which softens and becomes plastic or semi-fusible in the fire. This softening takes place at the temperature of incipient decomposition, and is attended with the escape of bubbles of gas. On increasing the heat, the volatile products which result from the ultimate decomposition are

and a coherent, grayish-black, cellular or friable mass (or part not volatile) varies from 50-85 p. c.

(b) Non-Caking Coal, Like the preceding in all external characters, and often the ultimate composition; but burning freely without softening or any appearance of incipient fusion. There are all gradations between caking and non-caking bituminous coals.

(c) Parrot Coal (Parrot Coal). A variety of bituminous coal, and often caking; but differing from the preceding in texture, and to some extent in composition, as shown by its products on distillation. It is compact, with little or no luster, and without any appearance of a banded structure; and it breaks with a conchoidal fracture and smooth surface; color dull black or grayish black. On distillation it affords, after drying, 40 to 50 p. c. of volatile matter, and the material volatilized includes a large proportion of burning aromatic oils, much larger than the above kinds of bituminous coal; whence it is extensively used for the manufacture of such oils. It is obtained from oil-pruding coaly shales, the most compact of which it much resembles. Torbanite is a variety of cannel coal of a dark brown color, from Torbane Hill, near Baibgale, Scotland; also called Boghead Cannel.

(d) Brown Coal (Bituminous Coal, Lignite). The principal characters of brown coals have already been mentioned, they are non-caking, but afford a large proportion of volatile matter; sometimes pitch-black, but often rather dull and brownish black. G. =

It is often culled lignite. But this term is sometimes restricted to those of coal which still retain the form of the original wood. Jet is a black variety of brown coal, compact in texture, and taking a good polish, whence its use in jewelry.

Coal occurs in beds. It is interstratified with shales, sandstones, and conglomerates, and sometimes limestones, forming distinct layers, which vary from a fraction of an inch to 30 feet or more in thickness. In the United States, the anthracites occur east of the Alleghany range, in rocks which have undergone great contortions and fracturings, while the bituminous coals are found extensively in many strata farther west, in rocks that have been less disturbed; and this fact and other observations have led geologists to the view that the anthracites have lost their bitumen by the action of heat. The origin of coal is mainly vegetable, though animal life has contributed somewhat to the result. The beds were once beds of vegetation, analogous, in most respects, in mode of formation to the peat beds of modern times, yet in mode of burial often of a very different character. This vegetable origin is proved only by the occurrence of the leaves, stems and logs of plants in the coal, but also by the presence throughout its texture, in many cases, of the form of the original vegetable; also by the direct observation that peat is a transition state between unaltered vegetable debris and brown coal, but sometimes found passing completely into true brown coal. Peat differs from true coal in want of homogeneity, it visibly containing vegetable fibers only partially altered; and wherever changed to a fine-textured homogeneous material, even though hardly consolidated, it may be true brown coal.

For an account of the chief coal fields, and of the geological relations of the different coal deposits, reference is made to works on Economic Geology.

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APPENDIX A.

ON THE DRAWING OF CRYSTAL FIGURES AND OF PROJECTIONS.

In the representation of crystals by drawing, the object may be either to show the entire form in perspective or to give simply a projection of the faces upon a single plane. The first of these cases is the more important, and must be treated here in some detail. Two points are to be noted in regard to it. In the first place, in the drawing of crystals the position of the eye is supposed to be at an indefinite distance, and it follows from this that all lines which are parallel on the crystal appear parallel in the drawing.

In the second place, in all ordinary cases, it is the complete ideal crystal which is represented, that is, the crystal with its full geometrical symmetry as explained on pp. 9 to 11 (cf. note on p. 11).

Projection of the Axes.

The projection of the axes of a crystal is the first step preliminary to the drawing of the

form of the crystal itself. The projection of the axes in the isometric system, which are equal and intersect at right angles, is here &rat given. The projection of the axes in the other systems, with the exception of the hexagonal, may be obtained by varying the lengths of the projected isometric axes, and also, when oblique, their inclinations, as shown beyond.

Isometric System.—When the eye is directly in front of a face of a cube, neither the sides nor top of the crystal are visible, nor the faces that may

be situated on the lateral edges. On turning the crystal a few degrees from right to left a side lateral face is brought in view, and by elevating the eye slightly the terminal face becomes apparent. In the following demonstration the angle of revolution is designated θ , and the angle of the elevation of the eye, e . Fig. 989 represents the normal position of the horizontal axes, supposing the eye to be in the direction of the axis BB: BB is seen as a mere point, while CC appears of its actual length. On revolving the whole through a number of degrees equal to θ the axes have the position exhibited in the dotted lines. The projection of the semiaxis MB is now lengthened to HN, and that of the semiaxis MC is shortened to MH. —

If the eye be elevated (at any angle, e), the lines B'N, BM, and C'H will be projected respectively below N, H, and H, and the lengths of these projections (which we may designate $b'N$, bM , and cH) will be directly proportional to the lengths of the lines B'N, BM, and CH.

It is usual to adopt such a revolution and such an elevation of the eye as may be ex-

pressed by a simple ratio between the projected axes. The ratio between the two axes, N : WU, as projected after the revolution, is designated by $1 : r$; and the ratio of A'N to MN by $1 : s$. Suppose r to equal 3 and s to equal 2, I then proceed as follows :

Draw two lines AA', HH' (Fig. 990), perpendicular to each other at right angles. Make $MH = MH' = b$. Divide HH' into 3 equal parts, and through the points, N, N', thus determined, draw perpendiculars to HH'. On the left hand vertical, set off, below U', a part U'R.

equal to

1 .

$r \cdot b'N$; and from R draw RM, and extend the same to the vertical N'. B'B

is the projection of the front horizontal axis.

Draw BS parallel with MH' and connect SM. From the point T in which SM intersects BN. (i.e. draw TC parallel with MH. A line (CC) drawn from C through M, and extended to the left vertical, is the projection of the ^de horizontal axis.

Lay off OD the vertical a part HQ equal to $\frac{1}{2}MH$, and make $HA = MA' = MQ$;

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APPENDIX A.

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where, $r = 8$, $ftDd = 2$, $theo 8 = 18^\circ \frac{1}{2}$. and $e = 9 \frac{1}{2}S'$. $lor \cot 6 = T$, and $\cot f = rt$.

It is desirable to go through the above construction with the help of all possible Biquadrants. The axes thus obtained, they will be represented, may be drawn on a suitable curvilinear paper by means of a ruler for (ultra use. Whenever needed, they may be readily transferred to a sheet of paper by passing through the lines. And finally, in the case of the axes so obtained, they may be tested to make sure that the respective semi-axes of the central point M are equal.

Tetragonal and Orthorhombic Systems.—The axes AA' , BB' . constructed in the manner described, are equal and the right angles to each other. The pro-

portions of the axes of a tetragonal crystal is obtained simply by using a scale of proportions. If MA and MA' taken as units, the value of the vertical axis (i^*) will be given by the species. Thus for $CaCO_3$, where $i = O'M$, we must lay off $\frac{1}{4}$ of MA above M and the same length below.

For an orthorhombic crystal, where the three axes are unequal, the length of i must be laid off above and below from M . and that of l in front of M . It is usual to make the side axis $MC = \frac{1}{2}i$.

Monoclinic System.—The axes a and b in the monoclinic system. other at an oblique angle ϕ . To project this system, and thus adapt the isometric axes to a monoclinic form. Lay off (Fig. 991) on the axes MA , $Ma = HA \times \cos \phi$. and on the axis BB' behind H , $TS = WB \times \sin \phi$. From the points b and a , draw lines parallel respectively with the axes AA' and BB' , and from their intersection D'' , draw through M . $D'D$. making $MD = MD'$. The line DD'' is the clinodiagonal, and the lines AA , CC . DD' represent the axes in a monoclinic solid in which $a = b = c = 1$. The points a and b and the position of the axis DIV will vary with the angle ϕ . The relative values of the axes may be given them as above explained; that is, if $J = 1$, lay off the direction of MA and MA' a distance equal to i , and in the direction of MD and MD' a distance equal to d , etc.

Triclinic System.—The vertical sections through the horizontal axes in the triclinic system are obliquely inclined also the inclination of the axis a in each axis b and c , is oblique. In the adaptation of the isometric axes to the triclinic forms, it is therefore necessary to

le length of J I ind behind H,<

B inclined to one a.

J^^

Brst place,

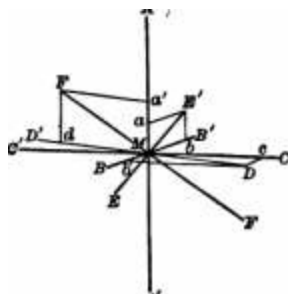
place, to give :liDattoii ofthi

.the requisite 992, obliquity to the mutual IncllDatton of the vertical eec-

lioDS, and afterwards to adapt tbe horizontal axes. The tncliualion of these socIIodb we may desijrnHte A. tind as heretofore, the angle between a and b, y, and a nod e, (8. BB' is the analogue of the brachydiiponal. and CC of the mncrodiagonal. An oblique inclinntion may be given the vertkiii sections, by varying the position of either of these sectiuDS. Permitting the bntchydiag-onal section ABA'B' to remain unallered. we may vary the other section as follows r

Lay off dig. 892)on MB, M'/ = MB X cos A. and on the axis CC (to the right or left of U. nccording as the angle A la to the right or lefti. Mc =. MC X sin A: completing the pandlelogram M£'D>!. and dntwin^ the diagonal HD, extending the same tn D' bo as to make HD' = MD. we nbuin the line DD'i Ibe vertical section isinit through this line lathe correct mncrod is con al ,A section. The Inclination of a to the new macro'iiag-

onal DIV is still a right angle: as also Ihe Inclination otatoi, their oblique Inclinalions may be given them aa follows : Lay off on HA (FI^



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APPENDIX A.

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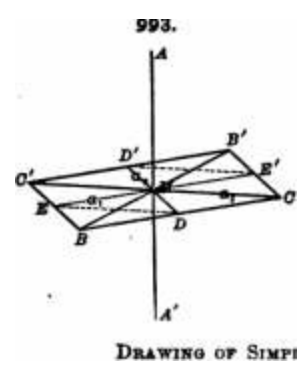
K cos o, and o

the projected bnchjiUagooal. Agtia iay oft on HA, Ha' = HA X cos n, and on HIX, U the left, Md = HD' X sia a. Draw tines from a' and d parallel to MD and HA; V, the 1d-

leiaectioD of tbeae linea, U one extremltT of the macrodUEOnal; and the Hoe FE*, in %vhch HF = HF', la the macrod[agonal. The TertlcaJ axis AA' and the borizooolal axes EB* (brachydtagonal) and FF* (macrodlagonal) thus obtained, are the axes la a tricllnlc form, in which a =6 = tf = 1. Diileient valuea maj be given these axea, accoidlDgioihe method heretofore illustrated.

Hsxagonal Bjrat«in.—The simplest method of obiaininK tbe axes for the hexagonal sjs-teui iB as follows : We start with tbe isometric axial crosi (Fig. o60) and change it so thait the front axis (MB, fig. 998) shall have a length equal io 1782 (= f^) times its original length. When the extremities of the litiee CC and BB' are Joined a rhomb will be formed having the angles at C and C equal to 120*. Now draw through the middle po!nIs of HB, HB' lines parallel to CC, and the resulting figure, CDEC'D'F', will be a regular hexagon. The lines Joiningits ancles, namely, CC', DD', BB', will be the lateral axes r^ quired; these are shown lu Fig. 90i, repeated from p. 60, It only remains lo ^ve the vertical axis the length required by the case In b&ud.

9M.



Jlis

ABD CUMPLBX CrTBTAU.

a beeu constructed tor tbe given species, the oi 090.

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mid la the other systems are obtained a joining the extremities of each of the lateral a: with those of the Terlicul axii This is illustrated for LLe Isometric system by Fi^. DBS Here, as in all cases, the lines whicb fall in front are drawn Btrongly, while tboee beliind are simply dotted.

For the diamtttral prism draw Ithrough B, B', C. C, of the priijected axes of any flpecfes (Fig. 9»5|, lines parallel 'o the axes CC, BB', uatll they meet; They make tbe pnrnllelnEram.

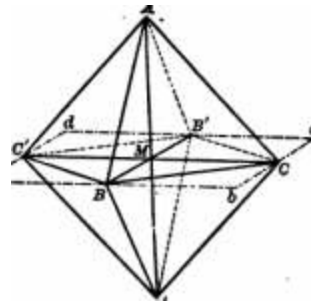
abed, which is a transverse section of the prism, parallel to the base. Through n, b, e, d draw lines parallel and equal to the vertical axis, making the parts above and below these points equal to the vertical semi-axes. Then, connect the extremities of these lines by lines parallel to ab, be, ed, da, and the figure

will be that of the diametral prism, corresponding to the axes projected.

In the case of the isometric system (Fig. 996), the diametral prism is the cube whose faces are represented by the letter a (100), Fig. TO, p. 54; in the tetragonal system it is the prism a (100), c (001), Fig. 1*9, p. 54; in the orthorhombic, monoclinic and triclinic systems the faces have the symbols a (100), b (010), c (001), as shown in Figs. 276, 812, 888.

The unit vertical prism in the tetragonal, orthorhombic and triclinic systems may be projected by drawing lines parallel to the vertical axis AA' through B, C, B', C, making

the prism in the tetragonal, orthorhombic and triclinic systems may be projected by drawing lines parallel to the vertical axis AA' through B, C, B', C, making



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APPENDIX A.

the prism above and below these points equal, and then connect the extremities of these lines parallel to BC, CB', B'C, and B. The plane BCB'C is a transverse section of the prism parallel to its base. This prism is bounded by the faces in (110), c (001). Cf. Figs. 160, 276, 813,

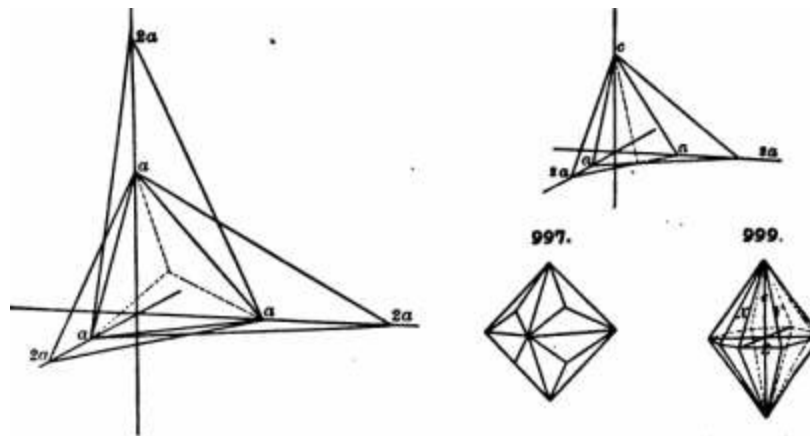
prisms in the orthorhombic, monoclinic and triclinic systems, as (310), (120), etc., are drawn in the same way, only that the lateral axes must be adapted to each case. Thus in (210) the axis b has double the unit value characteristic of the given species; for (100) it has half this value (or, otherwise expressed, the axis a is to be doubled), and so on.

In the tetragonal system the prisms (810), (310), etc., have eight faces (cf. Fig. 1B2, p. 64). Therefore, e.g. for (210), first draw the plane of the lateral axes from the unit length of each axis to the double of that axis (+ and -): the intersection of these lines are points through which lines must be drawn vertically (parallel to the vertical axis),

correspond to the edges lettered n in Fig. 1G2. For the hL-diagonal system the construction of the three prisms is so similar to those described as to need no further explanation.

Other simple forms are constructed on much the same principle. The points are first to be determined at which the respective faces cut the axes. Then light construction lines are drawn in each of the axial planes connecting these points, and at the same time the intersection-points noted for each pair of adjoining faces. From these last the actual intersection lines of the given form are obtained.

996. 998.



The diagrams will serve as illustrations. The form in Fig. 997 is the trisphenohedron (122) of the tetrahedron system (cf. Fig. 88, p. 88). For them the axial intersections are (see the explanation on pp. 28, 24) as follows :

Here the axes a, a, a, are, as always, taken in the order a, Fig. 896 shows each of the three planes projected limit-edges

in order (cf. Fig. 85, p. 88). Their respective intersection-points lettered a. The

The dotted lines are drawn in the

completion of the figure involves no further difficulty. It is to be noted, however, that it is not necessary to repeat the full construction for each octant, since the principle of symmetry with respect to the three axial planes, that is, along lines parallel to each of the axes, can be used to simplify the work. In practice, it is well to construct the front half of the form, and then use symmetry to complete the part behind, which is usually drawn in dotted lines. The symmetry. In any case, is a most valuable check on the accuracy of the work.

Another illustration is given by Figs. 998. 999 of the ditetragonal pyramid (122) of the tetrahedron system (cf. Fig. 166, p. 58) This requires no additional explanation.

When it is required to draw not only the faces of a single form, but also those of a number of forms modifying one

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AFPEXDIX A.

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■uotlier, ft somewhat different process U found desirable. It Is possible, indeed, 10 con-
slrucl a. complex figure iu thie way just mealtoned, e&ch plane belo); laid off on the giveu
Hxes, Udit its inlerraectiou-edges with adjoinlnji planes determined by two points, always
In Ibi: aziKl sectlous, which it has iu commou with each of them. In this way. however,
the Bgure will soou become so complex as to be extremely perplexing, ncd Ibus lead to
error •ud conaequeut loss of lime.

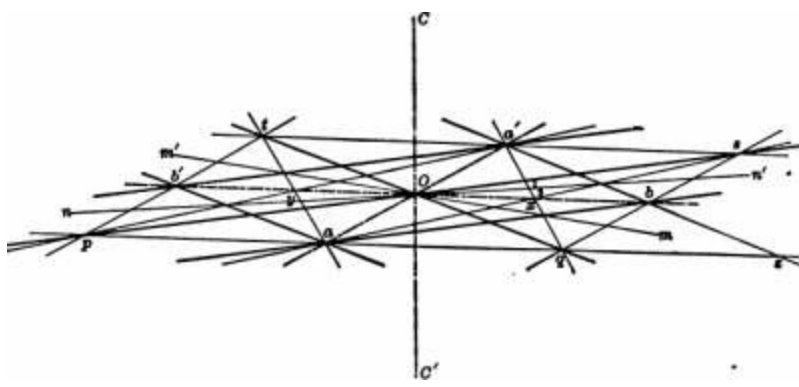
The process of determining the direction of the intersect!on-edges is much simplified If
The ezpreGSION giving the ratio of the iuterceptl on the axes for each plane ia so
tranaformed (by dividing by the coefflcieuL of the verticul axisi that the value lot the
vertical axis shall be unity (that fa, [□ general equal to ij. The extremity of the vertical axis
(-f or —) U then OM point of intersection for any two planes, and the other point will
always be in the plane of the lateral axes.

As an einmple of the traneformalion culled for, suppose 834 to be the symbol of a given
face, lis intercepts on the axes a. b, e would be (Art. 84) ia : Ifr : ie. This put Into thu form
required for the present purpose becomes (by muptiPLYug by 4), ia:2b:c.

It will be noted that the expression inus obtained is not ordinarily that which corr^
sponila lo the Naumaun symbols, Tor with them the length of a lateral axis (usually a) is
taken as the unit.

Example.—The explanation of lhn method' of drawing cryslala can beat be described by
use of a concrete example. Suppose Uiu it is reijuiied iu aruw an ortborhombic crystal
(HuduluBiK!) of pristuutic Lnbil (Fig. 1000), showing the fiicea a (tOO), e (001), m (110),
f (101), »(011), o(111). f (121). It Is evident that an ItideSuiie number lOOa

of tignres may 1>e niado, including the faces mentioned, tind yet of very different
appearance according lo the relative size of encb. It Is usually desirable, however to
represent the actuill appenrnnce of lbe crystaln i[i nature, only in idenl symmtlry, hencu it
is Important in all cases to bnve H sketch of the crystal to be represented, showing the
relative development of the different fnces. If this sketch Is made with a little care, so as
to exhibit also the parallelism of the Interaectlon-edges In the occurring zones, it will elve
material aid. The zones, it Is lo be noted, are a great help in (irawing figures o(crystals,
and they should be carefully studied, since the common direction of the Intersection-
edge, once determined for nay two faces in it, will answer for all others.



The first step is to take the Isometric axial cross-section already drawn on a suitable card, and of a conveniently large scale, and pierce it through on to the paper to be employed, and then adapt it to the requirements of the species to hand. For example, the axial ratio $U : u : v : i = O-M : 1 : O'TO$. Hence the vertical axis must be 0.7 of its isometric length, but (in this case) the lateral axes call for no change.

The next step is to draw on the plane of the lateral axes (Fig. 1001) a line giving the

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APPENDIX A.

trace of each face as it is needed. Each line is obtained from the expression of the azimuthal intercept transformed in the manner already explained, so that the value for i shall be unit. For the faces of the prism, $m(110)$, the required lines are pt and at (Fig. 1001). Obviously their intersection-edges are parallel to the vertical axis, and their intersections with the base $c(001)$ are parallel to pt and at . Similarly the mutual intersection-edges for the faces of the prism $k(110)$ and the prisms $a(100)$ will be parallel to CC .

Further, the corresponding hues for the upper faces of the macrodome, cr , $thsl$ Is , 101 and $10l$, are pq and It ; those of the faces of the microdome (Oil) are gt and pi ; those of the faces of the pyramids $o(111)$ are ab , ba' , $a'i$, $b'a$; those of the pyramid $e(121)$, whose intercepts in the required form are la : JA : le , areas, ga' , $a'p$, ta . Each of the planes mentioned passes also through C , the extremity of the vertical axis. Therefore, whatever intersection-edge is called for is given at once by the figure. For example, that for the faces $s(331)$ and $«(Oil)$ is the line joining C with g , since the lines at and g^* represent these planes intersect at g and all planes pass through C ; further, the same direction is that of the intersection-edge of $f(111)$, since the trace of $m(110)$ also passes through the point I . So also the intersection-edge of $o(111)$ and $s(121)$ will be the line joining C and a , and the same line is that required for $r(101)$ and $o(111)$; that of r and $r'(110)$ is given by the line joining C and i , since the trace of k'' , or fun' (Fig. 1001; cf. Fig. 1002), if produced, meets that of $r(101)$ at s .

The above explanation will show how the required intersection-edges can be obtained as needed. In the practical use of the method it is customary to begin with the pre-

dominating forui aud tbeu add tbe modifying faces Id succestlon. That li, Id the esse tn banil. tbe prism m would he drawn teiiuiunteu by e ; then perfasps tbe pinacold a added ; then tbe prUni J:(210), and afterward Ibe-IemilDal faces.

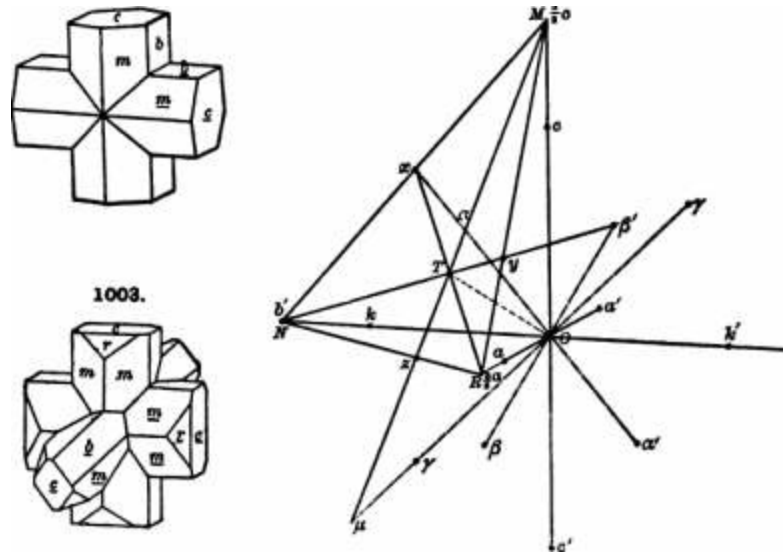
The actual figure U usually made on the shtet of paper below the axial projecilon (Fig. 1001), the directions of the Inter si ctionidgcs being carefully transferred by menns uf a trlauEle sliding agidust Ibe dde of nnothtr triangle or a ruler. The figure should be made cunufierabty larger lliaii as finally required, since only In this way Is sufficient accuracy nttaiunble. Tbe final figure of tbe required sise Is readily accomplabed by pholo-engraviug, if tbe drawing Is done clearly and iu black llinea.

To complete tbe figure beblnd, the Intersection-lines can be oblained Id the same way. It is to be nutbd, however, thai In tbe given case (of an orthorhombic crystal) tbe sj m> meuy can be used to simplify tbe process, or, if desired, merely to check tbe work duoe, since every point in front has a correspoDdlng eymmeirical point behind equally distant from tbe axial plane be, in the direction of ibeaxha, Tbe symmetry right and left relatively lo tlie axial plane ae also should be usid lo tesl the accuracy aa regards ibe two pities of tlie figure. When the upper half of llie figure (or. if preferred, the front half) Is completed, the symmetry, as noied, ia usually employed in the completion of Ihe remainder, since (in Ibis case) every liue in front above has one paiiallel and equal to It behind below. In the case of crystals of lower grade of symm etry tbe method Isiessiimple, and in some instance* considerable iugeauty Is culled for, hut pnictlce will serve to give facilliyy.

On TEtB Drawina of Twin Crtstals,

Twlsnad Oiyatals.— Id order to project a compound or twinned crvBial It Is generalIT necessary tn obtain first the axes of Ibe second individual, or semi-individual, In the position iu which Ihev ere brought by tbe revniullon of 180'. This is accomplished In Ihe following manner ■ ' In Fig. 1008 a compound crystal of stiuirolite Is represented, Iu which twinning has taken place (I) on an axis normnl lo 1082), aud In Fig. 1008 this is shown again with also twinnlnit on an axis normal to (2321, Tbe second method of tw[oiiing being the more ireneral. is of the greater importance foi- the sake of example. Iu Fle^ IftH r^, W. *.■ repre«enl the recUnguUr axes of slsurolUe d : !: i = 0-47 : 1 : 0-68). T- he t*vlu-nlmg-plane (2821 lias the position MNR. It is first necessary lo coustruct a normal from ibe center O to this plane. It perpendiculars be drawn from tbe center O to tbe llinea MN, NR, MB, they will meet rbem at tbe points *, a. i. dividing each line Into segments proportional to the squares of the adjacent axes ;• or Ni: Mi = ON' : OM". In this way the points x, y I are fixed and lines drawn from nnv two of them to Uie opposite angles R. N. or U will fljc the point T, A line joining T nod O is norma! to the plane iMNR = 28-2). Furlber-more it is nbvious that if a revolution of 180° about TO take place, that every point in the plane MNR will remain equally distant from T. Thus tho point M will take the place

gtnU u, ff, St, and the common centtir O will be the new axea correit»oDdiDg to HO, NO, O. In order to obtain the unit axes coireipondlog tah.h, d It b merely neceawry to-dniw through e a line parallel to MTu, meeting xO at y, then yOr' 1b the near Teril-



cal Hxla correponding to eOe*; also, fiOfi'carresponAi to bOV, and oOa' correponda to aOa!. TlieBc tlireeiiiKes. then, are the axea for tlit xpcond inillvidiml Id Its twiunvd posUioD : upon tlieiii, 111 llie usual way, the new figure iiiiiy be coustructed and then ttsnaferred to Ita proper positioD with reference Ui the iioTiinil crystal.

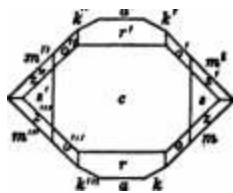
For the second melhodoI iwianiae. wheu the iixIb ia normn) to (033\ the construction Ih more ^iiuple. It Is obvious that the axis is the line Oz, and using this, as before, thv new axun nn Fiiund ; kOk correaponds lo eO<f (sensibly colncidlug with W), since e (OO)A (m = 45° 41'. and so on.

Ill muuy cases the simplest method Is to cocatract first the normal crystal, then dt«w through its center the twinning-pkne and the axia of revolulfon, and determine the angular points of ilie reversed crystal on the principle alluded to above : that by the revolution every point remains at the same distance from the axta, measured In a plane at right anele to the axia. Tlius in li'ig. 884, p. 120. nheo the scalenohedron has been drawn, since liie miiiuing-plana is the b^l plane, each angular point, by the revolution of 180*, obtains a poaitlon equidistant from this plane and directly below it. In thia way each angtdar point Is determined, aud the compound crystal ia completed in a '

Dbawino of Projections.

HorizontaI PreJeoUoiM.—It la often conveoient to have, Instead of a drawing of a complete crystal. B .ip-called horizontaI proje'ction of its faces upon a single plane. The plane selected is usually thai normal to the prtiiialir zone—in other words. that parallel to the base (if nl! crystals except those of the clinomctiic system ; Figs. 80, 83. 84. 8«. 39, on pp. 15, 16. are ezamplei. O'Iiit planes itiny also be taken with advantage in certain cases, ai, for example, the clinodUgoiial axial plane with monocliolo crystals. Fig. 836, p. 108, represents such a projection of a crystal of epidote.

A.PPKNDIX A.



IS ua'. W repret'til the lutarsl

t or llie pyramid, i (121). lUe

In Fig. loot) tbe dotted Hoc

:s (beiu aensibly eciuaJ). Tlii

"•(110) o

k (210). luB imea 7i7. i litius <M, u'j. u'p< uiid ^ .

The intersettlon-edgtii for aoy two planes wfl be jriven by the line drawn through e (the eslremily of Uie vertlckl axis a* projected, coniiiiun to all plaues), and also through the polot where the Uues represeDting the (jiven planesiaieraect. Thus the r&uesB &oil t are repreened iti the projecliou by the Ifoea i; hence the intersection-dlreclion Coi the two

The projection shown in Fig. 1005. when each face of a irjlsal is represented by a lin« on a given plane, iiBUtily lliAl u'lrnal to the pmmatic zone (vertical snisl, was proposed by Quenstedt, and ia known ns the (iutmUdt Pi-irietion; It was made by hiin the buU of ui eiaboiftte system of cryBtalloRrapliy, It will be noticed lliat wioeR here are represented by poinU, since all the lines repr.'sintirijr i lie faces of a given zone must pass through tlie same point in the plane of projection ; this is slUl true malhematically of lines which ate parnliel.

Spherical PrcJ«cUoik.—Vnrions itiotbods Imve been aiiggesled aud are in use for representing the positions nf tbe fucts of a crystil. especially with rtspect to their zonsl rela-tions, tlie angles between them, file,: these do not, however, iitteitipl ui give a picture of the crystal itself. One of tthese—the Quenstedt pnijection—represented by Fig. 1006. has nlremly been spoken of. Another vnlnable method is thst of GoldscUmiMt. which lie makes the base o(his ingenious and practical system (see references on p. 117).

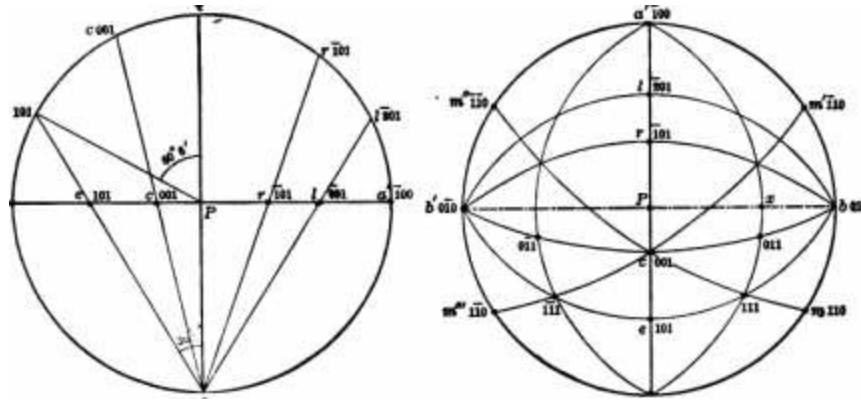
The most serviceable melhoii and the only one used in the present book Is the spiieric*! projection desciihed in Arls. 33 to 42. and of which examples are given in the paees follow. fng (see Fig. 110, p. 41; 171. p 58, eic). The method of construction needed hu been •xplained in Art. 40, sii far as ll applies to the orthomelric systems.

For the monoclinic system it is to be noted (cf. Fig. 827. p. 108, and Fig. 1008) that while the poles of the prismatic faces still lie on the circumference of the circle and can be fixed at once with a protractor, and while all the orthodomes (AOO He in the diameter from n (100) to a' (100) at 00° from b (010), the base e (001) is not the middle point of the circle, as in the projections of the orthometric system B. The position of e can, however, be readily fixed, since the complement of the angle at (100 A 001 = 13) gives its angular distance from the middle point of the diameter oa' (P in Fig. 1008); and this distance of the

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APPENDIX A.

projection is perpendicular to the tangent of half the angle in terms of the given radii. So it is also the projection of any orthodome whose angular distance from a (100), o' (100), or e (001), is known can be fixed on the sphere of projection in the same way.



Instead of using the tangent of the half angles with a scale of proportional parts the distances may be readily constructed without a scale. Thus for epidote we have as. $100 A 101 = 38^\circ 54'$; ae. $100 A (101 = 84'' 87''$:

tt'r, $100 A 101 = ar 41'$;

a'i. $100 A 101 = 25-57'$.

Draw a circle with the given radius (Fig. 1007) and measure oB on the circumference from a (100) and if (100) the successive angles given. Then join the points m fixed with the point o, the opposite extremity of the diameter at right angles to aPa'. which is supposed to be the position of the eye in this method of projection. The points where these lines intersect the diameter aPa' give the position of g (101), e (001), r (101), l (201), to be used for the projection of Fig. 1008.

When the positions of the poles of the orthodomes of the given crystal are fixed in the diameter no' (Fig. 1008), and also those of the prisms in the circumference of the circle, it is obvious that the pole of any other face can be determined by drawing the proper circular arcs. Thus the arc 010, 101, 010 and that 110, 001, 110 together determine the

position of 111, and so on for any pyramid. Again, the arc 100. Ill, 100 and 010, 001, 0i0 determine the position of Oil. It is to be noted, however, that for sake of accuracy it may be better to fix the pole of Oil. or any given clinodome, independently. It (Fig. 1008) a diameter be drawn connecting b (010) and Zi* (010), it is clear that the angular distance between b and hence Pz ($P^* = 90^\circ$) can be obtained at once by the solution of a simple right-angled triangle in which the hypotenuse (010 a Oil or in general 001 a Okl) is known and also the side at S—the last being equal to $90^\circ - \alpha$. When the point x is known the circular arc axn" determines the position of Oil, and so in general of any clinodome. Fig. 1007 makes it clear at once why, as stated in Art. 40, the tangent of half the angle from the central point of the sphere gives the position of any pole. The radius of the circle being taken as unity. It is seen here (Pic 1007) that the actual angle, for example $8'$. for T ? (101) at the center is one-half the corresponding angle $80^\circ 8'$ the circumference; ; further, the distance P- is proportional to $\tan 30^\circ$ (tan $80^\circ 6'$), where the radius aP is equal to unity. *

In the triangle $ty \gg ttm$ the spherical projection can be drawn by an extension of the method just described. It is assumed that the angles between the pinacoids are known, (A, B, C); also the angles of the triangle abc which are the supplements of the axial angles α, β, γ . In the first place the positions of a (100), b (010), and of the prisms are fixed by the protractor on the circumference of the circle. How to draw through aa' k

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1 and a'. A circular arc drawn through BB' and e (001) will be tangent to the circle at point E, whose distance from the center P can be obtained in a right-angled spherical triangle $! \square$ in which the hypotenuse (oe) and the adjacent angles at a are known. Hence the position of K is fixed. Again, draw the diameter bb', and draw right angles to it as a diameter, AA'. If L is the point on bb' where it is cut by the circular arc through bel/, the distance PL can be similarly calculated. With K and L both fixed, it is easy to draw oeo' and beV, and thus determine the position of e. After the zone circles aea' and M/ have been drawn, the position of any dome (M or OkC) can be similarly found by a single calculation, and as the prisms have been fixed it is easy to locate any required pyramid.

On the general method of projection applicable to triclinic crystals, see Qroh, Phys., . Kryst., p. 679, «t ttg., and Lieblich. QMmelr. Kryit., p. 10, el Mg. *

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APPKJSDIX B,

TABLES TO BE USED IN THE DETERMINATION OF MINERALS.

This Appendix contains a table of crystals, more or less complete, of minerals arranged according to certain prominent crystallographic or physical characters. These, it is

will be of aerrice not only to the Htudenl. but also to the Bkilled miueralogiM. Table I., ofMJnenli arranged accora lug to Sjritem of Crystal llzatk>n, is intended to Mabraca all null-recoguized ipeciea, IbouKh those of relatvelj greater importance, especially >• regard* occurrence, are Indicated By being printed In heavier type.

The other tables make no claim to completeoesa, being limited often to common aيتد important species.

For an exhaustive system of Determtnative Tablea based paitlcularly upon blowpipe and chemical characters, the student Is referred to the work of Professors Briufa and Penfield, meniloned'oD p. 206.

I. MINERALS ARRANGED ACCORDING TO THEIR SYSTEM OF CRYSTALLIZATION.

The following lists are Intended to ioclude all well-recognlEad species, whose crystalllz-allou is known, arranged according to the system to which they belong, and furh^ classifled by thelr luster and specific grBTliy ; tbe hardneas is also given In each case.

I. CRYSTALLIZATION ISOMETRIC*

A LUBTBR UhHETALLIC.

ShI Ammoniac (p. 819)

EnliQJte(p.53i))

Fauiiuite |p. 460)

Bylvltto (p. 319f

HaUUlp 818)

Hydri-phflile (|>. 821).

Sodalite (p. 41?)

Analcite (p. 460)

NosaUte (p. 418}

Niirtljiiiiite (p. 864)... Hallynite (p 4IS)....

Leucit«(p, 881)

Lazurite (p. 418)

Sulphohalite (p. 531).. Itatouite (p. 838)....

Voliaite(p. 537)

Langbeinite(p 533)...

Zinnite (p. 414)

Ferrihydrite(p,882)

Boracite (p. 518)

Pharmacoite (p. 818) Nitrobarite (p. 61) Fluorite (p. 830). Helvite(p. 414).. Olanite(p, 415).. Uvaldeite(p. 618) Davianite (p. 414). Hauzite (p. 801), Diamond [p. 371] Spinel (p. 388)... Periclase (p. 831).

Arsenolite(p. 880)... Hübnerite (p. 419). Hercynite (p. 33B)... Sphalerite (p. 291)...
Nanokite (p. 817)...

Marahite (p. 317)

Akimotoite [p. 2B2).. Perovskite (p. 4R7)... Berzeliite; (p. 405).... Qandite (p. 189)
Pyrochlore(p. 480)... Koppite(p. 409)

Zirconite<|. 846)

Halobellite <p. 489).

Lfwipite (p. 516)

Alopite(p. 516)

Percyite. Boleite(p. 822) Hübnerite (p. 516) .. Manganoite (p. 833).

Embolite (p. 810)...! Oerargyrite (p. 319)

Miersite(p. 31B)

Microcline(p. 489) ... Inhibromite (p. 319). Brimleyite(p. 819)...

Ouprite(p. 881)

Synchroite (p. 414) ... Buntsielite<p. 882).., Honnigsteinite(p.49a),.

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APFENDIX B. B. LUOTBE HZTALUC (AMD StrBHBTUXIO).

Hauerlte (p. 801)

BpbalerlU! (p. 2»1)

Alabaodite (p. aes)

aut)uiite(p. 297)

Dyunalyte (p. 468]

StwiiiUe(p. 81G)

Ohromita(p. 841)

Biunite(p. 8P8)

TeiUMuUta(p. 318)

Tetrahadxlte (p. 81S).. Magae»io(ernii: (p. 841) PsntloDdit* (p. 398)... Polydymile (p. SWJ),...

JaCDbElte(p. 341)

Syt^liDodymite(p. 298),

Umiwlte (p. 297)

Carrollite (p. 287)

Ulxbyltetp. 843)

PyTlU{p. 800)

FnuiUiaita(p. 341)....

Magnetite (p. 889)

Boniit« (p. 297)

Oersdorfflie (p. 303)

Onprito(p. 381)

BrouRDiardiM (p. 30B)..

CorynUe (p. 803)

Argyrodite (p. 816)

OobalUte (p. 801)

Cairfeldite (p. 816)

Umannite (p. 301)... SnultiU, ChloaDthlW

(p. 301)

Skuiierudite (p. 807)... WillyHmite (p. 302)... Polynrgyrite (p. 316)...

Lnurite (p. 802)

Argentite (p. 288)

Iron (p. 281)

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HoBsiie (p. 289)

Ooppor (p. 278)

Uruinito (p. 521)

SUvorip. (p. 278)

Silverite (p. 302)

Lead (p. 279)

Palladium (p. 281)

Amalgam (p. 279)

PlaUnum (p. 280)

Oold(p.276)

Iridium (p. 380)

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6-2-a-7 I 5-5-5

6-4-6 U , 5-5-6

6-7-6 86 ■ 6

6-87 I 6-5

6-97 2-5

7-0 7-5

73-7-36 2-2-5

7-8-7-8 4-5

7-4-7-6 1 2 5-8

8-2-8 5 I 2 5

8i\ -8-9 I a-5-8

8 8-8 9 2-.-.-8

9-9-7 5-5

101-111 2-5-3

113-11-8 13 7-11-1 14-19 ' 15-6-19-3 I 32-6-22 8

II. CRYSTALLIZATION TETRAGONAL.

A. Luster Unmetallic.

Torbendte(p. 516)., I Trippkeitefp. 516)..

Oct^edrite (p. 346).

Rutile Ip. 345}

, Xanotlma (p. 494)... ' P.'welli(e (p, 541) ...

Thorite (p. 480)

P«rR.l8<inite <|.. 490) ,

Ziiooa(p. 4281

Itomeile(p. 516)

Sipyliiu (p. 490)

Oi>immnliie<p. 408)..

SoheeUte (p. 540)....

PhoBgeoite (p. 364)..

OalomeKp. 817)

Wulfemto(p. 541)...

OoMiteritv (p. 344)...

Hallockile ip. 322)...

Taplolite(p. 493)....

Siolslie

8-8-a-D5 I

4-l&-4-3.'>;

4-46-4-56 4'S3 4-4-5 4 4-4-5-8

4-68-47

,ab,GoOgIc

AFPBKDIX B. B. liDBTBR HiTAij,ic (akd Submbtaluc).

Ohalcopyrlt* (p. S»7)...

Rnm«(p.M6j

t'eigUMJuitB (p. 490)...., Hauamaimita (p. 342)..., BrauniU (p. 348)

PolUDlU(p. 845)

iteluit«(p. 5«3)

Haucbecornite (p. 3B6)

TapioHle (p. 4Ba)

PlHltuedte (p. 84G)....

ni, CRTSTALIZATION HEXAGONAL.*

Rhom bo belli 111 speclea are dUdDgtilshed by a iMter R.

A. LCBTBR Umicbtalijc.

Jarosite (p. 587) R

Ifaimondlte (p. 5S6)...

Hamliulte(p. 608)R...

Pyiochioite (p. B61) R.

jKreniBjidvitetp. 61B)...

DlopUM (p. 4S4) R...

STBiibertfitc (p. 61fl) K.

Ci-onBtedlile (p. 475) R..

HemnloKte (p. 1507] R,

Counellit! {p. 5B0)

MesltHe (p. 368) H

RhodochTOilts iHC9) R.

8vabile(p. 601)

81dm1te{p. 35«) R

RliBbclopliBulie(p.GOG|R

Wiirtziie (p. 2B5)

Conindam (p. S3.1) R..

WlUemita (p. 452) R...

Si)l'a>rocofaa1t!le (Sfil) R.

MeluDOCBiite (p. 407) R.

Trilomite (p. 407) R...

NordeuBkiOldiie |518) R

Csiyocerile (p. 407)6...

Parisite(p. 864)

Smithiontte (p. 3601 R.. BeudanUte (p. 516) B. .. I Pluinboguramitef(n.ai4} ' Cappeleoile
(p. 407).., ai«enocUU (p. 294).. Hematite (p. 334) R... XaolbocoDite (p. 815)1

Zinoitelp. S3S)

FrouiUte ip. Sll) R...

lodyrite (p, 819)

Fluocerite (p. 323)

PyrMgyrita (p. 811) R.

Penfleldlieip. 323)

BaryeiHtefp. 408)

Ty»'.nile(p. S3I)

Pyromorphlta (p. 499) T>iiadliiita(p. 500)....

Mimetlts(p. 500)

Oinnabar (p. 393) R...

■ Some pseudo-hexlcon*! ipeciei are Included.

830 I 3-0-8-6

885 : 8iie 1

8-38-8-42 8-46-3 60 I

B-52 3 B3-3 88 8-94-4 01

4-9-5 0 4-9-5-3 5-5-2

5-4-5-7

t-6-5 3-8-6 5-5-6-5

6-66-6-86 7-0-7-25 I 8-0-8-2

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APPBKDIX B.

B. LUSTBB HBTAI.UO (abd Scbiotallio).

Tesobemncberfte (p. 864) Thermotiatrlte (p. 886).,

Ctiroallite (p. 822)

BtniTtlB(p. 607)

£pBomlLe(p. 688}

Ma»cagaite (p. 628)

KesqueboDite <p. MO},,..

OiMlarite(p. 688)

Moieaoiite {p. 68S)

Bnlphnr (p. »78} ,

Lltidackerite(p. 610)....

Newberrite (p. SIO)

Niter (p. 617)

eideronatriite (p. 587)....

Pliielllte (p. 882) ,

H«troUt« (p. 461)

Okeiiiie?(p, 438)

Fe!Bftliauji[e(p. 587)... ThomBoalte (p. 402'....

WavelUte (p. 518|

Hnmbergiie (p. 518)

Pirssodlte (p. 806)

8>iirobori(e(p. 531)

PJMilierile (p. 512)

I'egiiiiile (p. 612)

Elpidileip. 407)

HowlUe?(p. 519}

Fraluiit« (p. 442)

ADhydrits(p. 538)

AraKonite (p. 301)

Spodioalte? (p. 508)

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Danbotil* (p. 450)

T\ -n)lite(p, 611) ,

HnrKtiglte (p. 442).... .

Rfdiliiigit.-(p. 506)

Bertmncliie (p. 440)....

Lftiiliiiiiw(p. 866)

loUto(p. 4071

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Hopeile (p. 507)

Pliospostderlte (p. 510)

CRYSTALLIZATION ORTHORHOMBIC. A. Luei'BR Unmbtaixio.

T«lo{p.476}

Berylloolte (p. 496)

Hftidtagerlte (p. 610}

8treDgite{p. 610)

lAwoDlte (p. 447}

Hnlnte (p. 448)

Aathophyllta (p. 896)..

Andalniite (p. ^)

Bnrtatit* (p. 384)

Antaiiite(p. 515)

MoDtlneUte (p. 482}....

CbifdieatM (p. 618)

SUUmudt* (p. 483)

Soorodita (p. 609)

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A.PPBNDIX B.

IL CRYSTALLINE HABIT,

I. ISOMETRIC SYSTEM.

In the following list some species are enumerated whose crystal form is often so distinctive as to be a characteristic.

Oxides.— Metallic Oxides: Quartz; Pyrite,

Unkmetallic: Fluorite; Cuprite; Cerussite; Halite; Sylvite; Borate; Barite; Lithium salts. Also Perchlorate; Cerargyrite; Perovskite.

Cubic form occurs with the following: Apophyllite (tetragonal); Cryolite (monoclinic), Almandine with its rhombohedral species: Chabazite; Alunite; Calcite rarely Quartz; and Ulexite.

Octahedra.— Metallic and Oxides: Haguetite; Frankite; Chromite; Unkmetallic. Also sometimes, Quartz; Pyrite; Lithium; Dysanite.

(In the list of Minerals: Beryl (includes Emerald and Garnet); Cuprite; Diamond; Pyrochlore and Melanite; Rhabdophane; Alunite.

Forms which are characteristic of regular octahedra occur with the following species, as Beryl; Ulexite; Chalcocite; Zircon, etc; also with some rhombohedral

species, as Dolomite.

Dodecahedron.— Metallic Lustre : Magnetite; Amalgam,

Nonmetallic Lustre : (Lignite; Cuprite; Sodalite.

Trapezohedron.— Unmetallic Lustre ; Garnet: Leucite; Analcite.

Pyritohedron.— Metallic Lustre ; Pyrite; Cobaltite. Also Garnet; Hauelite (submetallic)

Tetrahedron.— Metallic Lustre : Tetrahedrite.

Unmetallic Lustre : Spinel; Biotite; Helvite; Eulytite; Dolomite.

The tetragonal pseudocubes of Calcopyrite may resemble leucite.

II. TETRAGONAL SYSTEM.

Square Pyramidal.— Submetallic Lustre : Bismuth; Hannayite.

Unmetallic Lustre : Zircon; Wulfenite; Vesuvianite; Octahedrite; Xenotime.

Square Prismatic.— Unmetallic Lustre : Zircon, Vesuvianite; Scapolite; Apophyllite; Phosphite.

Square tabular crystals occur with Apophyllite; Wulfenite; Torberlite.

Prisms nearly isometric are noted with a number of orthorhombic species, e.g., Topaz; Andalusite; Danburite.

III. HEXAGONAL SERIES.

Hexagonal Prismatic.— Unmetallic Lustre : Beryl; Apatite; Pyromorphite; Vanadinite;

Malachite (usually indistinct rounded form). Also Nephelite; Hilairite; Tyrochroite, and others.

Hexagonal prisms are also common with the rhombohedral. Quartz, Calcite; Tourmaline; Willemite; Phenacite; Diopside, etc. Again, with the Malachite, etc. Numerous rare species could be included here.

Many orthorhombic (or monoclinic) prisms exhibit a prismatic angle of about 60° (and in some cases the form both in natural crystals and artificially as the result of twinning). Thus, Aragonite; Strontianite; Leadhillite; Lignite. It is also to be noted that the rhombohedral dodecahedron, e.g., of Garnet, has often the form of a hexagonal pyramid with

trihedral lernilnntlons fcf Plfr. 443. p. 184).

Tahulnr hexacfonal prism" arp noled with various «peeel!». Thua. Metallic Lotteb : Hematlle; Tmenlfe: Pyrrliotltfi. TInmetallic Ltjbter : Tridymlta.

Hexarnnal Pvramlda.—Apntile: Corundum (rhombohedral); Quarts (rbombphedral-trappKnlipdrall: Hankslle.

This form la often almulated by Tarloiiis orOior kombU aperies. In part as the reauill_ of twionlne. For pxamnle. mktalt.ic lubtbr : Chalcoclte; Btephanile; PolybB«Ite; Jordanile, etc Alan Brooklte (Fie. SOI n M)

■Unicktallic LcffTEN : Wliherlte; Bromllte; Ceruaslte: loHte.

^Mgonal Prian.—Tourmalide.

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AFPENDIS- B. fioft

RhombolMdroiM.—Angle 1&' (nod IOS"); Oaldle; Dolomite; Bldeilte; Jtbodochroslte. Aoglu not far froiD SO": Cliabazlte; A.1uDite,

Sonlenohodroit*.—Calcite aod allied Carbonatea; Prousttte.

IV. ORTHORHOMBIC, MONOCLINIC AND TRICLINIC SYSTEMS.

Prlimatlo Oryatala.— Mbtallic LDBTBit; SUBnile; Arsenopyrite; BournODlte; Maoga-Dilei OOtliite, etc.

Unmi'.Xaluc Luster : (w-l/wrAomM:) Topaz; Siaurolite; Andalusite; BarttQ; Celeitlte; Danburiie. Also (monotiinic) Pvroxeeae; Amphlbole; Orthoclase. and many otbera.

Epidote cry^talb are orten prfsmBlic in uapect (Fig. 8So, p. 488).

Tabulmr Orjrrtali.—Bnrilc; Ceruulle; Calamine; Dlaspore; Wollastonlte,

Acionlar Oryitala.— Hktaluc Lubtbr : Stibnlle; Blamuthlnfte: Hillerite; JameRODite; Aikiuie, auil oitiFt species,

UmietalijcLiibtbu: Fectolite; Nalrollle; Scolecite; Thomsonile, and other Zeolites. Also Arugouile; Slronliuuite; leu often Caiclte. Also many oilier species.

Twin Orystals,—Tlie habit of the twins occurring with many species Is very cliaracler-istlc. Relerence is made to pp. 118 lo 130 and tbe accompanylni; figures for a presentHtion of tills subject.

III. STRUCTURE OF MASSIVE MINERALS.

Fibrous.— Fibrous leucophaea: Asbestos (serpentine); also the similar asbestos-like variety of serpentine (chrysotile): Crocidolite (color blue).

Fibrous leucophaea*, chiefly in the form of : Calcite; Gypsum. Also Anhydrite; Barite; Celestine; Apatite; Brucite; Enstatite; Wollastonite; Dufrenoyite; Vivianite. See also Columnar below.

Fibrous Radial.— Water-soluble; Thortvanderite; Nitrate; Silicate, and other Zeolites; Gypsum; Urea-soluble.

Columnar.— Metallic Minerals : Stibnite; Hematite; Jamesonite; Zinkenite, etc.

Non-metallic Minerals : Limonite; Goethite; Actinolite; Amphibole (tremolite, actinolite, etc.); Epidote; Zoisite; Tourmaline; Sillimanite; Nepheline and other Zeolites; Strontianite; Wilberite; Topaz.

Cyanite has often a radial structure.

Fibrous and columnar varieties pass into one another.

Radial.— Gypsum in Pyrophyllite.

Foliated,— Metallic Minerals : Graphite; Ulybdenite; Tetradymite; Sternbergite; Natronite

Non-metallic Minerals : Tellurite; Orpiment; Gypsum; Pyrophyllite; Serpentine; Gypsum.

Microscopic.— The Minerals, p. 468: also the Brittle Micaceous, p. 470. and the Chlorites, p. 472. Also Bismutite; Orpiment; Talc; Torbernite; Anhydrite.

Oranlar.— Metallic Minerals : Galena; Hematite. Many sulphides, sulpharsenites, etc., have varieties which are fine-granular to compact and impalpable.

Thinly crystalline Minerals : Pyroxene (cocolite); Gmelin; Calcite; Borate, etc.

Botryoidal, mammillary, reniform, etc.— Metallic Minerals : Hematite; Arsenic; Allamonicite. TM _ i j n

Non-metallic Minerals : Malachite; Prehnite; Calamine; Smithsonite; Chalcedony; Hyalite; Sphalerite, etc.

Stalactitic.— Metallic Minerals : Limonite; Psilomelane; Marcasite.

Non-metallic Minerals : Calcite; Aragonite; Gibbsite; Chalcedony.

IV. PHYSICAL CHARACTERS.

I. CLEAVAGE.

OuHo.— Mbtalmo Luent: Galena, , „ „—

UmntTALLic LuaTEH : Halite: SylTiio. The cleavage of Anhydrite (alio of Cryolite) simulates this. Cf. also Corundum, p, 838.

.dByGoOglC

OoUhsdral.—Fluorile; Diamond. Magnetite (aba Fraoklinlte) bu often dlitincl octft-faedral parting..

Dodeoahadrtl.—Spbalerite.

tUioinbahedral.~Calcite and other species of the same group, pp. 854-860.

Square PrlamatlC (90°).—Scapolite; Ruiile; XeDOlime.

Priuaatlo.—Barlte; Celestlte; Amphlbole (64° and 136°), etc.

Batal.— Hktallig Luster: Qrapliite;Molybdenite.

Unmetallic LtrSTBH ; Apoplijllite; Topaz: Talc; the Micu and Cfalorites; Chalcophjl-lte, etc. Pyroxene ofteu shovrs narked basal pariog.

Plnacoid«L— Metallic Lt'BTER : Slibolte.

UsHBTALLic LuBTEit: Gypsum; Orplmeut; Euclase; Dlaspore, etc.

II. HARDNESS.

1. Soft Mlnorala.—Tbfl followiDg minerals or less; They hence have » grtiity 'eel. (See furtlier ihe Tal

Metaxuc Luster: Qiapblte; Molybdenite; Tetradyi Kat^agite; some of the Naliva Metals (Lead. etc.).

Ummetallic Luster: Talc; Pyrophyllili CtuDsbaj Sulphur; Gypsum.

Also Calomel, ArseaoUte, and many hydn

2. Hard Mlnarala.—Mlaenils whose hardui Tbe followlog minerals are here included:

LnBTRa Unmstallio.

Sternbergite ; Argentite; Brucite; Tyrolite; Orpiment; Cerargyrite; Biilpbates, phosphates, etc. ia equal to or greater than 7 (Quartz s Ty.

Quartz (p. 8M) 7

Triilymto(p. S38) 7

BrnryliteCp. 408) 7

Dumortierite (p. 44S) 7

Daiiburito (p. 480) 7-7-25

Boracite (p. 619) 7

Ziiliyili: (p. 416) 7

Ctanite (p: 494) 5-7 as

TouRMAUNB (p. 447) 7-75

Garnet (p. 415) 6-6-7-6

loLiTE (p. 407) 7-7 6

Staurolite (p. 450) 7-7'5

8c)iorlomlte(p. 419) 7-7'5

Sapphirine (p. 451) ,^7-6

Euclase (p. 466) 7'5

Hambergite (p, 518) 7-8

ZiRCOK (p. 428) 7-6

Ahdalusiib (p. 433) 7*5

Bertl(p 405) 7.5-8

Lawsonite (p. 447) 75-8

Pheiiacite (p. 428). 7-6-8

Gahnite (p. 880) 7-5-8

HercyDlie (p. 889) 7-6-8

Spinei, (p. 888) 8

Topaz (p. 431) 8

Hhmliziie (p. 518) 8

Chbvsobbtl (p. 842) 8'6

CoRUHDUu (p. 838) B

DiAHOKD (p. 271} 10

The following minerals have hardness equal to 6 to 7. or fl-6—7.

Luster Metallic : Iridosmloe (p. 280); Iridium (p. 280): Bperryllte (p. 802).

LtlBTBR UNKBTALLtc: Ardennlle (p. 446); Axlulte (p. 441); Betrandlte (p. 446):

III. SPECIFIC GRAVITY.

AttflDttontscalted to the remarks In Art. 280 (p. 158), on the relation of specific gravity to chemical composition. Also to the Bnttemenis ic Art. S81 as to the average specific gravity among minerals of metallic and unmetallic luster respectively. The species in each of the separate parts of Table I. of minerals classified with reference to crystallization are arranged according to ascending specific gravity. Hence the lists give at a glance minerals distinguished by both low and high density.

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APPENDIX B,

IV. LUaTER. (See Art. US, p. 168.)

MetalUo.—NaUte meUlH; most Sulphides; some Oxides, those containing Iron, manganese, lead, etc

SDbmetallo.'-EeTe belongs chiefly certain iron compounds, as Ilmenite; Ilvaite; Cuscutin; Tsilalite (and allied species; Wolfenite, etc. Also Unmlnlte, etc.

Adamantine. —(a) Some of the most dense: Diamond; Corundum; Cassiterite; Zircon; Rutile. (&) Heavy compounds of lead, also of silver, copper, mercury. Thallium, Cerussite. Auelesite, Plumbite, etc.; Cerargyrite: Cuprite; some Cinnabar, etc (c) Also certain varieties of Bismuthite, Titanite and Oxidite.

Metallic—Adarnaitina —Pyrrargyrite; some varieties of the following: Ouprita, Cenurite, Oclabedrite, Butile.

Resinous or Waxlike.—Bismutite; Elreolite; many Phosphates.

Tibsona.—Quartz and mica, as Garnet, Beryl.

PMf.—The foliated species: Talc, Brucite, Pyrophyllite. Also (on cleavage surfaces) conspicuously the following: Apophyllite, gillite, Heulandite. Also, less prominent Barite; Celestine; some Feldspar, and others.

BIU7.—Borne fibrous minerals, as Opal, Calcite; also Apatite.

V. COLOR

The following lists may be of some use in the way of suggestion. It is to be noted, however, that especially in the case of metallic minerals a slight surface change may alter the effect of color. Further, among minerals of unmetallic luster particularly, no sharp line can be drawn between colors slightly different, and many transitions of shade occur in the case of a single species. For these reasons no list, unless continually extended, could make any claim to completeness.

(a) Metallic Minerals.

Silver-white, Tin-white.—Native silver; Native Antimony, Arsenic and Tellurium; Amal: Arsenopyrite and Loupingtonite: several sulphides, arsenides, etc., of cobalt or nickel, as Cobaltite (reddish); some Tellurides. No sharp line can be drawn between these and the following group.

Steel-gray.—Platinum: Manganite; Chalcocite; Sphalerite; Bornite.

Bluish-gray.—Molybdenite.

Lead-gray.—Many sulphides, as Galena (bluish); Sphalerite; many Sulphates, etc., as Malachite, Dufrenoyite, etc.

Iron-black.—Graphite: Tetrahedrite; Polybosite; Stephanite; Enargite; Pyrrhotite: Magnetite; Ilmenite; Frankelite,

Black (with submetallic luster).—Ilmenite; Limonite; Columbite; Vanadinite, etc.; Wolfenite; Uvaite; Uraninite, etc. The following are usually brownish black: Bismutite; Hausmannite.

Copper-red.—Native copper.

Bronze-red.—Bismutite (quickly tarnished); Niccolite.

Bronze-yellow.—Pyrrhollite ; Pentlandite ; Brelthauptite.

Bronze-yellow.—Chalcopyrite ; Hillierite (bronze.) Pale bronze-yellow : ^rrltai Ha>-casite (whiter than Pyrite).

Gold-yellow.—Native gold.

Streak.—The following minerals of metallic luster are notable for the color of their streak :

White : Pyrrhollite. Red : Malargyrite. Dark Red : Hematite (also Cuprite).
Black : Cinnabar (usually uniaxial).

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Dark Brown : Manganoite ; Franklandite ; Chlorite.

Yellow : Limonite.

White.—The following are conspicuous for their bright or variegated tarnish :
Chalcopyrite ; Tetrahedrite ; some Limonite.

{b} Umibitaluc LnBTEB.

Colorless.—In Crystals : Quartz ; Calcite ; Anorthite ; Orthoclase ; Albite ;
Albite : Barite : Adularia ; Topaz : Apophyllite ; Natrolite and other Zeolites ;
Diopside ; Nepheline ; Meibomite ; Calamine ; Cyanite ; Phenacite, etc.

White : Quartz ; Calcite ; Gypsum ; Hyalite (botryoidal).

White,—Cuvierite : Amphibole (tremolite) : Pyrope (diopside, usually greenish).

White : Calcite ; Milky Quartz ; Feldspars, especially Albite ; Barite ; Ceolite ;
Scapolite ; Talc ; Meibomite ; Magnesite ; Kaolinite ; Amblygonite, etc.

Blue.—Bastnaesite : Azurite ; Crocidolite.

Indigo-white : Indicolite (Tourmaline) : Uralite.

Azurite-Bastnaesite : Lazurite ; Azurite ; Lapis Lazuli ; Turquoise.

White-Bastnaesite ; Sapphire ; Cyanite ; Lignite ; Asurite ; Chalcocite and others ; copper compounds.

Skt-blub. HouMTAiM-fiLoE : Bervl; Celeatlte.

YIOLET-BLDB : Amethyst; Fluorite,

Grebnibh Blue : AmazoD-stooe ; CbryBocolla ; Calamine ; 8mi[hgonlt« ; some Tui> quois ; Beryl.

Oreen.— Blackibb Oreen : Epidote ; SerpeDtue ; Pyroxeeo : Ampblbole.

Ekbuuld-uherh : Beryl (Ememld) ; Malachite; Dloptase ; Atacamite j and maay Other copper compounds ; BpodumeDe (liiddenite); Pyroxene (rare) ; Oahulle ; Jadeit« and Jitde.

BLriBH Orbbb : Beryl ; Apatite ; Fluorite ; Amnxon-BlODE ; Frebnfte ; Calamine; Smilbaouille ; Clirygocolla ; Chlorile : tome Turqnois.

Mountain oreen ; Beryl (aquamarlue) ; Euclase.

ApPLE-SBRBn 1 Talc ; Garnet ; CbryBoprnae ; Wllemite ; Garnierlte; Pyrophyllle j ■ome MuscoTJI« ; Jadeite and Jade ; Pyropliylite.

PlBTACBio-QBBEN: Epidote.

OaABS-OKEEH : Pyromorphite ; Wnveilile ; Varlacte ; Chrysoberrl.

Gratibh Greer : Amphibole and Pyroxene, many commou kinda ; Jaaper; Jade.

Yellow-oreen lo Olite-Orebn : Beryl : Apatite ; Cbrysoberyl; Chrysolite (oliT»> green); Chlorite ; Serpentine ; Titiinite ; Datolite; OHvenite ; VesuTianile.

Tollow.—StTLPHUR-TEELLOW r fiilpbur : some VesuTianLte.

ORANaB-YELLOW : Orpimeit; Wulfenile : Mlmelite.

Straw tbi.low, also WiNB-TELow, Waitellow : Topaz; Bulphur; Fluorite; Can-crlnite ; Wulfenits ; Vanadinite ; Willemlte ; Calcile ; Bsnie; Chrysolite ; Choadroldte; Titaiiite; Datolite. etc.

Bruwhish Yrliaw : Much Sphalerite : giderlte ; GOIhfte.

OCHBR-TEi,ix>w: G&tbite ; Yellow ocer (limonlte).

Red.— Rubt-red: Ruby (coruudum); Ruby spinel; much Garnet; Proustlte; Vanadinite ; Sphalerite : Chondrodite.

CocniNEAL-RBD: Ciiprile ; Cinuabar.

Hyacinth-red. —Zircon.

Ohangbred: Zincite.

Okahgb-rbd to aurora-hed : Realgar ; Wulfenite.

CitiMMiN-RRD : Tourmaline (rubellite) ; Spinel, Fluorite.

Scaribt-red : Cinnabar.

BiiiCK-RBD: Some Hematite (red ocher).

Rose red to Pink ; Itnse quartz ; Rliodonite ; Rhodochroalte i Erytbrlle; some Scapolite Apophylllle and Zoiaite ; Eudialyte ; Pelallp : Margarite.

PEACH-BLoasoM Rbd to LiLAC I Lepidollte ; Rubellite.

Flbbh-rbd: Some Onhoclase ; Willemite (the variety troostlte); some Chabaztle ; Stit bite and Heulandlte ; Apatite ; rarely Cakice ; Polylialiie.

Bbownisb Red : Jasper ; Limonile ; Garnet ; Sphalerite ; Slderite ; Rutlle.

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Brown.— Reddibh BROvn : Some Oarnet; Mme SphaleiiU ; Stanrollte ; CaBsiterite ; ButllB.

Clote-bbowr : Aiiioite : Zlraxm : Pyromorphite.

Tellowibu Brown : Slderite and routed carbonates ; Sphalerite ; Ja«per; Llmonlte; GCtliite ; Toiirmaliue; VemiTiaDite; Chondnxlite ; Staurolite.

Blackish Brown : TiUnite ; tome Sldtrite ; Bpbalerlte.

Bhokv Bbowh : Qiiariz.

Black; TouTmaliue ; black Garnet (raelauite); wme Hlca feapeclally biotite); alio lome Amphibule, Pyroxene and Epidote (tbese are moatly greenisli or bruwiilsb black); fiiTtber. some Bphalerile auiJ some kln dg of Quart/ (varylne from smoky browu lo black) ; also Allinite; Bamarsklte. Some black mluemls with aubmetallc luBl«r are meutioed on p. 483.

Streak.— T\ie»treakiBtohK noted in Ibe case of nome mjnerala with un metal lie luster. By far llie mujoiity bave, even when deeply colored In the mass (e.g. Tourmallue}, a streak differing but little from wblle. The following may be mentioned:

ORANOB-rKLLow: Zincite, Crocoite.

CoCBiNEAL-RBD : Pyrargyrite and pTOUBtite.

ScARLBT RED : Cionabir.

Brownish kbd : Cuprite ; Hematite.

Brown : Limonite.

The streak of the various copper, green and blue minerals, as Malachite, Azurite, etc. is about the same as the color of the mineral itself, though often a little paler.

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